

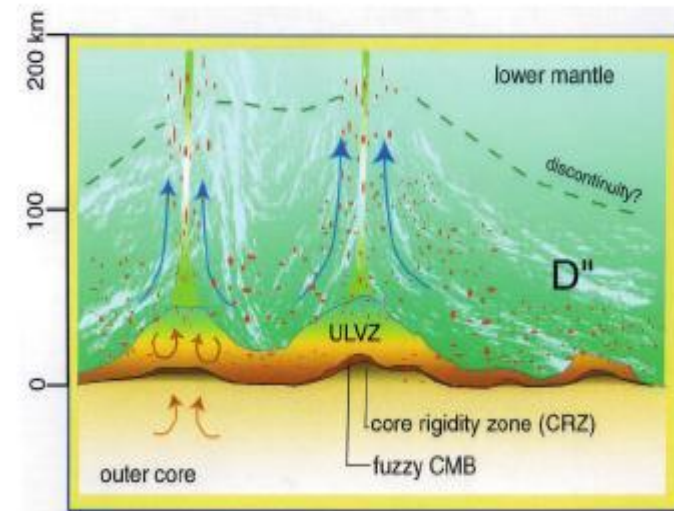
Melting at the Core-Mantle Boundary and the Topography of the Ultralow Velocity Zone

Sash Hier-Majumder

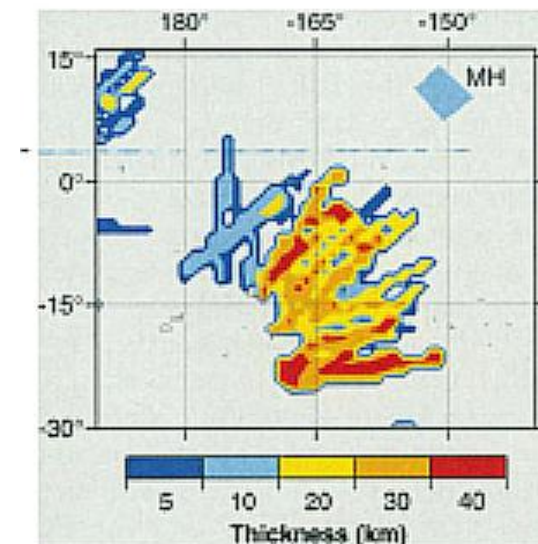
CIDER workshop, June, 2008

Ultralow velocity zone at the core-mantle boundary

- Thickness varies between 5-40 km
- Estimated melt volume fraction between 5-30%
- Logarithmic ratio of seismic velocities vary between 2.5-3.2



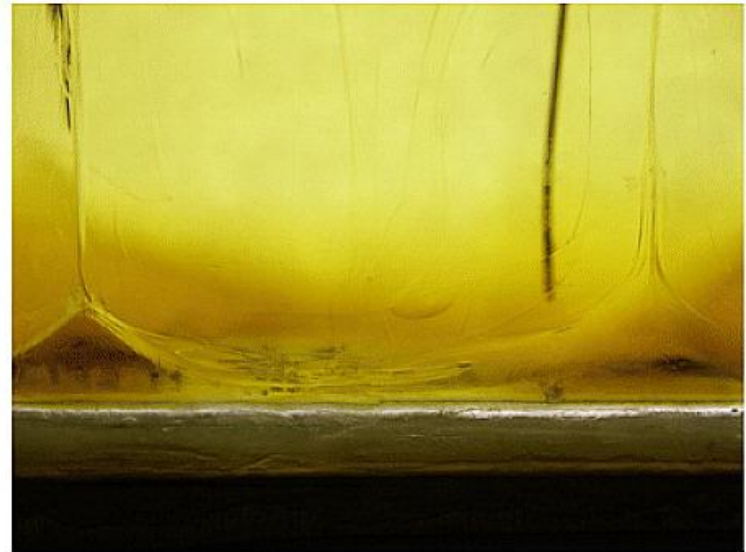
Garnero [2004]



Williams and Garnero [1996]

Structure of the basal layer

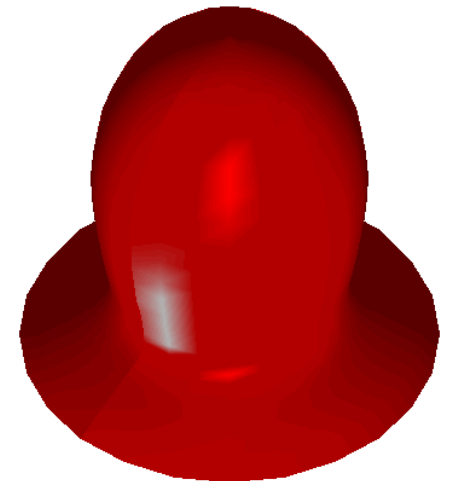
- Lateral variation in topography
- Lower viscosity compared to the mantle
- Low seismic velocity



Jellinek and Manga [2004]

Outline

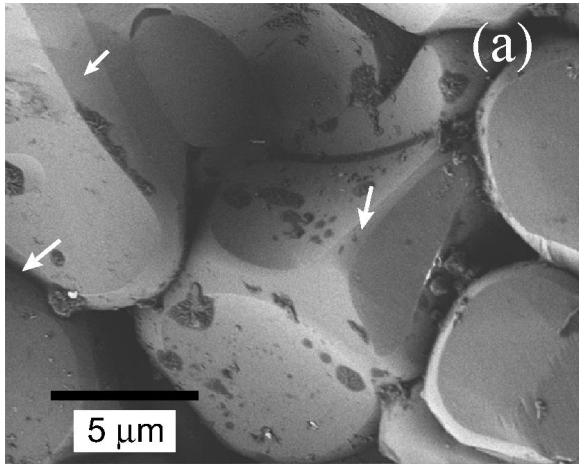
- A new, microstructure based, model for seismic velocities
- Estimated melt volume fraction at the ULVZ
- Influence of surface tension on the topography of the ULVZ
- Influence of viscosity contrast on the topography of the ULVZ



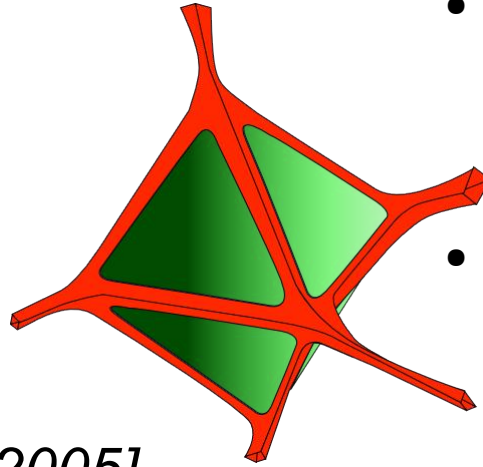
Microstructure modeling

- Elastic wave velocities in two-phase media is controlled by the area of contact or contiguity between adjacent particles [*Takei, 1998, 2000*].
- Steady-state contiguity is achieved by viscous sintering of the multiphase aggregate

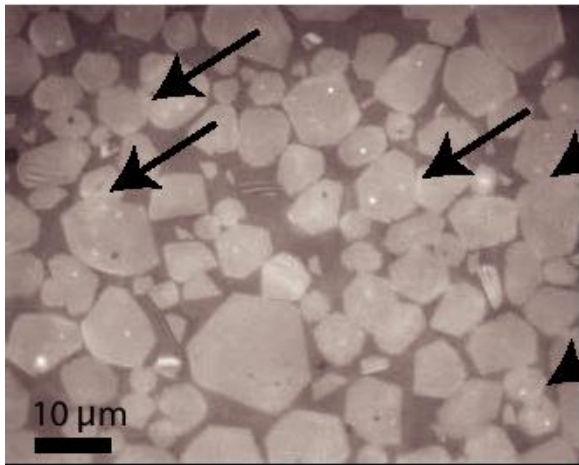
Disaggregation



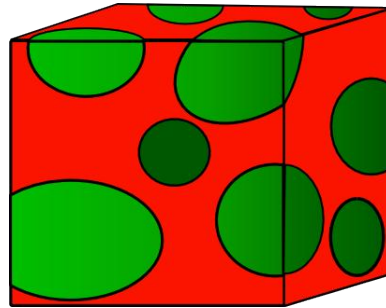
Hier-Majumder *et al.* [2005]



- Grains are contiguous at melt fractions below the disaggregation melt fraction
- At higher melt fractions, grains are suspended in the melt



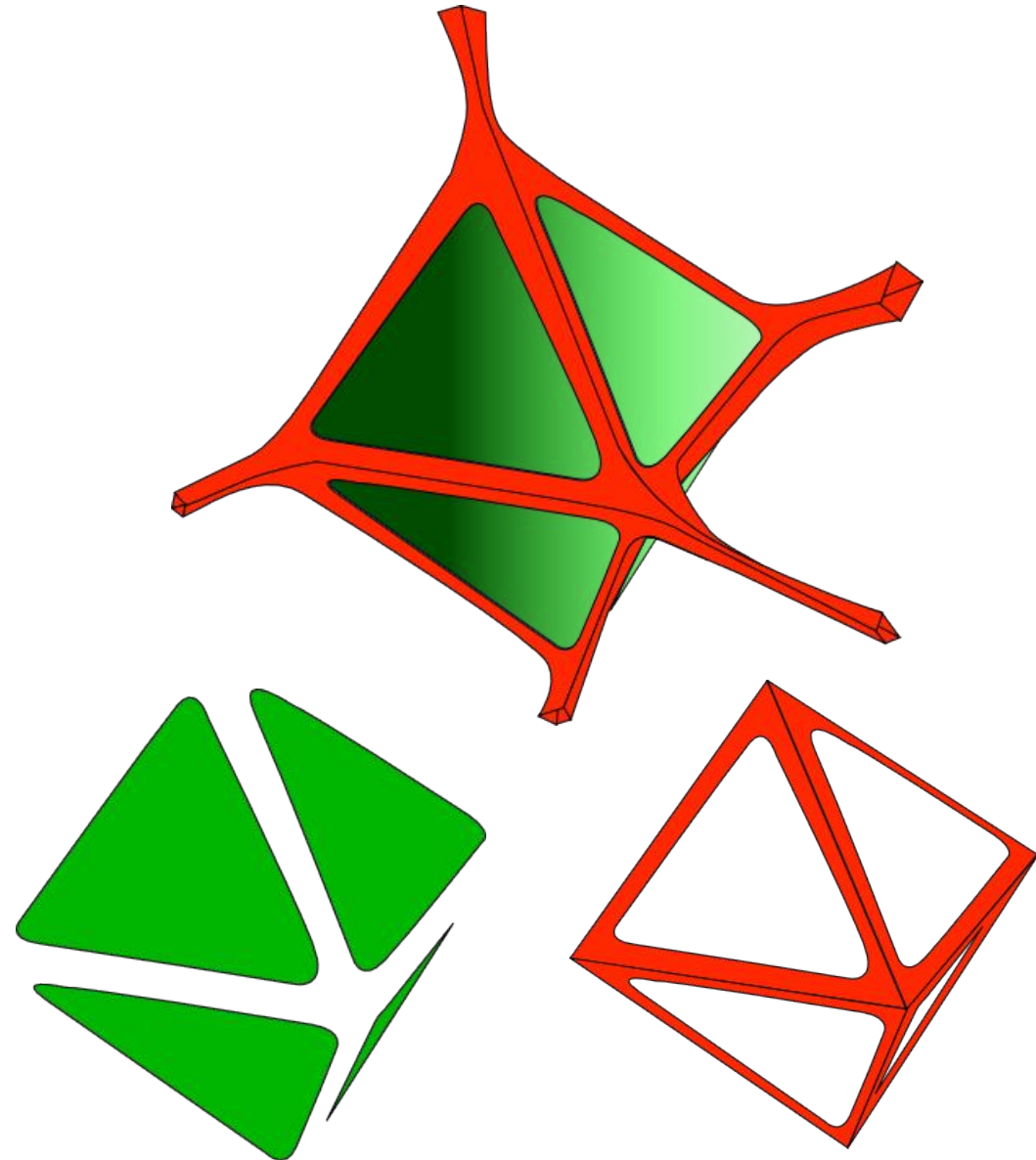
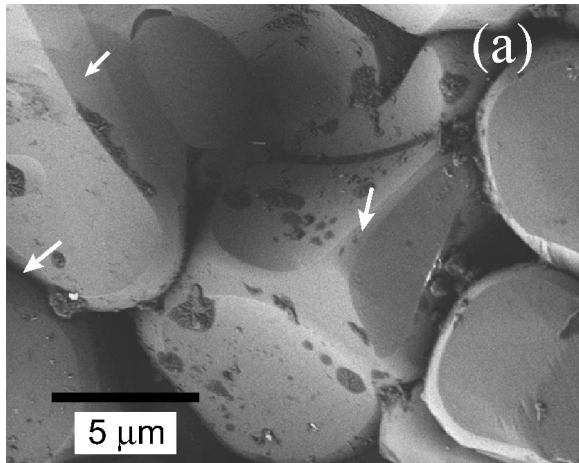
Scott and Kohlstedt [2006]



