



From Dust to Planets:
**Volatile element depletion in the
inner Solar System**

Dynamics and Evolution of Earth-like Planets
Feb. 12 , 2015

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Outlines

- 1. Nature of planetary building blocks (chondrites min/pet)**
- 2. Volatile depleted nature and its current budget (chemistry)**
- 3. Earth atmosphere is leftover from erosion (isotopes)**

Two paths to enlightenment

How do I learn how planets and solar systems form?



meteoriticist



astronomer

-PLETCH-

-PLETCH-

Courtesy of S. Desch via S. Stewart

**You must find solar systems
and planets forming now,
and observe them thusly**



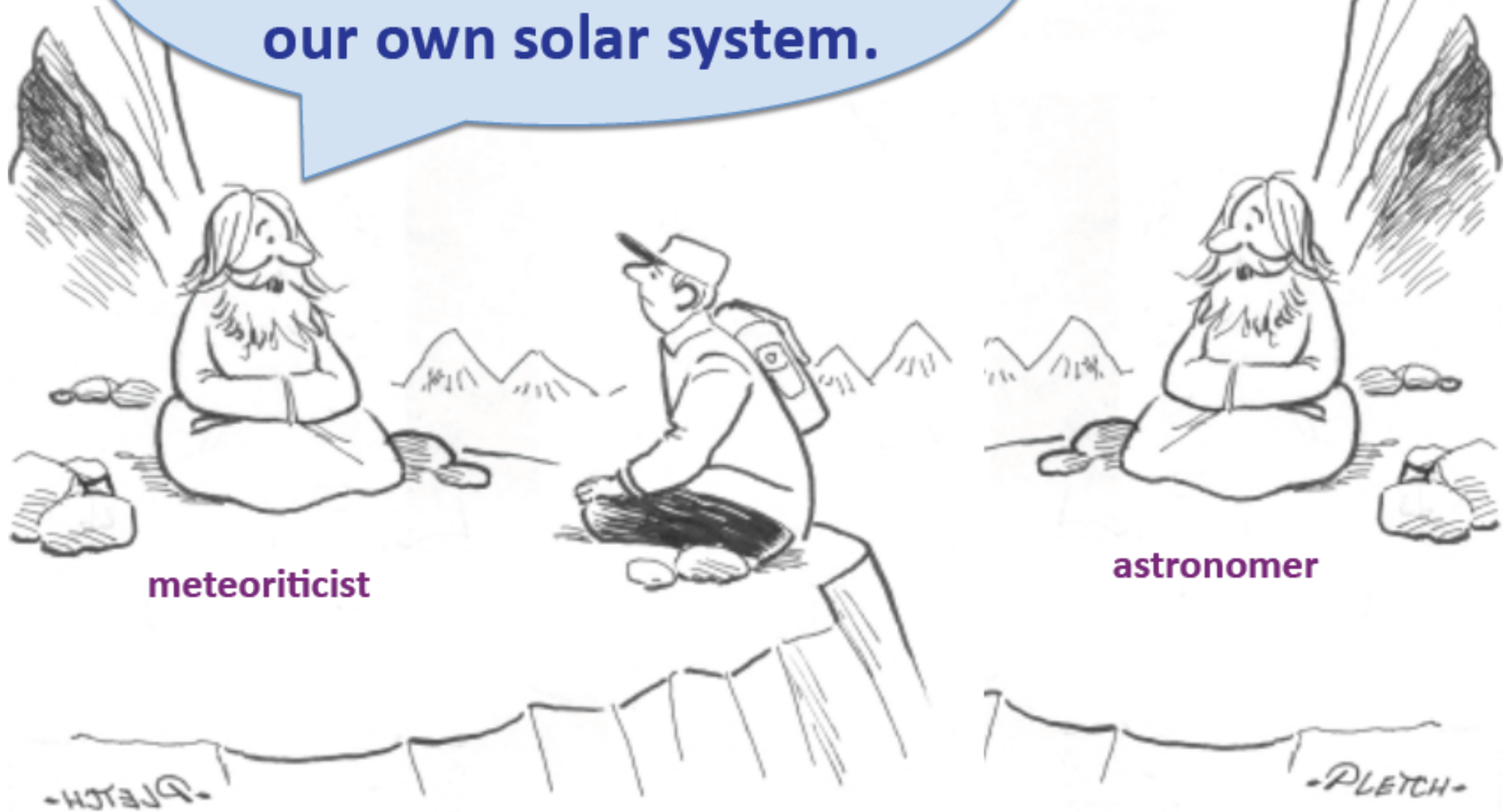
Courtesy of S. Desch via S. Stewart

Astronomer's View

(as understood by a meteoriticist)

- **Super-Earths and Mini-Neptunes are everywhere, Jupiter-likes far less common.**
- **Disk run out of gas in only a few Ma.**
- **In this timeframe, dust grains grow from mm to Mars-sizes.**
- **Planets must form eventually because we see them everywhere, but we can't see how it was done nor where they formed.**
- **Migration requires gas, but if gas is around, runaway growth would make Jupiter instead of super-Earths and mini-Neptunes**

You must read the ancient record of the formation of our own solar system.



meteoriticist

astronomer

Galactic Chemical Evolution

$X (H) = 71.10\%$

$Y (He) = 27.41\%$

Z (Astronomers "metal": Li-U) = 1.49%

*C, N, O, and Ne dominate more than half of Z,
and C and O are most abundant "metals":*

C: 0.247%

O: 0.657%

*CO is second most abundant gas next to H₂
C/O < 1 dominates galactic chemical evolution*

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Concept

Primary Solids: Dust inherited from ISM.

