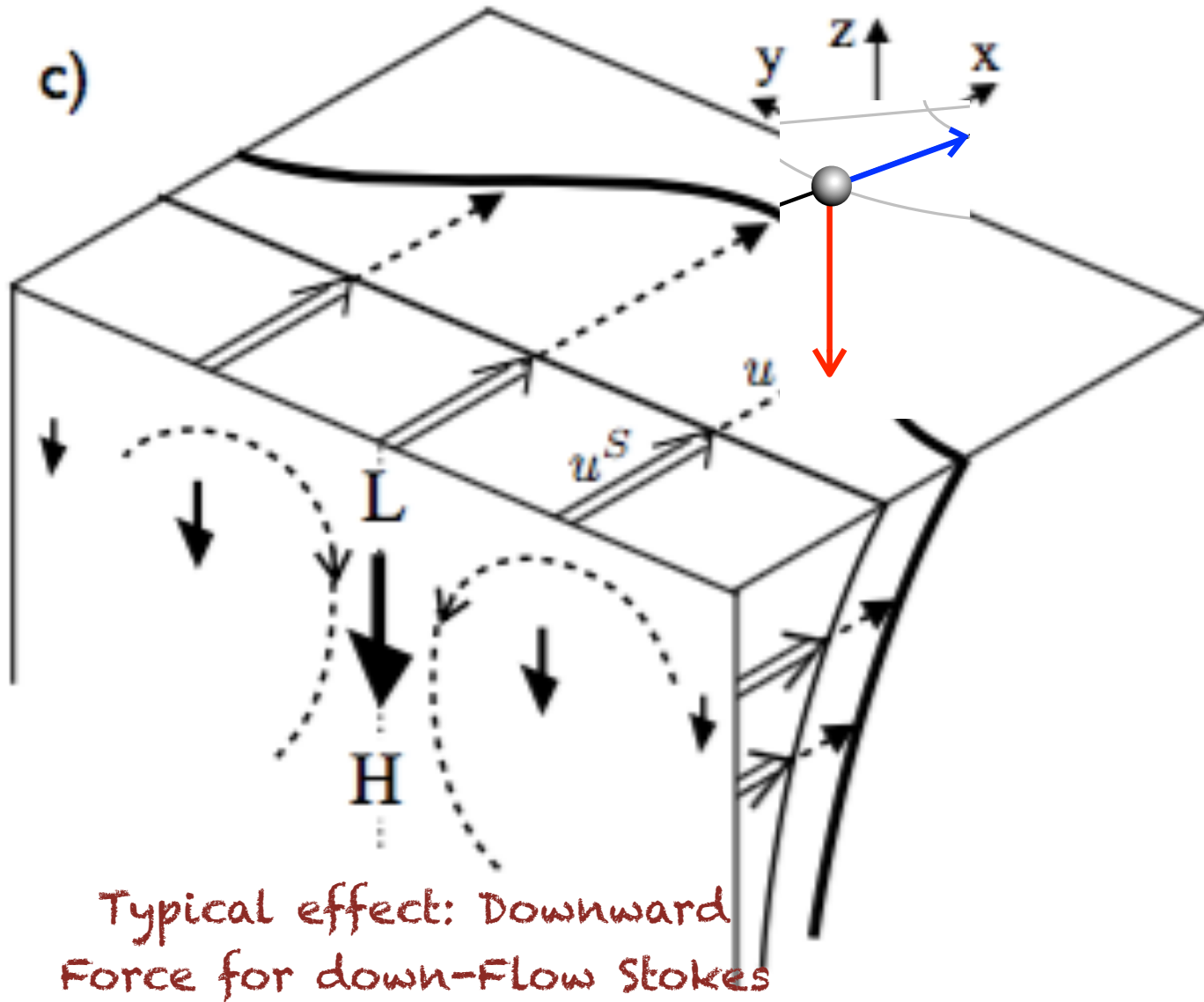


im
Thor

c)



Typical effect: Downward
Force for down-Flow Stokes

$$M_{Ro} \equiv \max(1, Ro)$$

$$Re = \frac{UL}{\nu}$$

$$Ro = \frac{U}{fL}$$

$$Ri = \frac{N^2}{(U,z)^2}$$

$$\alpha = H/L$$

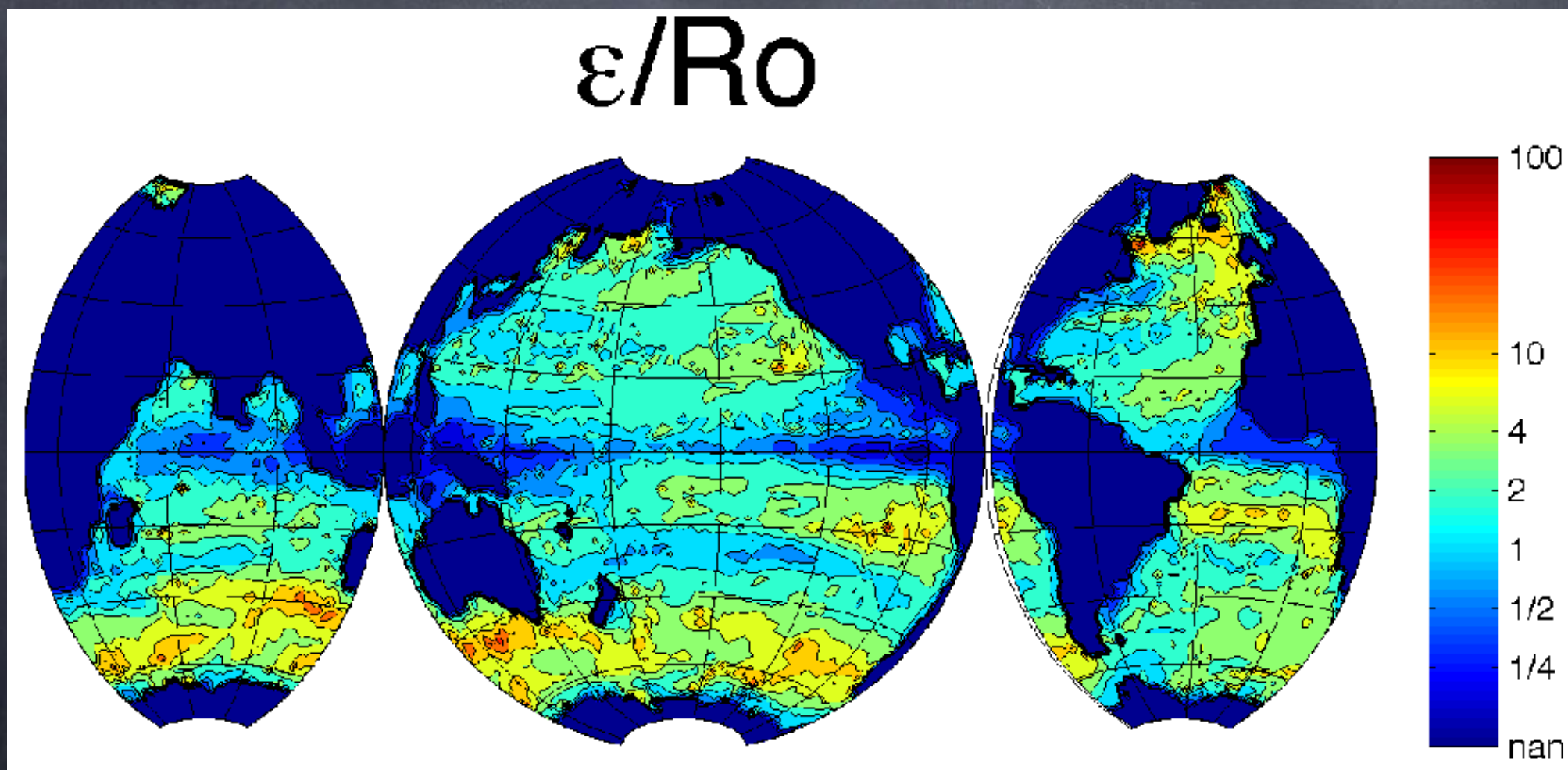
$$\epsilon = \frac{V^s H}{fLH_s}$$

"wavy hydrostatic" if $\epsilon \gg 1$

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b - \epsilon v_j^L v_{j,z}^s} + \frac{\alpha^2}{Re Ri} w_{,jj}$$

N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. Journal of Geophysical Research-Oceans, 121:1-18, 2016.

Do Stokes force directly affect larger scales?



“wavy hydrostatic”

$$\epsilon \gg 1$$

$$\epsilon = \frac{V^s H}{f L H_s}$$

$$Ro = \frac{U}{f L}$$

J. C. McWilliams and BFK. Oceanic wave-balanced surface fronts and filaments. *Journal of Fluid Mechanics*, 730:464-490, 2013.

Do Stokes forces affect (sub)Meso-Scales?

Movie: P. Hamlington

LES of Langmuir turbulence with a submesoscale temperature front

Use NCAR LES model to solve Wave-Averaged Eqns.

