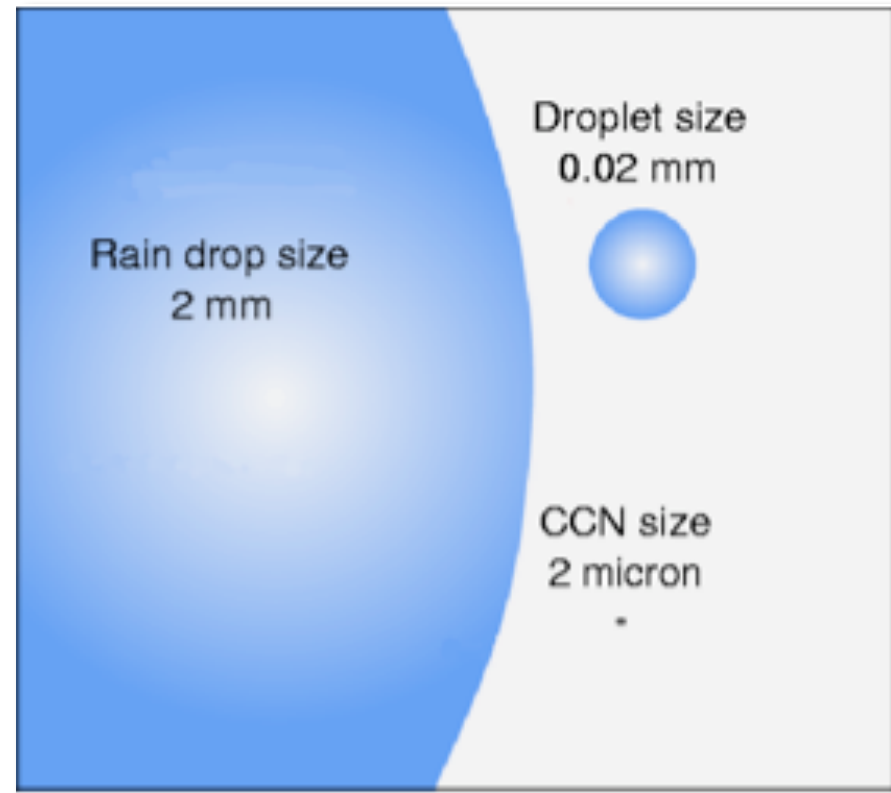


# Condensation of cloud microdroplets in homogeneous isotropic turbulence

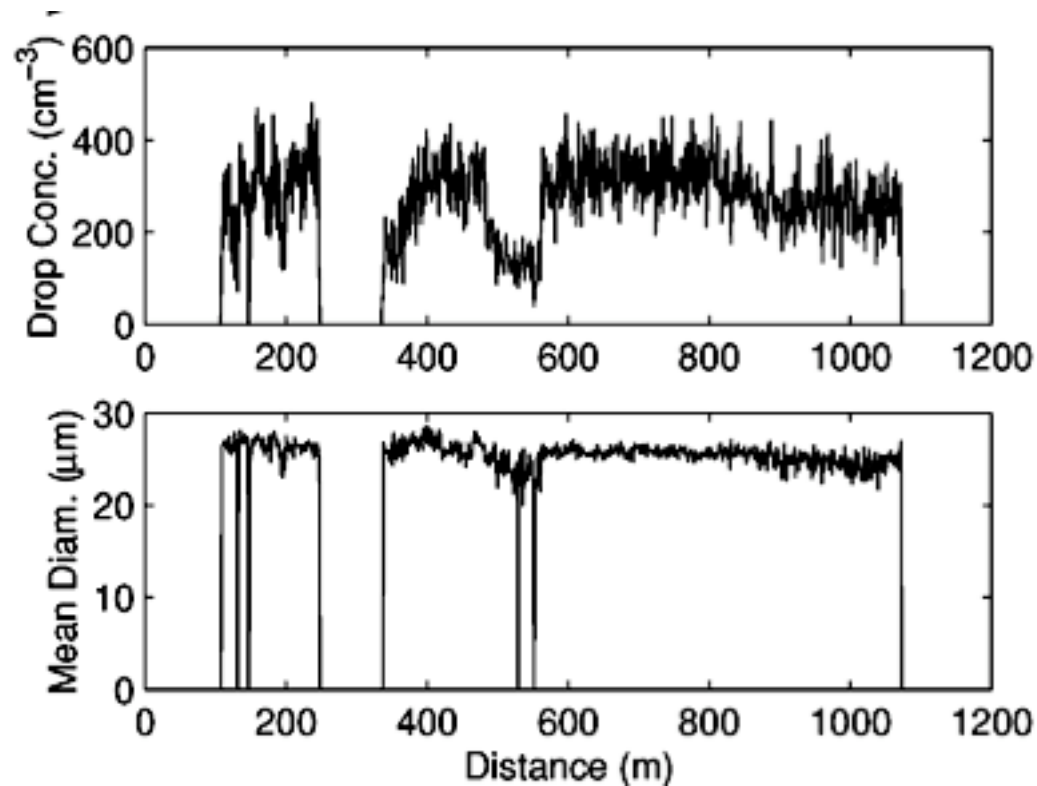
*Alessandra Lanotte*  
*CNR ISAC, Rome, Italy*



