

Biosphere Atmosphere Interactions: Ingredients for Toy Models

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(1) Who: Name Calling

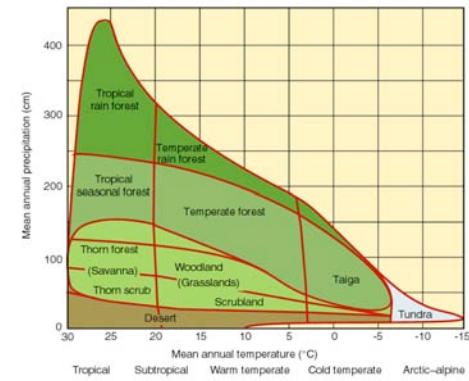
1. Taxa/Species
2. **Physionomic/structural:** Biomes e.g. desert, rainforest - albedo, roughness
3. **Bioclimatic + CO₂ +...**
4. **Functional:** based on satellite photosynthesis index:



$$\{\% \text{tree} / \text{shrub}\} \times \begin{cases} \text{deciduous} \\ \text{evergreen} \end{cases} \times \begin{cases} \text{broadleaved} \\ \text{needleleaved} \end{cases}$$

$$\{\% \text{herbaceous} / \text{crops}\} \times \begin{cases} \text{C3} \\ \text{C4} \end{cases}$$

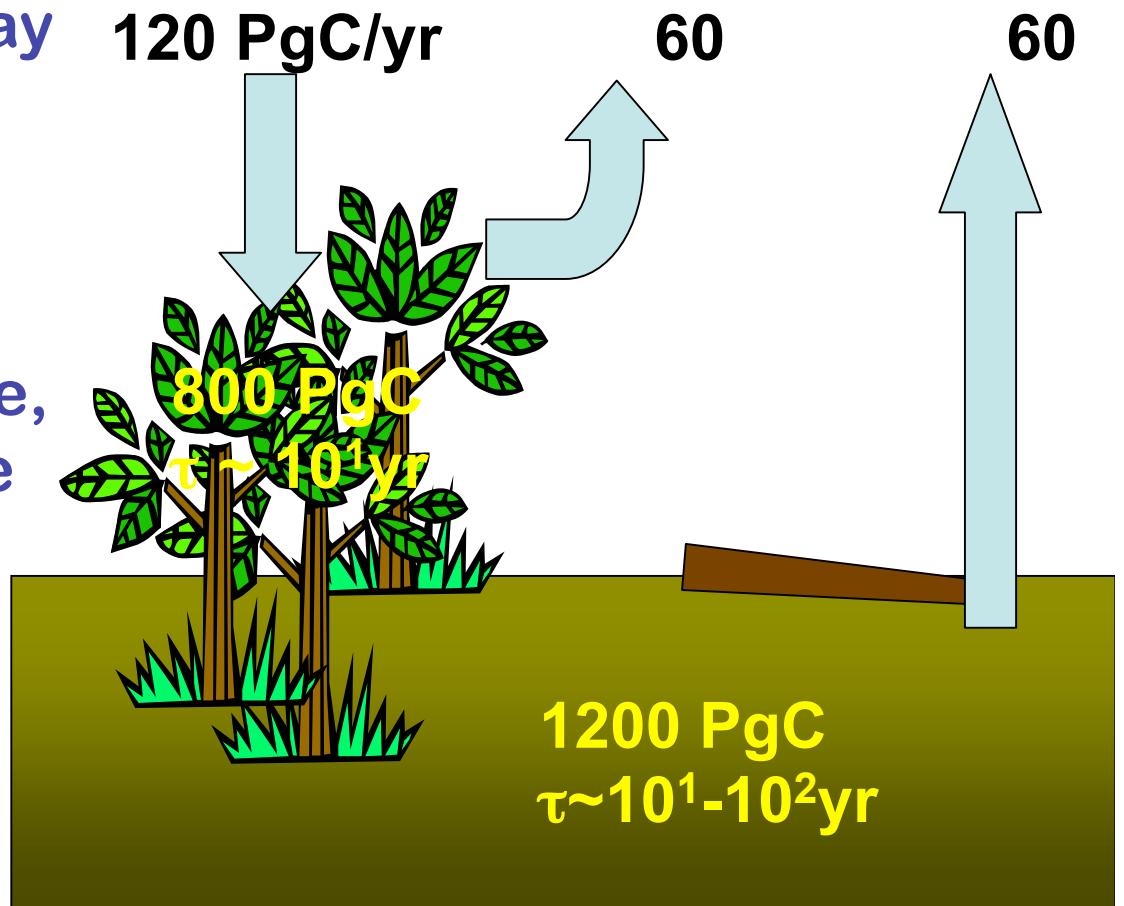
$$\{\% \text{bare}\}$$



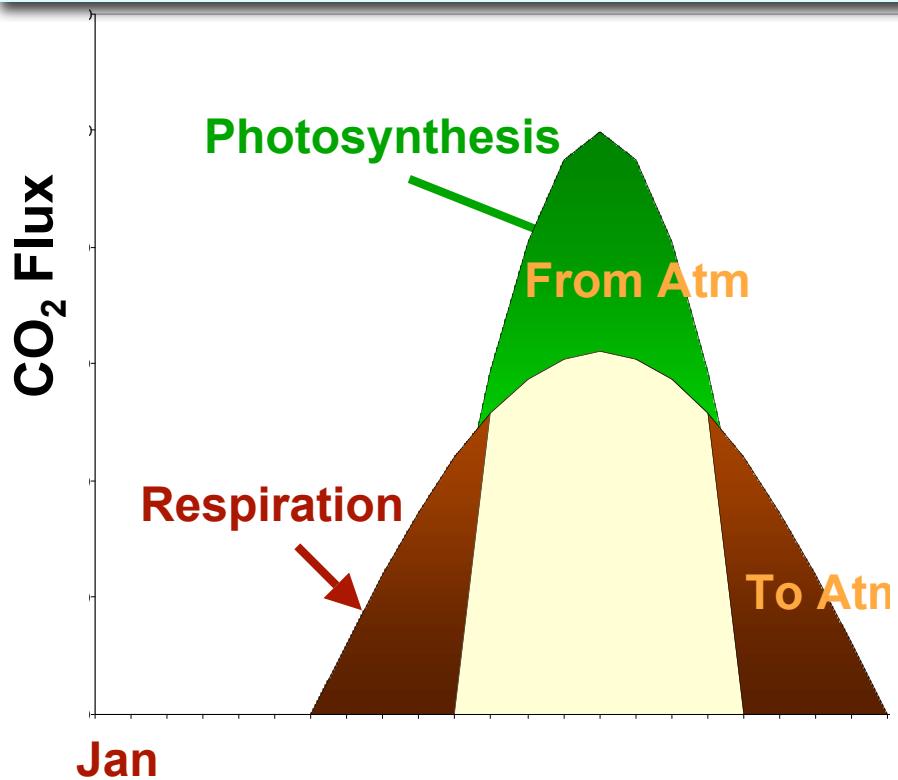
Static land cover: typically 4 --> lookup table of structural properties and BGC characteristics. Prognostic land cover: 3+

(2) What? [lifecycle traced by C cycle]

- Growth, mortality, decay 120 PgC/yr
- Population: {ages}
- Photosynthesis (climate, CO_2 , soil H_2O , resource limitation)
- Decay (T , soil H_2O ,...)



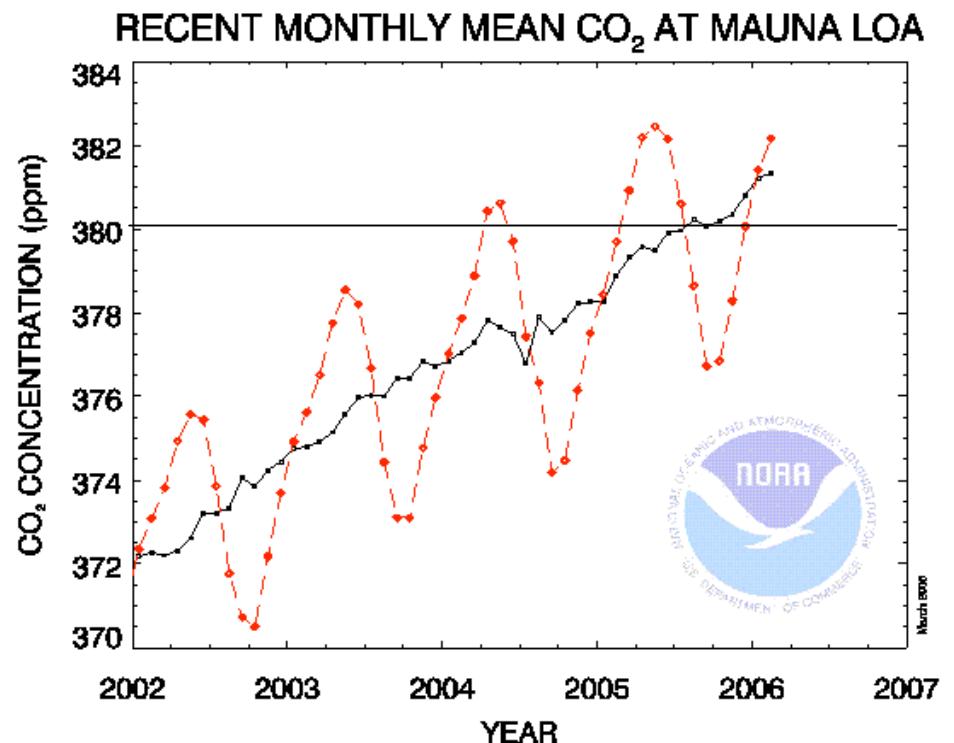
Two major fluxes - seasonal cycle



- **Seasonal asynchrony between photosynthesis and decomposition**

- net fluxes of CO₂ to and from atm
- seasonal cycle of CO₂ in atm

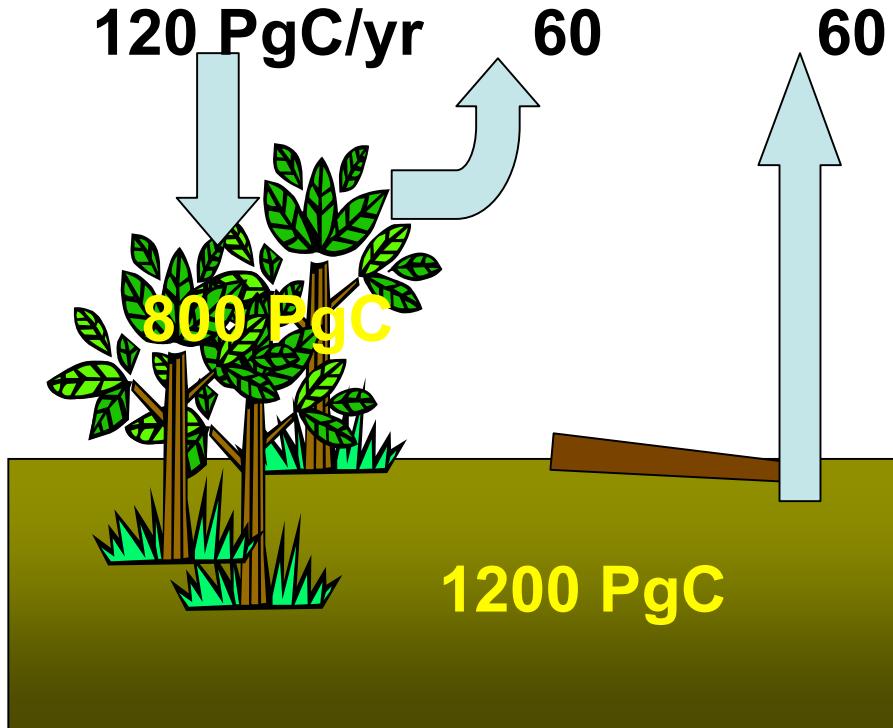
- **Annual imbalance** → carbon source/sink



Fung et al JGR 1983, 1987:
Simple prescriptions of
seasonality. 1987 based on
satellite obs + temperature

(3.1) How?