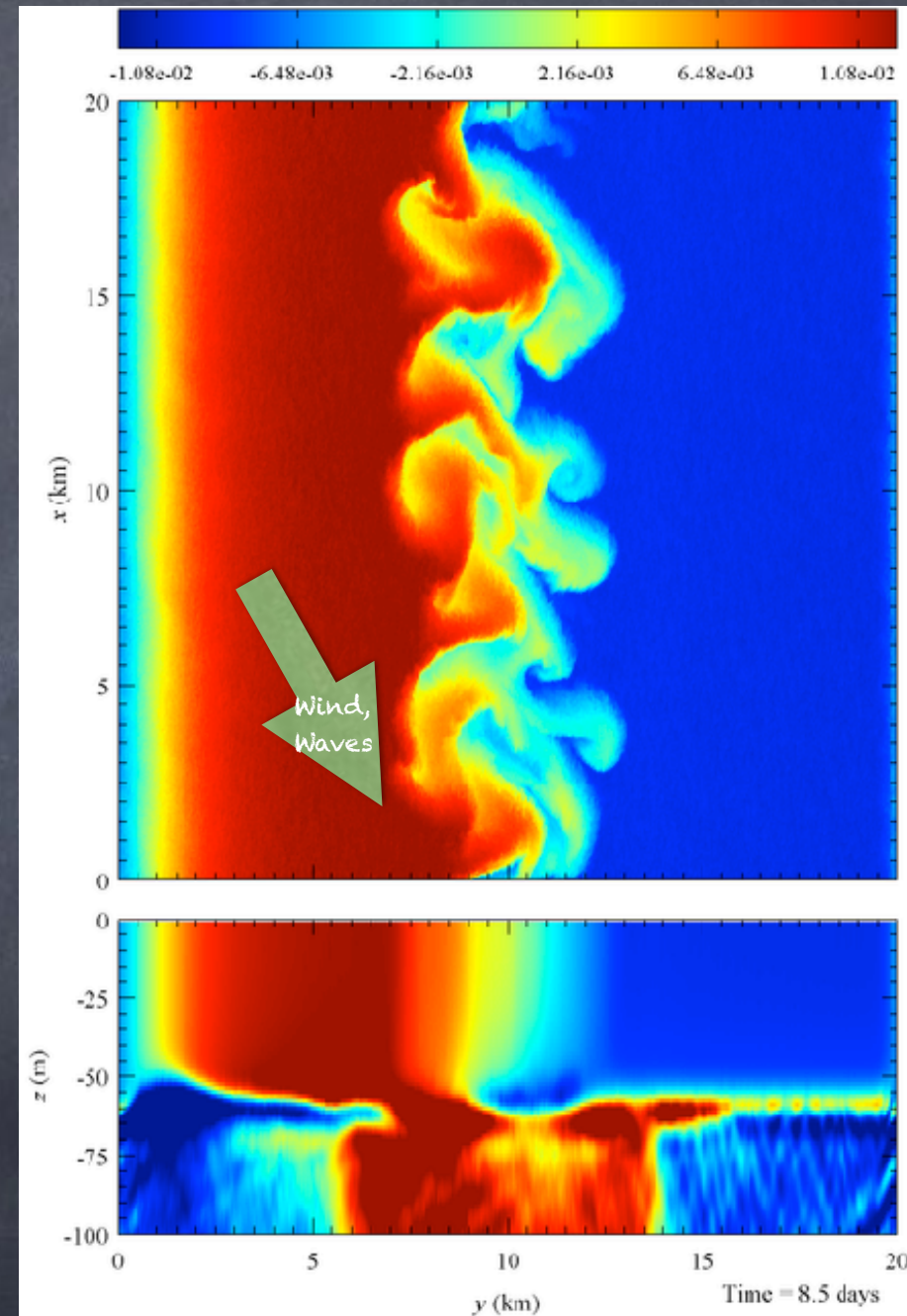
2 Versions: 1 With Waves & Winds
1 With only Winds

Computational parameters:

Domain size: 20km x 20km x -160m

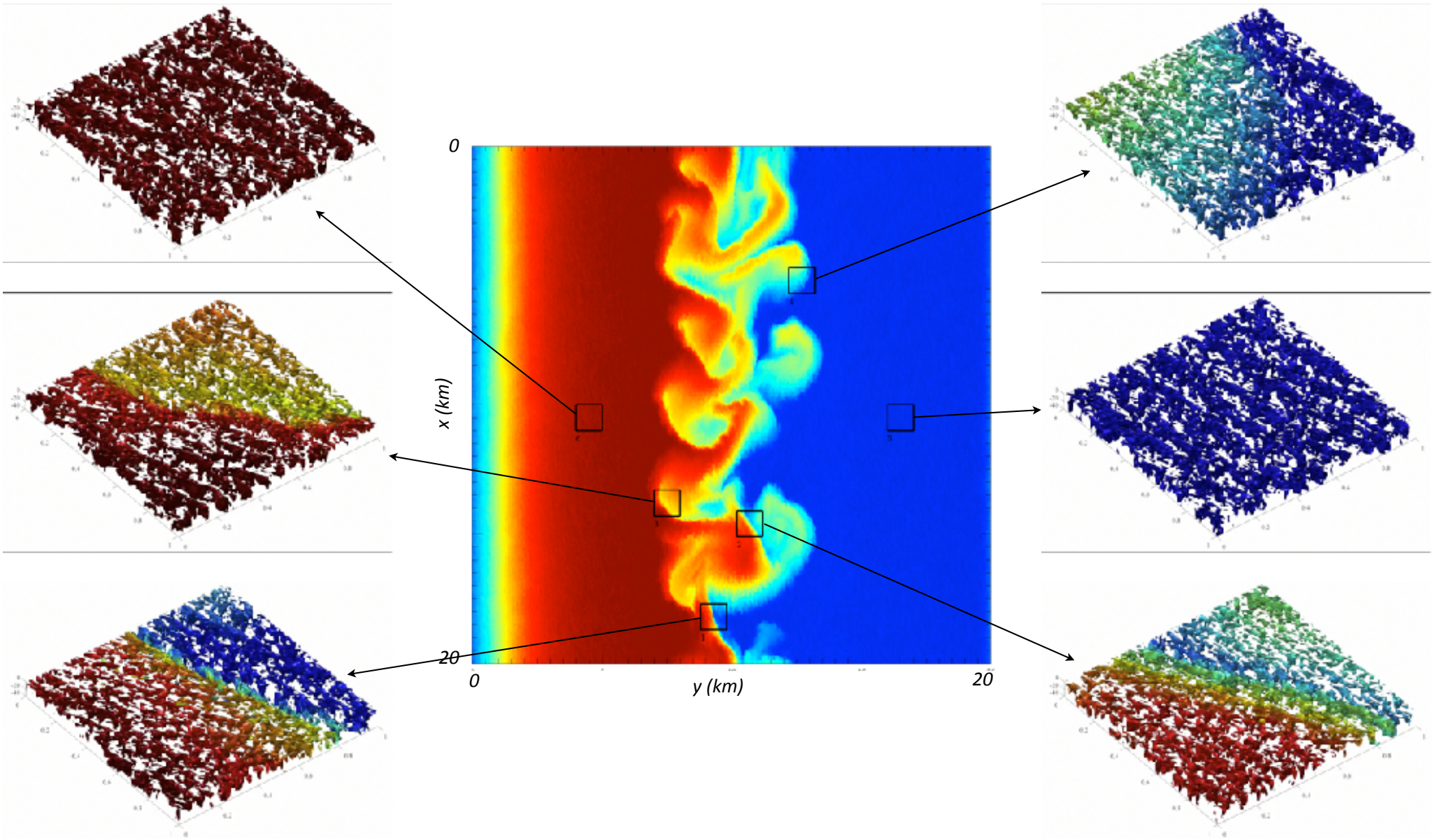
Grid points: 4096 x 4096 x 128

Resolution: 5m x 5m x -1.25m



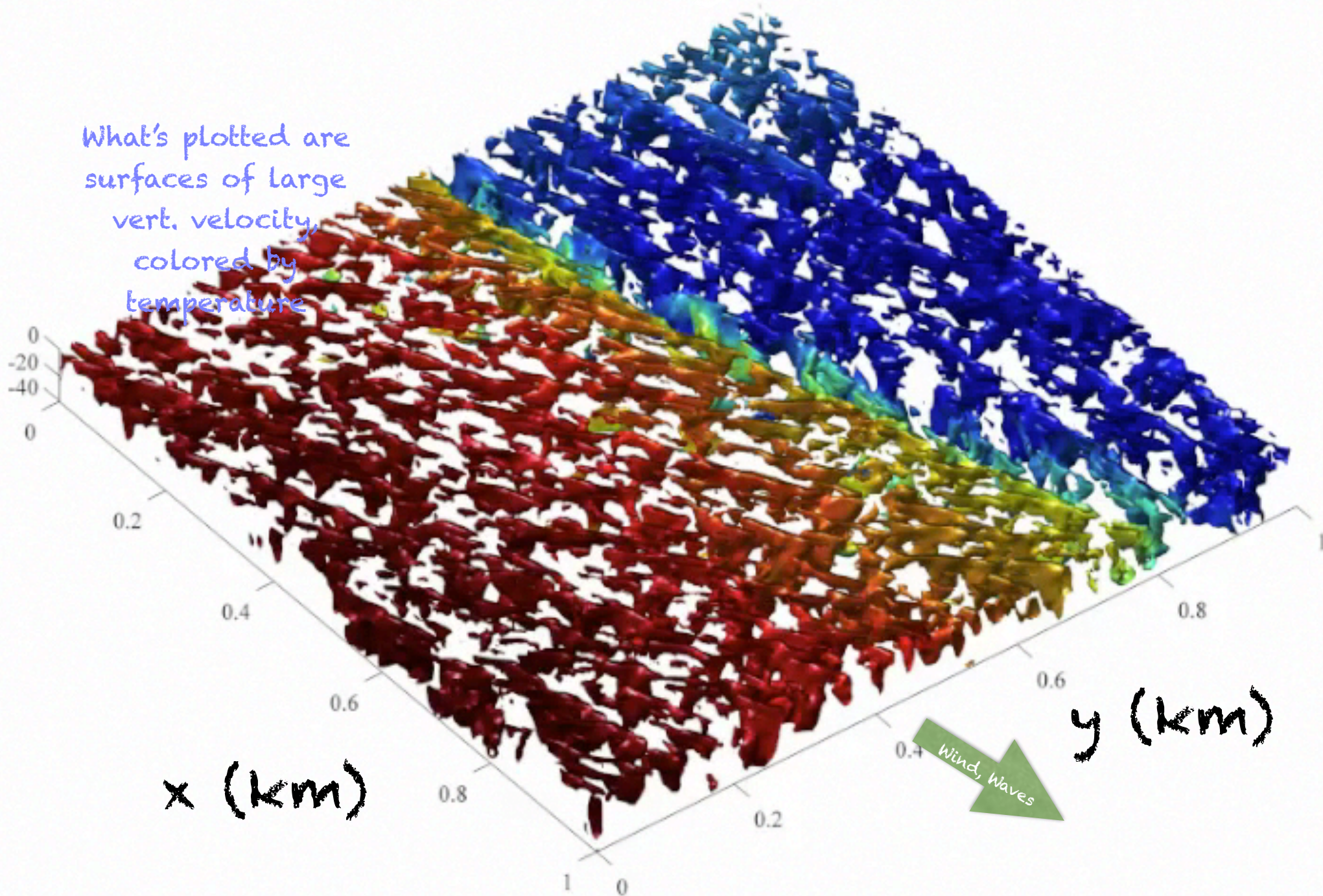
P. E. Hamlington, L. P. Van Roekel, BFK, K. Julien, and G. P. Chini. Langmuir-submesoscale interactions: Descriptive analysis of multiscale frontal spin-down simulations. *Journal of Physical Oceanography*, 44(9): 2249-2272, September 2014.

Diverse types of interaction: Stronger Langmuir (small) Turbulence



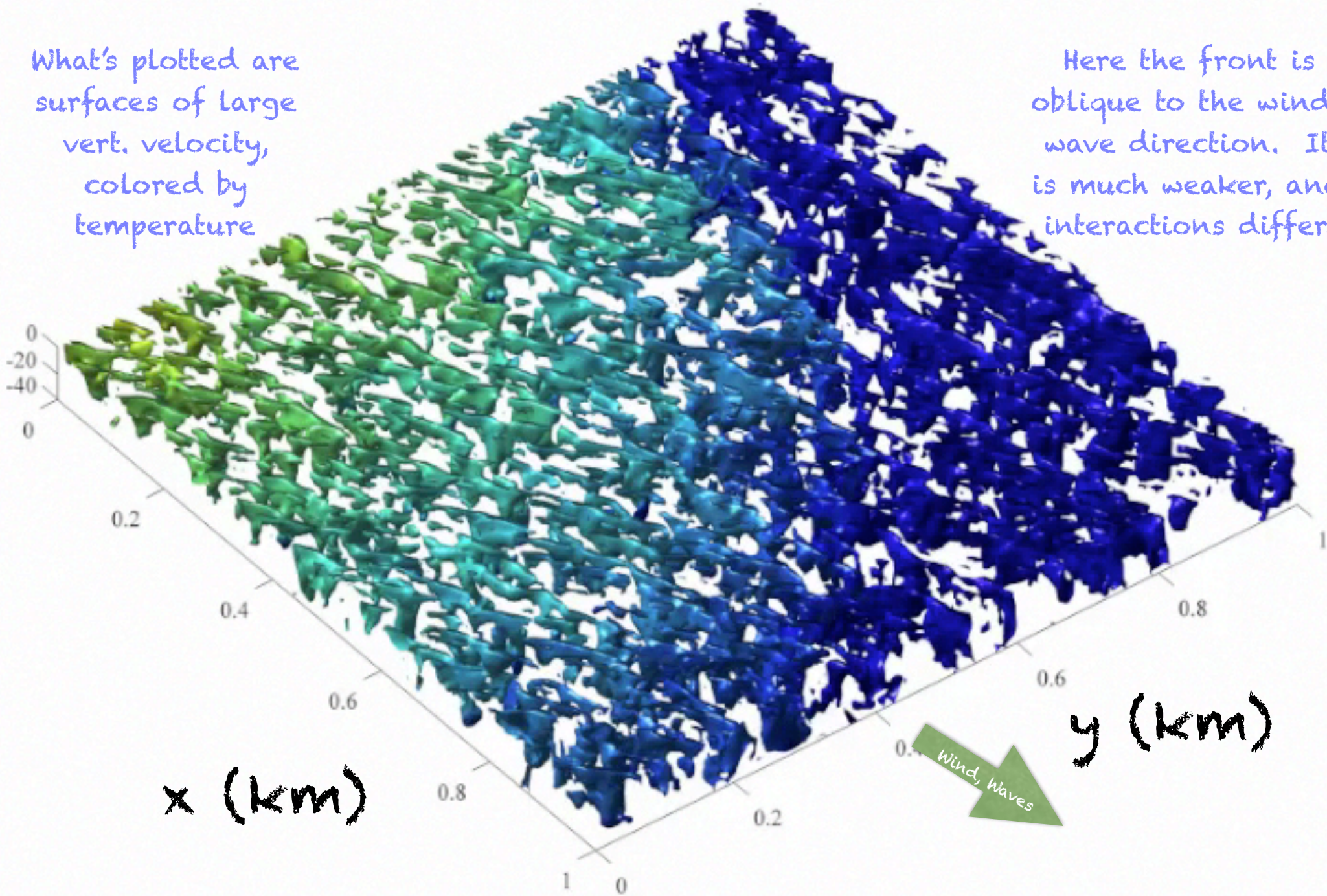
P. E. Hamlington, L. P. Van Roekel, BFK, K. Julien, and G. P. Chini. Langmuir-submesoscale interactions: Descriptive analysis of multiscale frontal spin-down simulations. *Journal of Physical Oceanography*, 44(9): 2249-2272, September 2014.

Zoom: Submeso-Langmuir Interaction!



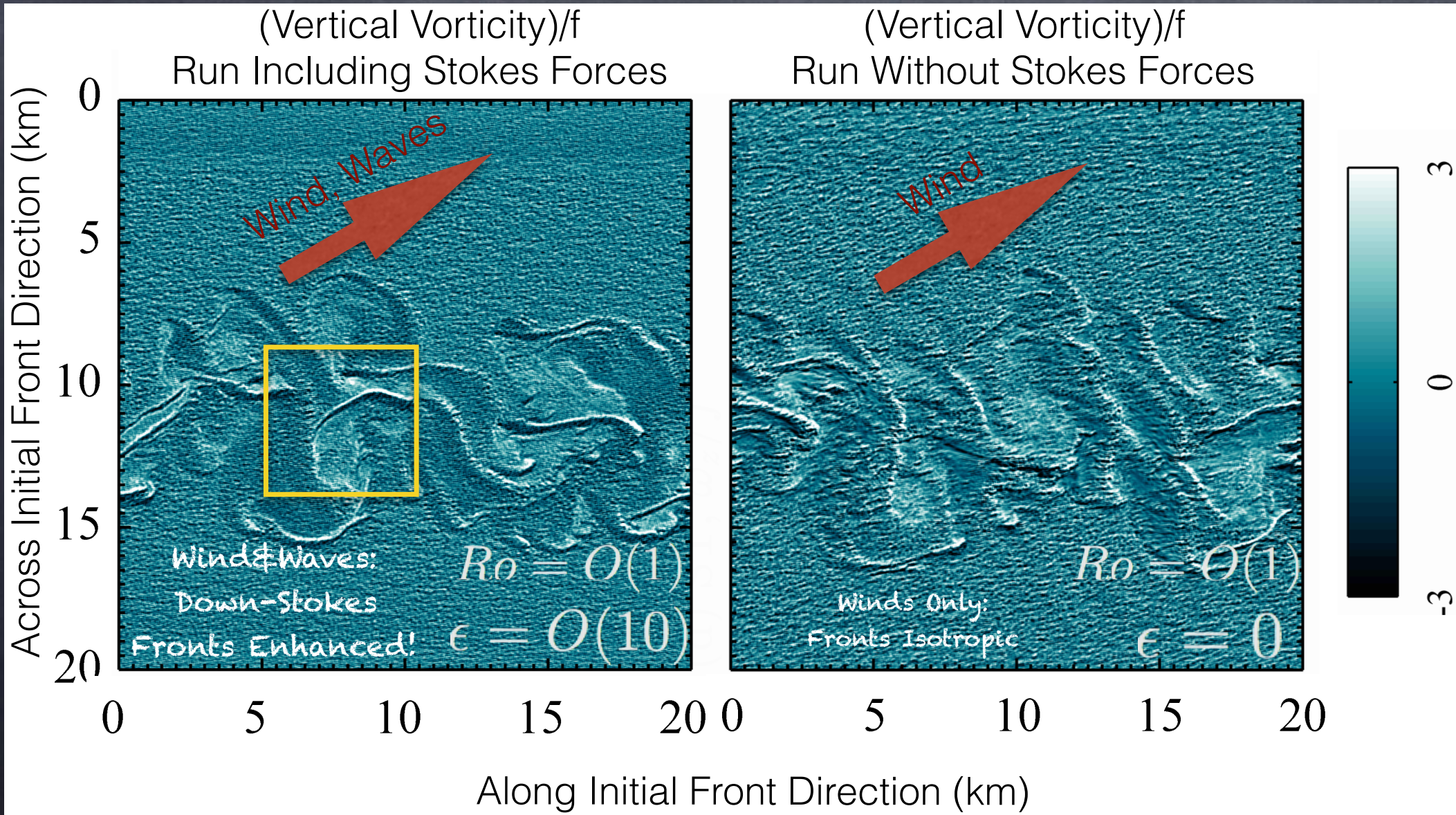
What's plotted are surfaces of large vert. velocity, colored by temperature