*Informal KITP talk*

# A traverse through a warm cloud

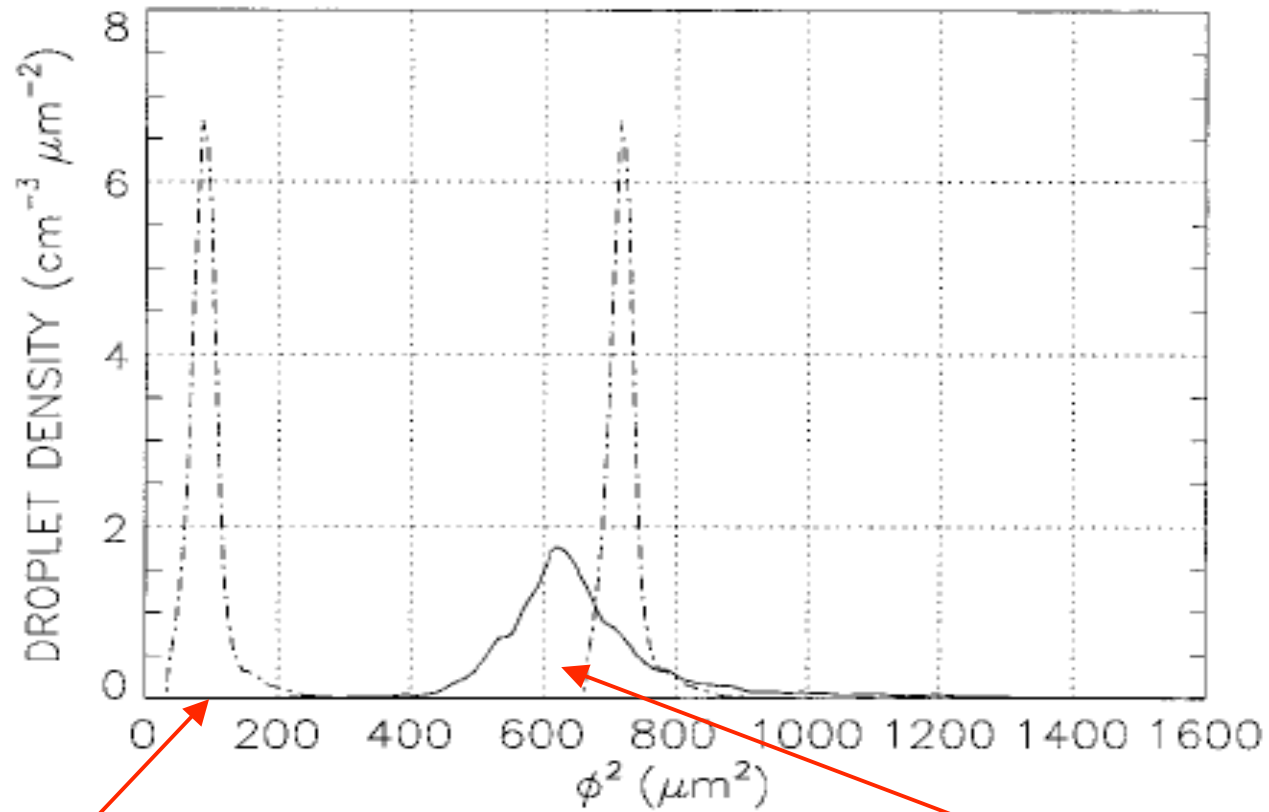
From R. Shaw, Ann. Rev. Fluid Mech 2003



- Drop concentration is not at all homogeneous in space.
- Fluctuations of mean radius down to zero means that dry and cloudy air are segregated during mixing.