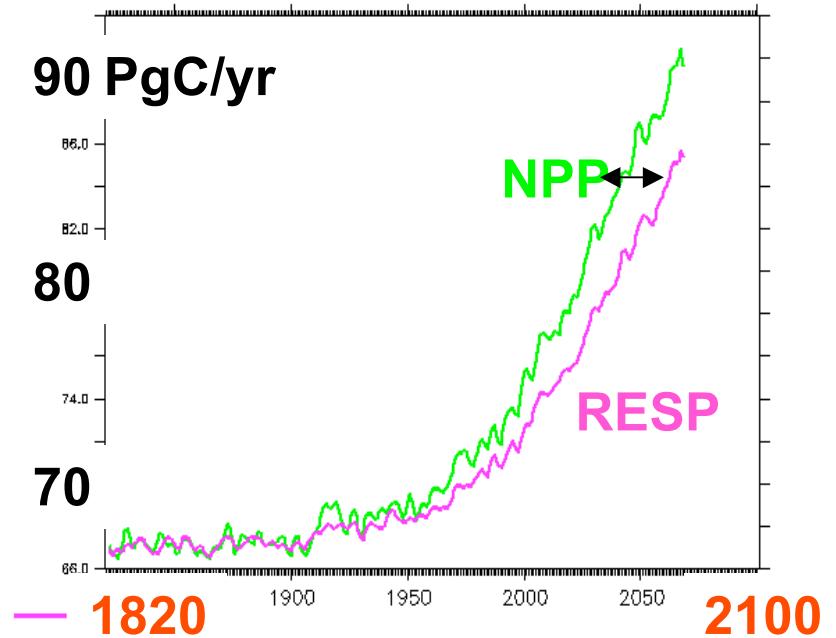
First simple toy model: carbon only



$$\tau_{veg} \sim 15 \text{ yr}$$

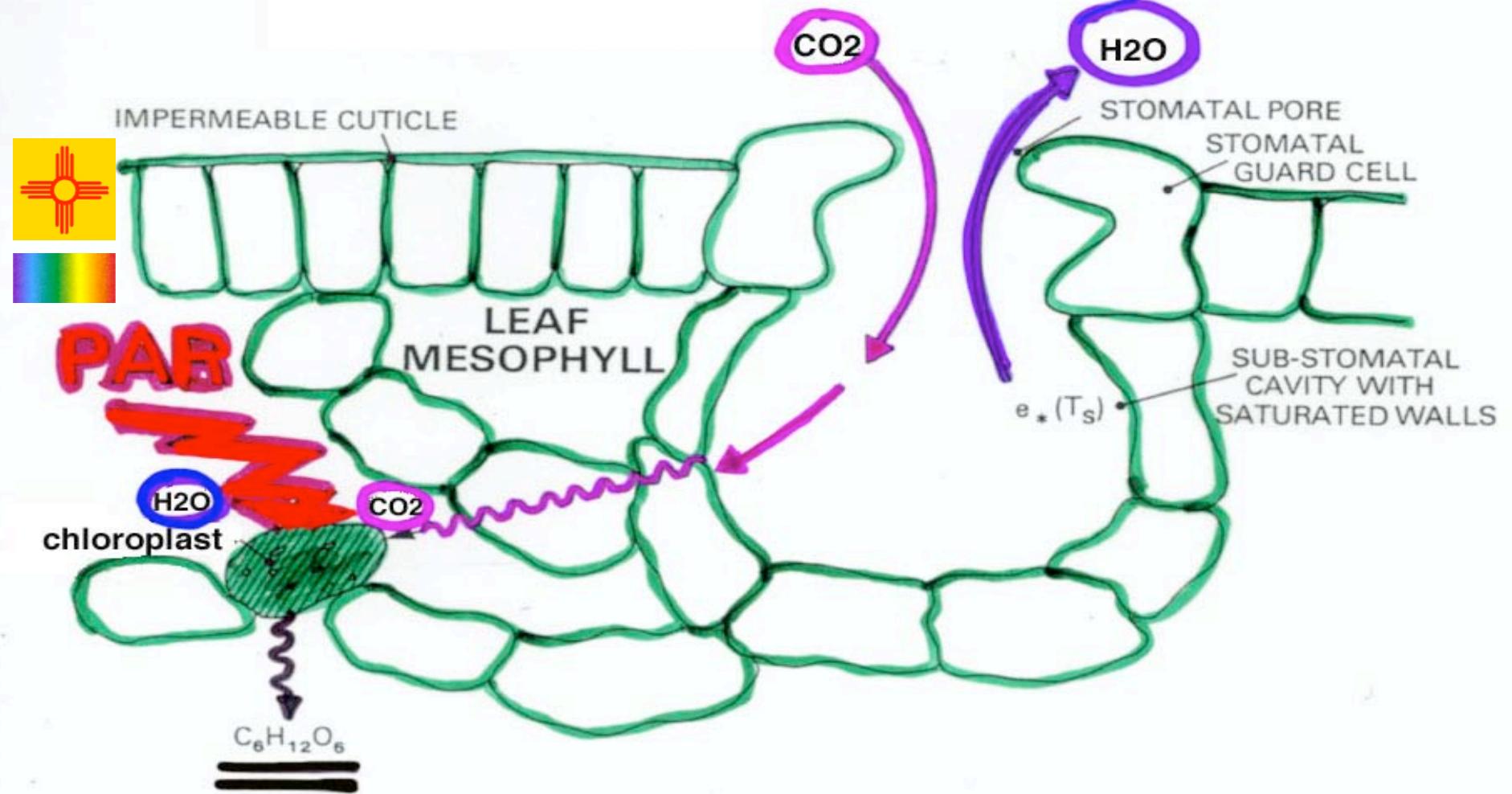
$$\tau_{soil} \sim 25 \text{ yr}$$

$$\frac{\partial M_{veg}}{\partial t} = NPP - \frac{M_{veg}}{\tau_{veg}}$$
$$\frac{\partial M_{soil}}{\partial t} = \frac{M_{veg}}{\tau_{veg}} - \frac{M_{soil}}{\tau_{soil}}$$



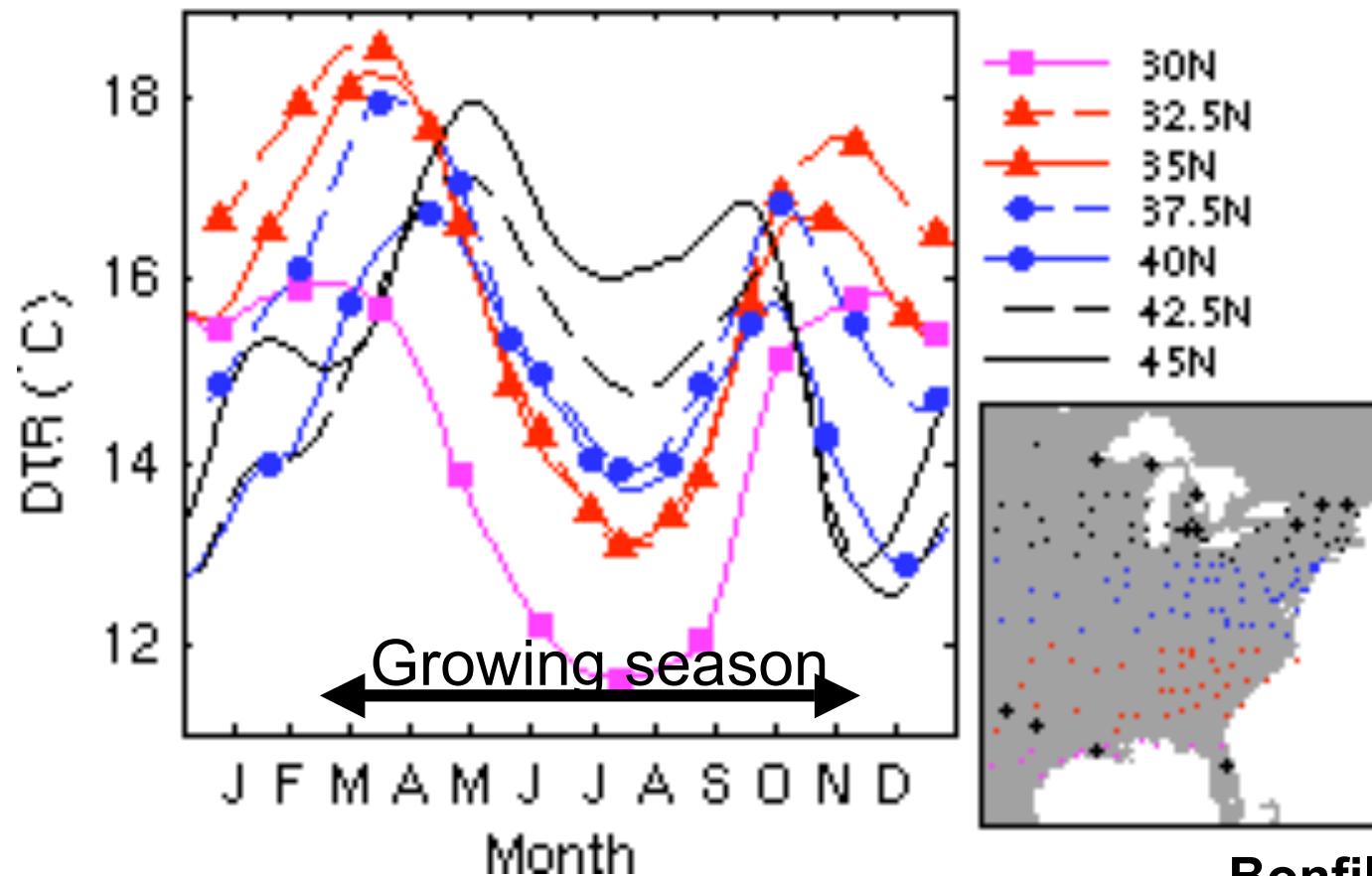
(3.2.1) Leaf Photosynthesis

Piers Seller's PAR Diagram



Diurnal Temperature Range: coupling of energy-water-carbon fluxes

$$C \frac{\partial T_g}{\partial t} = SW \uparrow + LW \uparrow - \underbrace{SH}_{\text{warms PBL}} - \underbrace{LH}_{\text{cools PBL}}$$