Interfaces in rocks

$$\chi = A_{gg}\gamma_{gg} + A_{gm}\gamma_{gm}$$

$$\psi = \frac{A_{gg}}{A_{gg} + A_{gm}}$$



Microstructure

Within the grains and the melt

$$0 = \nabla \cdot \mathbf{u}$$

$$0 = \nabla \cdot \mathbf{T}$$

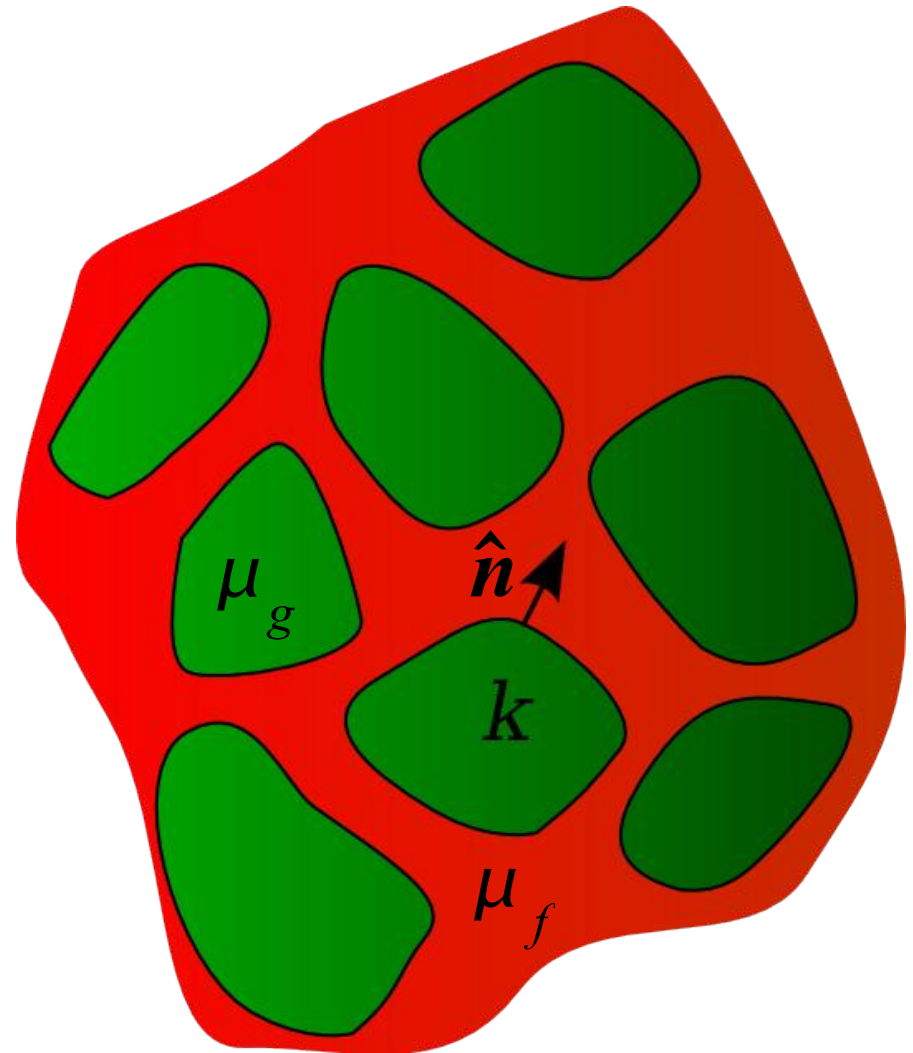
Boundary condition

$$\Delta \mathbf{T}^k \cdot \mathbf{n}^k = \gamma^k \left(\nabla \cdot \mathbf{n}^k \right) \mathbf{n}^k - \tilde{\nabla} \gamma^k$$

$$0 = \Delta \mathbf{u}^k$$

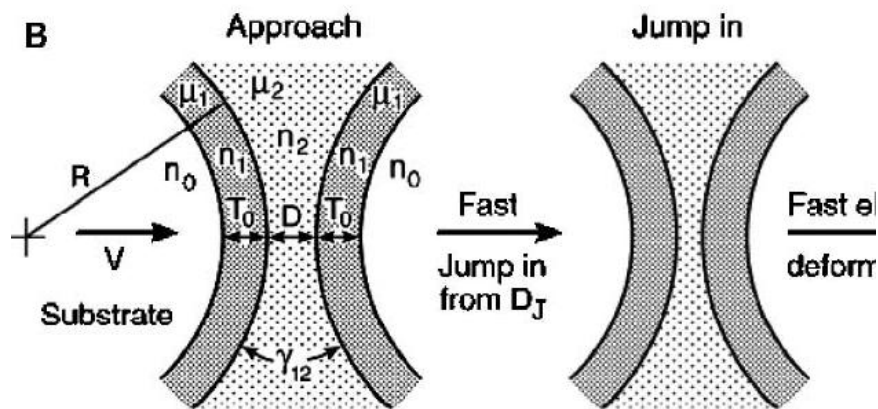
Kinematic relation

$$0 = \frac{\partial F^k}{\partial t} + \mathbf{u}^k \cdot \nabla F^k$$

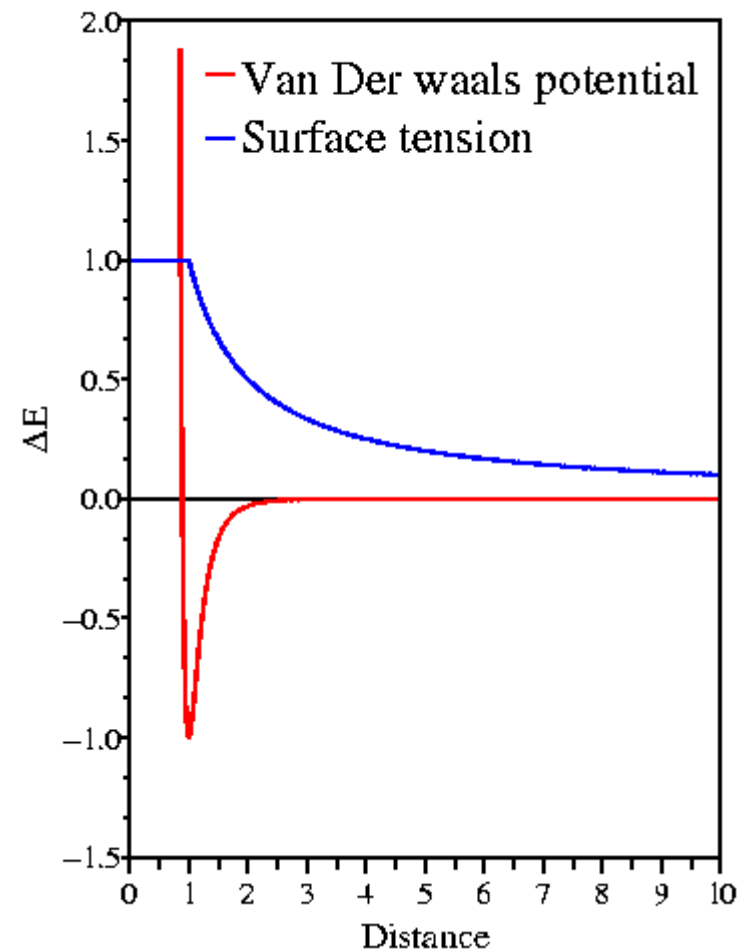


Midrange forces

Interaction between grains
gives rise to variation in
surface tension



Chen et al.[2004]



Steady-state grain shape

$$\begin{aligned}
 \mathbf{u}^g \cdot \mathbf{n} &= 0 \\
 \mathbf{u}^m \cdot \mathbf{n} &= 0 && \text{No-slip boundary condition} \\
 \mathbf{u}^g \cdot \hat{\boldsymbol{\theta}} - \mathbf{u}^m \cdot \hat{\boldsymbol{\theta}} &= 0 \\
 \hat{\boldsymbol{\theta}} \cdot \Delta \mathbf{T} \cdot \mathbf{n} + \hat{\boldsymbol{\theta}} \cdot \tilde{\nabla} \gamma &= 0 && \text{Marangoni condition}
 \end{aligned}$$

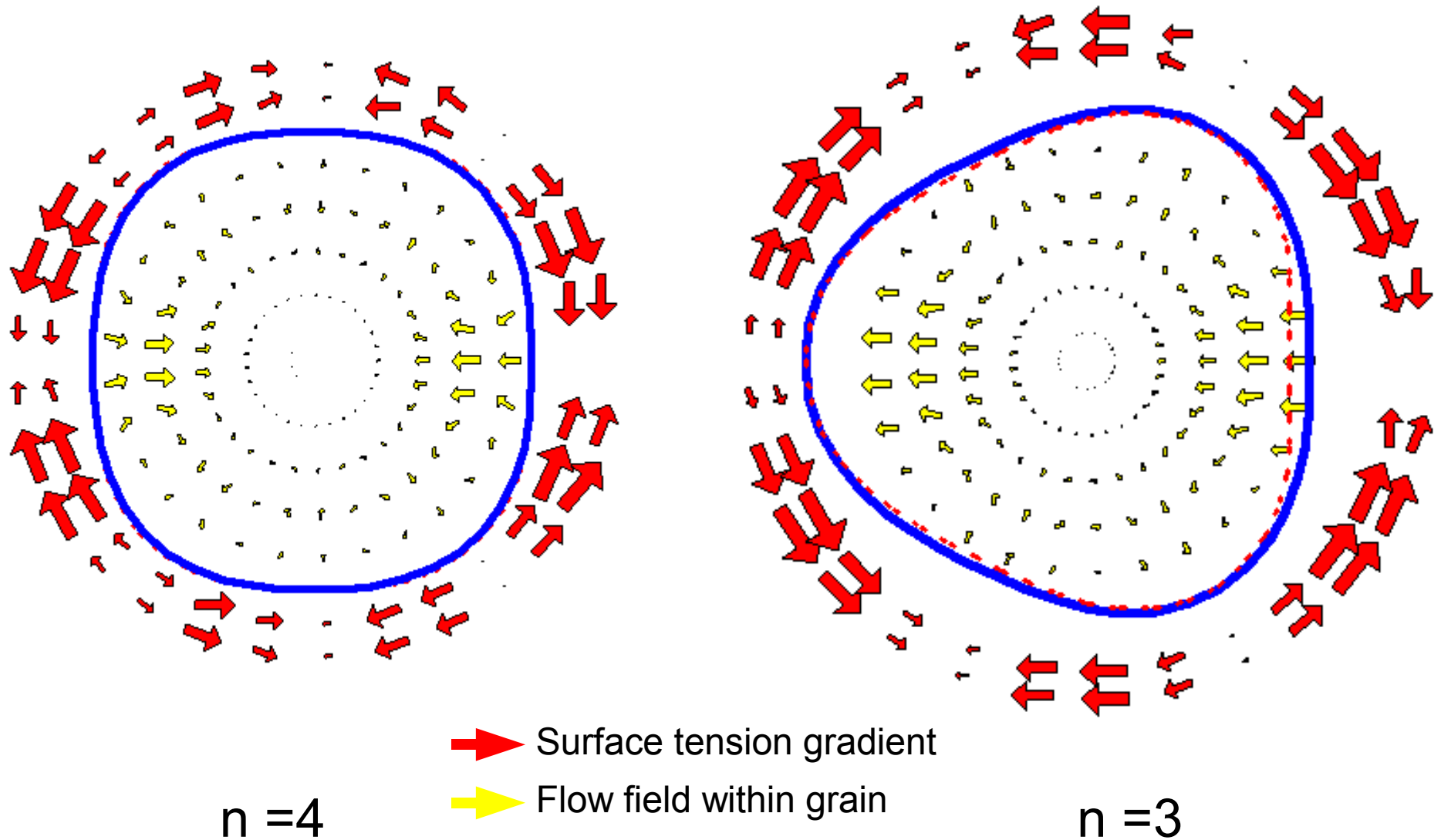
$$\gamma = \gamma_0 + \sum_n \gamma_n P_n(\cos \theta)$$

$$F(r, \theta) = r - a - \epsilon \sum_n f_n(\theta)$$

Steady-state grain shape

$$\frac{d}{d\theta} \left(\sin \theta \frac{d f_n}{d \theta} \right) = P_n(\cos \theta) \sin \theta \left(2a - 3Ca \frac{\gamma_n a^2 [(n+1)\lambda + n]}{(1+\lambda)(2n+1)} \right)$$