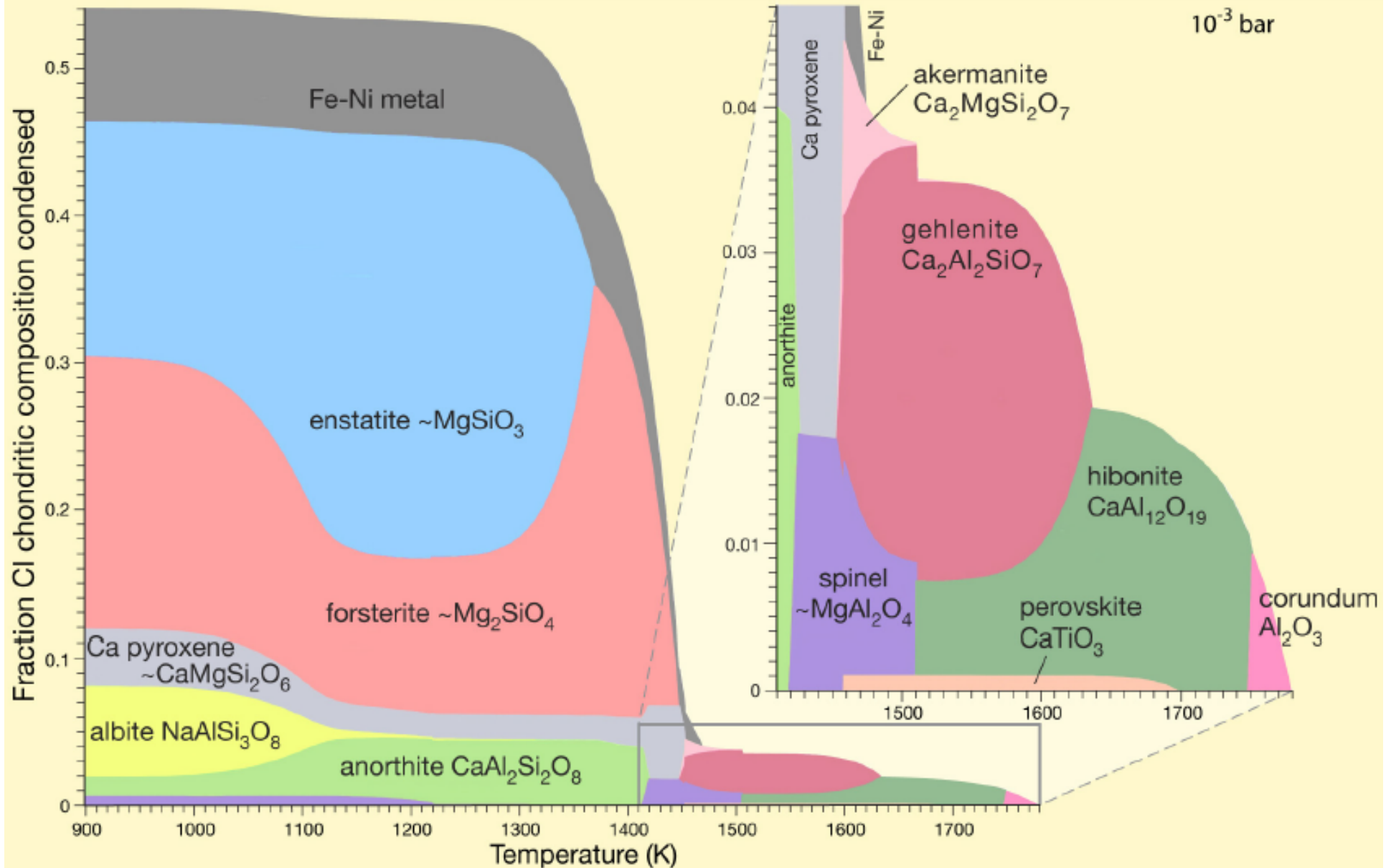
Secondary Solids: Condensates, evaporation residue, molten chondrules, etc.

Results of further processing of the primary solids within solar nebula.

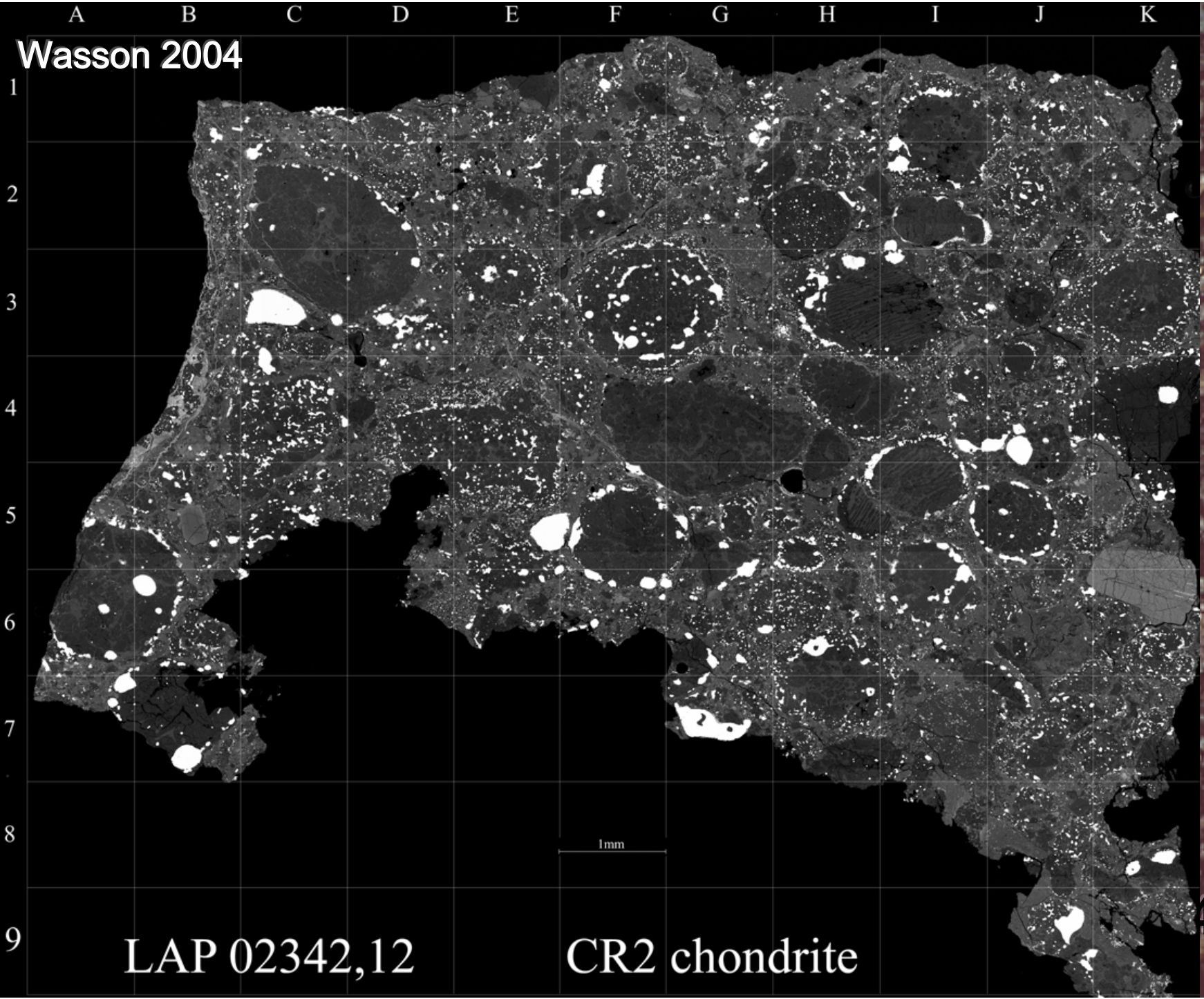


Mineral stability in solar nebula

(courtesy of Andy Davis)