Spring chill
when
growing
season
starts

Bonfils et al. GRL 2005

(3.2.1) Rates of Carbon Assimilation and Transpiration

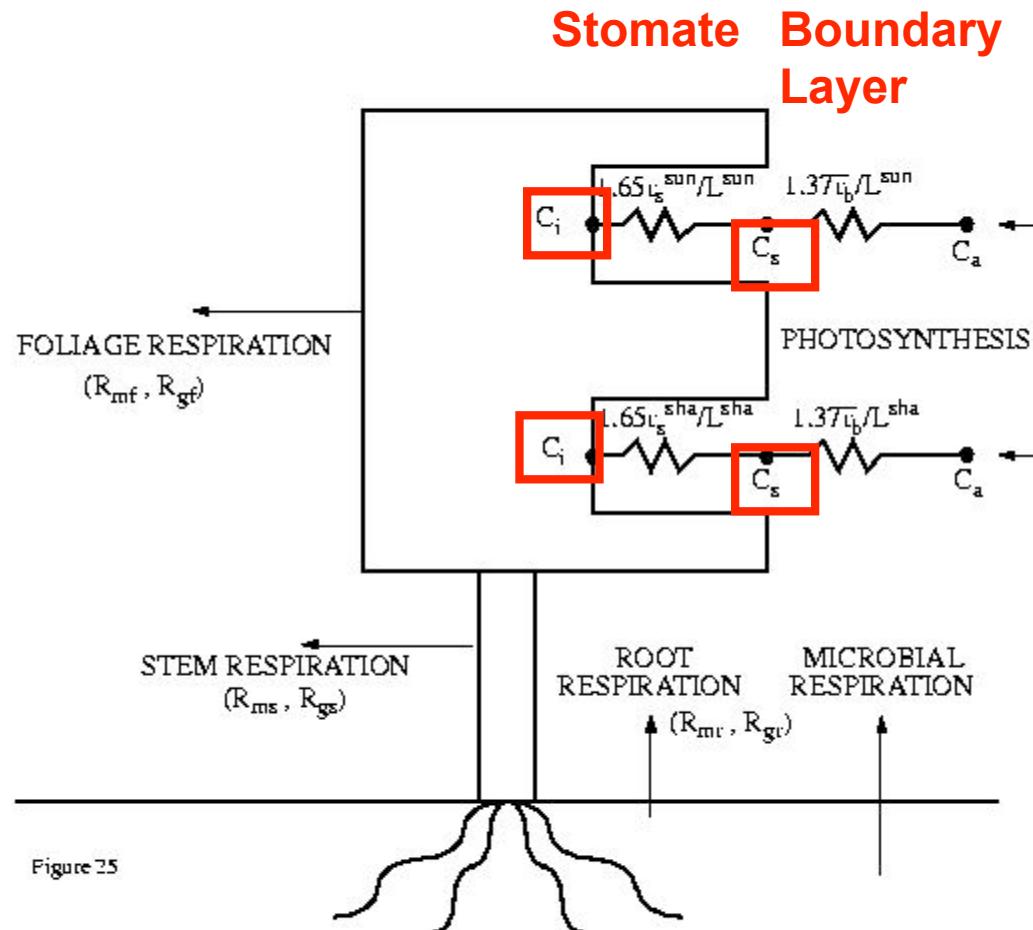


Figure 25

Farquhar:
$$A = g_s (C_a - C_i)$$

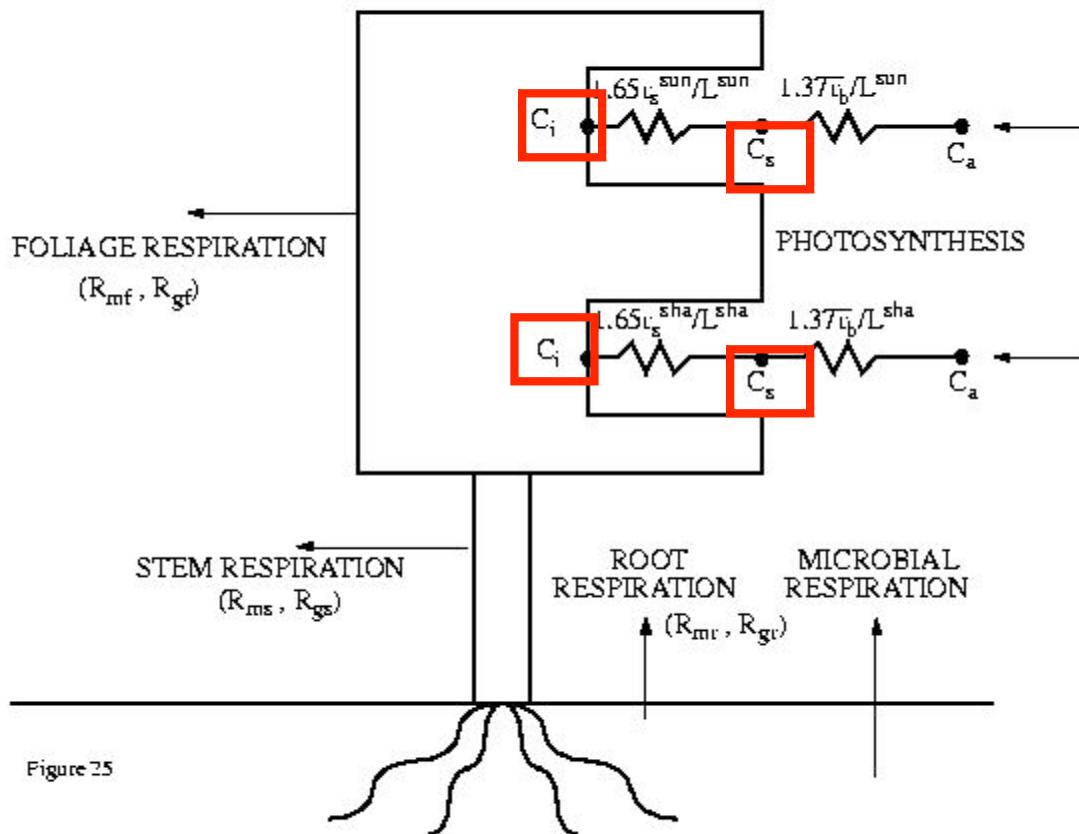
Collatz: leaf boundary layer
$$A = g_l (C_a - C_s) \\ = g_s (C_s - C_i)$$

Need to determine
stomatal conductances g_s
and C_i

Transpiration
 $Tr = g_w (1 - RH_a)$

Assimilation Rates of Sun and Shade Leaves

(Sellers et al. J Climate 1996)



Per area of leaf sfc:

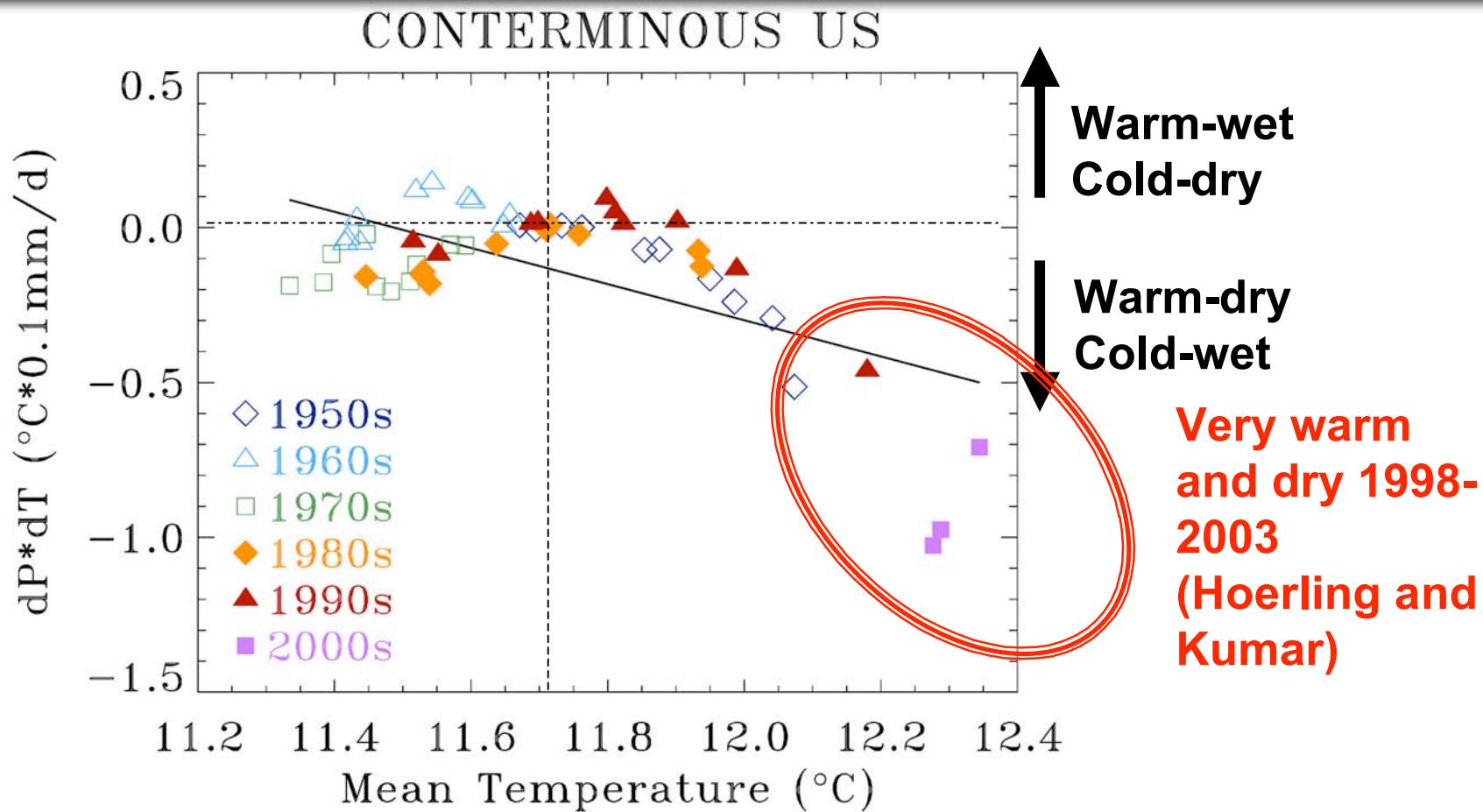
$A_{\text{sun},\text{shade}} = \min \{$
 PEP-Carboxylase
 Light_sun,shade
 Export/utilization }
limited rates of assimilation