# Droplet size spectrum

From JL. Brenguier & L. Chaumat, JAS 2001

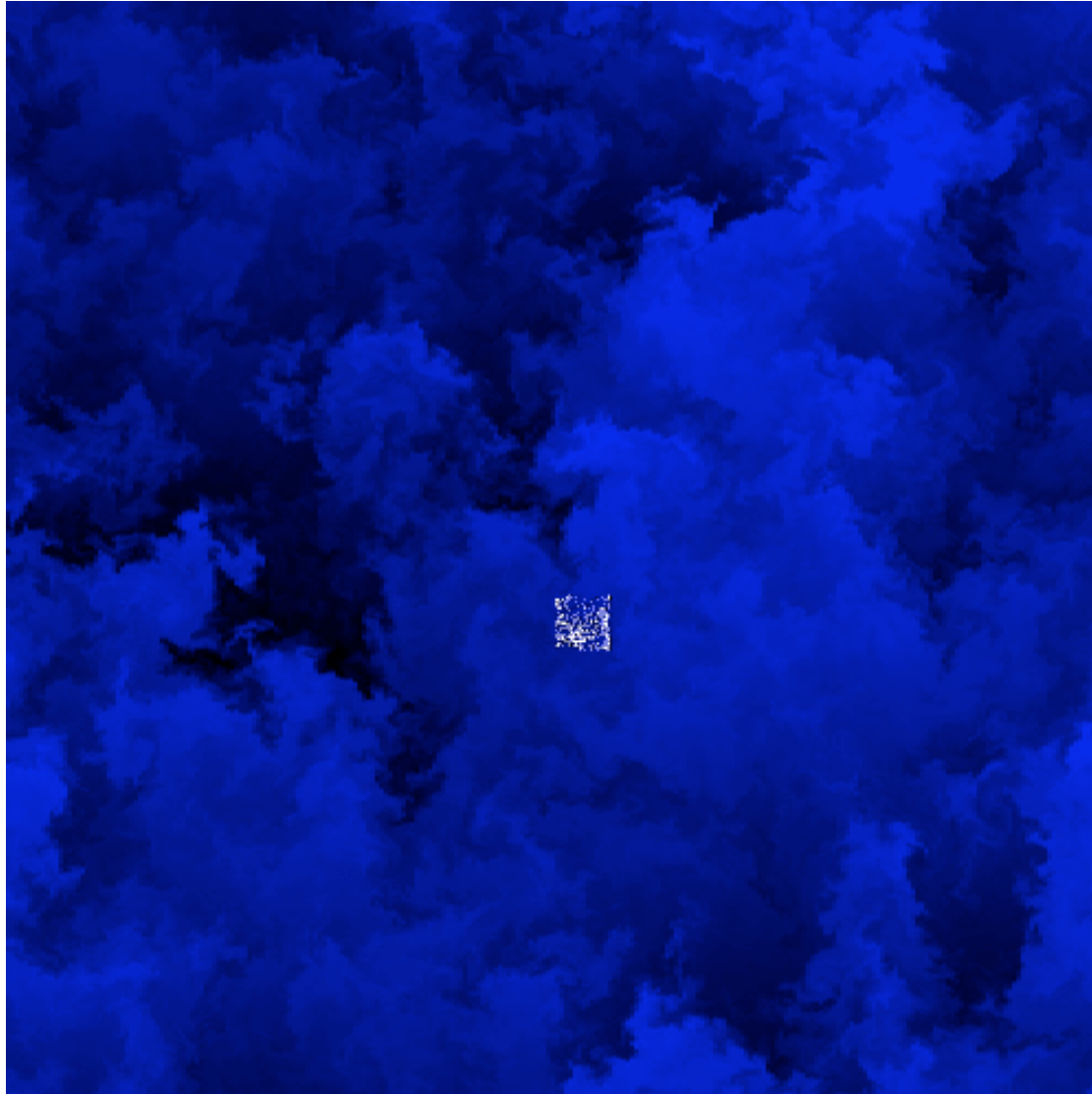


Measured at cloud base h=700m

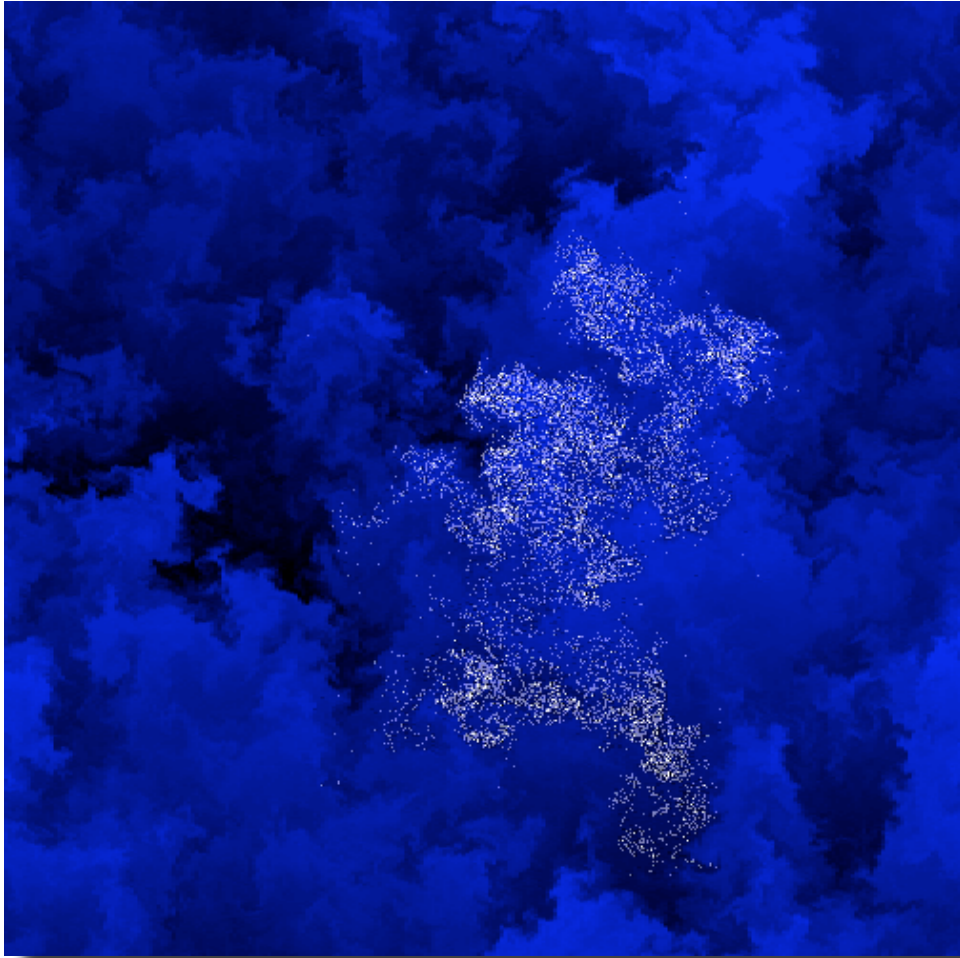
Measured at h=2200m

average LWC= 2.5 gr/m<sup>3</sup>; average #drops= 250 drops/cm<sup>3</sup>

# Parcel Model



# Beyond air parcel model



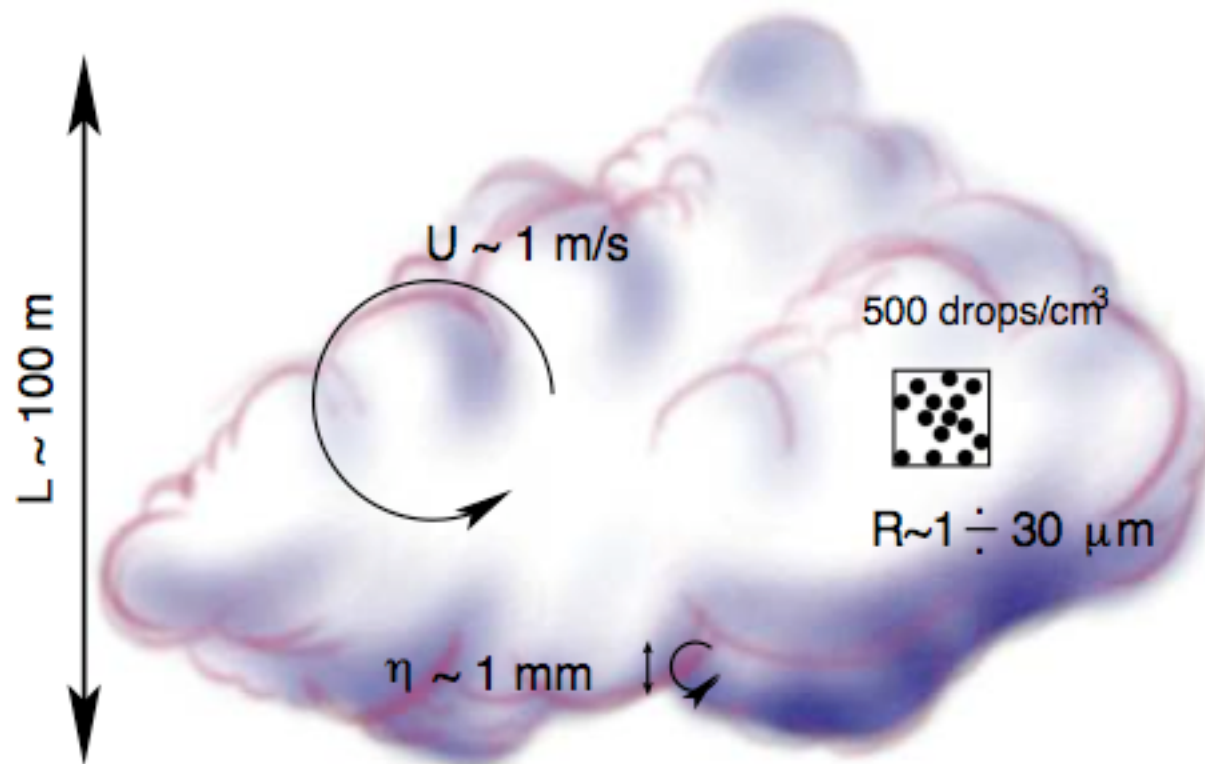
Role of flow & vapour  
turbulence  
for droplet spectrum  
broadening  
during condensation

$$\langle (\mathbf{X}_1(t) - \mathbf{X}_2(t))^2 \rangle \simeq \epsilon t^3$$

$$\rightarrow \langle r^2 \rangle^{1/2} \sim (3 - 100)m, t = (10 - 100)s, \epsilon = 10^{-2} m^2/s^3$$

# Warm non-precipitating cloud

$\tau_\eta = 0.1$  s;  $\varepsilon = 10^{-3} - 10^{-2} \text{m}^2/\text{s}^3$ ;  $\text{Re} = 10^7$ ; concentration 150-500 drops/cm<sup>3</sup>



**Note:** Typical Stokes numbers:  $\text{St} \approx 0.001 \div 0.9$  for  $R \approx 1 \div 30$   $\mu\text{m}$

# Generalized Twomey's model equations

(Twomey, Geofis. Pura Appl. 43, 1959)

$$s(\mathbf{x}, t) = e/e_s(\mathbf{x}, t) - 1$$

- $\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\nabla p + \nu \Delta \mathbf{u} + f$

$$\partial_t s + \mathbf{u} \cdot \nabla s = A_1 u_z - \frac{s}{\tau_s} + \kappa \Delta s$$