Wasson 2004



LAP 02342,12

CR2 chondrite

4

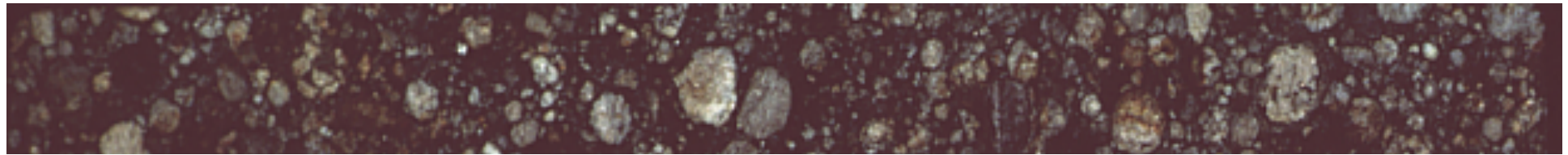
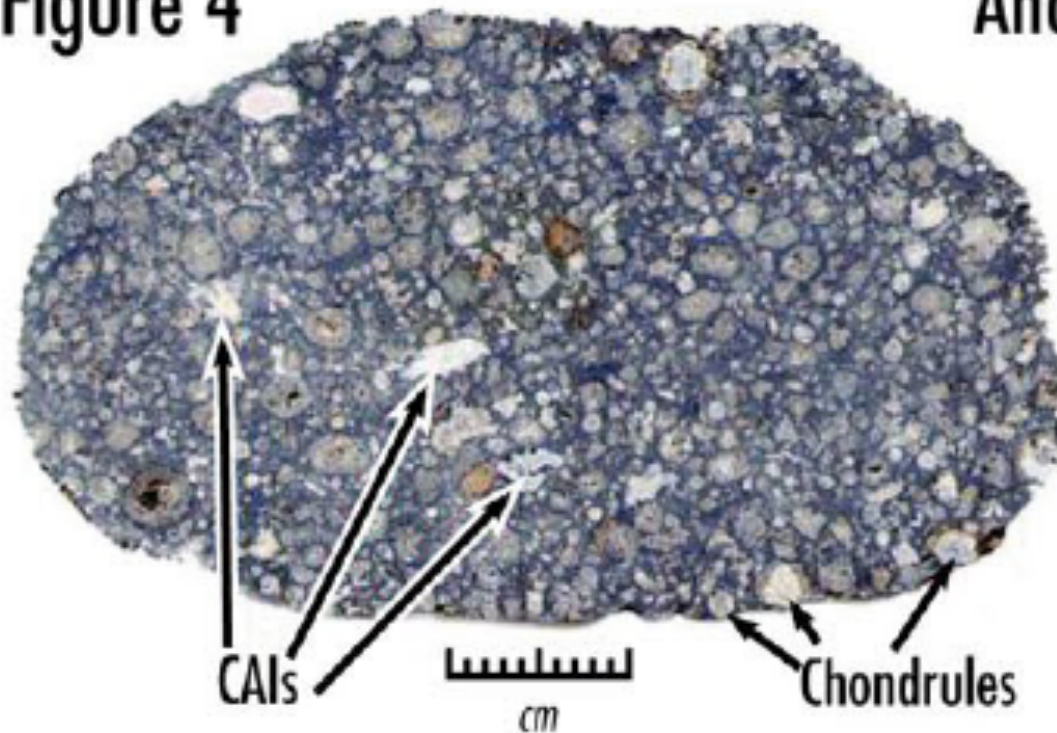


Figure 4



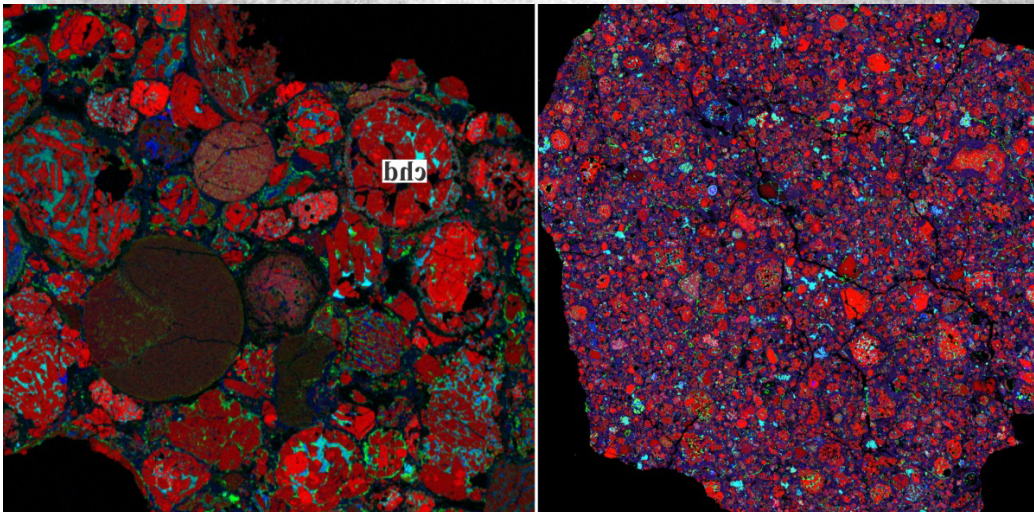
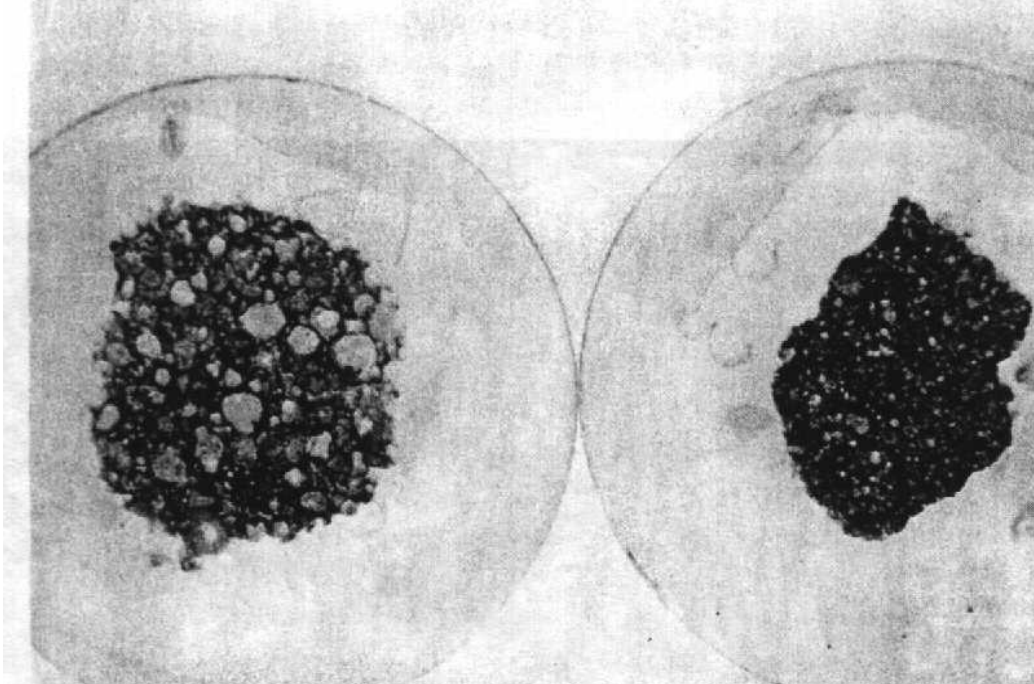
Anatomy of cosmic sediments:

1. High temperature refractory bits and pieces (CAIs, chondrules)
2. Low temperature fine grained dark matrix with pre-solar grains, organic matter and water

A key question is: *How and when do these diverse components come together?*



J. A. Wood (1985)

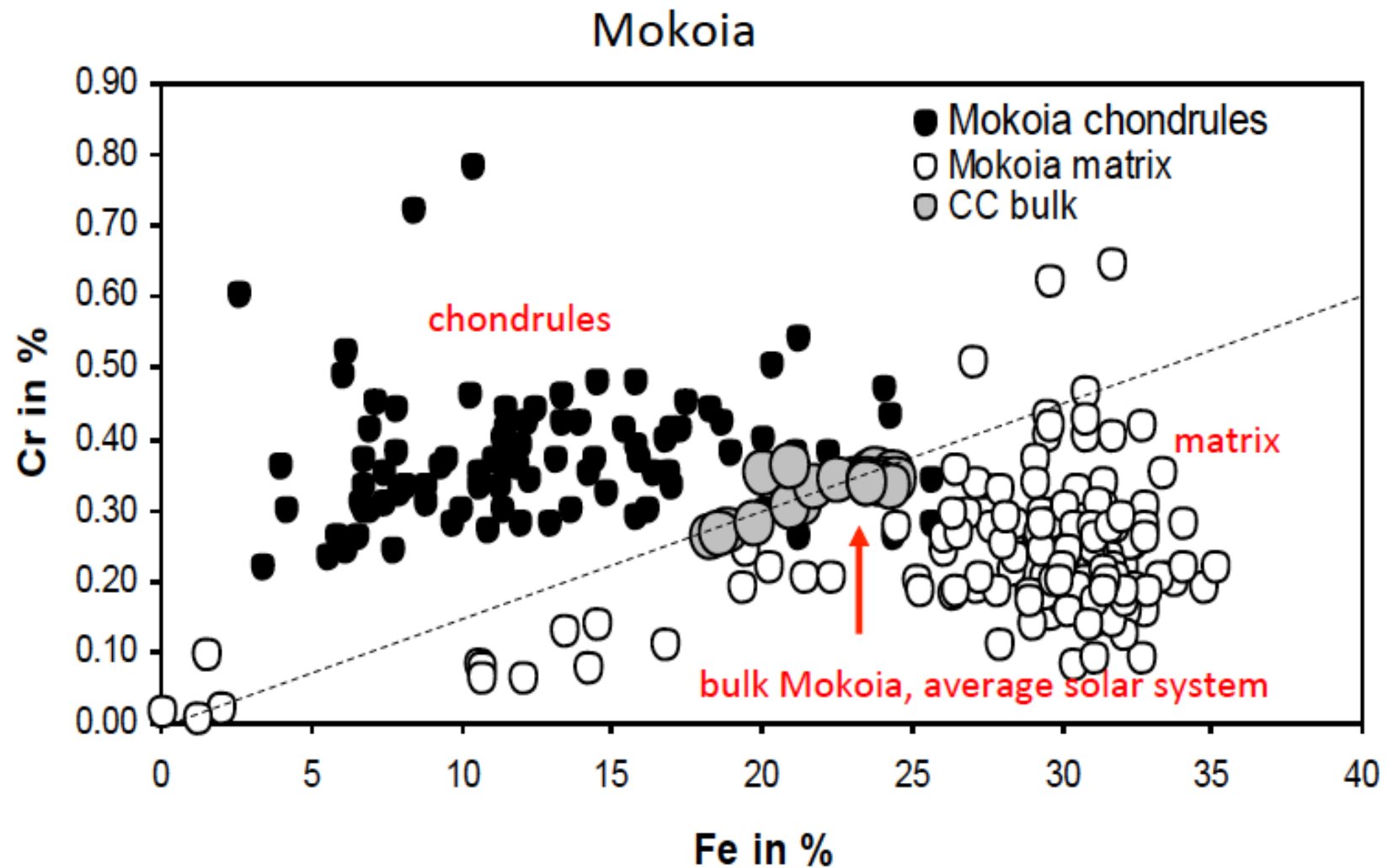


“Fredriksson’s Paradox”

- Different petrography;
- Variable proportions of chondrules to matrix;
- Identical bulk chemical composition