Varies with plant functional types

Normalize canopy

$$\begin{aligned} GPP = & A_{\text{sun}} * f_{\text{sun}} * \text{LAI} \\ & + A_{\text{shade}} * f_{\text{shade}} * \text{LAI} \end{aligned}$$

Recent Drought: Co-Variations of ΔT and ΔPrecip

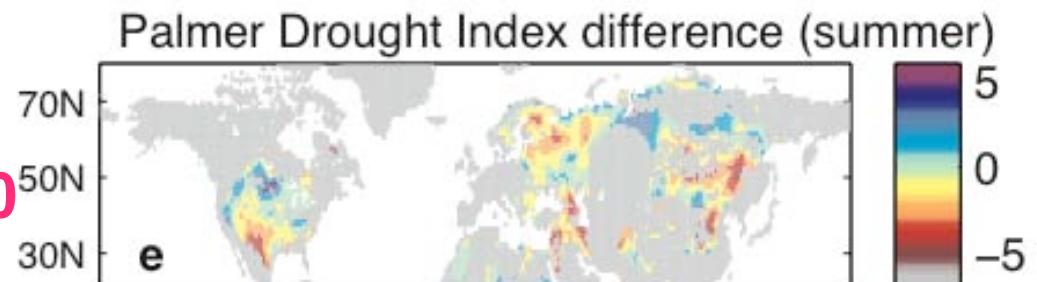
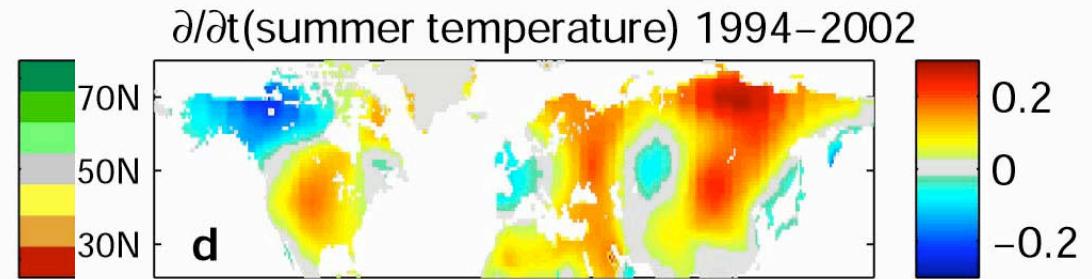
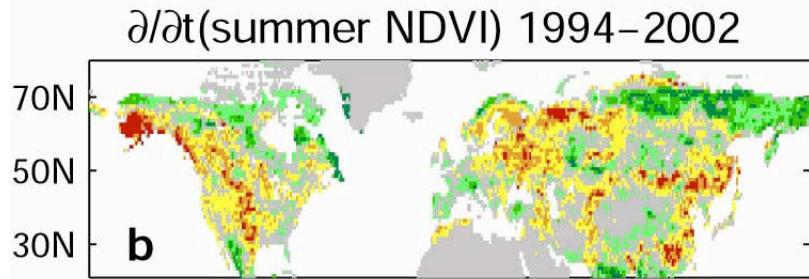
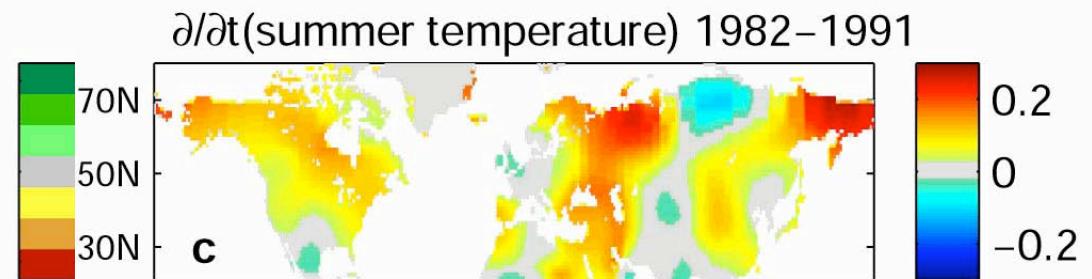
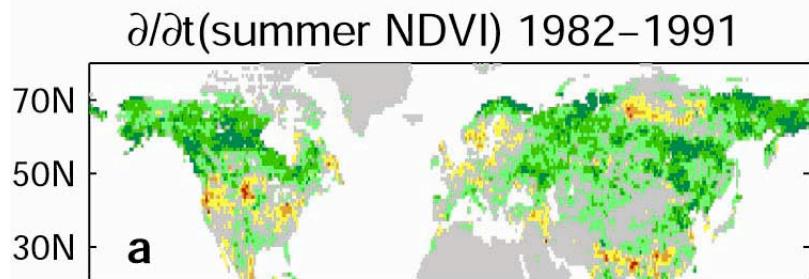


Buermann et al. PNAS 2007

Global distribution of droughts:

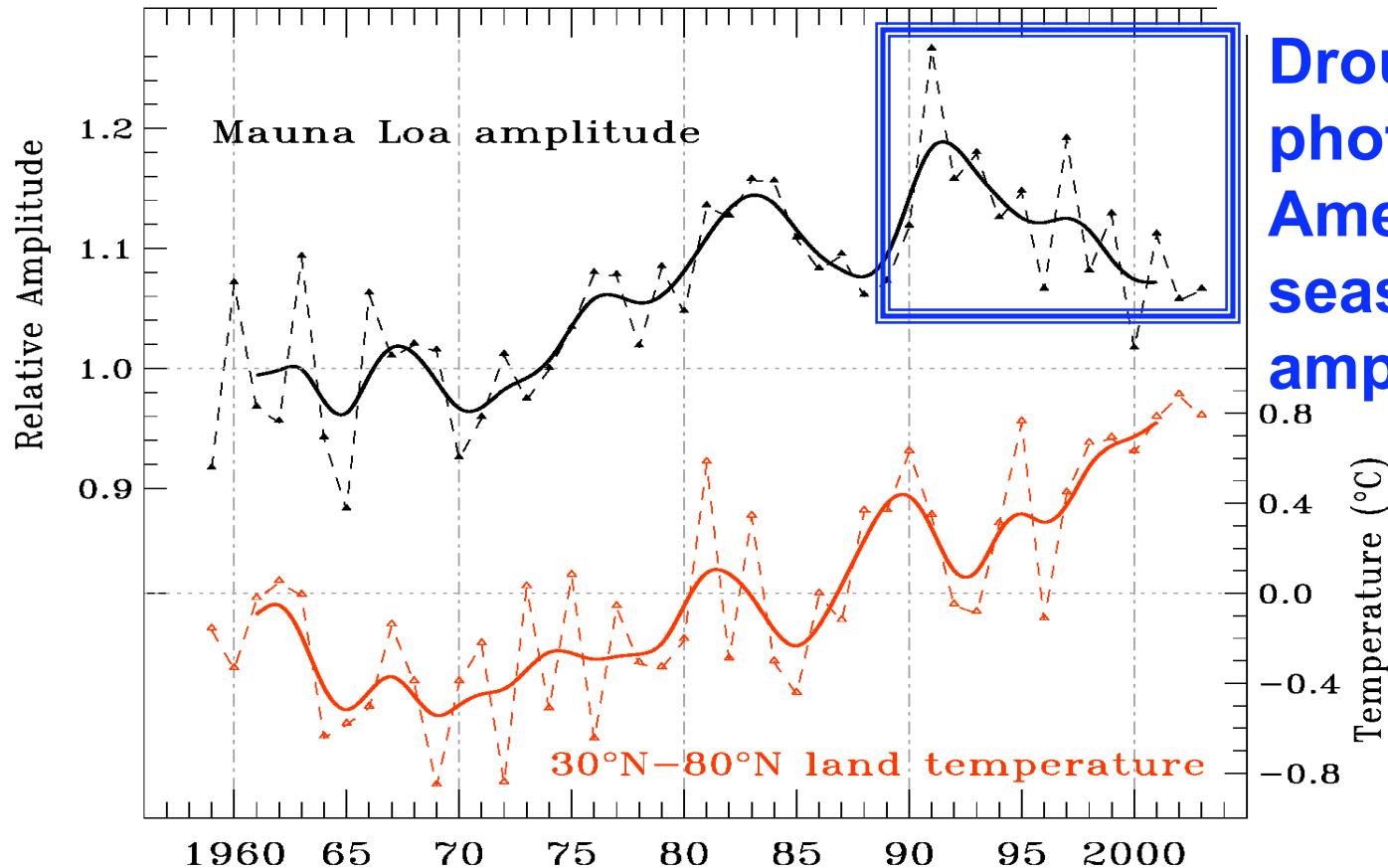
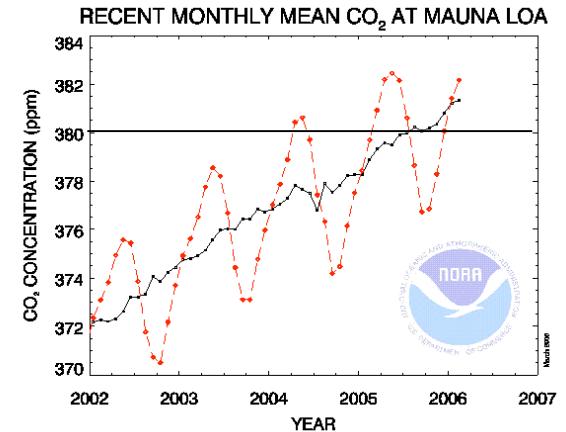
Dai, A., K. E. Trenberth, and T. Qian "A global data set of Palmer Drought Severity Index for 1870-2002: Relationship with soil moisture and effects of surface warming" *J. Hydrometeorology*, 2005.