$\tau_s^{-1} \sim n \langle R \rangle$

supersaturation absorption time

$$\frac{dX}{dt} = V(t)$$

- $\frac{dV}{dt} = -\frac{V(t) - u(X(t), t)}{\tau_d} + g$

$\tau_d \sim R^2 / (\nu \beta)$

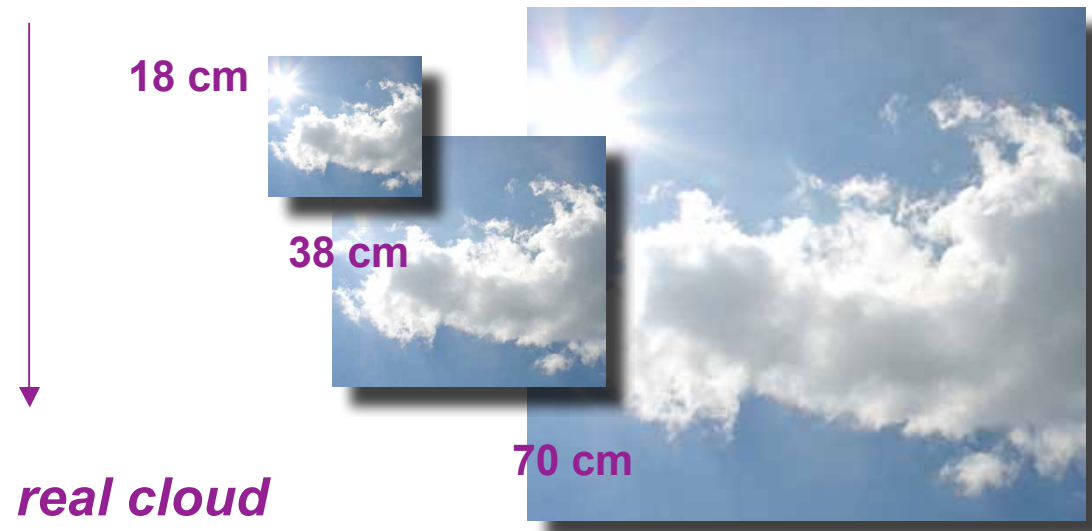
particle Stokes time

$$\frac{dR^2}{dt} = A_3 s(X(t), t)$$

# DNS for droplet condensation in turbulent flows

# = 130 drops/cm<sup>3</sup>,  $\varepsilon = 0.001 \text{ m}^2/\text{s}^3$ ;  $v_{\text{rms}}=0.1 \text{ m/s}$ ;  $\eta = 1\text{mm}$   
 LWC= 0.07; 1.2 g/m<sup>3</sup>;  $R_0 = 5$ ;  $13 \mu\text{m}$ ;  $St = 0.005$ ;  $0.035$

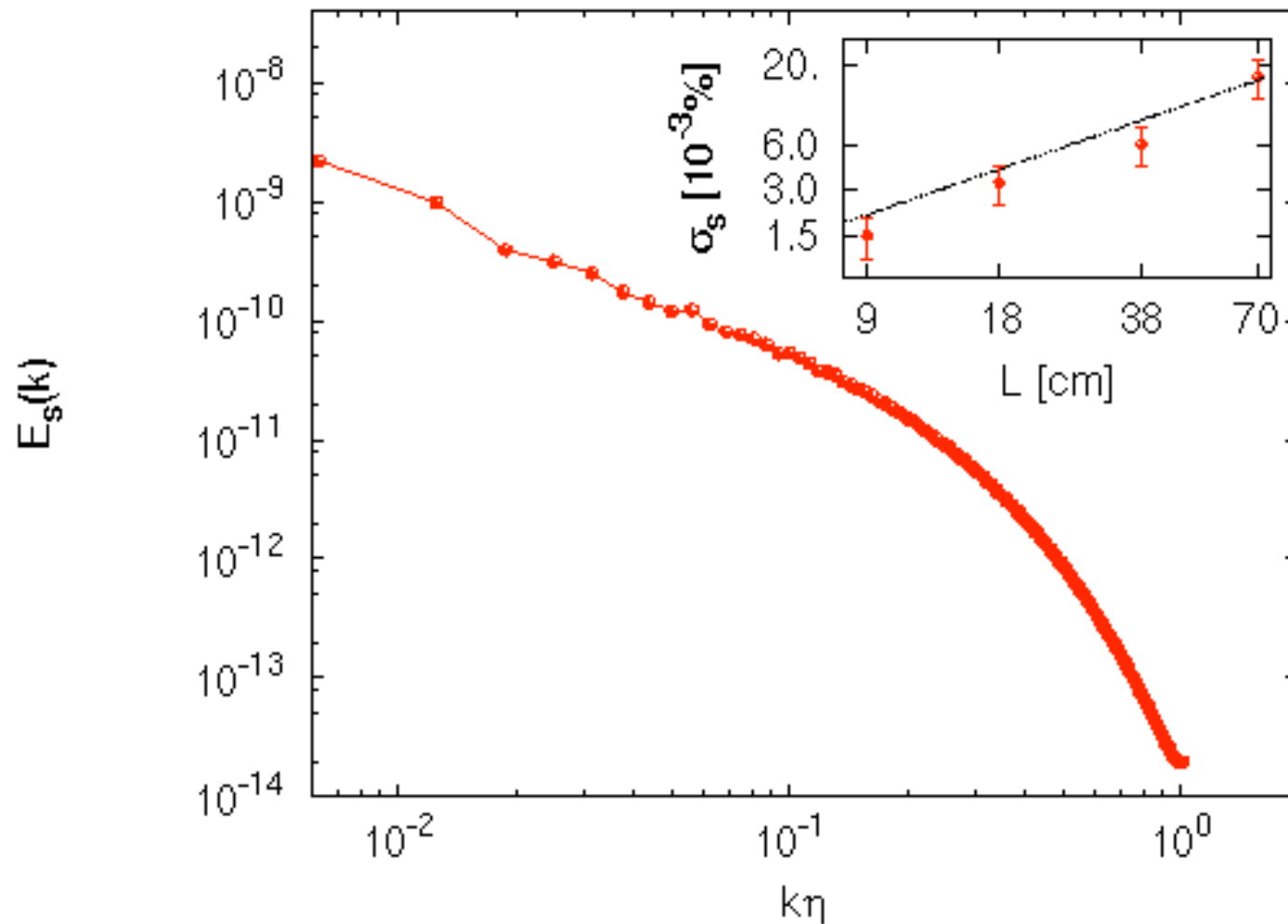
| $N^3$   | L     | $Re_\lambda$ |
|---------|-------|--------------|
| $64^3$  | 0.09m | 40           |
| $128^3$ | 0.18m | 65           |
| $256^3$ | 0.38m | 105          |
| $512^3$ | 0.70m | 185          |
| ....    | ....  | ....         |
| ...     | 100m  | $\sim 10^3$  |