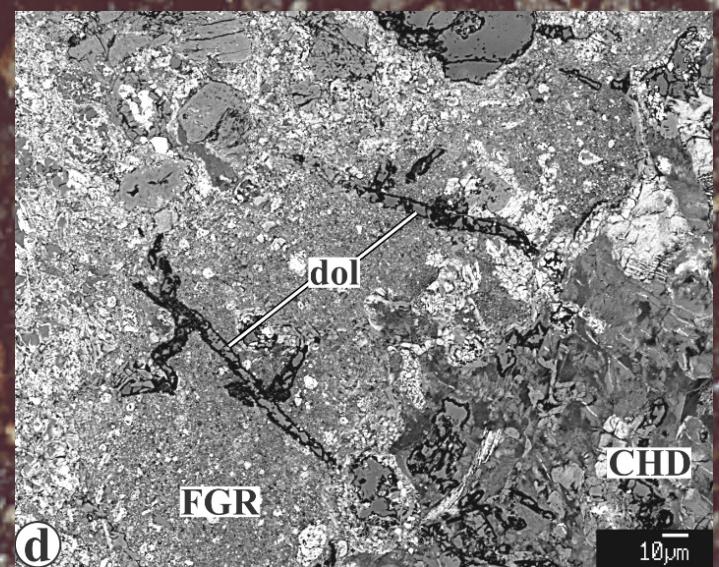
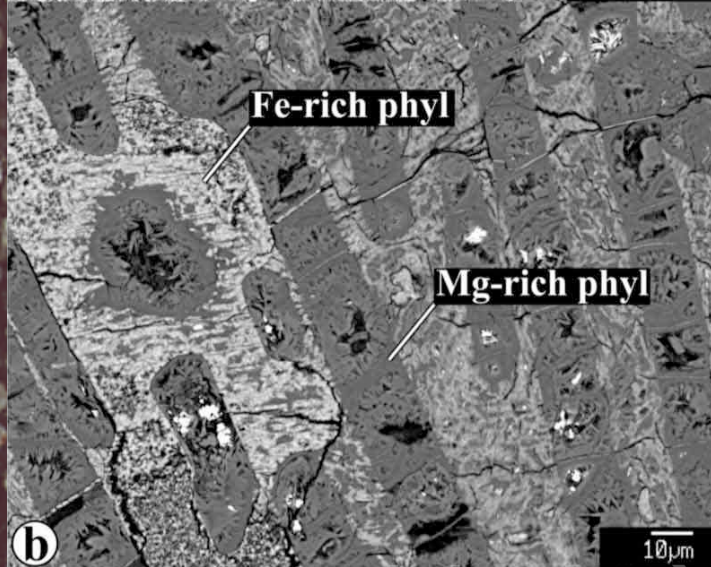
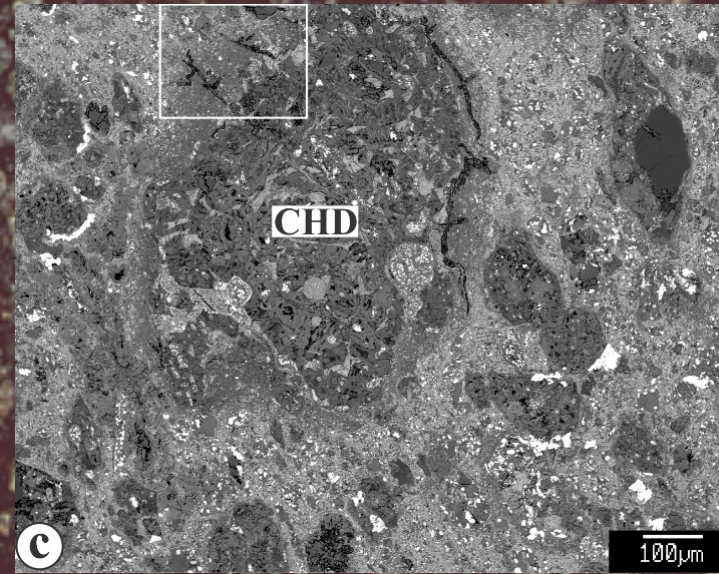
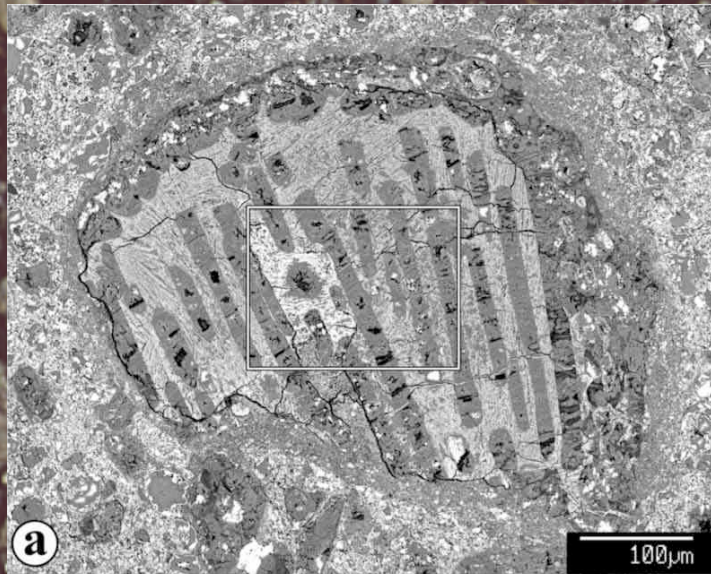
Implying:

- Prompt aggregation with the matrix after chondrules and CAIs formation (\ll orbital period).
- If chondrules continued to orbit for thousands of years, expect aerodynamically sorted, thoroughly mixed morphology, textural identities lost!
- Local production, reflect local precursor reservoir chemistry

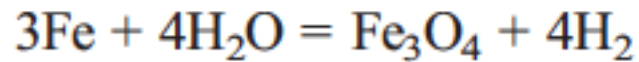
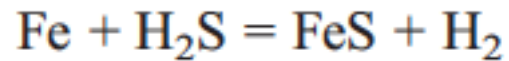


Parent body processes cannot produce forsteritic olivine.
 Aqueous alteration may produce FeO-rich olivine.

Sutter's Mill (CM chondrites)



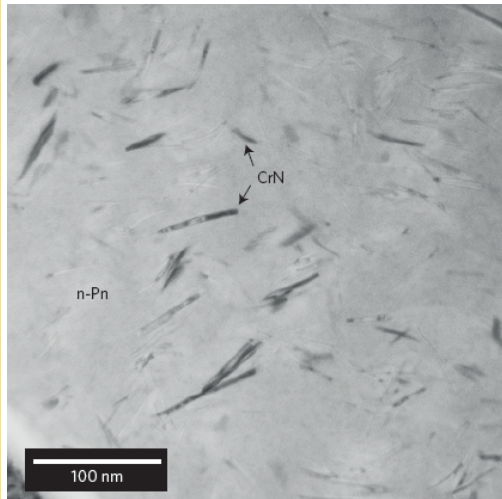
Jenniskens, Yin, Krot et al (2012)



Remnants of the Early Solar System Water Enriched in Heavy Oxygen Isotopes

Naoya Sakamoto,¹ Yusuke Seto,¹ Shoichi Itoh,¹ Kiyoshi Kuramoto,² Kiyoshi Fujino,¹ Kazuhide Nagashima,³ Alexander N. Krot,³ Hisayoshi Yurimoto^{1,4*}

Science 2007



Carsbergite (CrN) a secondary mineral named after beer

Reactive ammonia in the solar protoplanetary disk and the origin of Earth's nitrogen