Halting of the Summer Greening Trend: slowing of northern hemisphere land sink



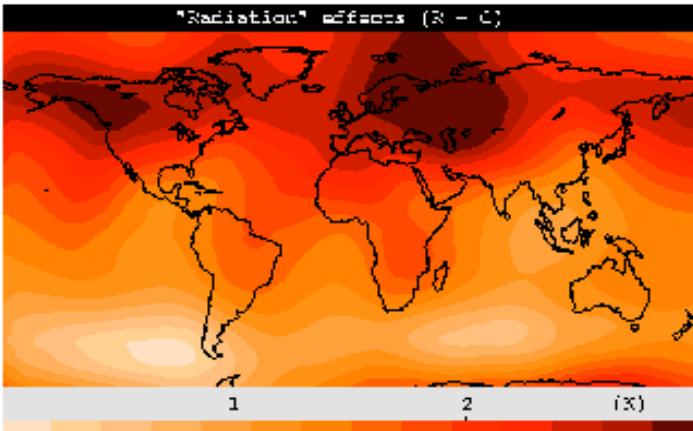
Where
 $\partial/\partial t(\text{NDVI}) < 0$ and $\partial/\partial t(T) > 0$
water stress

Evidence for slowing NH land sink: Changes in MLO Amplitude since 1990

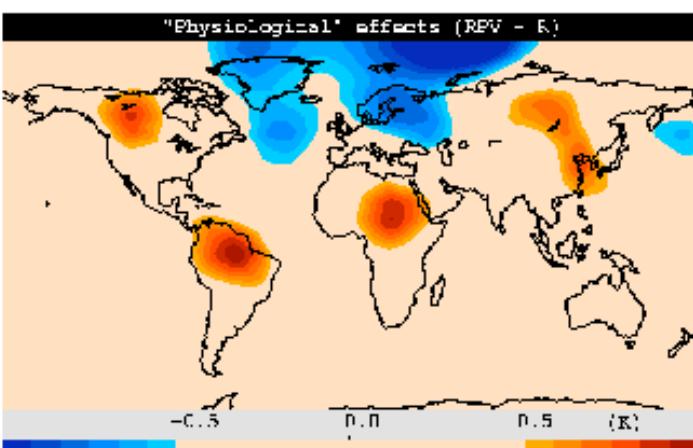


Droughts reduced photosyn in N America and seasonal CO₂ amplitude at MLO

2xCO₂: Additional Feedbacks due to Plant Physiology in GCM with biophysics.



Warming
due to
radiation



Additional
Warming
due to
plant physiology

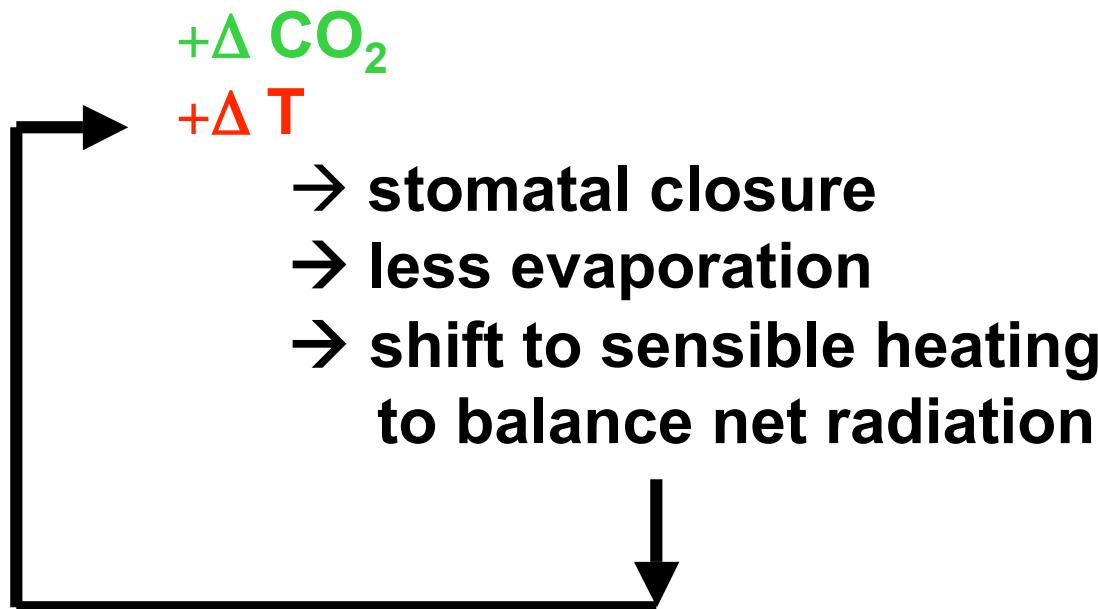
In the tropics:

- Nutrient limitation of photosynthesis
- Stomatal closure at high water stress
- Reduces transpiration and
- Causes net radiation to be balanced by sensible rather than latent heating

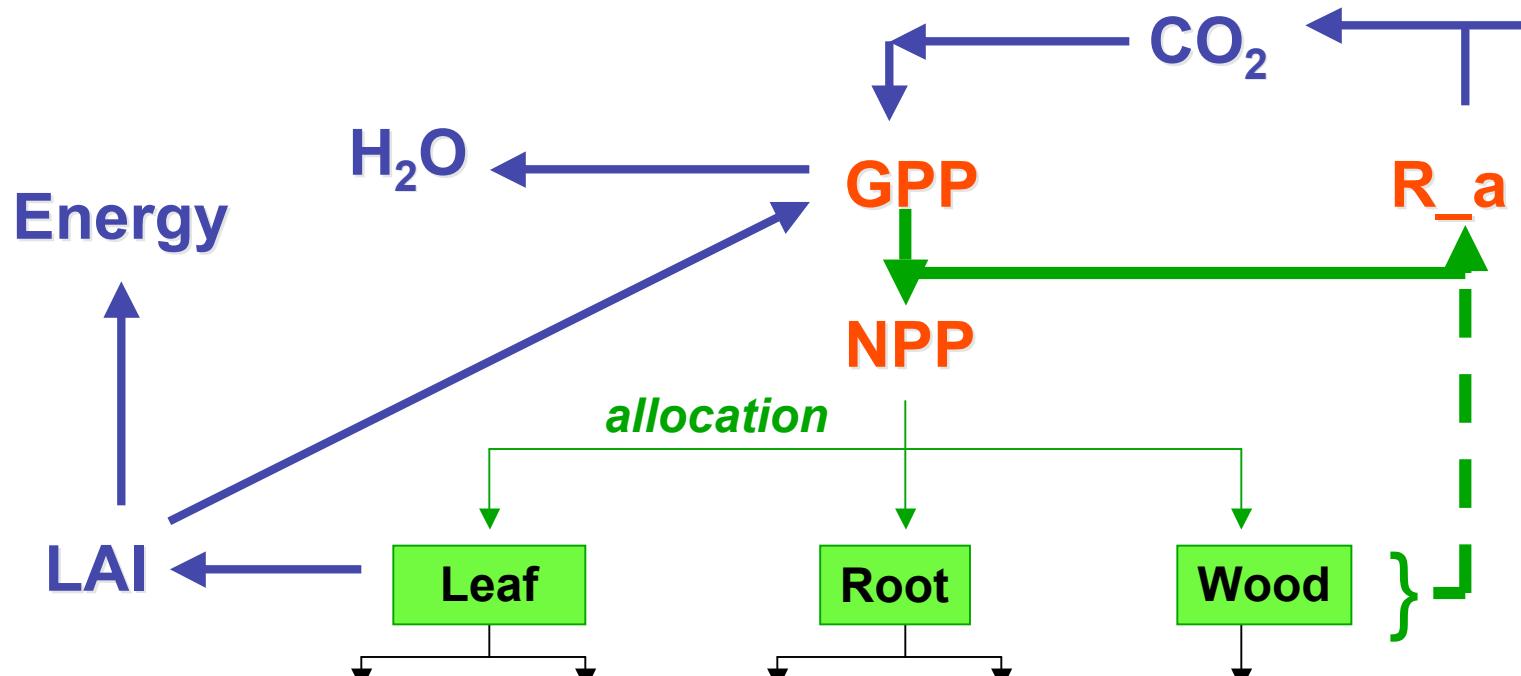
Sellers et al. Science 1996

Stomatal Suicide

$$C \frac{\partial T_g}{\partial t} = SW \uparrow + LW \uparrow - \underbrace{SH}_{warms PBL} - \underbrace{LH}_{cools PBL}$$



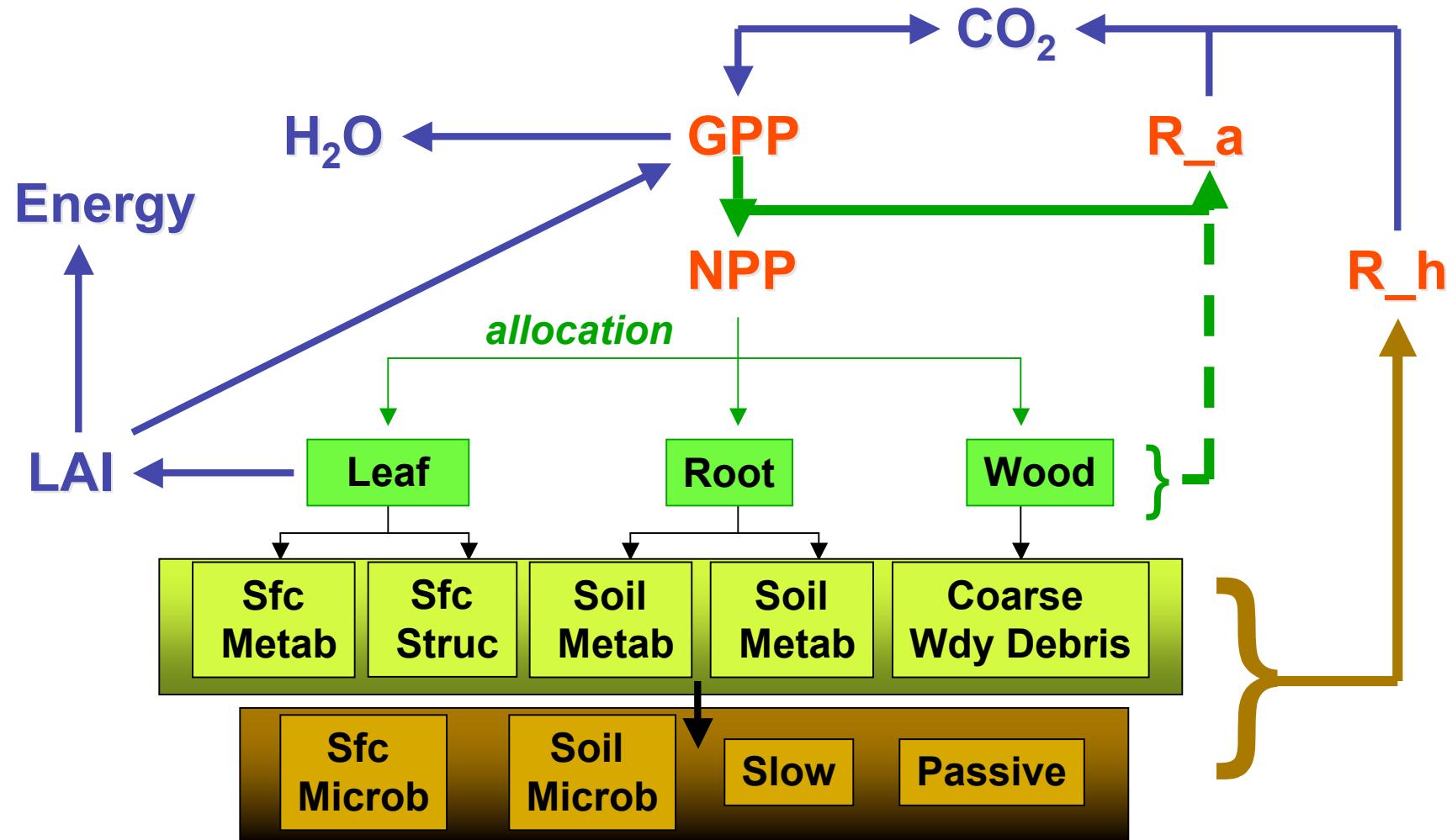
(3.2.2) Photosynthesis --> Growth



Strategy for survival: allocate new C to
Roots when water and nutrient limited
Wood when light limited