Here the front is oblique to the wind/wave direction. It is much weaker, and interactions differ.



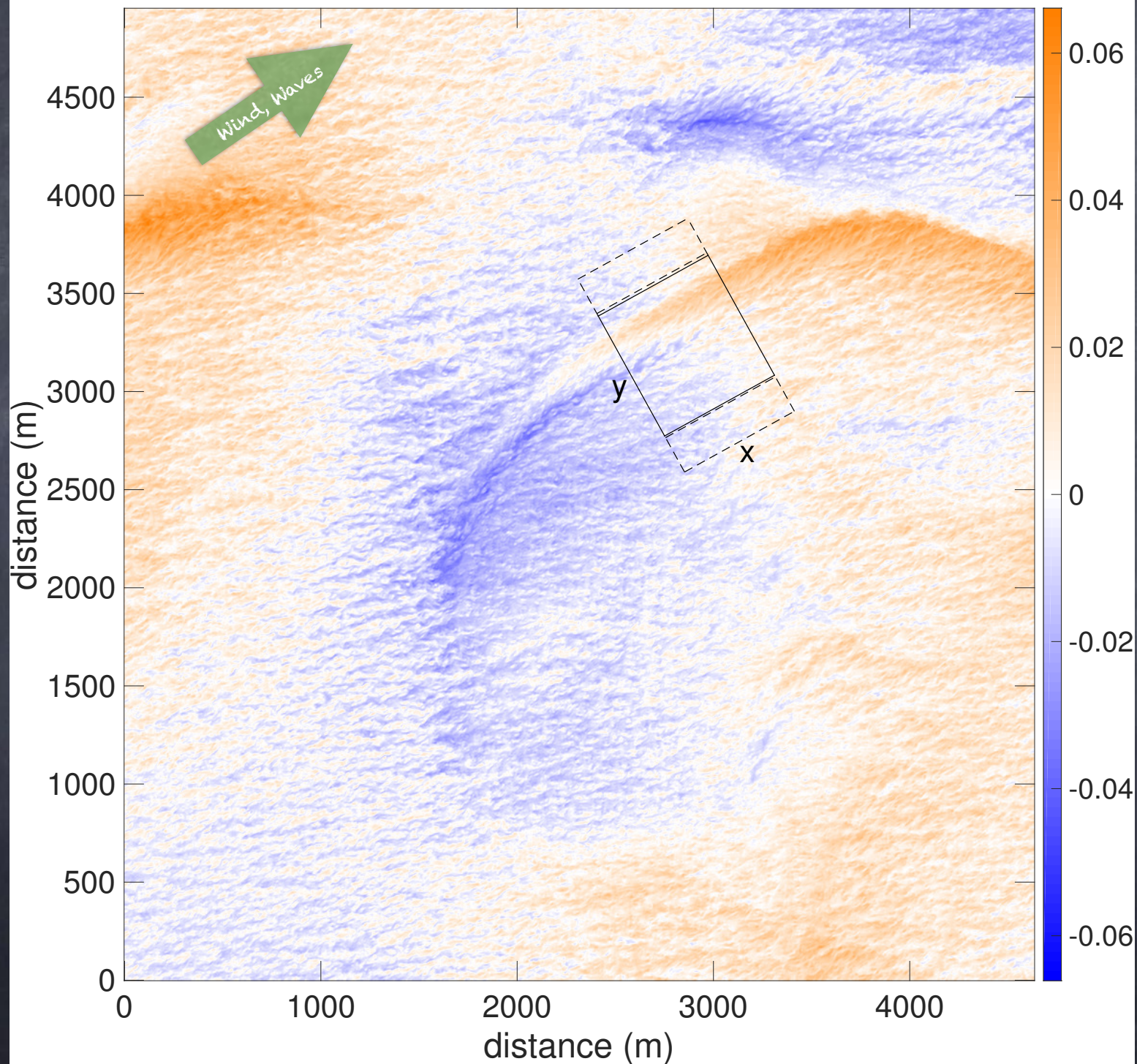
Are Fronts and Filaments different with Stokes shear force?

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = -\pi_{,z} + b - \varepsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{Re Ri} w_{,jj}$$

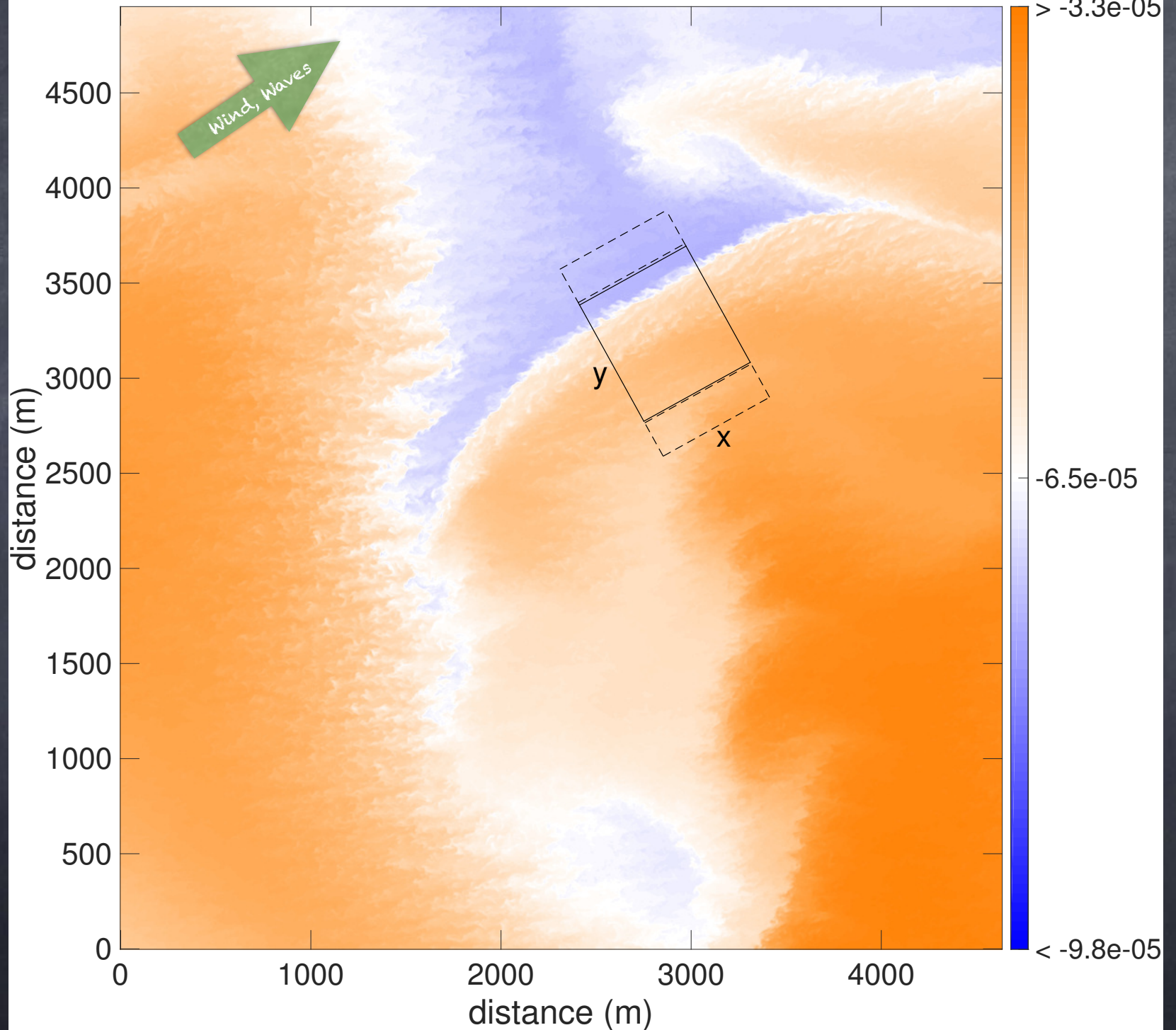


- N. Suzuki, BFK, P. E. Hamlington, and L. P. Van Roekel. Surface waves affect frontogenesis. *Journal of Geophysical Research-Oceans*, 121:1-28, 2016.
- N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. *Journal of Geophysical Research-Oceans*, 121:1-18, 2016.
- J. C. McWilliams and BFK. Oceanic wave-balanced surface fronts and filaments. *Journal of Fluid Mechanics*, 730:464-490, 2013.