Faceting during sintering



Numerical solution

$$\begin{aligned} \mathbf{u}^i(\mathbf{r}_0) &= \frac{2}{1+\lambda} \left[\mathbf{u}^\infty(\mathbf{r}_0) - \frac{1}{4\pi Ca} \sum_k \int_{\Gamma^k} \Delta \mathbf{f}^k(\mathbf{r}) \cdot \mathbf{J}(\mathbf{r}, \mathbf{r}_0) d\Gamma^k \right] \\ &\quad + \frac{(1-\lambda)}{\pi(1+\lambda)} \sum_k \int_{\Gamma^k} \mathbf{n}^k \cdot \mathbf{K}(\mathbf{r}, \mathbf{r}_0) \cdot \mathbf{u}^k(\mathbf{r}) d\Gamma^k \end{aligned}$$

$$0 = \frac{\partial F^k}{\partial t} + \mathbf{u}^k \cdot \nabla F^k$$

$$\mathbf{J}(\mathbf{r}, \mathbf{r}_0) = -I \ln(|\mathbf{r} - \mathbf{r}_0|) + \frac{(\mathbf{r} - \mathbf{r}_0)(\mathbf{r} - \mathbf{r}_0)}{(|\mathbf{r} - \mathbf{r}_0|)^2}$$

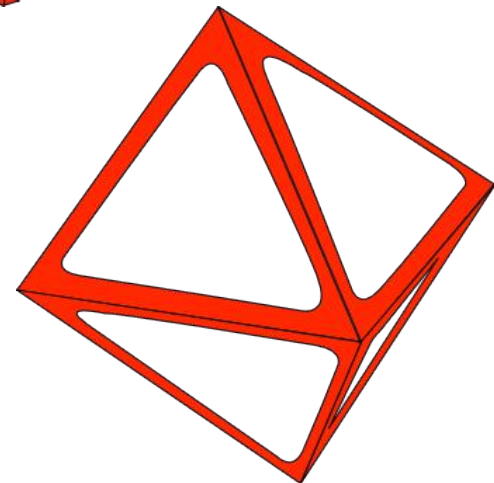
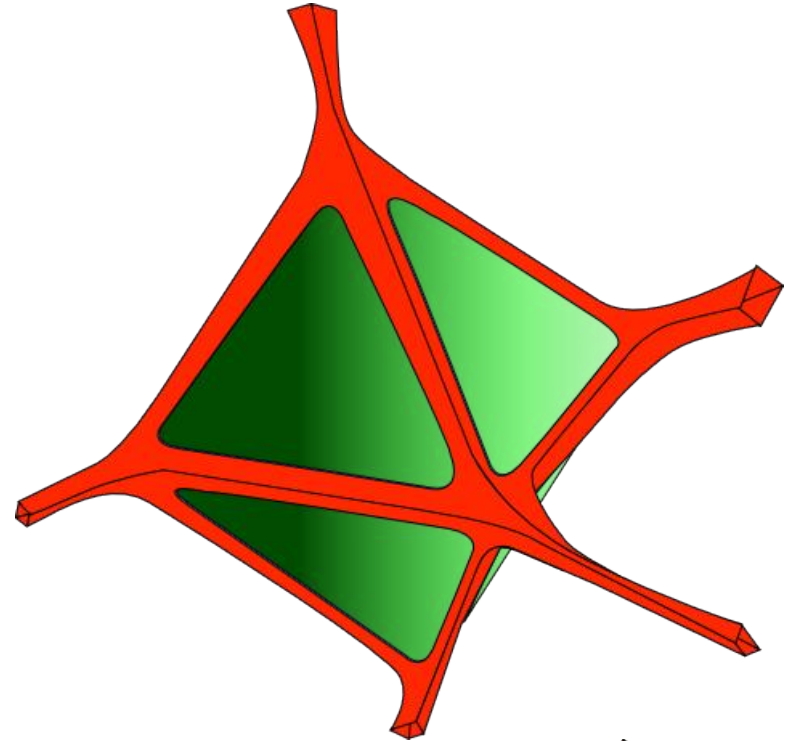
$$\mathbf{K}(\mathbf{r}, \mathbf{r}_0) = -4 \frac{(\mathbf{r} - \mathbf{r}_0)(\mathbf{r} - \mathbf{r}_0)(\mathbf{r} - \mathbf{r}_0)}{(|\mathbf{r} - \mathbf{r}_0|)^4}$$

Nondimensional parameters

$$\lambda = \frac{\mu_g}{\mu_m}$$

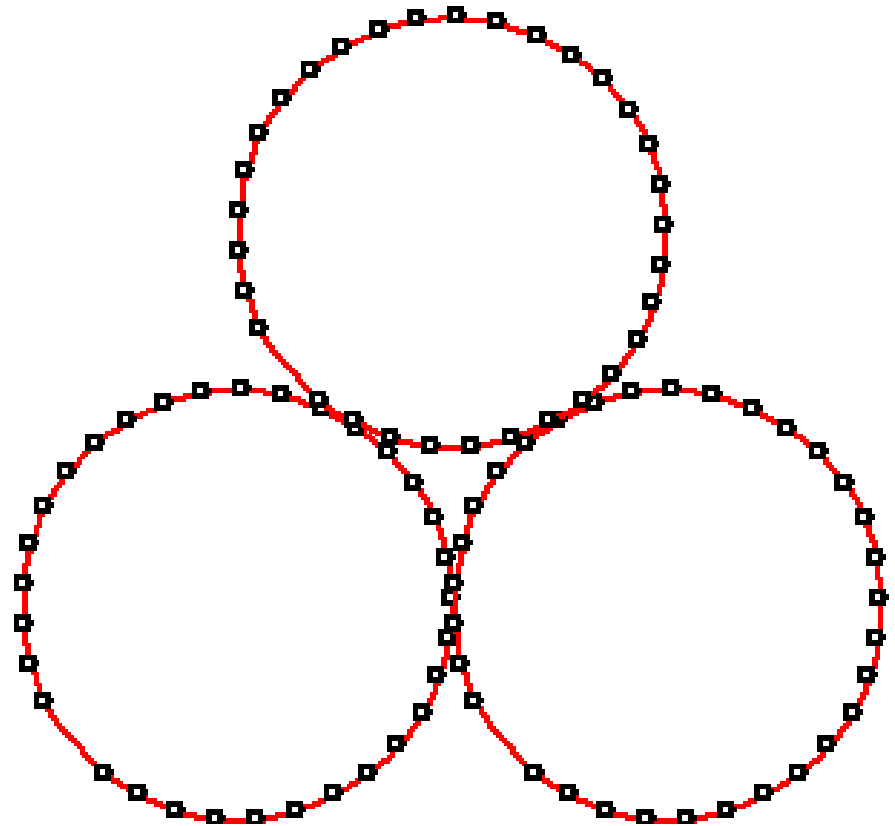
$$Ca = \frac{\mu_m u_0}{\gamma_0}$$

$$\psi = \frac{A_{gg}}{A_{gg} + A_{gm}}$$



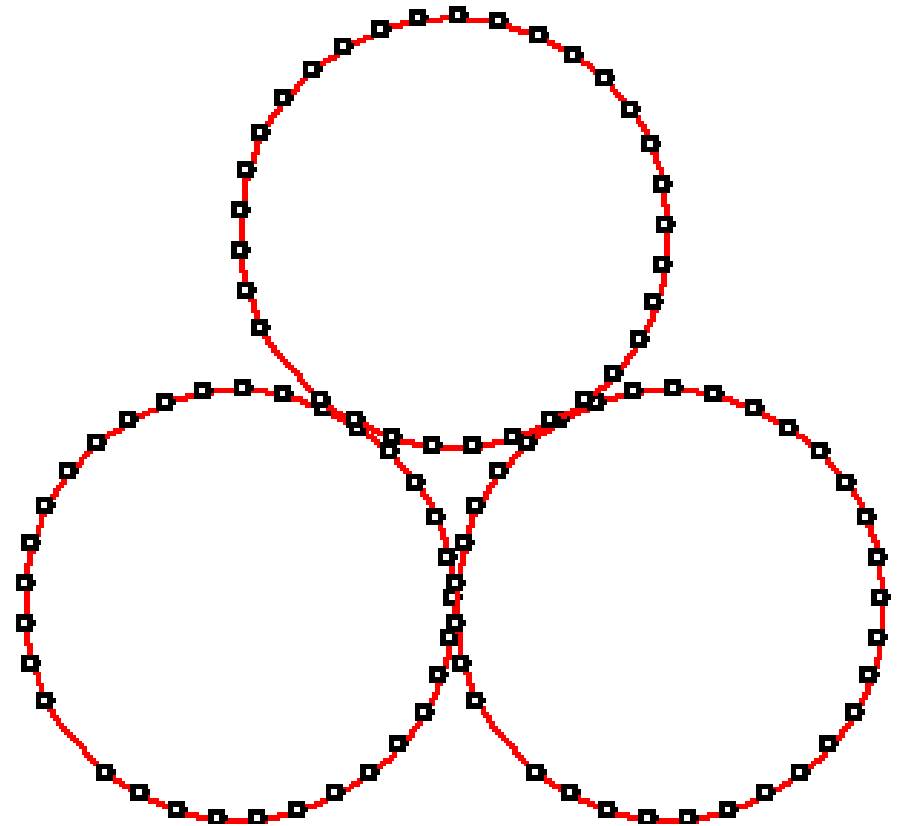
Boundary Elements Model

- Discretizing the BIE
- Evaluating singular integrals
- Updating the position of Lagrangian marker points



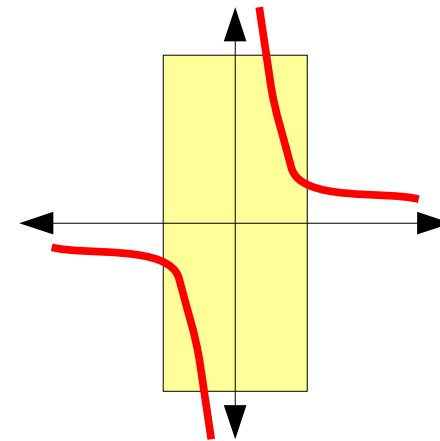
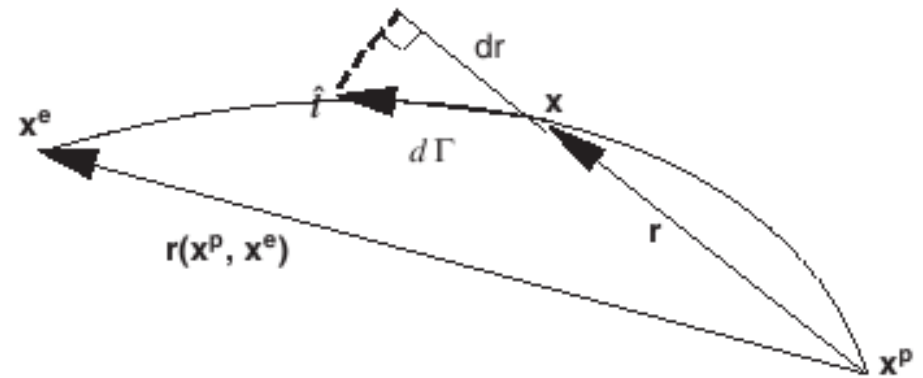
Discretization