Friedlingstein et al. GCB 1995

(3.2.3) Photosynthesis, Growth, Mortality & Decay



(3.2) Prognostic Carbon Cycle (with transpiration and dynamic albedo)

Atm

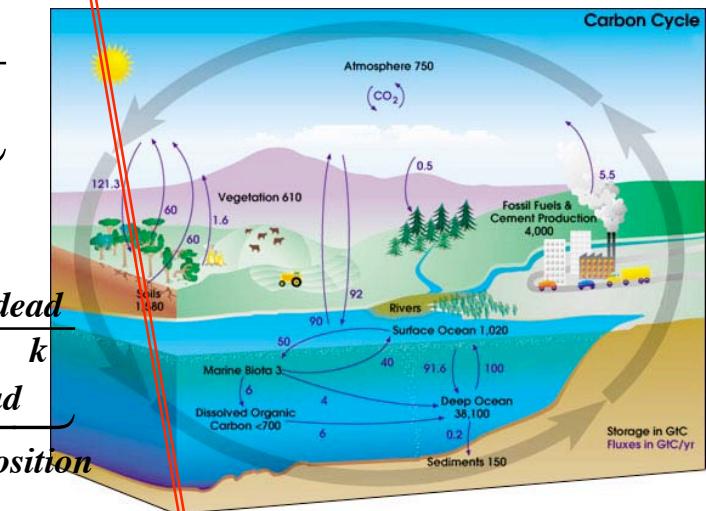
$$\frac{DC_a}{Dt} = (FF + Def + \underbrace{F_{oa} \uparrow\downarrow}_{air-sea_flux} + \underbrace{F_{ba} \uparrow\downarrow}_{atm-land_flux})_0 + \Im(C_a)$$

Land-live

$$\frac{\partial C_{b_live}^k}{\partial t} = -\alpha^k \underbrace{F_{ab} \downarrow}_{photosynthesis}_0 - \frac{C_{b_live}^k}{\tau_{live}^k} \underbrace{\phantom{F_{ab}}}_{mortality}$$

Land-dead

$$\frac{\partial C_{b_dead}^k}{\partial t} = \frac{C_{b_live}^k}{\tau_{live}^k} + \sum_j F_{jk} - \frac{C_{b_dead}^k}{\tau_{dead}^k} \underbrace{\phantom{F_{jk}}}_{decomposition}$$

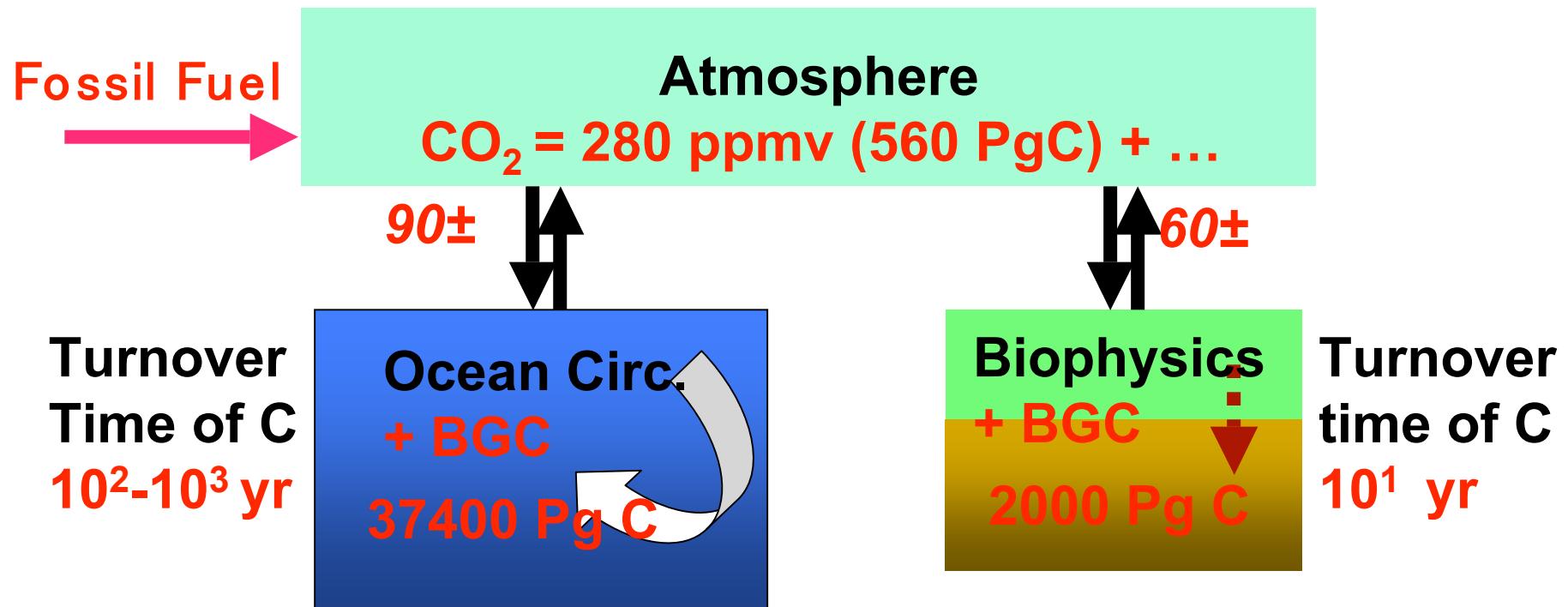


Changes climate

Responds to
climate change

(4) Will the warming accelerate the warming? CO₂-Climate Feedbacks

Technical Goal: include interactive carbon dynamics in the climate model. Specify Emission (t), not atm CO₂ (t)



Challenge: validation of model

How would CO₂ and climate co-vary?

Suppose there is warming...

Atm CO₂ would increase because:

- Warming may enhance decomposition
- Increased ocean stratification → more carbon in mixed layer → reduced air-to-sea flux
-

Atm CO₂ would decrease because:

- warming may enhance photosynthesis
- Enhanced marine productivity and export

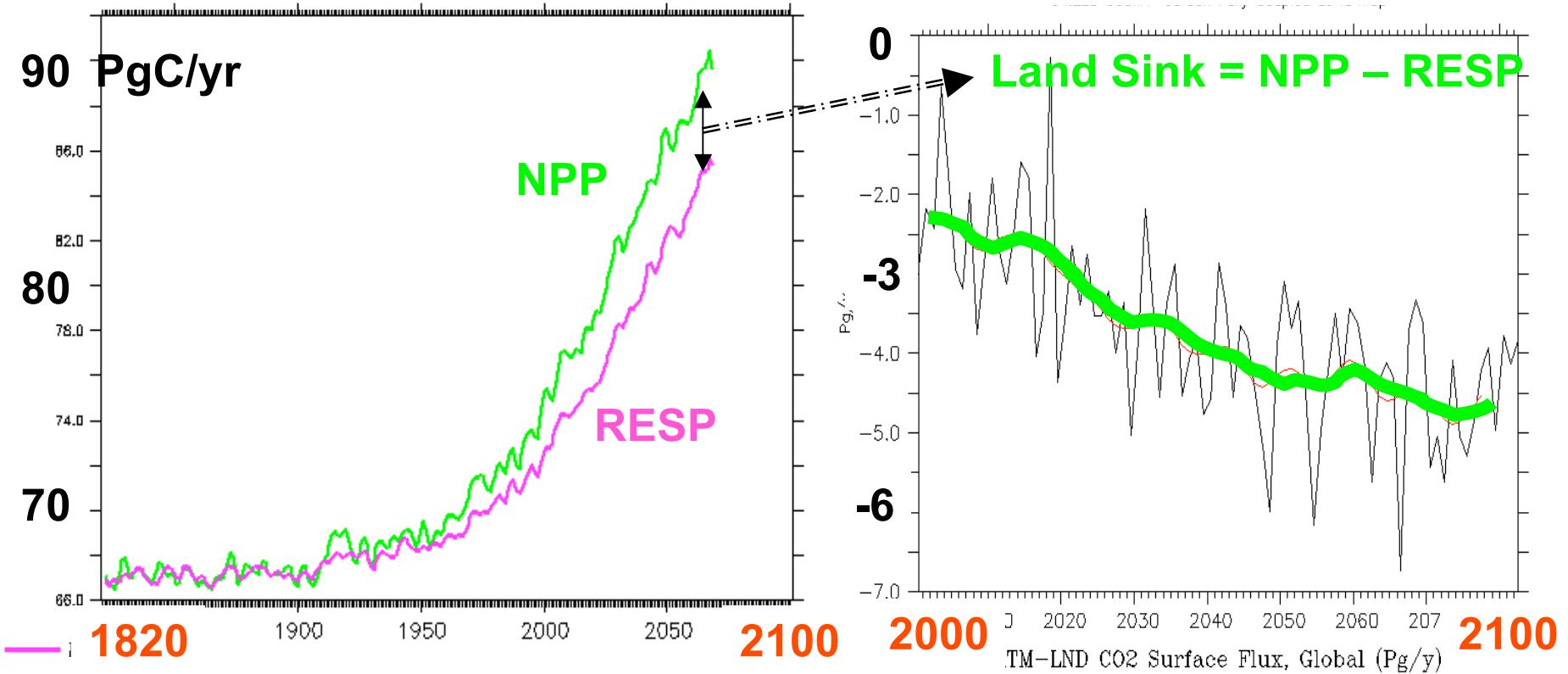
In model, three flavors of CO₂:

- CO₂_tracer(x,y,z,t)
- CO₂_bgc=CO₂_tracer(x,y,**lowest layer**,t)
- CO₂_rad=CO₂_tracer(x,y,**column**,t)