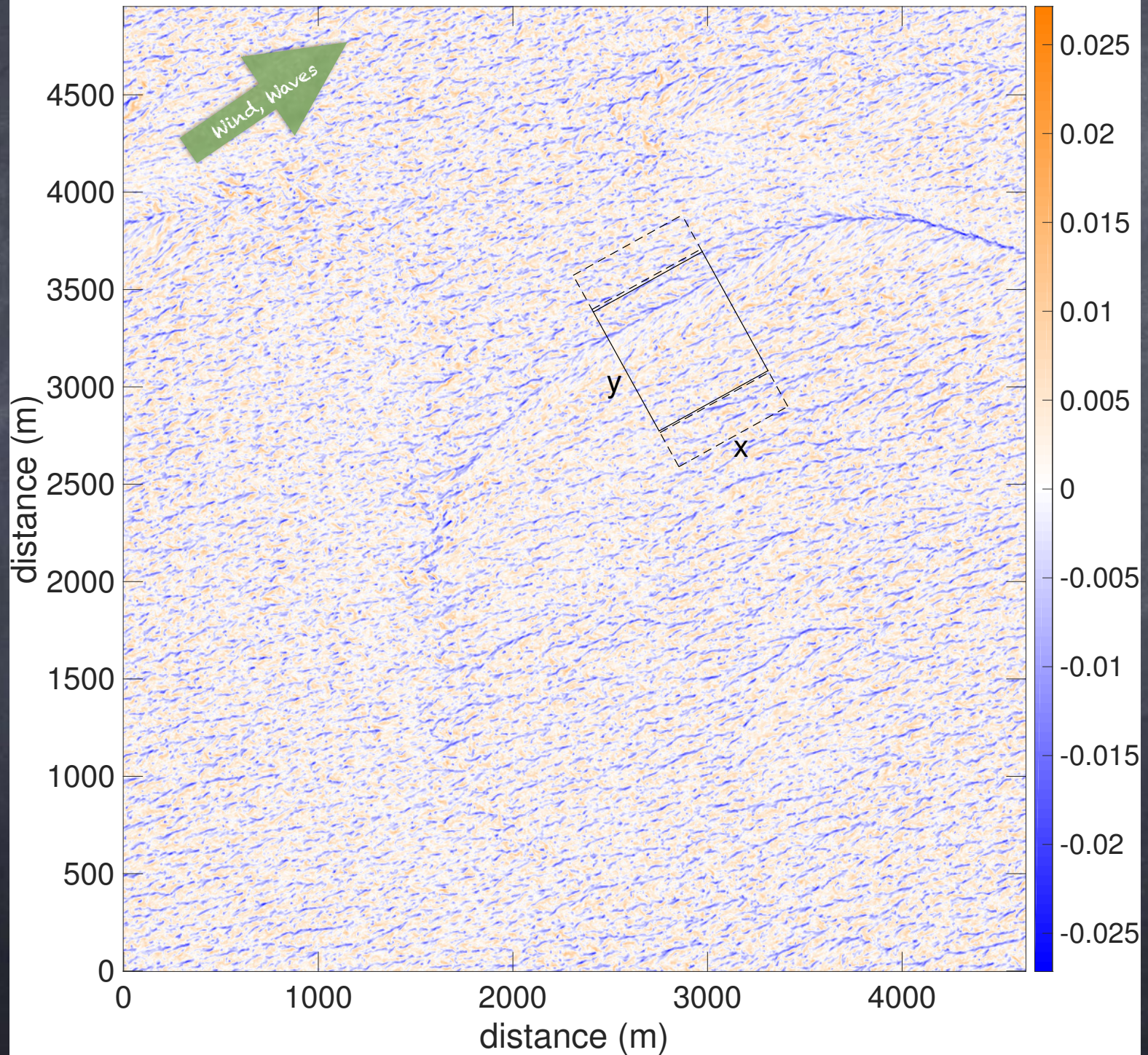
velocity in the x-direction - the horizontal mean (ms^{-1}) at $z = -11.25\text{m}$



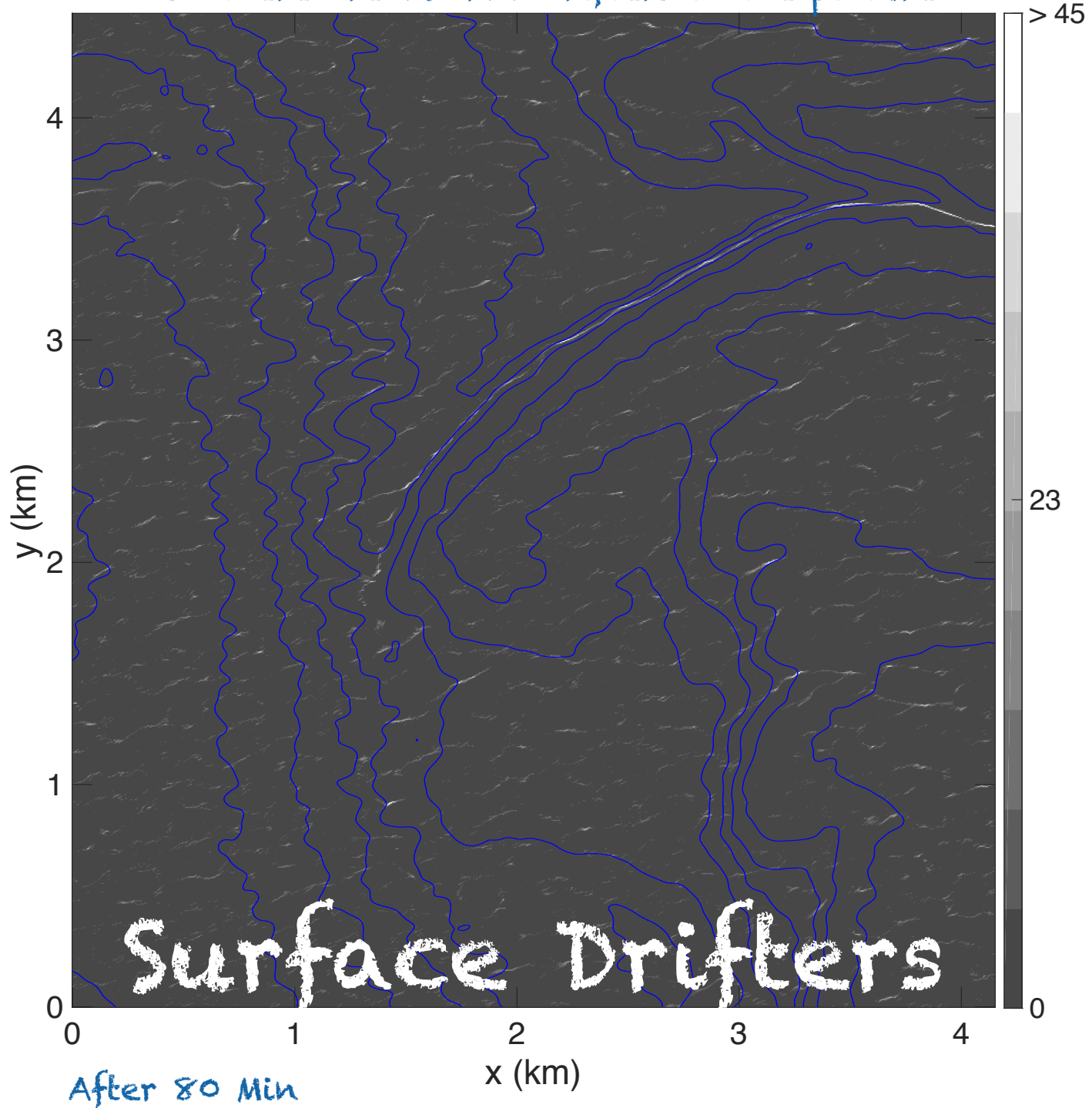
buoyancy - the horizontal mean (ms^{-2}) at $z = -11.25\text{m}$