- Collocating the poles at nodes
- Cubic spline interpolation of position, allows analytical evaluation of tangent, normal, and curvature
- Linear interpolation of the velocity field



Integral evaluation

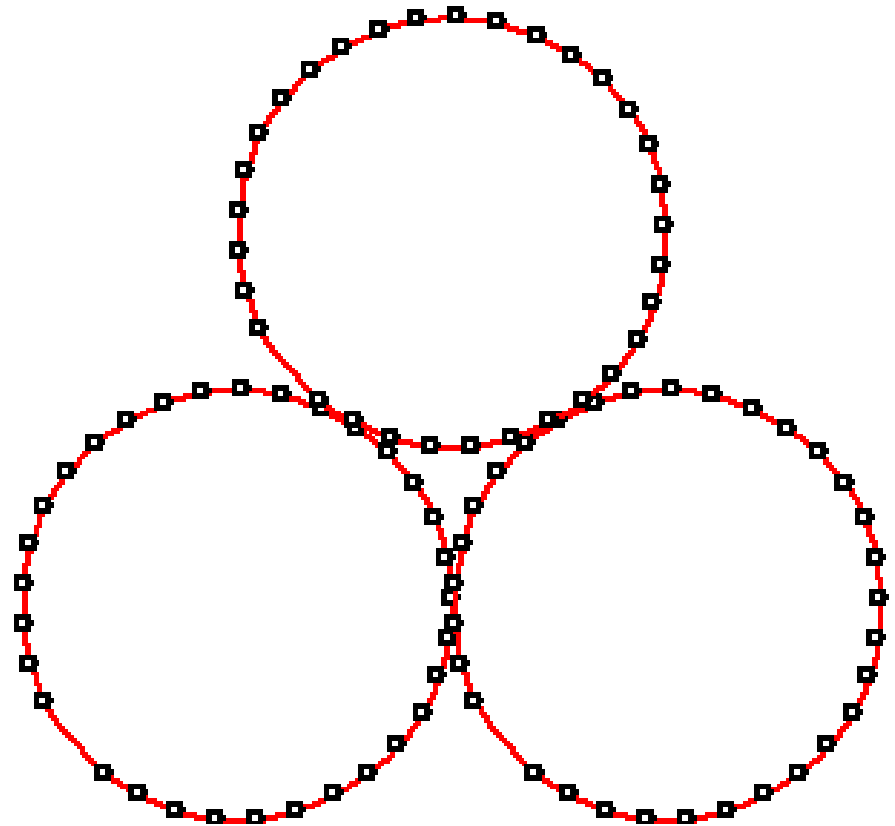
- Gaussian quadrature within nonsingular elements
- Radial integration method [Gao, 2006], within singular elements
- Velocity optimized using a quadratic optimization scheme to ensure



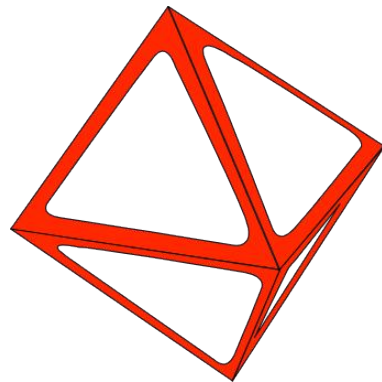
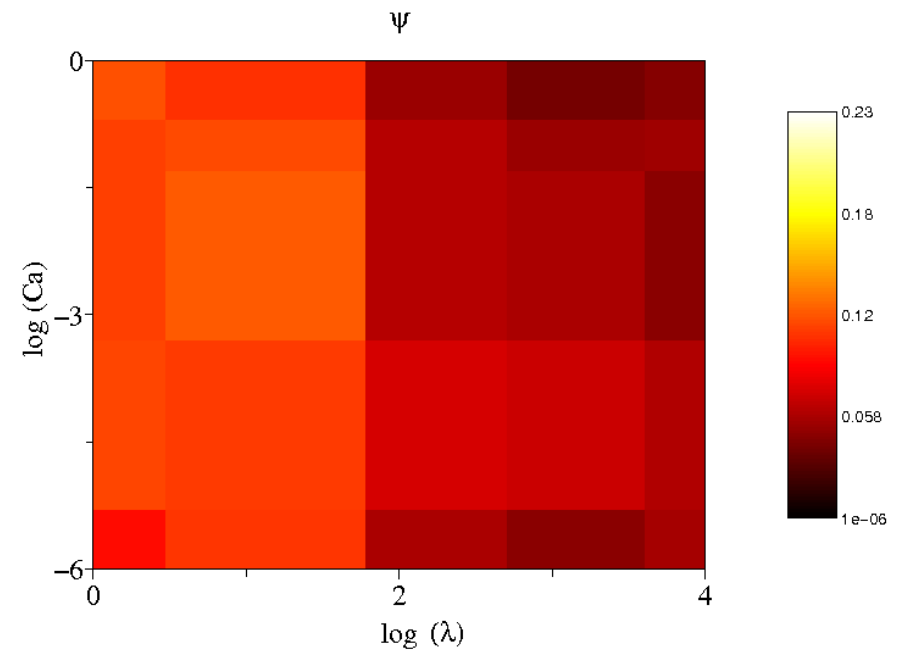
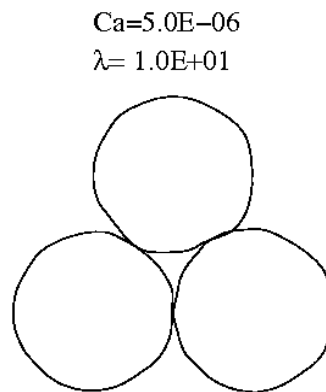
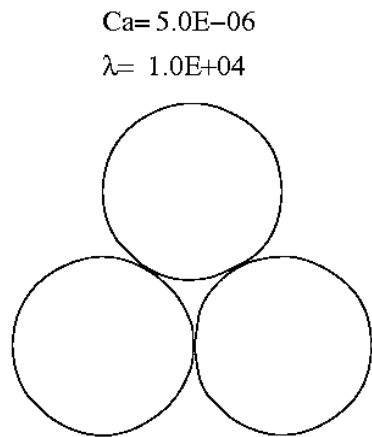
$$\mathbf{u} \cdot \mathbf{n} = 0$$

Updating shape

- Update nodal coordinates using forward Eulerian integration
- Remove high curvature nodes
- Marker position updated until convergence is attained



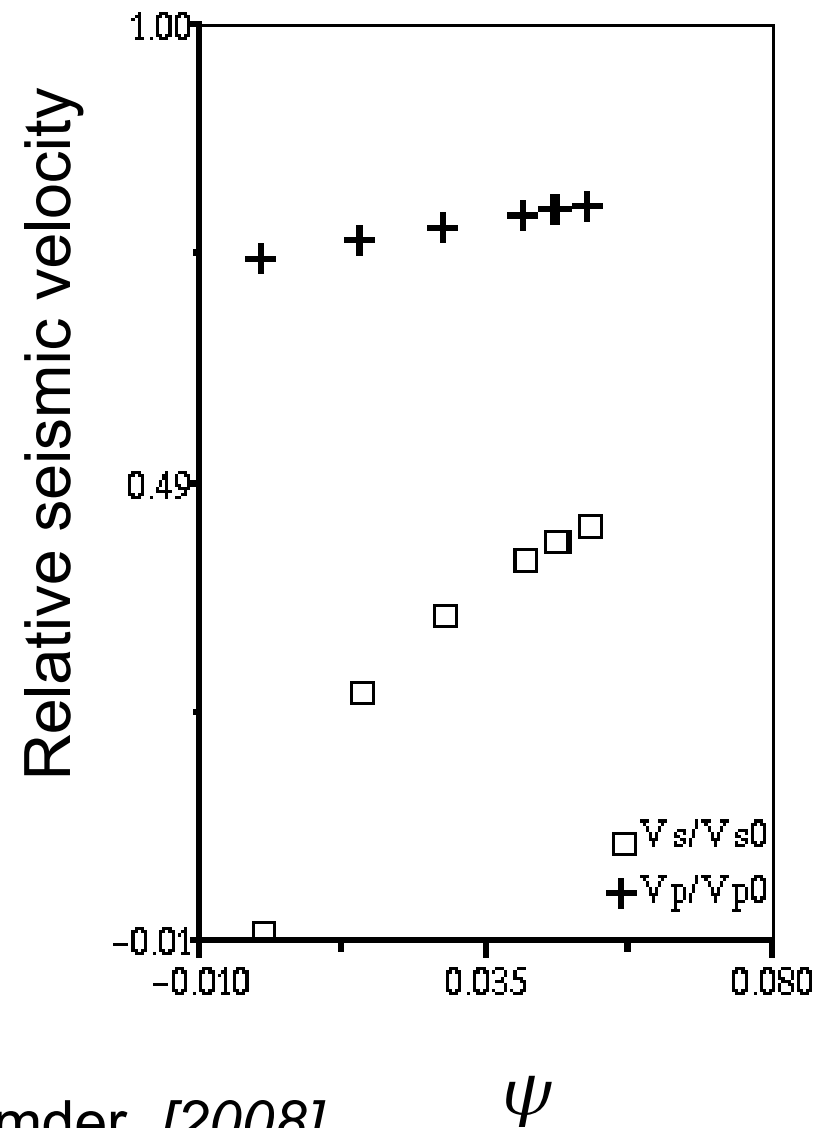
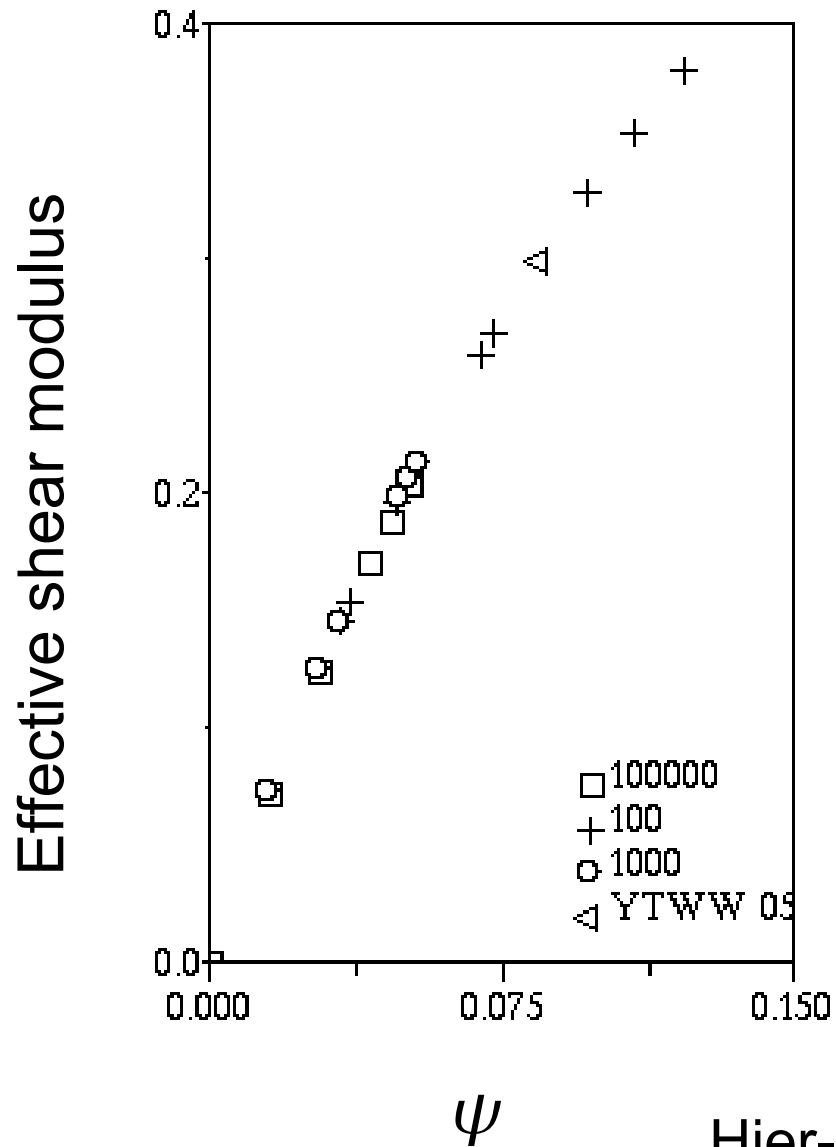
Contiguity