Dennis Harries^{1,2*}, Peter Hoppe³ and Falko Langenhorst¹

nature
geoscience

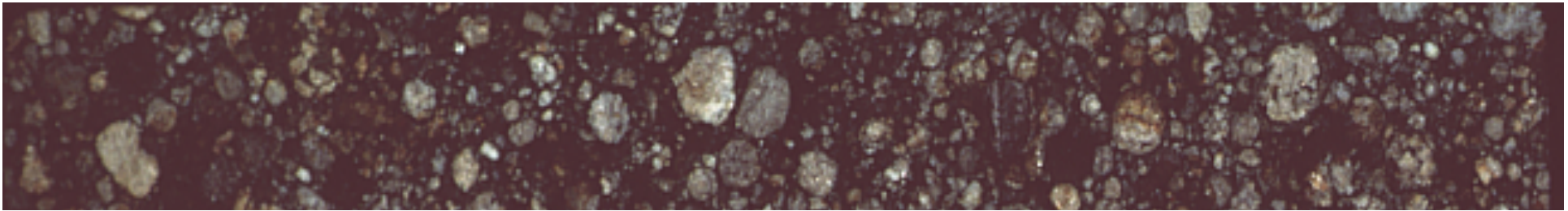
2015

“Sponges” that store volatiles

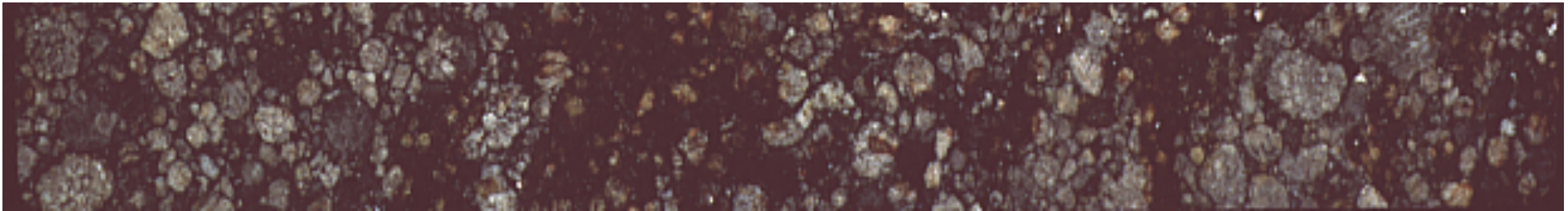
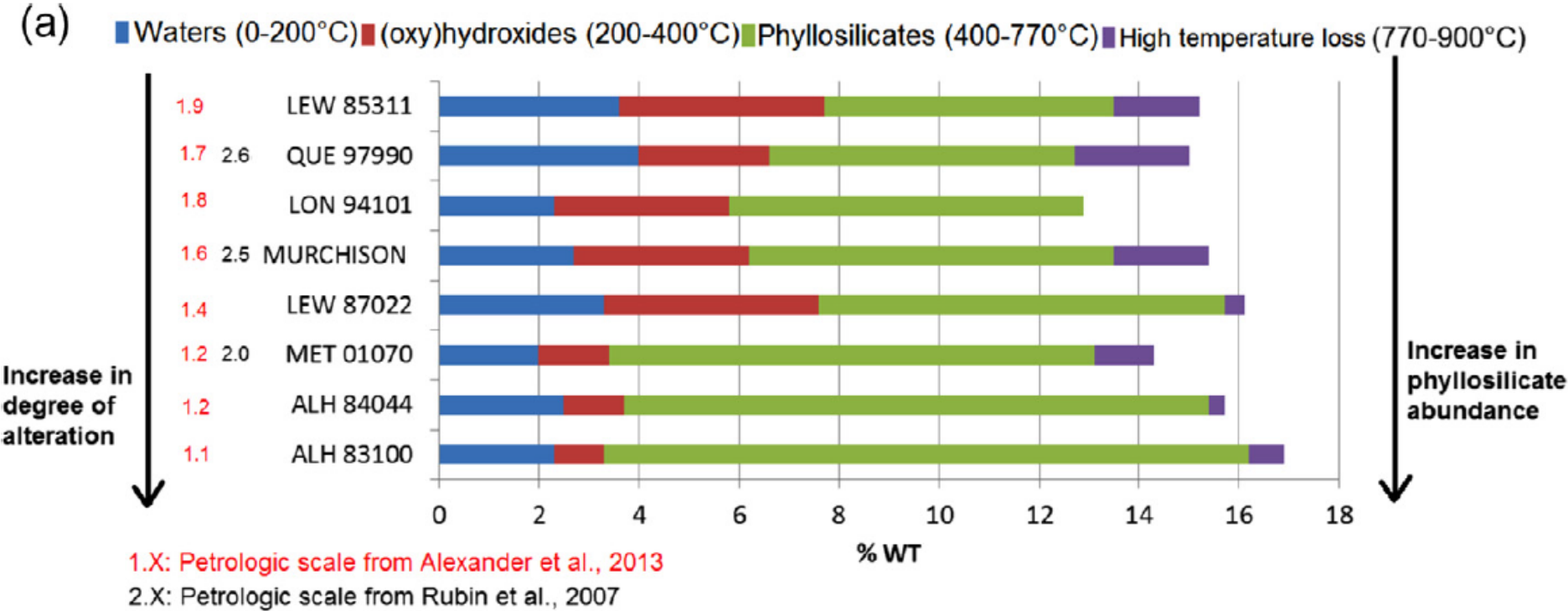
Table 1. Aqueously formed minerals in CI, CM, CR, CV, CO, R, and ordinary chondrites.

CI, CM, and CR chondrites

serpentine	$\text{Mg}_3\text{Si}_2\text{O}_5(\text{OH})_4$
cronstedtite	$\text{Fe}_2^{2+}\text{Fe}^{3+}(\text{Si},\text{Fe}^{3+})_2\text{O}_5(\text{OH})_4$
saponite	$\text{Ca}_{0.25}(\text{Mg},\text{Fe})_3(\text{Si}_4\text{O}_{10})(\text{OH})_2 \cdot n\text{H}_2\text{O}$
tochilinite	$2[(\text{Fe},\text{Mg},\text{Cu},\text{Ni})\text{S}] \cdot 1.57\text{--}1.85[(\text{Mg},\text{Fe},\text{Ni},\text{Al},\text{Ca})(\text{OH})_2]$
calcite	CaCO_3 (trigonal)
aragonite	CaCO_3 (orthorhombic)
dolomite	$\text{CaMg}(\text{CO}_3)_2$
siderite	FeCO_3
breunnerite	$(\text{Mg},\text{Fe},\text{Mn})\text{CO}_3$
troilite	FeS
pyrrhotite	Fe_{1-x}S
pentlandite	$(\text{Fe},\text{Ni})_9\text{S}_8$
magnetite	Fe_3O_4
cubanite	CuFe_2S_3
sphalerite	$(\text{Fe},\text{Zn})\text{S}$