Models expts:
BGC coupling,
Radiative coupling

Land-Atm Fluxes

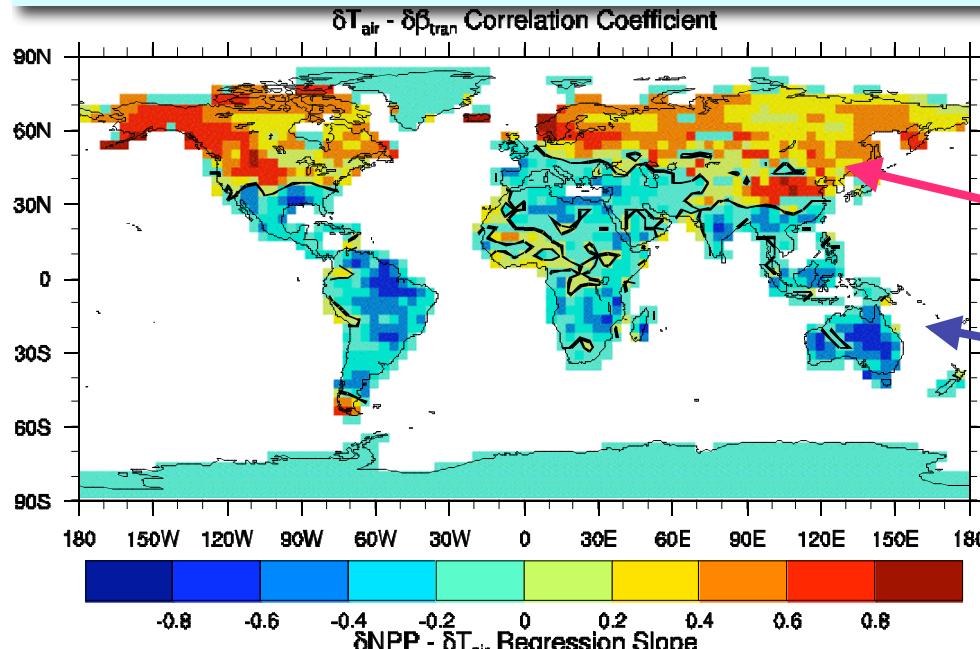
NCAR CCSM1-Carbon: FF



NPP incr with atm CO₂ + ΔT
RESP incr with incr biomass
+ ΔT

Slowing down of land sink
around 2050 despite continued
increase in atm CO₂ and
temperature

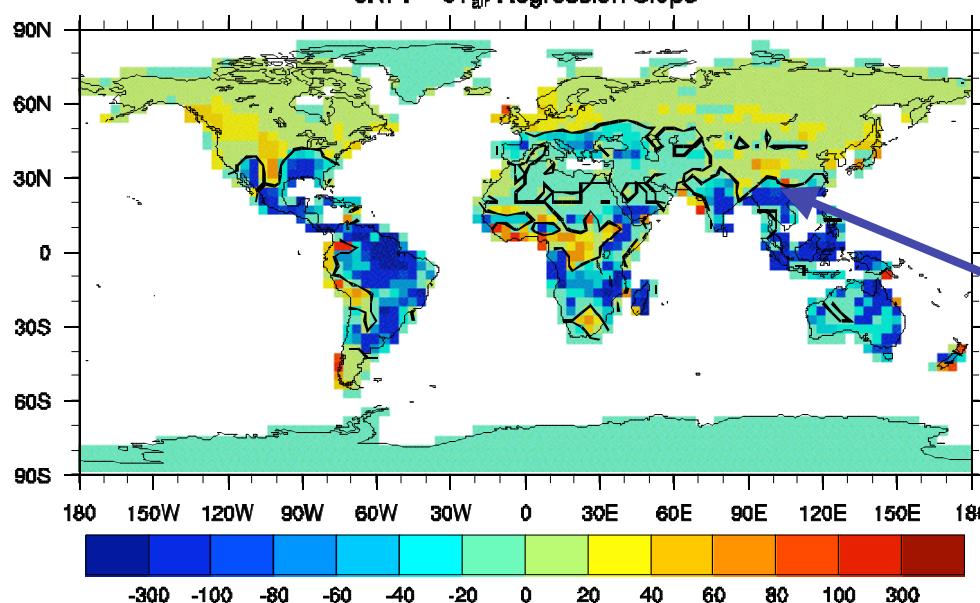
21st Century Correlations & Regressions: FF= SRES A2 ; δ = Coupled minus Uncoupled



{ δT , δ Soil Moisture Index}

W arm-wet

W arm-dry



Regression of δNPP vs δT

NPP decreases with carbon-climate coupling

Fung et al. Evolution of carbon sinks in a changing climate. PNAS 2005

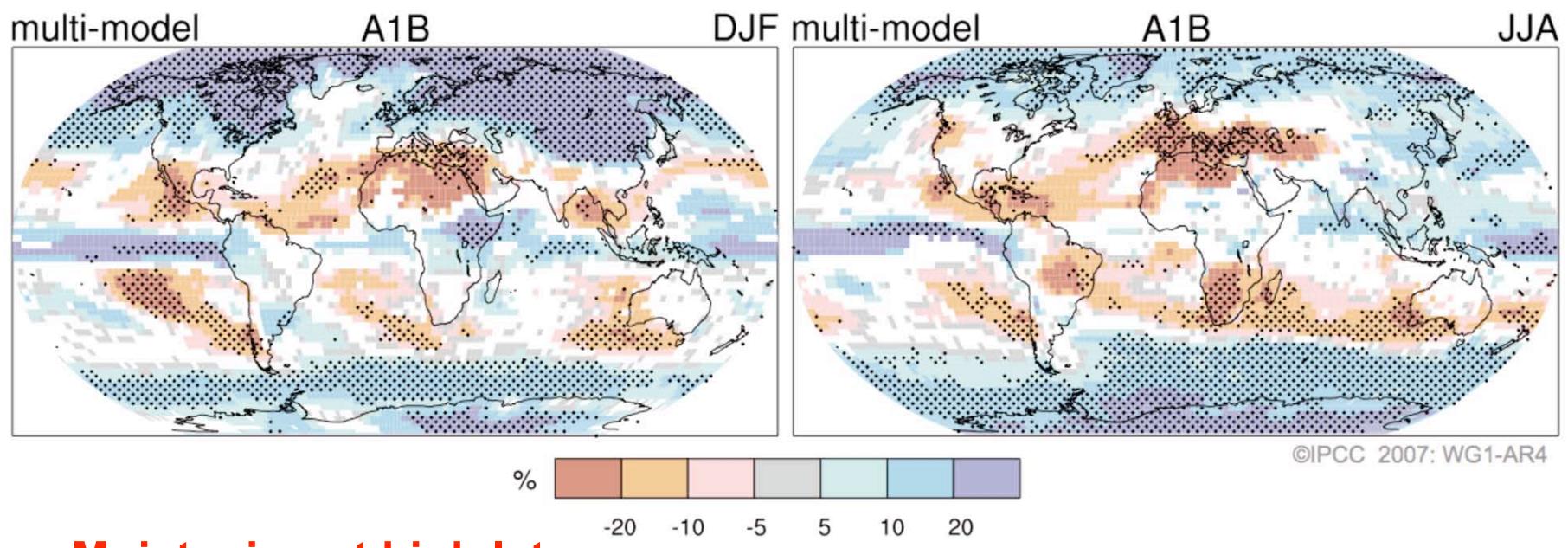
C4MIP: BYOModel

| | Atm | Ocean | | Land C | Ocn C |
|-------------|-------------------------------|-------------------------------------|--|--|-------------------------------|
| Hadley | HADCM3 2.5°x3.75°, L19 | 2.5°x3.75°, L20 flux-adjusted | | MOSES/TRIFFID – DynVeg, stomates , GCM soil moisture, 1 soil pool | HadOCC: NPZD, DIC, TALK |
| IPSL-CM2 | LMD5 64x50, L19 | OPA, no flux-adj | | LUE*APAR, 4 soil pool, 1 water bucket | OCMIP' |
| LLNL | PCM, 2.8°x2.8°, L18 | POP 0.6 ° x0.6°, L40 | | DynVeg, IBIS- CENTURY | OCMIP' |
| NCAR-CSM1.4 | CSM1, T31 L18 | NCOM3.6x3 | | Stomates, 9 soil pools, LSM 6-layer water | OCMIP'+Fe patch |
| MPI | ECHAM | MPI-OM | | JSBACH | HAMOCC5 |
| FRCGC | CCSR/NIES/ FRCGC T42L20 | COCO; No flux-adj, (0.5-1.4)x1.4 | | Sim-CYCLE | NPZD |
| UVIC | 1-layer Energy Balance | MOM-2.2 | | TRIFFID DynVeg, stomates, 1 soil bucket | DIC-abiotic |
| UMD | QTCM | Mixed layer-Qflux | | VEGAS DynVeg, 3 soil pools | OCMIP-abiotic |
| CLIMBER | 2.5D stat-dynam 10°x51° | X-avg,2.5° lat, 3 basins | | LPJ | NPZD |
| Bern-CC | EBM 2.5°x3.75° | HILDA box-diffusion model | | LPJ | perturbation |

IPCC WG1-AR4: regional differences in precipitation change

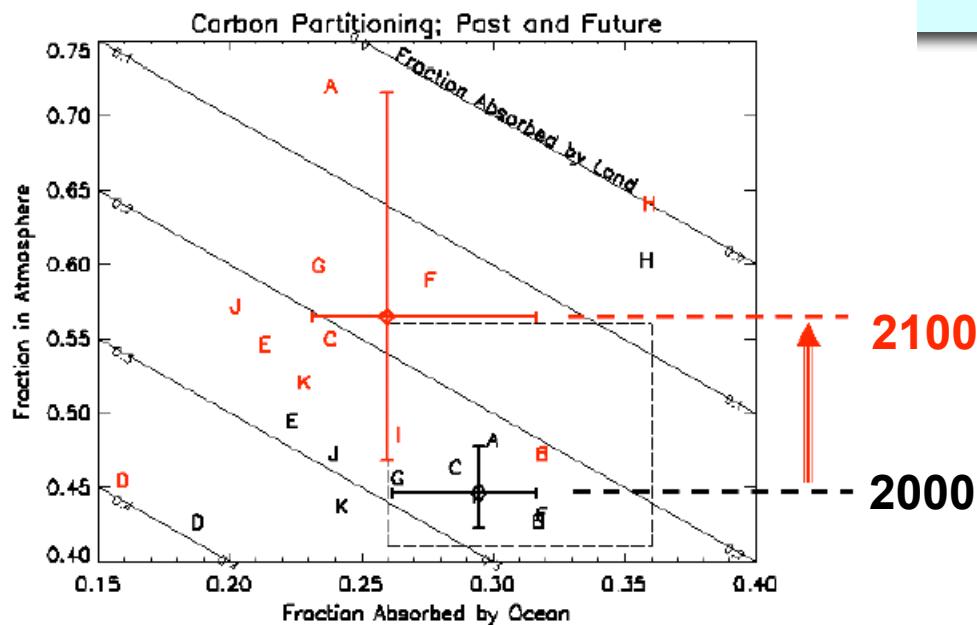
White area: < 66% of models agree in the sign of change
Stippled area: > 90% of models agree in the sign of change