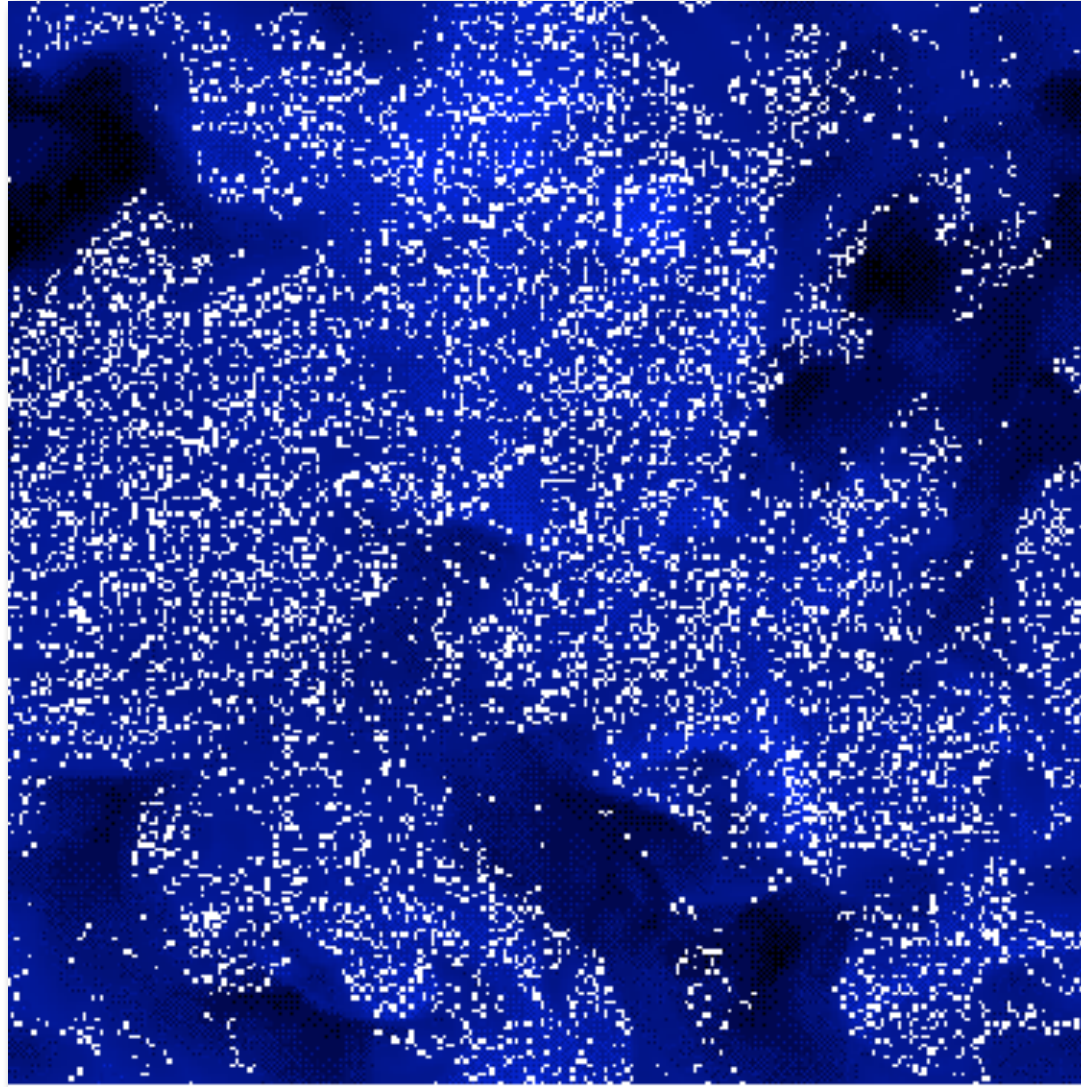
$L=70 \text{ cm} \rightarrow 32 \text{ Millions of droplets starting from a delta-function distribution in size}$