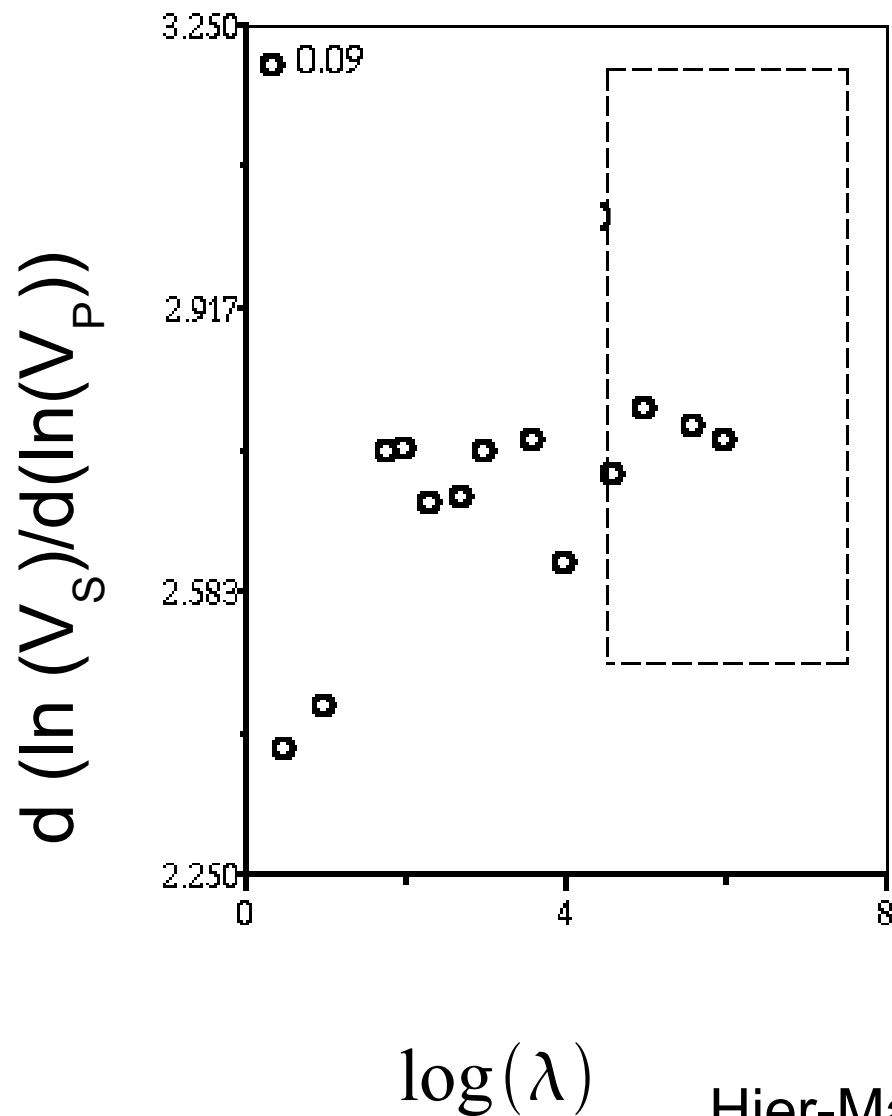
$$\psi = \frac{A_{gg}}{A_{gg} + A_{gm}}$$

Hier-Majumder [2008]

Contiguity and elastic properties

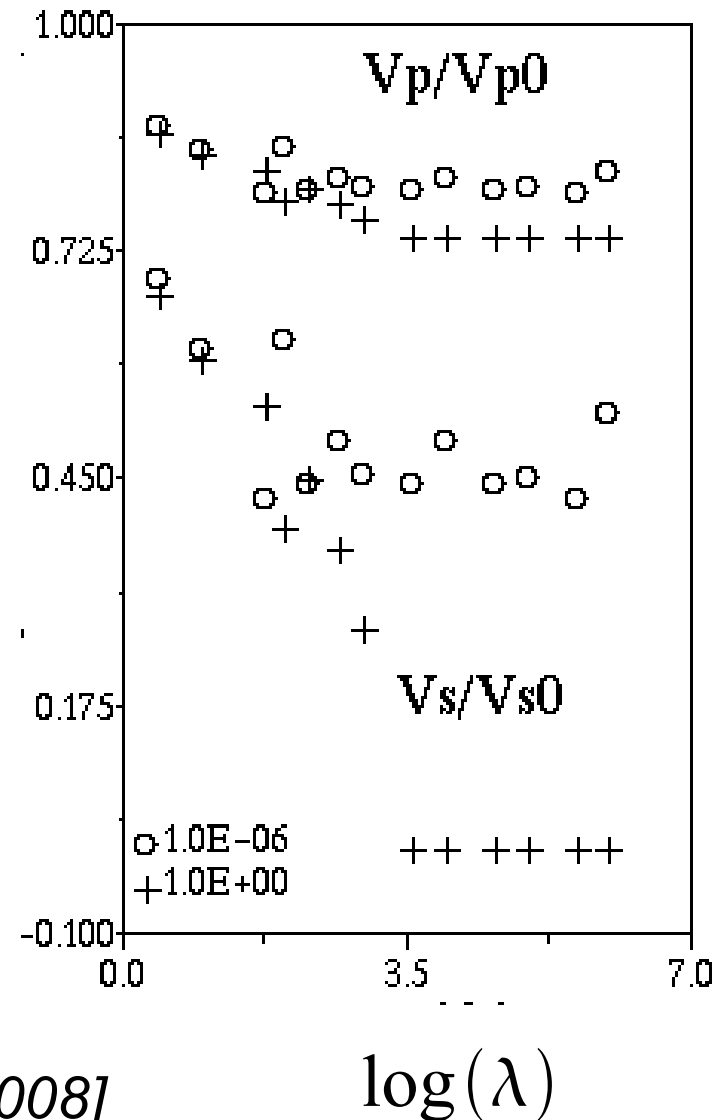


Viscosity and elastic properties



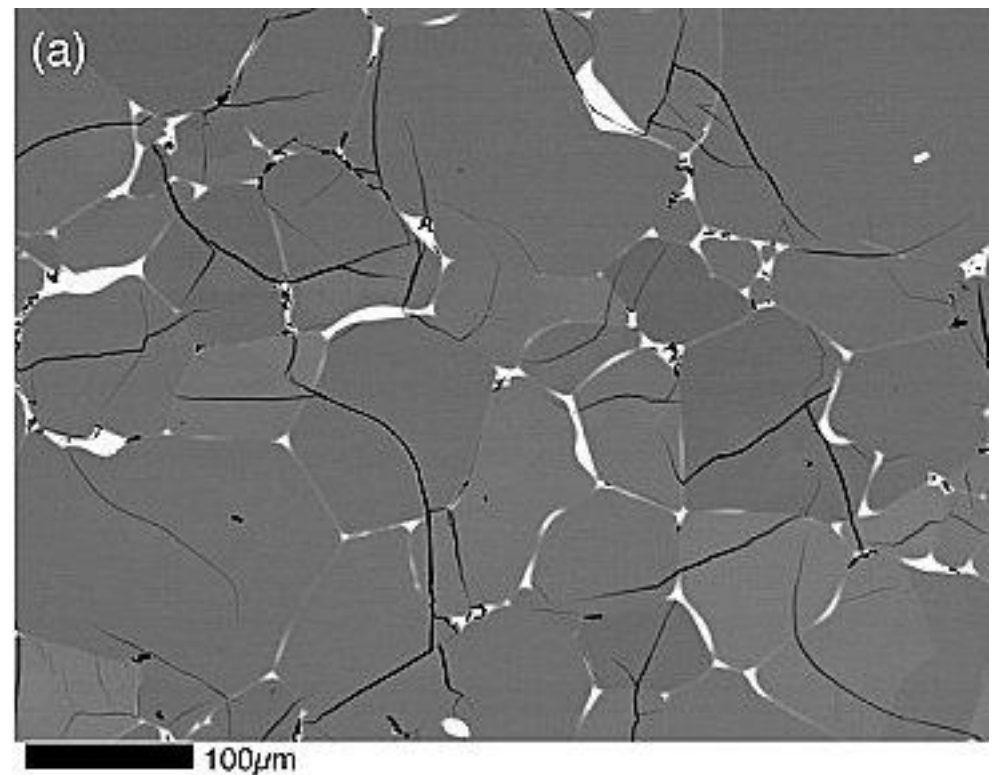
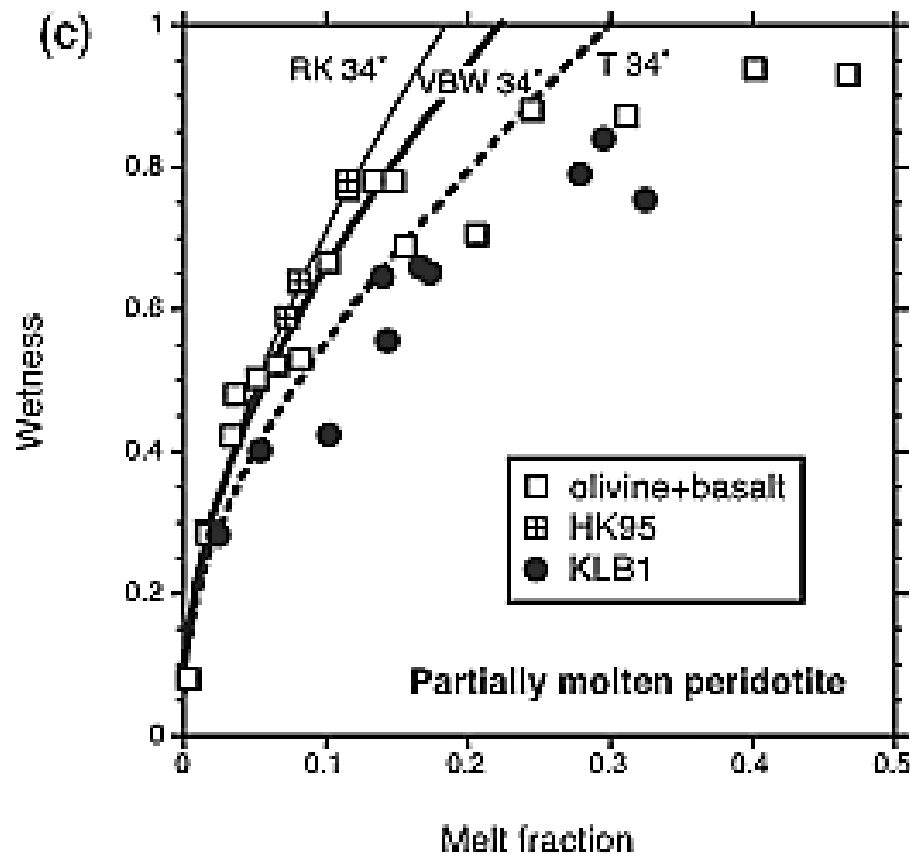
Hier-Majumder [2008]