A. Garenne et al. / *Geochimica et Cosmochimica Acta* 137 (2014) 93–112



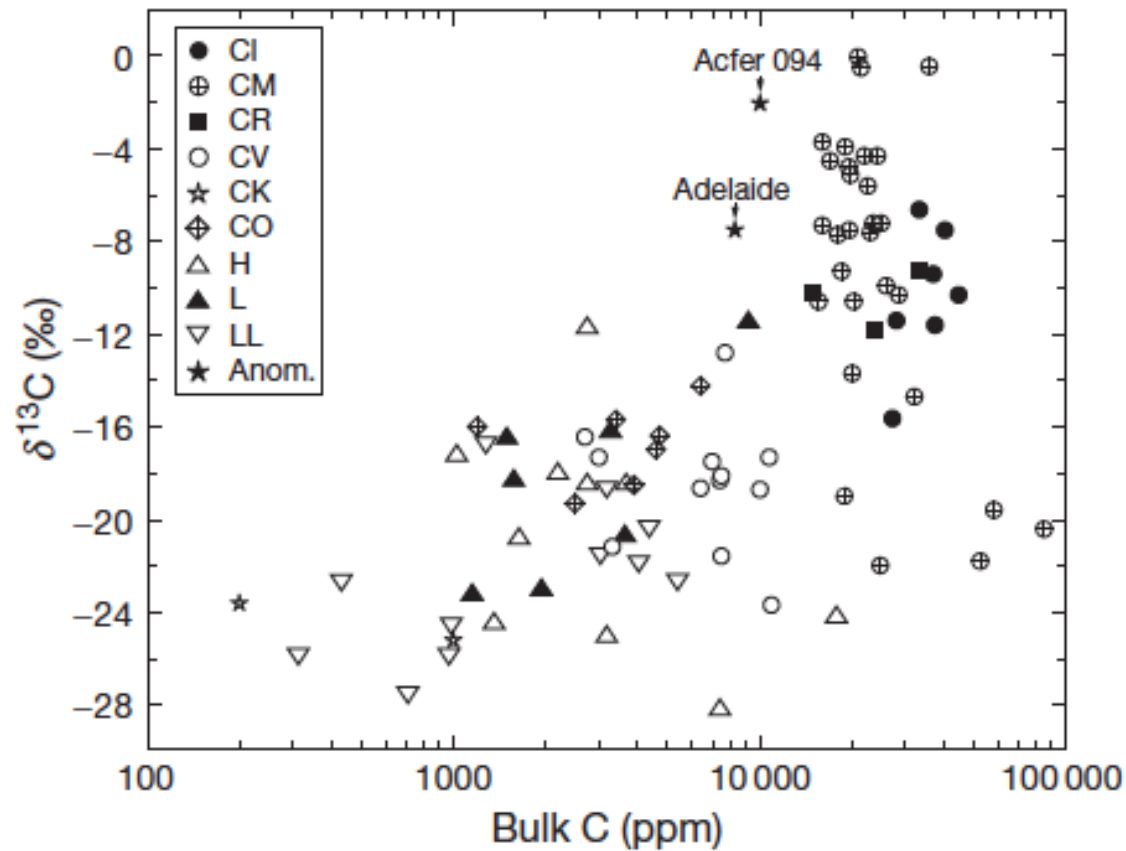
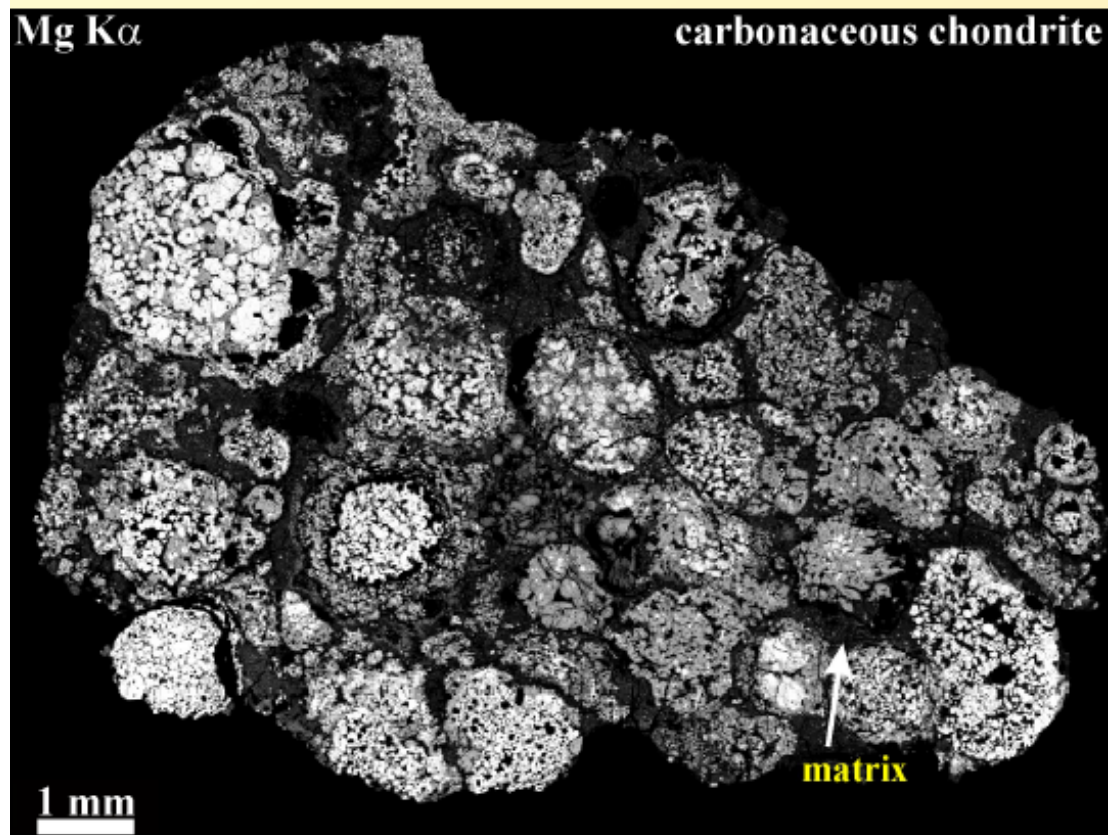


Figure 5 Bulk carbon contents (ppm) versus $\delta^{13}\text{C}$ (‰) in chondrites (Sources: Grady and Pillinger, 1986; Kerridge, 1985).