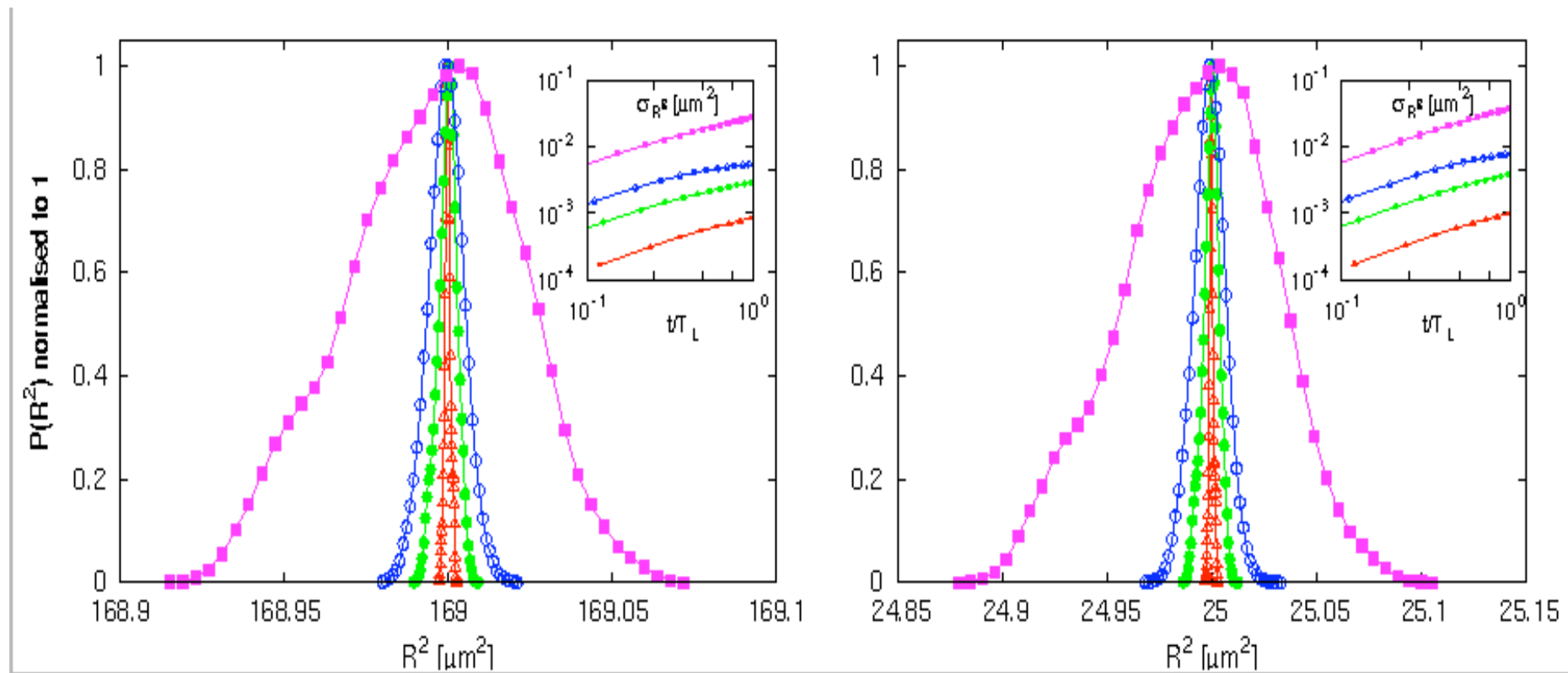
# Supersaturation fluctuations



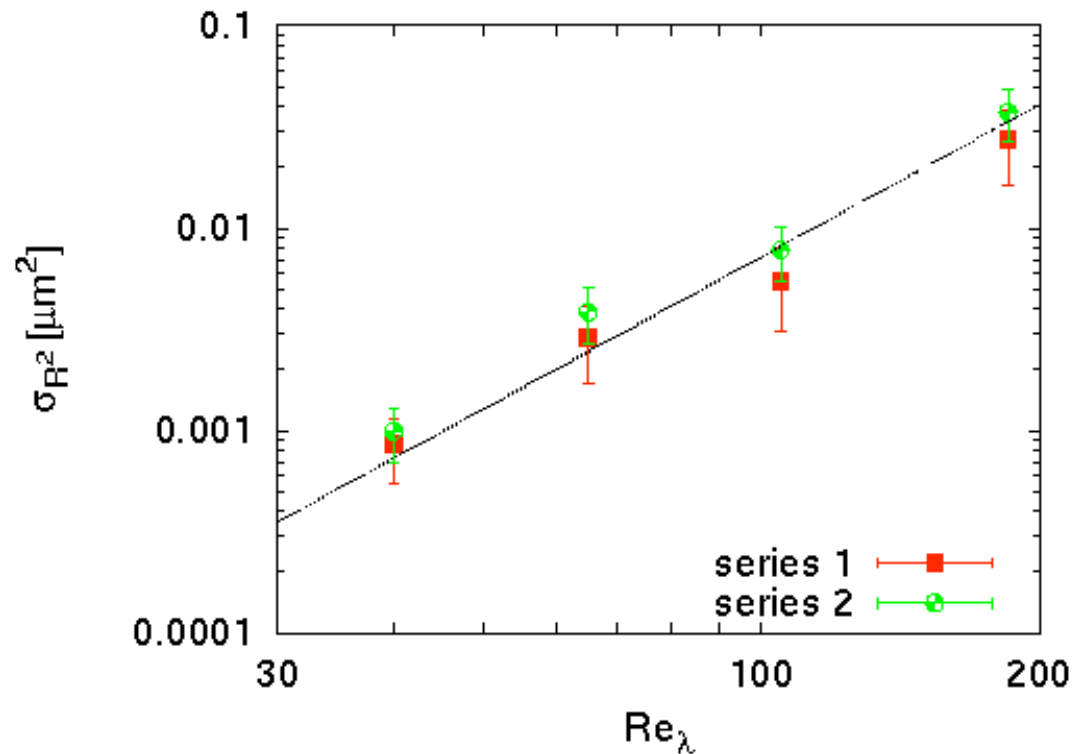
# Instantaneous drops distribution in space