Effective seismic velocities



Contiguity and melt fraction

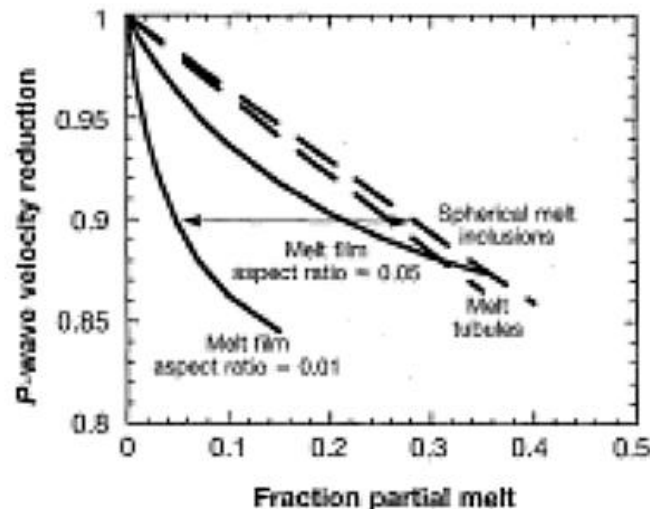
$$\psi = 1 - A \sqrt{\phi} \quad (\text{at a given viscosity ratio})$$



Yoshino *et al* [2005]

Extent of melting in the ULVZ

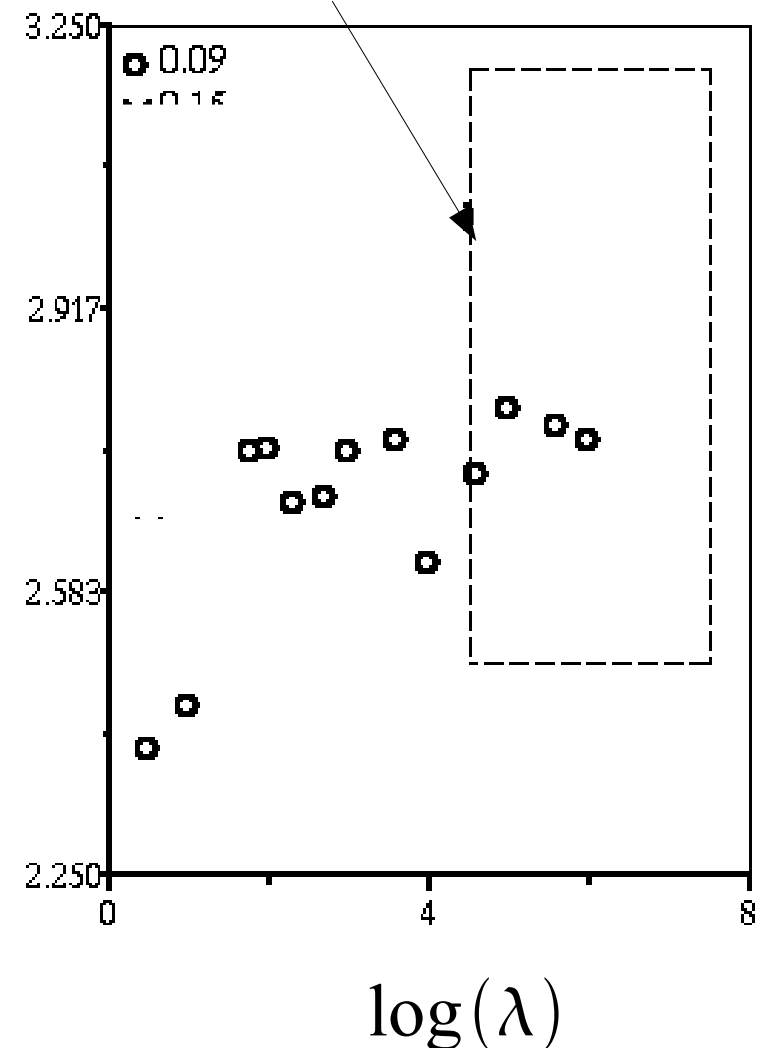
- Estimate from inclusion models 5-30 volume % melt for 3:1 reduction of shear and P wave velocities



Williams and Garnero [1996]

Masters *et al* [2000]

$$d(\ln(V_S)/d(\ln(V_P)))$$

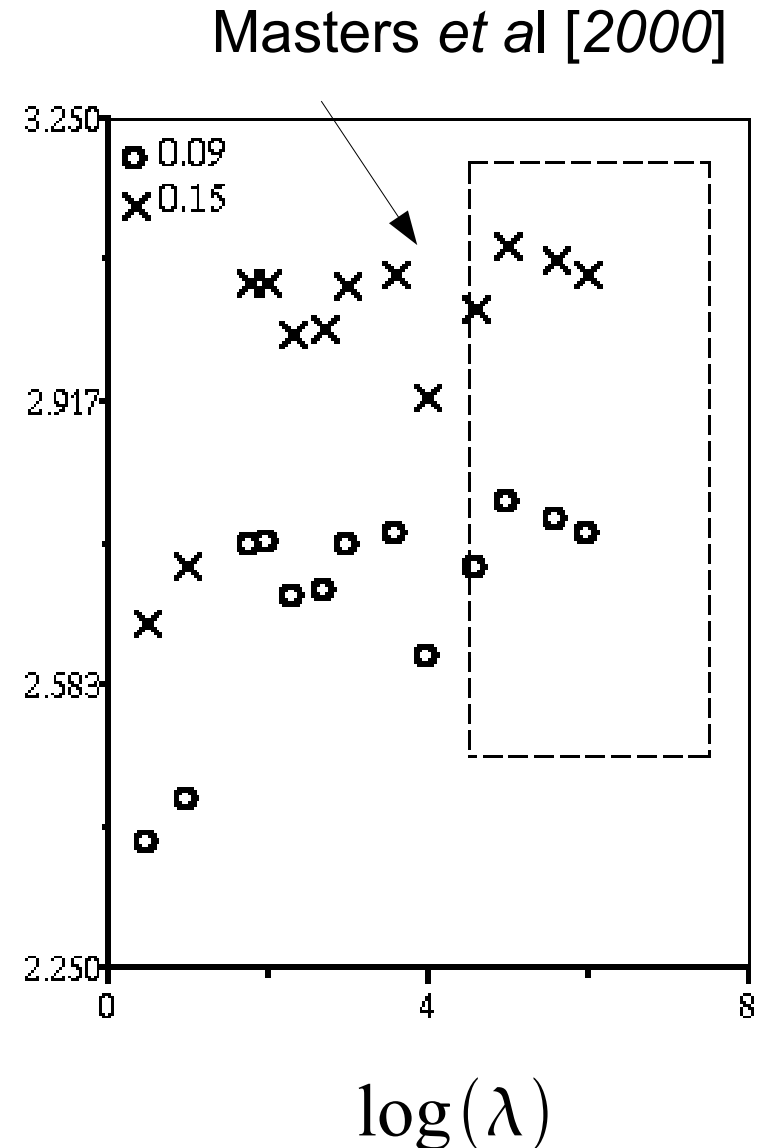


Hier-Majumder [2008]

Extent of melting in the ULVZ

- Current model indicates between 10-15% melting for 3:1 reduction of S and P wave velocities
- The ULVZ is molten, but probably not disaggregated

$d(\ln(V_S)/d(\ln(V_P)))$

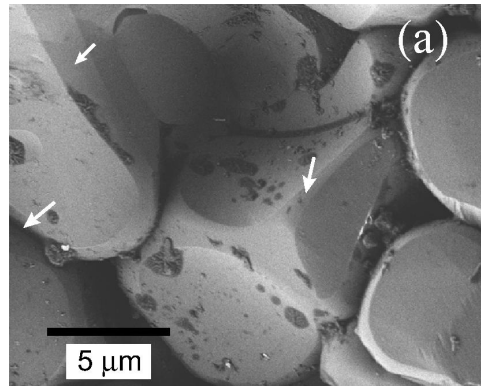
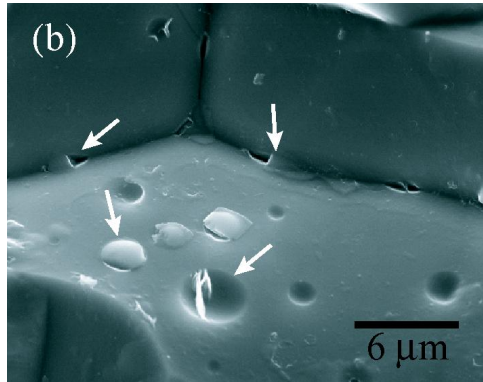


Hier-Majumder [2008]

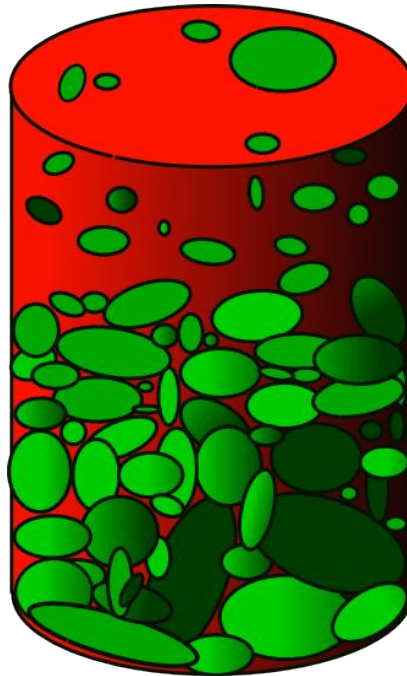
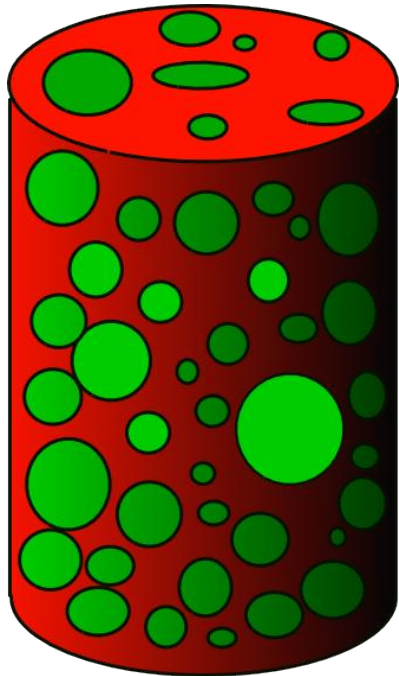
Influence of melting on the topography of the ULVZ

- Coupled effect of surface tension, compaction, buoyancy, and deformation
- Mantle flow and viscosity contrast between the ULVZ and the ambient mantle

Two-phase theory



- Coupled viscous flow of the melt and the matrix
- Surface tension balance pressure, viscous stress, and body forces
- Incorporates both melt geometry and disaggregation.