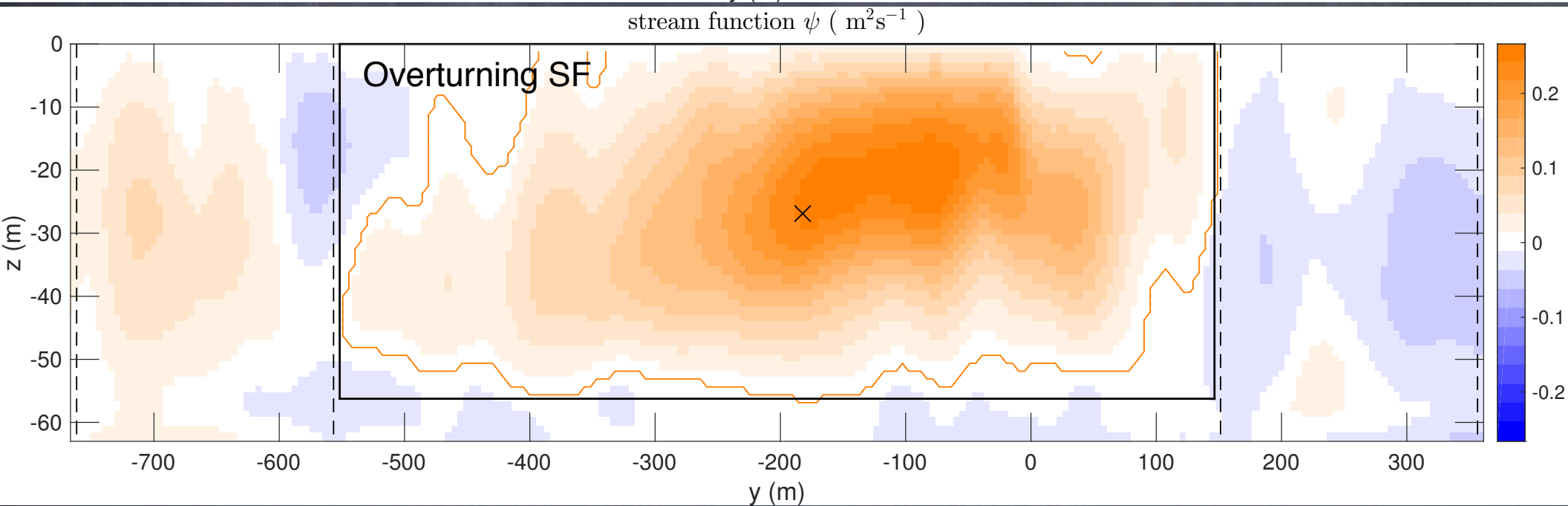
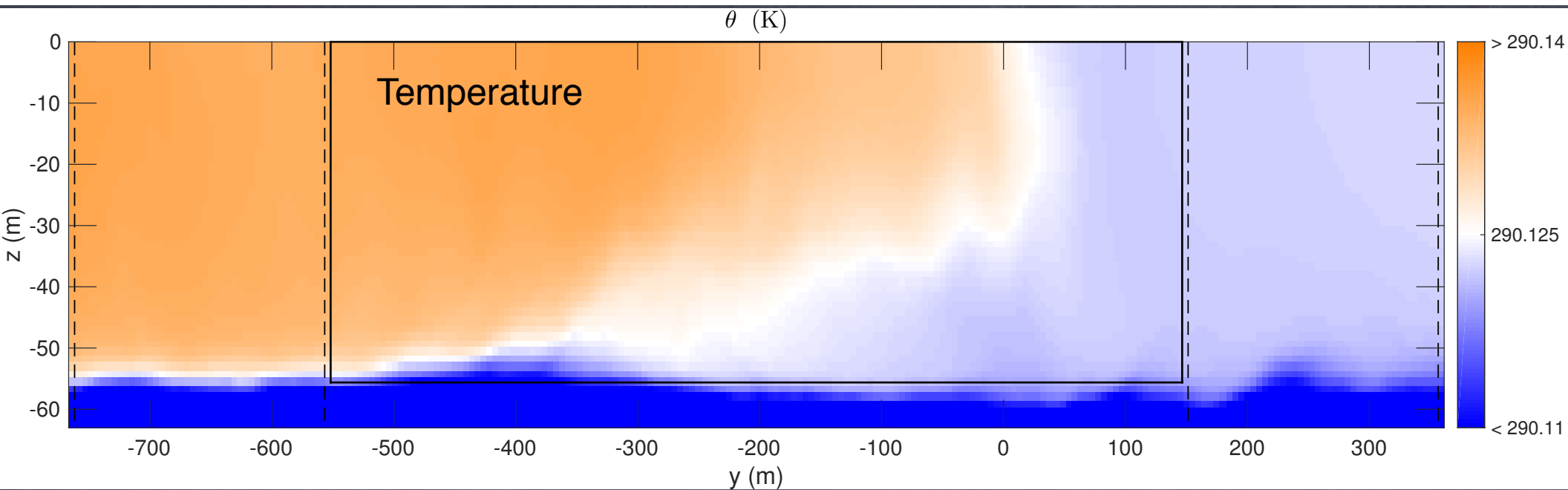


vertical velocity (ms^{-1}) at $z = -11.25\text{m}$



Initially every surface node has 1 drifter, so there are 851796 drifters in the picture





N. Suzuki, BFK, P. E. Hamlington, and L. P. Van Roekel. Surface waves affect frontogenesis. *Journal of Geophysical Research-Oceans*, 121:1-28, 2016.

Do (wavy hydrostatic) Stokes Forces Matter?

Yes! At Leading Order (in LES)

Table 3. Integrated Budget for Overturning Vorticity^a

Responsible Force	Relative Value
<i>Relative Tendency of Overturning Circulation along the Cell Boundary</i>	
Net tendency	11 ± 8%
Sources	
Buoyancy anomaly	100%
Stokes shear force anomaly	44 ± 4%
Interaction with v^H	44 ± 8%
Frontal anomaly in pressure gradient	6 ± 9%
Nonlinear interaction with v^B :	2 ± 1%
Sinks	
Frontal turbulence anomaly (mostly, imbalance in wavy Ekman relation)	-82 ± 11%
Coriolis on along-front jet	-66 ± 2%
Lagrangian advection of (v^ψ, w^ψ)	-36 ± 7%



N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. *Journal of Geophysical Research-Oceans*, 121:1-18, April 2016.

N. Suzuki, BFK, P. E. Hamlington, and L. P. Van Roekel. Surface waves affect frontogenesis. *Journal of Geophysical Research-Oceans*, 121:1-28, May 2016.

~~$Ri < 1 \Rightarrow SI$~~

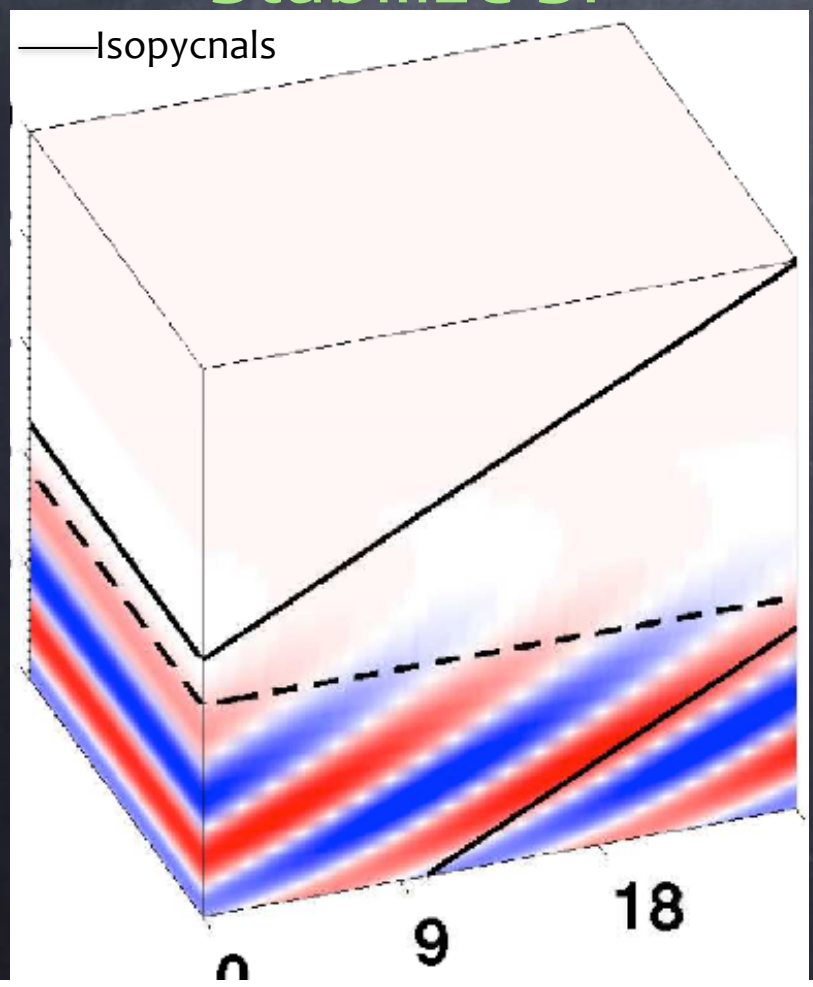
Wavy Submesoscale
Instability Different:
Symmetric Instability