# Square radius spectrum



# Spreading vs Reynolds number



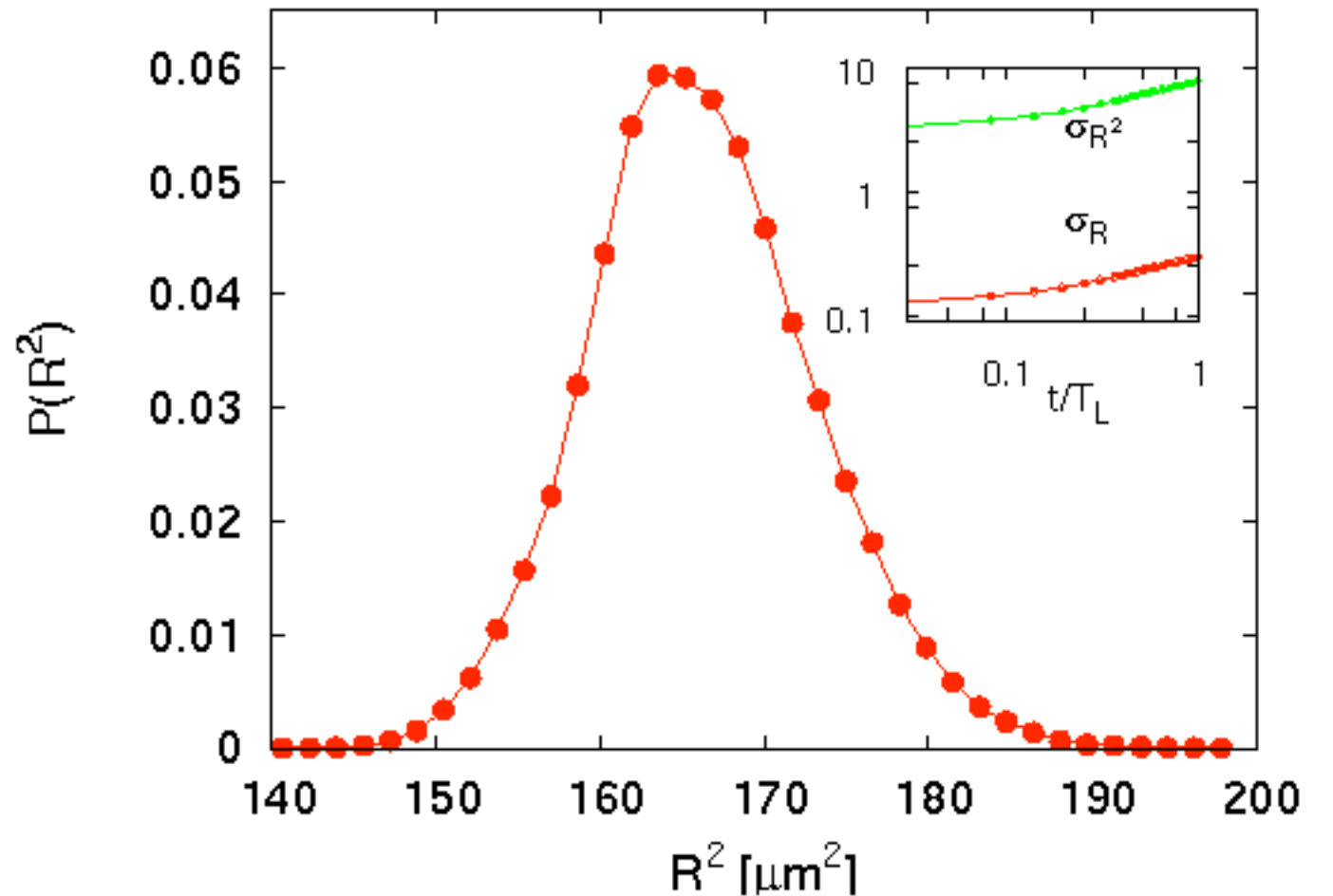
From dimensional argument:

$$\sigma (R^2) \sim c_1 \text{Re}_\lambda^{5/2}$$

---> With droplets feedback on supersaturation field,  
the dimensional prediction becomes  $\sigma (R^2) \sim c_2 \text{Re}_\lambda^{3/2}$

E.g. for physical parameters ( $L=100\text{m}$ ;  $\text{Re}_\lambda \sim 7000$ ):  $\sigma (R^2) \sim (4 - 8) \mu\text{m}^2$

# Evidence from a large scale simulation



with

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CNR-IAC, Rome, Italy

Preprint at:

<http://climate08.wikispaces.com/References> --> Clouds section