Governing equations (1D)

Conservation of Mass

$$\frac{\partial \phi}{\partial t} = \frac{\partial}{\partial y} \left((1 - \phi) w \right)$$

Force balance ('action-reaction')

$$\xi (1 - \phi) \frac{\partial^2 \chi}{\partial \phi^2} \left(\frac{\partial \phi}{\partial y} \right) + \frac{4}{3} \frac{\partial}{\partial y} \left(\frac{1 - \phi^2}{\phi} \frac{\partial w}{\partial y} \right) - R (1 - \phi) - \frac{4}{3} \frac{w}{\phi^2} = 0$$

Stress drop condition

$$\frac{\partial \chi}{\partial \phi} = -\Delta P + B \frac{D_m \phi}{D t}$$

Bercovici *et al.* [2001]

Some parameters

$$\xi = \frac{\gamma_{gm} \alpha_0}{\rho_s g \delta_s}$$

Modulates surface tension

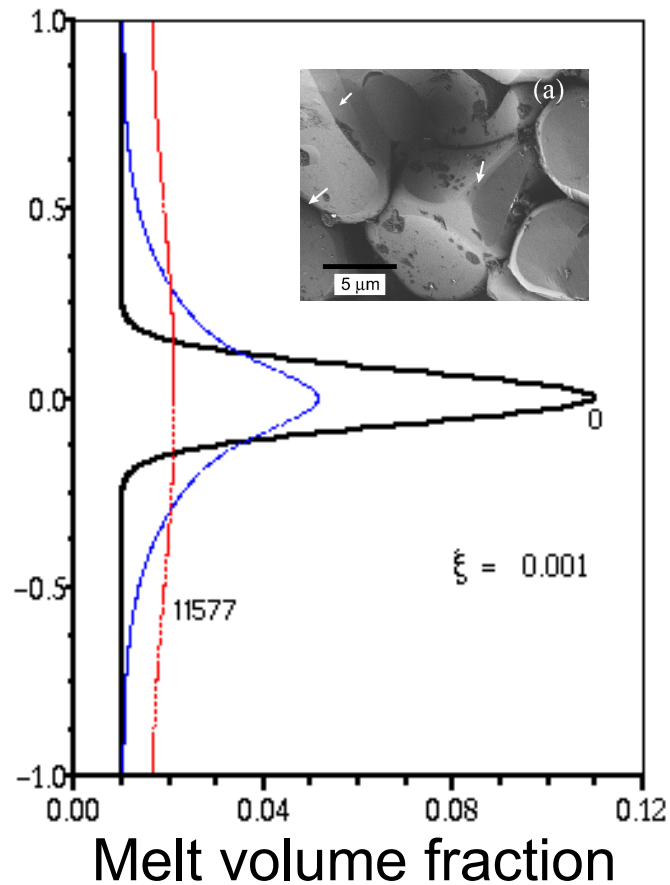
$$R = \frac{\Delta \rho}{\rho_s}$$

Controls buoyancy

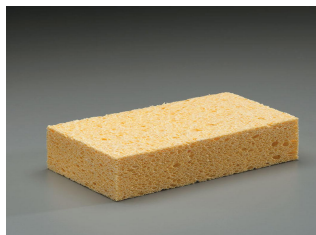
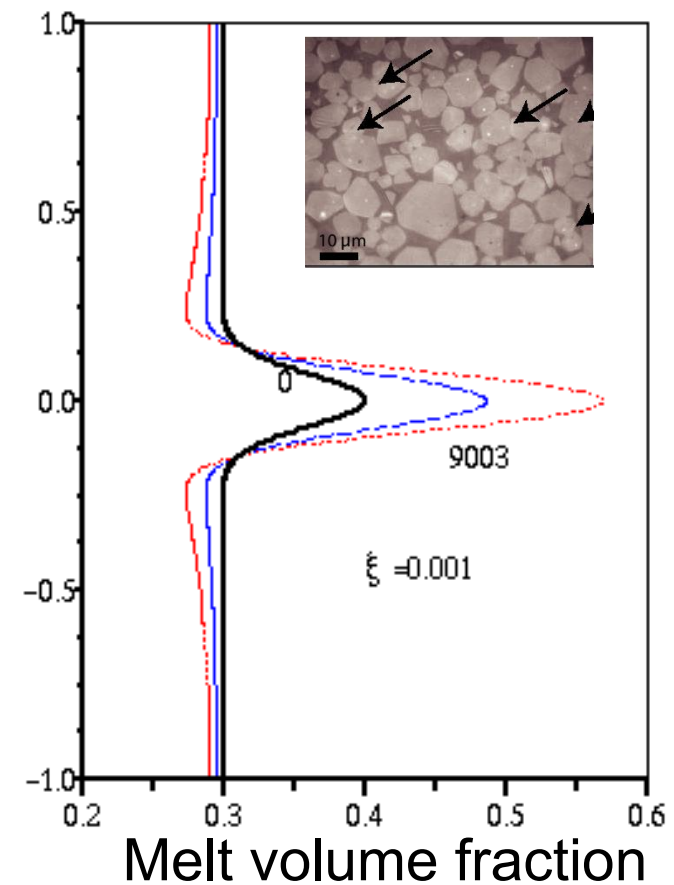
$$\delta_s = \sqrt{\frac{4 \mu_s}{3 c}}$$

Compaction length of the solid

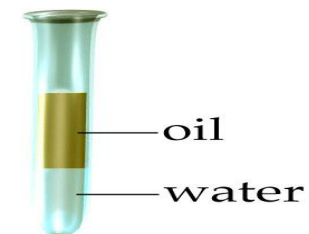
Distribution of a neutrally buoyant melt layer



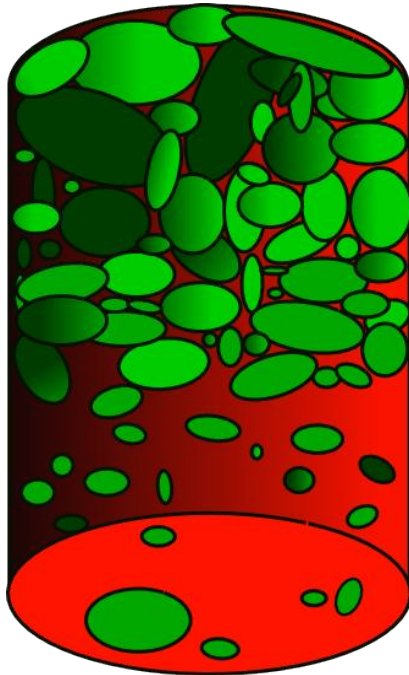
Height



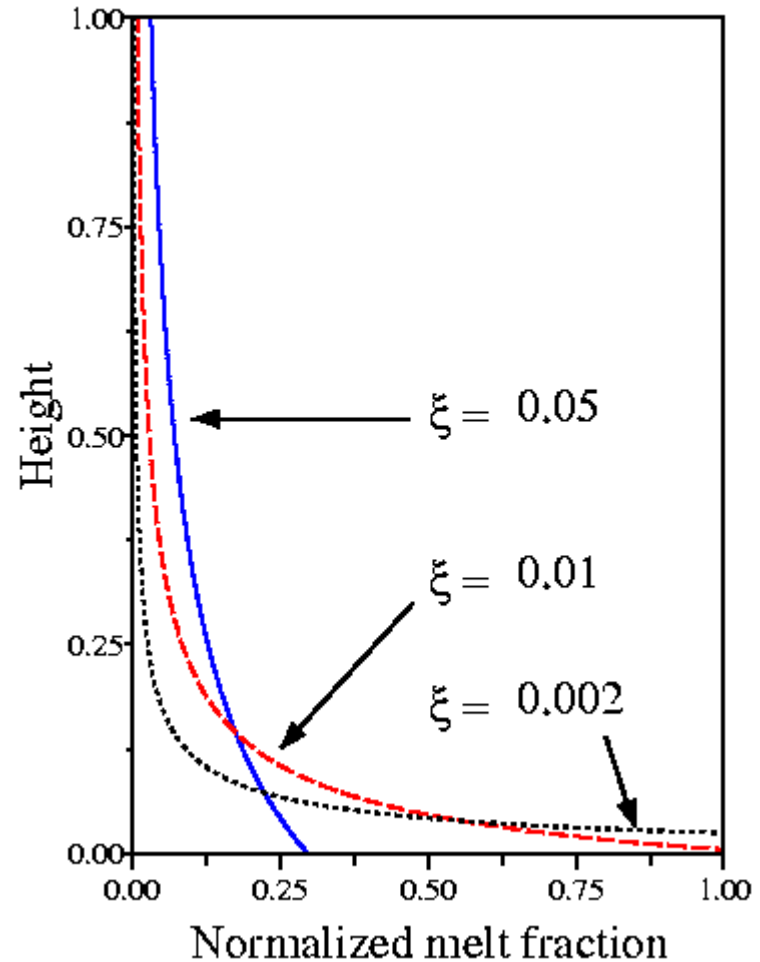
Hier-Majumder *et al* [2006]



Storage of dense melt

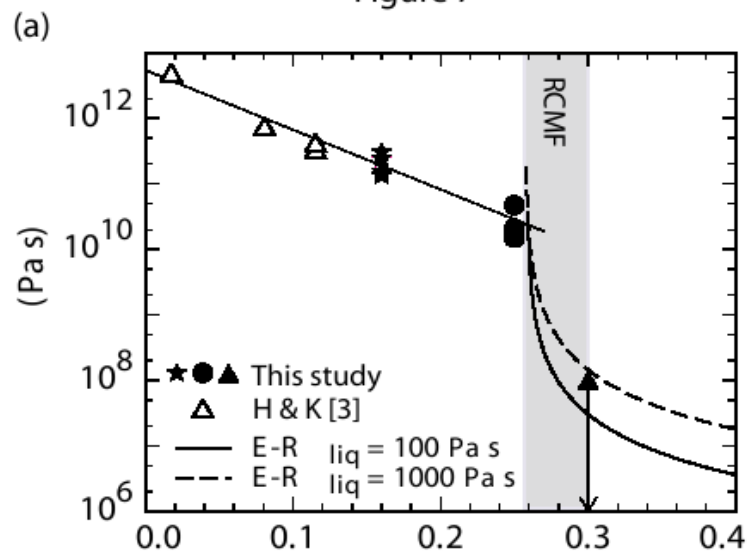
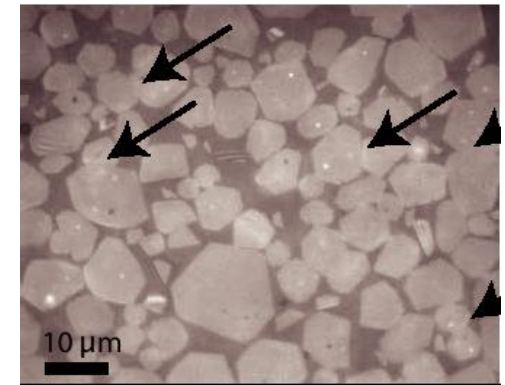
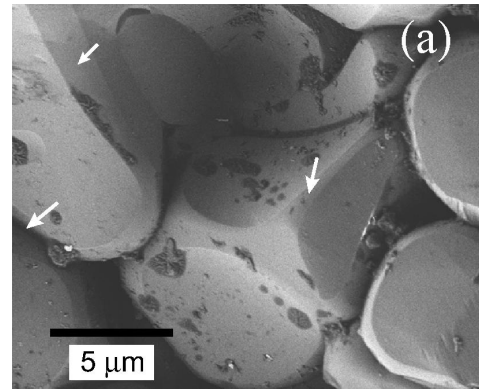
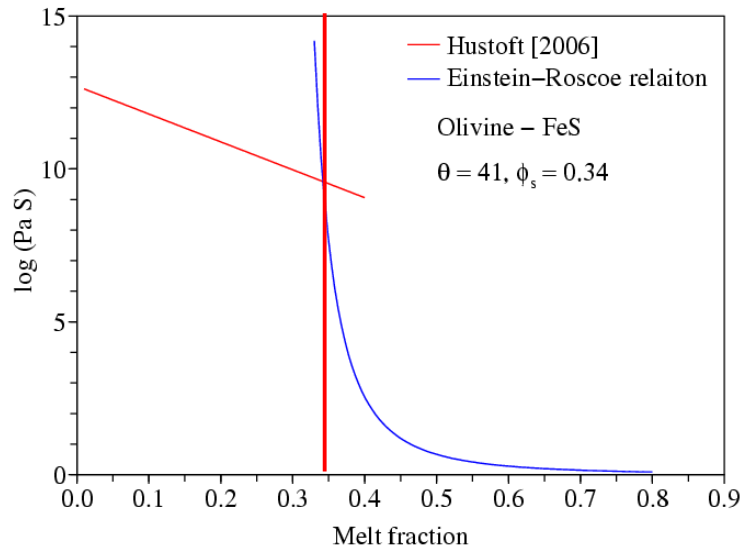


$$\xi = \frac{\gamma_{gm} \alpha_0}{\rho_s g \delta_s} \quad \delta_s = \sqrt{\frac{4 \mu_s}{3 c}}$$

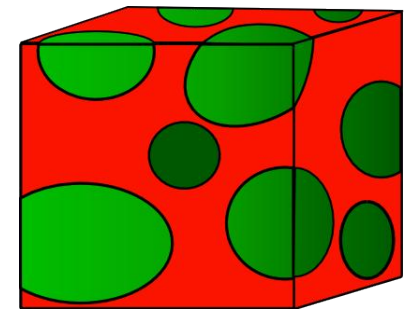
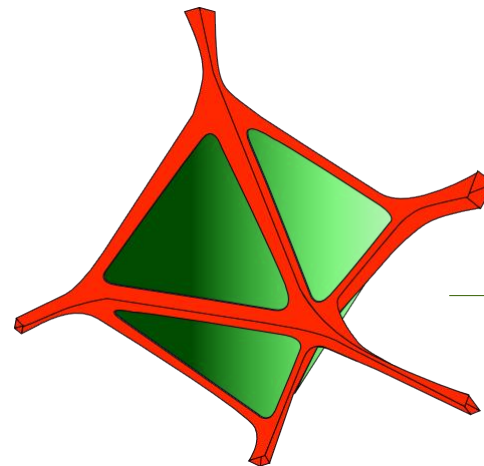