Projected Patterns of Precipitation Changes



- **Moistening at high lat,**
- **??? N American summer,**
- **??? In tropics**

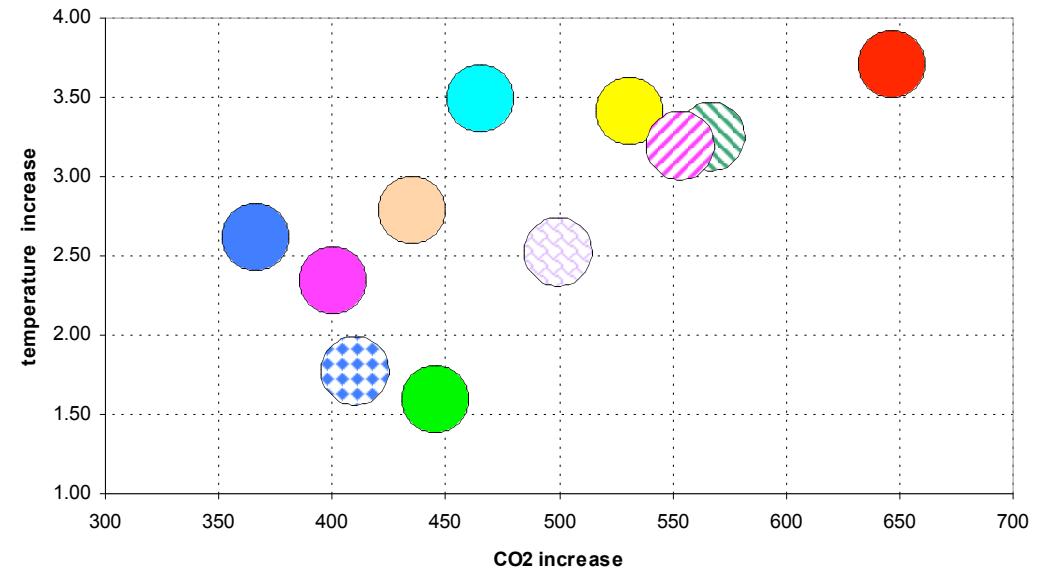
Summary: C⁴MIP Robust Result



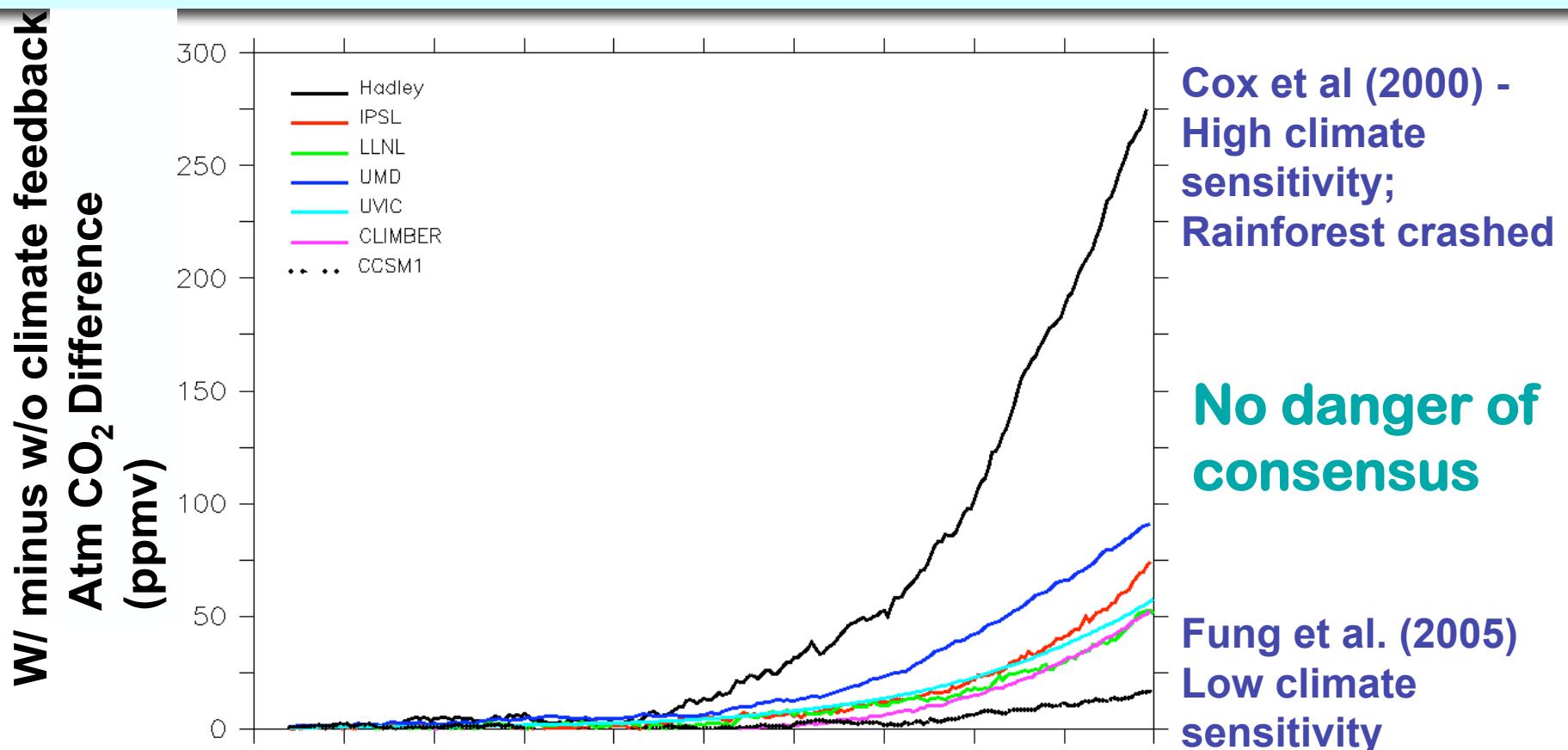
IPCC AR4 WG1 CH7

Coupled Carbon-climate
models: 2100 minus 2000:
350-650 ppmv increase in
CO₂

Carbon-climate
feedback
accelerates
warming



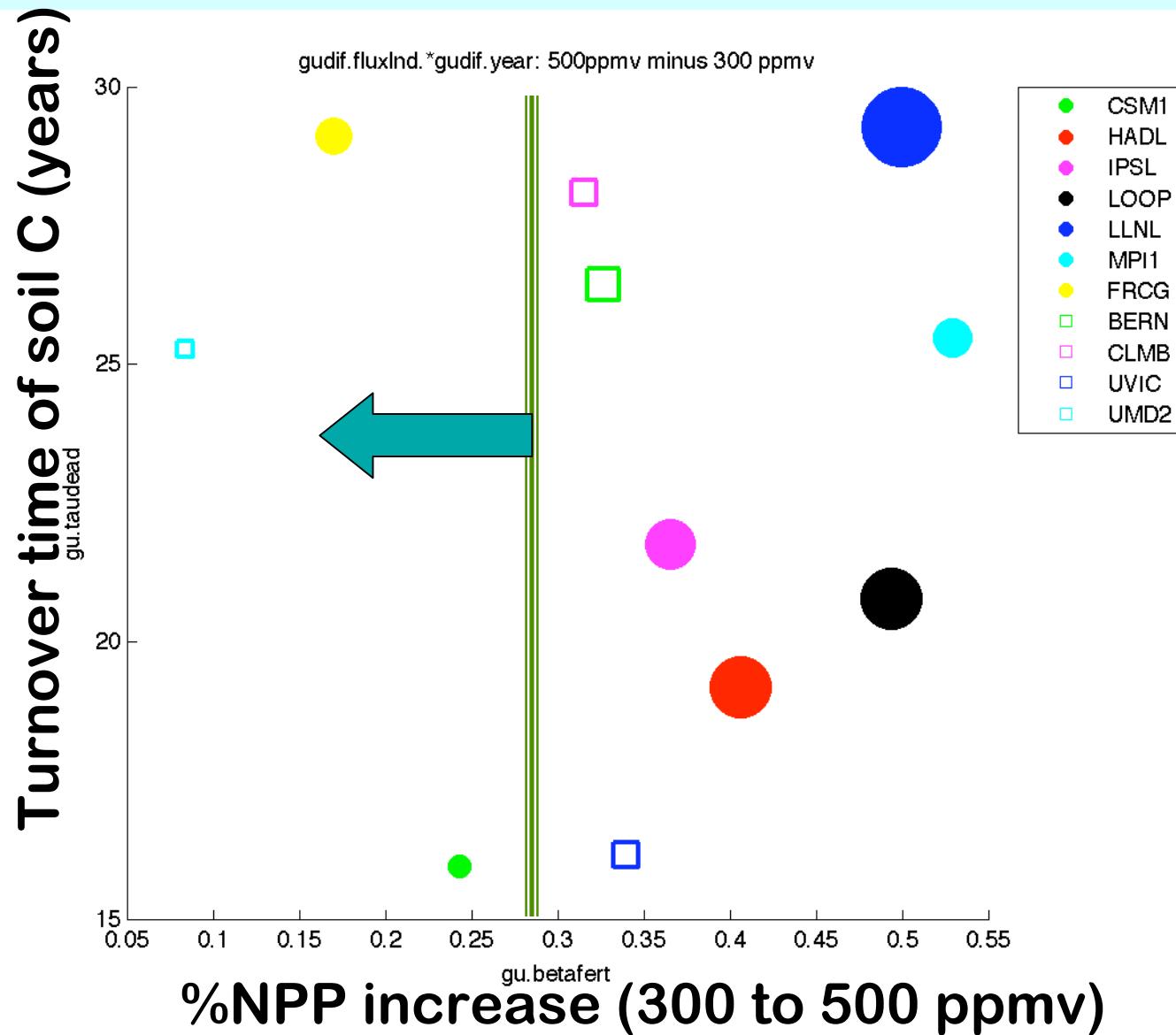
WCRP-IGBP Coupled Carbon Cycle Climate (C⁴MIP): FF=historical + SRES A2, BYOM



All models: with C-climate feedback → higher CO₂, warmer

Friedlingstein et al. (J Clim, 2006)

Control Climate: Land Uptake (300-500 ppmv) as a function of land parameters



Challenge

- C4MIP - ecosystem function dependent on plant-soil moisture coupling -
 - soil moisture(z);
 - how do plants access soil moisture(z) (deep roots, tap roots...) [Lee et al. PNAS 2005]
- Resource limitation (N, P,...) on CO₂ fertilization
- No clue about time constants for ecosystem dieback and recovery [Moorecroft et al. - ED]
- No clue about soil microbe population in new climate + biogeochem space [Treseder]
- No clue about ecosystem function in a new climate + biogeochem space