★ $fQ < 0 \Rightarrow SI$

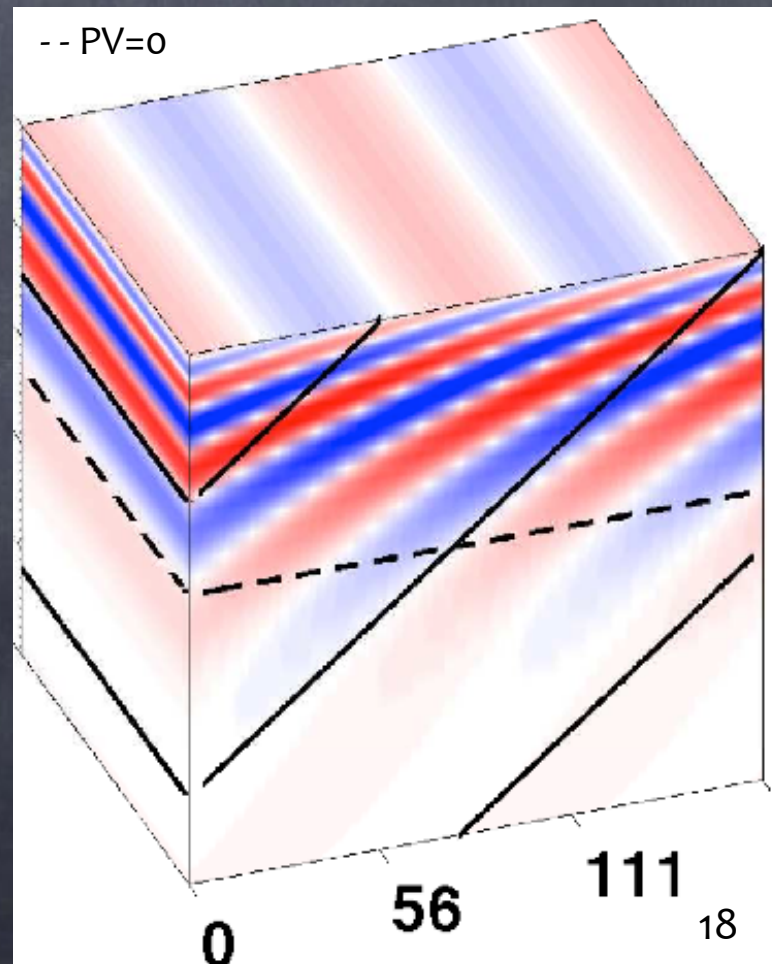
$Ri = 0.5$
Stokes Forces
Stabilize SI

Cross front
velocity for
the fastest
growing
mode

$Ri = 2$
Stokes Forces
Destabilize SI



S. Haney, BFK,
K. Julien, and A.
Webb.
Symmetric and
geostrophic
instabilities in the
wave-forced
ocean mixed
layer. JPO
45:3033-3056,
2015.



Conclusions

- The Stokes vortex or Stokes shear force simplifies the wave-mean interactions, simple enough that simulations are becoming common
- Stokes forces, as treated here, can be included in hydrostatic models like GCMs (wavy hydrostatic)
- Stokes forces affect Langmuir turbulence, but also (sub)mesoscale fronts (more energy, anisotropy) and submesoscale instabilities.
Need to assess climate & environmental impact!

Wave-Averaged

Equations following Lane et al. (07),

McWilliams & F-K (13)

$$\underbrace{v_j^L}_{\text{Lagrangian}} = \underbrace{v_j}_{\text{Eulerian}} + \underbrace{v_j^S}_{\text{Stokes}}$$

Boundary conditions,

plus:

$$Ro [v_{i,t} + v_j^L v_{i,j}] + \frac{M_{Ro}}{Ri} w w_{i,z} + \boxed{\epsilon_{izj} v_j^L} = \text{(Lagrangian) geostrophic} \quad -M_{Ro} \pi_{,i} + \frac{Ro}{Re} v_{i,jj}$$

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b} - \epsilon v_j^L v_{j,z}^S + \frac{\alpha^2}{Re Ri} w_{,jj}$$

hydrostatic

$$b_t + v_j^L b_{,j} + \frac{M_{Ro}}{Ro Ri} w b_z = \frac{1}{Pe} b_{,jj}$$

$$v_{j,j} + \frac{M_{Ro}}{Ro Ri} w_z = 0$$

$$\epsilon = \frac{V^s H}{f L H_s}$$

$$Re = \frac{UL}{\nu} \quad Ro = \frac{U}{fL} \quad Ri = \frac{N^2}{(U_{,z})^2} \quad \alpha = H/L \quad M_{Ro} \equiv \max(1, Ro)$$

J. C. McWilliams and BFK. Oceanic wave-balanced surface fronts and filaments. *Journal of Fluid Mechanics*, 730:464-490, 2013.

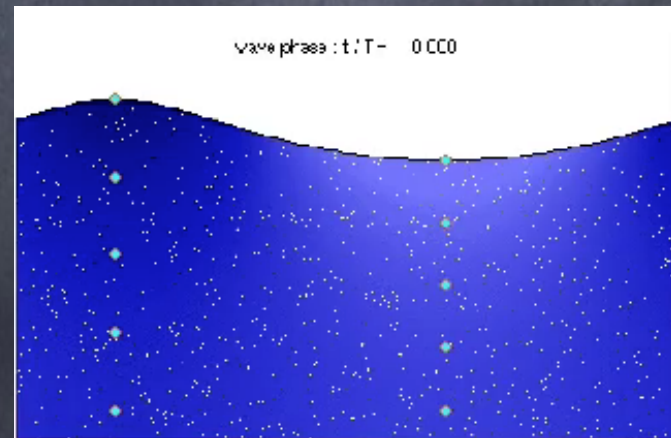
N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. *Journal of Geophysical Research-Oceans*, 121:1-18, 2016.

3 Wave Effects, 1: Lagrangian Advection: Particles, tracers, momentum flow with Lagrangian, not Eulerian flow

$$\begin{aligned}
 Ro [v_{i,t} + v_j^L v_{i,j}] + \frac{M_{Ro}}{Ri} w v_{i,z} + \epsilon_{izj} v_j^L &= -M_{Ro} \pi_{,i} + \frac{Ro}{Re} v_{i,jj} \\
 \frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] &= -\pi_{,z} + b - \epsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{Re Ri} w_{,jj} \\
 b_t + v_j^L b_{,j} + \frac{M_{Ro}}{Ro Ri} w b_z &= \frac{1}{Pe} b_{,jj}
 \end{aligned}$$

Adding a Stokes advection
 term converts
 total to Lagrangian advection

$$\underbrace{v_j^L}_{\text{Lagrangian}} = \underbrace{v_j}_{\text{Eulerian}} + \underbrace{v_j^S}_{\text{Stokes}}$$



3 Wave Effects, 2: Lagrangian Coriolis: Particles, tracers, momentum flow with Lagrangian, not Eulerian flow—Experience Coriolis force during this motion

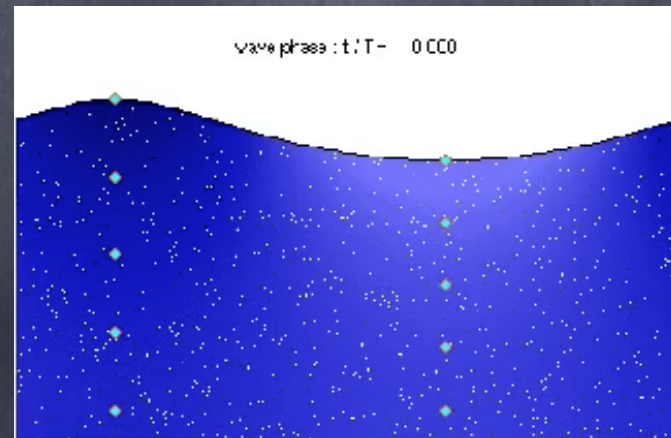
$$Ro [v_{i,t} + v_j^L v_{i,j}] + \frac{M_{Ro}}{Ri} w v_{i,z} + \boxed{\epsilon_{izj} v_j^L} = -M_{Ro} \pi_{,i} + \frac{Ro}{Re} v_{i,jj}$$

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = -\pi_{,z} + b - \epsilon v_j^L v_{j,z}^S + \frac{\alpha^2}{Re Ri} w_{,jj}$$

$$b_t + v_j^L b_{,j} + \frac{M_{Ro}}{Ro Ri} w b_z = \frac{1}{Pe} b_{,jj}$$

Adding a Stokes Coriolis
 term converts total to
 Lagrangian

$$\underbrace{v_j^L}_{\text{Lagrangian}} = \underbrace{v_j}_{\text{Eulerian}} + \underbrace{v_j^S}_{\text{Stokes}}$$



Imperfect Frontogenesis: Governing equations with Eddy Viscosity and Diffusivity

x momentum equation
(horizontal cross front)

$$\frac{Du}{Dt} - fv = \alpha u - \frac{1}{\rho_0} \frac{\partial p}{\partial x} + \boxed{\nu \nabla^2 u}$$

(Viscosity)

y momentum equation
(horizontal along front)

$$\frac{Dv}{Dt} + fu - \alpha v + \boxed{\nu \nabla^2 v}$$

z momentum equation
(vertical)

$$0 = b - \frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

Thermodynamic equation

$$\frac{Db}{Dt} = \boxed{\kappa \nabla^2 b}$$

(Diffusivity)