Hier-Majumder *et al.* [2006]

Melt fraction and viscosity

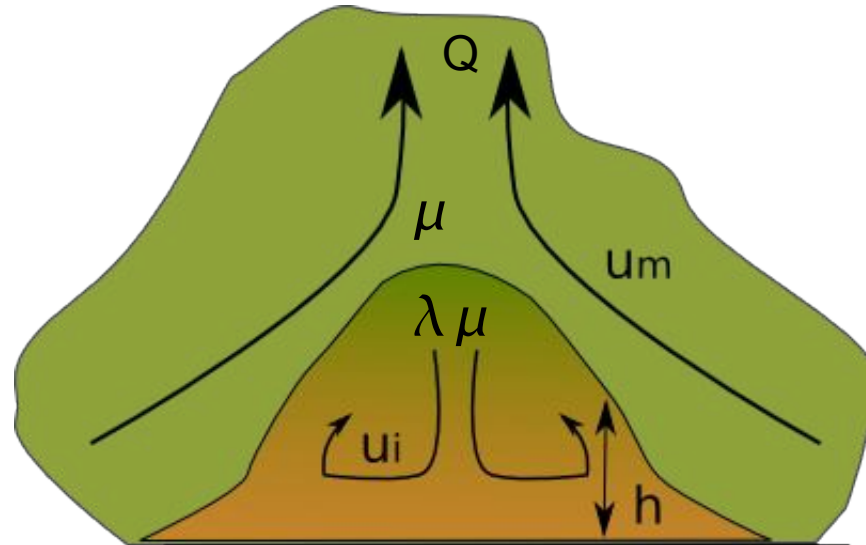


Scott and Kohlstedt, [2006].



Low viscosity dense layer

Q = Sink strength



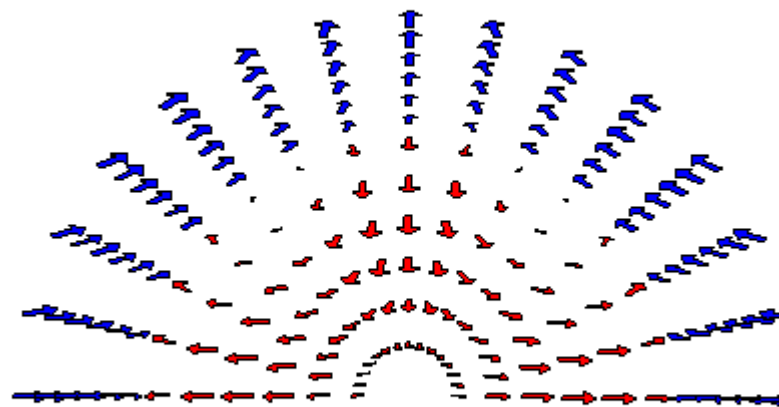
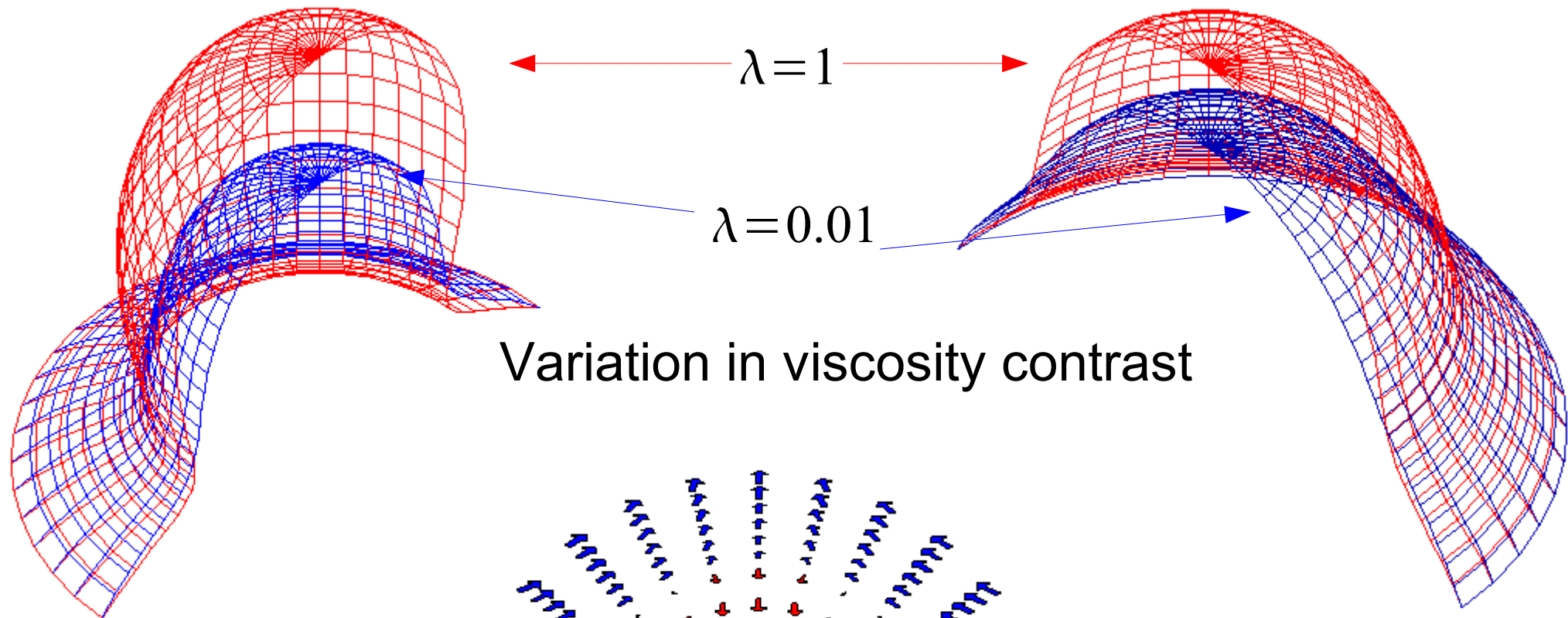
$$\tau_{r\theta}^i = \tau_{r\theta}^m$$

Boundary conditions

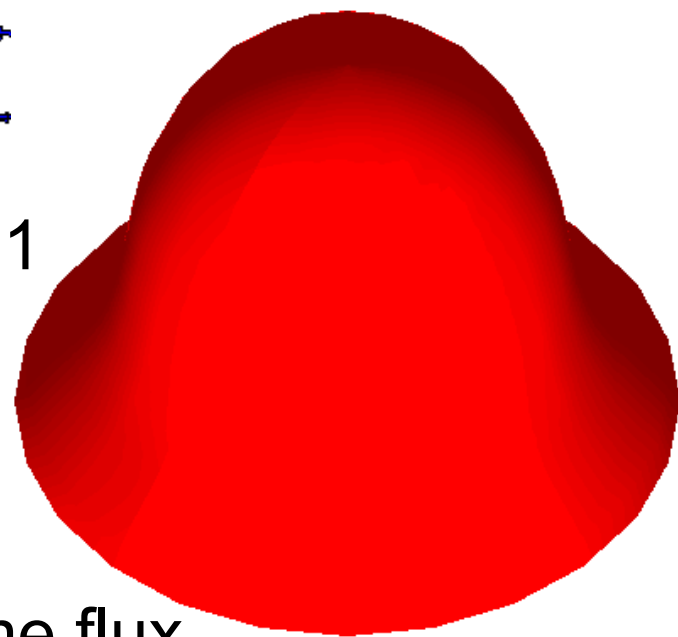
$$\mathbf{u}^i \cdot \hat{\boldsymbol{\theta}} = \mathbf{u}^m \cdot \hat{\boldsymbol{\theta}}$$

Topography relation

$$\hat{\mathbf{z}} \cdot \Delta \mathbf{T} \cdot \mathbf{n} = \Delta \rho g h$$



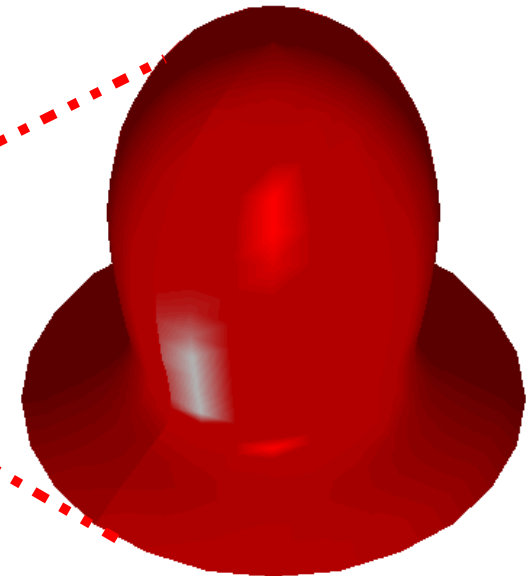
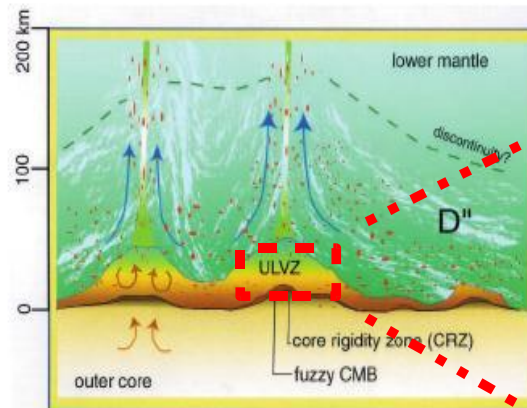
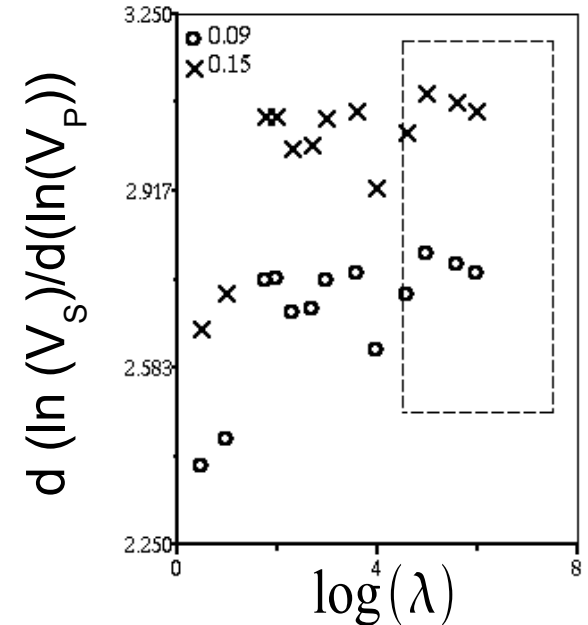
$Q = 1$



Variation in normalized plume flux

Conclusions

- Observed range of S and P-wave velocity changes can be explained by 10-15% melting
- The ULVZ is likely molten but not disaggregated
- Topography of the ULVZ is likely influenced by
 - Tension on grain boundaries
 - Plume flux
 - Density and viscosity contrast



Lamb's solution for a single particle

$$\begin{aligned}
 P(r, \theta) &= \sum_n p_n \\
 u(r, \theta) &= \sum_n \left(\frac{(n+3)r^2 \nabla p_n}{2\mu(n+1)(2n+3)} - \frac{n \mathbf{r} p_n}{\mu(n+1)(2n+3)} + \nabla \Phi_n \right) \\
 p_n &= a_n r^n P_n(\cos \theta) \\
 \Phi_n &= b_n r^n P_n(\cos \theta) \\
 p_{-n-1} &= \frac{A_n}{r^{n+1}} P_n(\cos \theta) \\
 \Phi_{-n-1} &= \frac{B_n}{r^{n+1}} P_n(\cos \theta)
 \end{aligned}$$

Dihedral angle