Conservation of mass

$$\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0$$

u - Cross front velocity

v - Along front velocity

w - Vertical velocity

p - Pressure

b - Buoyancy

ρ_0 - Density

f - Coriolis

α - Strain field

$\frac{D}{Dt}$ - Lagrangian derivative

κ - Diffusivity

ν - viscosity

Perturbation analysis

Assuming eddy viscosity and diffusivity are small corrections to the conventional strain induced frontogenesis equations.

$$\begin{aligned}U &= \bar{U} + u = \varepsilon^0(\bar{U} + u^0) + \varepsilon^1 u^1 + O(\varepsilon^2) \\V &= \bar{V} + v = \varepsilon^0(\bar{V} + v^0) + \varepsilon^1 v^1 + O(\varepsilon^2) \\W &= w = \varepsilon^0 w^0 + \varepsilon^1 w^1 + O(\varepsilon^2)\end{aligned}$$

Note, zeroth order contains both geostrophic and ageostrophic flows—but only *inviscid* flows

$$\varepsilon = \frac{1}{Ek} = \frac{\nu}{fL^2}$$

Eddy viscosity
(Ek – Ekman number)

$$\varepsilon = \frac{1}{Pe} = \frac{\kappa}{fL^2}$$

Eddy diffusivity
(Pe - Peclet-ish number)

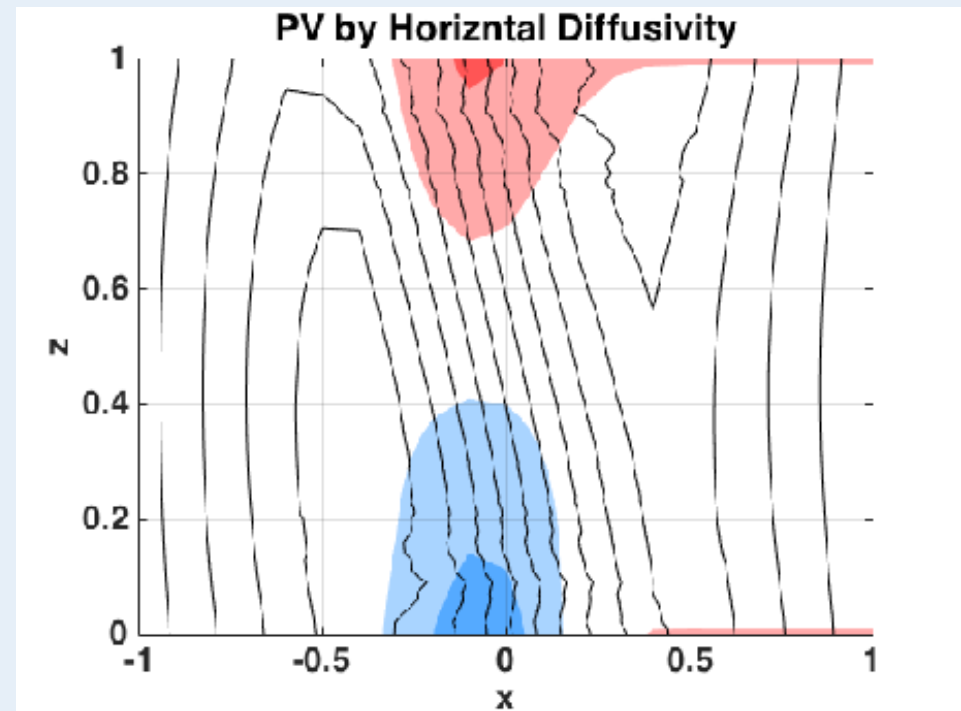
**Zeroth order –
known
solution.**

**First order –
addition due to
forcing terms**

At first order, an inviscid, diffusive PV effect is imposed,
and (e.g. Singh, Veronis & Taylor, 2013)

Localized potential vorticity -> Green's fct. analysis

- Studies (e.g. Thomas 2005) show frontal PV is localized due to friction.
- Zeroth-order perturbation solution is inviscid, zero-PV.
- Most of the turbulence and mixing (i.e., PV source) is at the front.
- Thus the potential vorticity can be approximated as a **point source/sink** at the front.



Solid lines represent buoyancy, red- positive PV, blue- negative PV (Thomas 05)

First order solution— Amenable to Green's Fct. treatment due to point PV sources

PV equation nontrivial at first order:

Viscous forcing

Diffusion

First order PV
using zeroth order
terms

$$\frac{D}{Dt_0} Q^1 = (\nabla \times \mathbf{D}_u^1) \cdot \nabla b^0 + \omega_a^0 \cdot (\nabla D_b^1)$$

$$Q^1 = \frac{1}{f^2} \frac{\partial^2 \phi^0}{\partial z^2} \frac{\partial^2 \phi^1}{\partial x^2} - \frac{2}{f^2} \frac{\partial^2 \phi^0}{\partial x \partial z} \frac{\partial^2 \phi^1}{\partial x \partial z} + \left(1 + \frac{1}{f^2} \frac{\partial^2 \phi^0}{\partial x^2} \right) \frac{\partial^2 \phi^1}{\partial z^2}$$

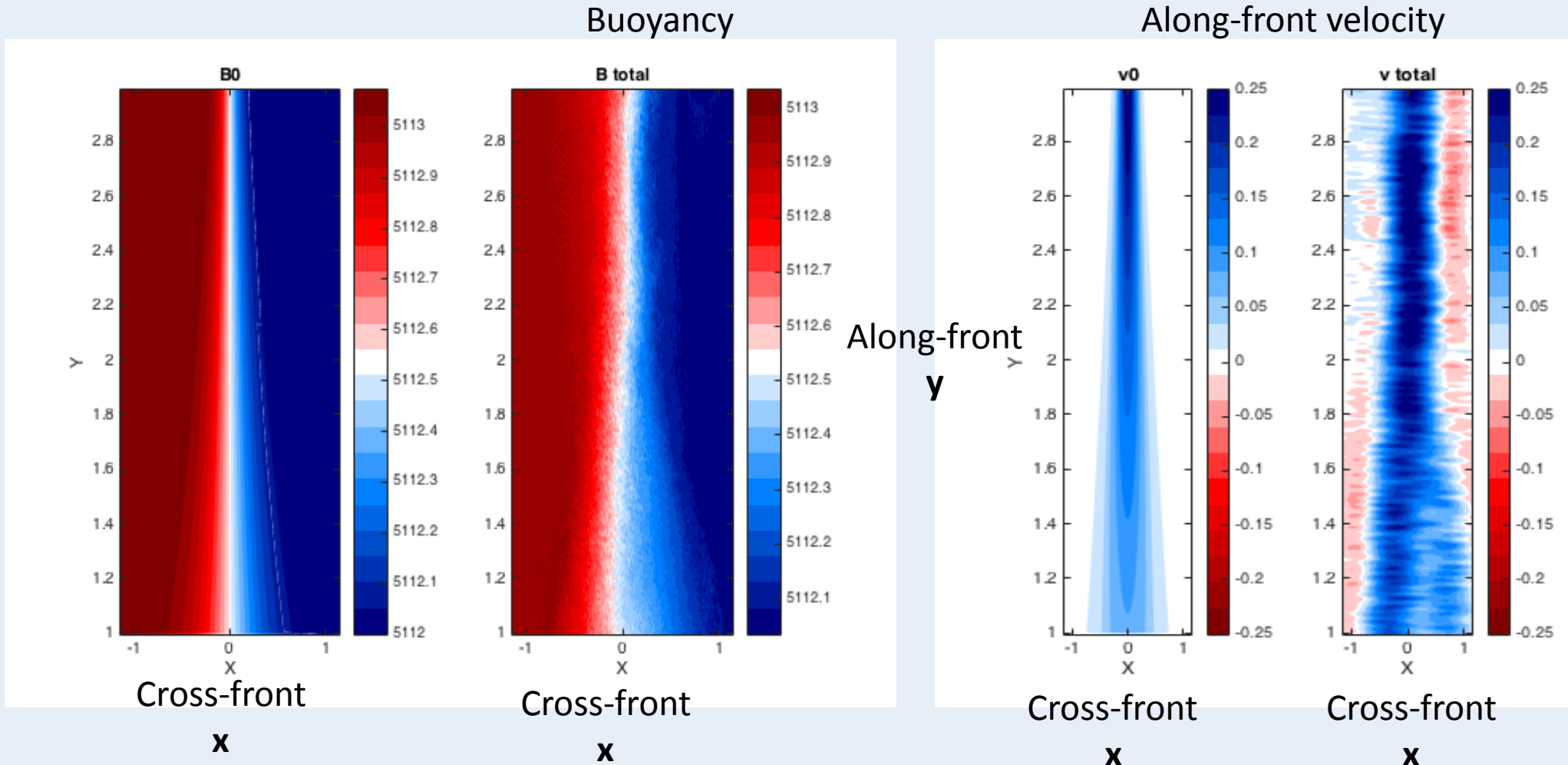
↑

$\delta(x - x_{front})$

Second order linear PDE in first-order
pressure correction.

Analytically invertible if PV is a point source (delta fct.)

Perturbation analysis guides analysis of frontal Large Eddy Simulations



First order solution

- Perturbation parameter ε for both viscosity and diffusivity is $O(0.1)$
- The largest contributor is horizontal diffusivity.
- Perturbation theory appears to be valid.
- We can find first order velocities.