Krot et al (2014a ToG)

Matrix material

- Mixture of grains 10 nm - 5 μm in size with \sim solar bulk composition which rims chondrules, CAIs, metal grains, etc.
- Only 4 out of 4,000 chondrites have pristine matrices (Acfer 094, ALHA77307, Adelaide, Kakangari).



- Pristine C chondrite matrices contain forsterite Mg_2SiO_4 , enstatite MgSiO_3 , and amorphous Fe-Mg-Si-O containing Fe-Ni metal or Fe-sulfides with admixture of carbonaceous material, refractory grains, and stardust ($\sim 0.001\%$).
- Complex mixture of grains that formed at different times and places; acquired when components accreted.

Comets and chondritic porous IDPs

- Chondritic porous IDPs closely resemble matrices of pristine carbonaceous chondrites.
- Both contain Mn-rich forsterite, rapidly cooled enstatite, and amorphous silicate material.
- Forsterite, enstatite, and amorphous silicate also present in comas of long-period comets: up to 30-50% crystalline (Wooden et al., 2004).
- Major fraction of silicate in comets formed in the disk at high temperatures, probably as condensates.
- Thermally processed dust was ubiquitous in the disk.



Ed Scott 2004

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Cosmochemical Classification of the Elements

Table 1. *Cosmochemical classification of the elements*

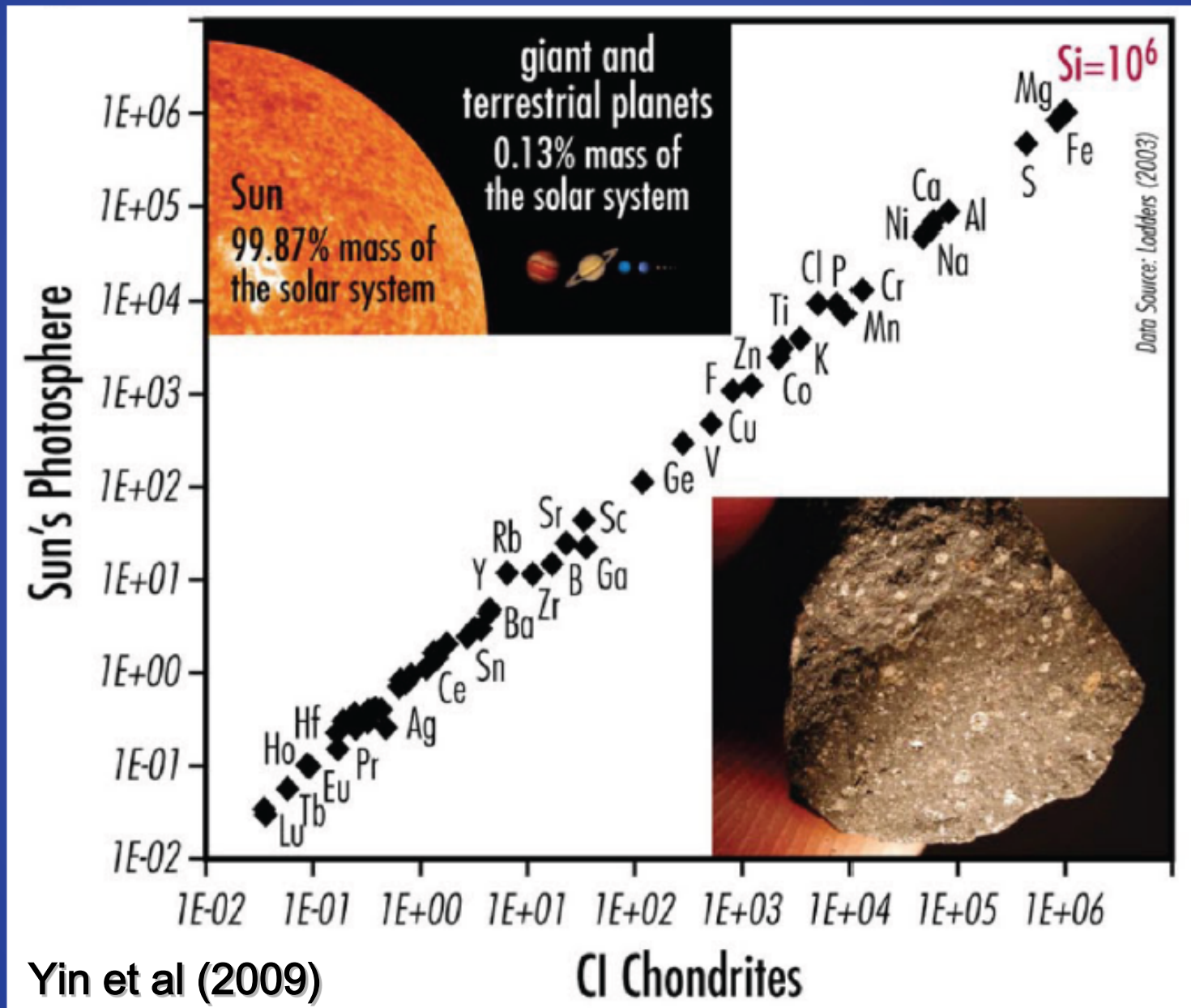
	elements	
	lithophile (silicate)	siderophile + chalcophile (sulphide + metal)
refractory	$T_c = 1850\text{--}1400\text{ K}$ Al, Ca, Ti, Be, Ba, Sc, V, Sr, Y, Zr, Nb, Ba, REE, Hf, Ta, Th, U, Pu	Mo, Ru, W, Re, Os, Ir, Pt
main component	$T_c = 1350\text{--}1250\text{ K}$ Mg, Si, Cr, Li	Fe, Ni, Co, Pd
moderately volatile	$T_c = 1230\text{--}640\text{ K}$ Mn, P, Na, Rb, K, F, Zn	Au, Cu, Ag, Ga, Sb, Ge, Sn, Se, Te, S
highly volatile	$T_c < 640\text{ K}$ B, Cl, Br, J, Cs, Tl, C, O, Ne, Ar, Kr, Xe	In, Bi, Pb, Hg, H

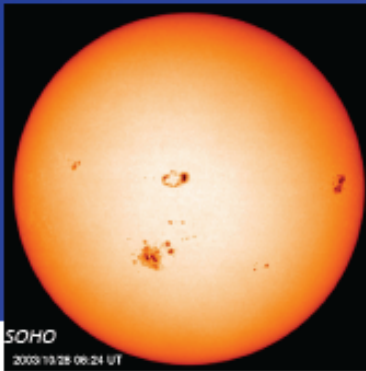
T_c denotes condensation temperatures at a pressure of 10^{-4} bar.

H. Palme (2001)

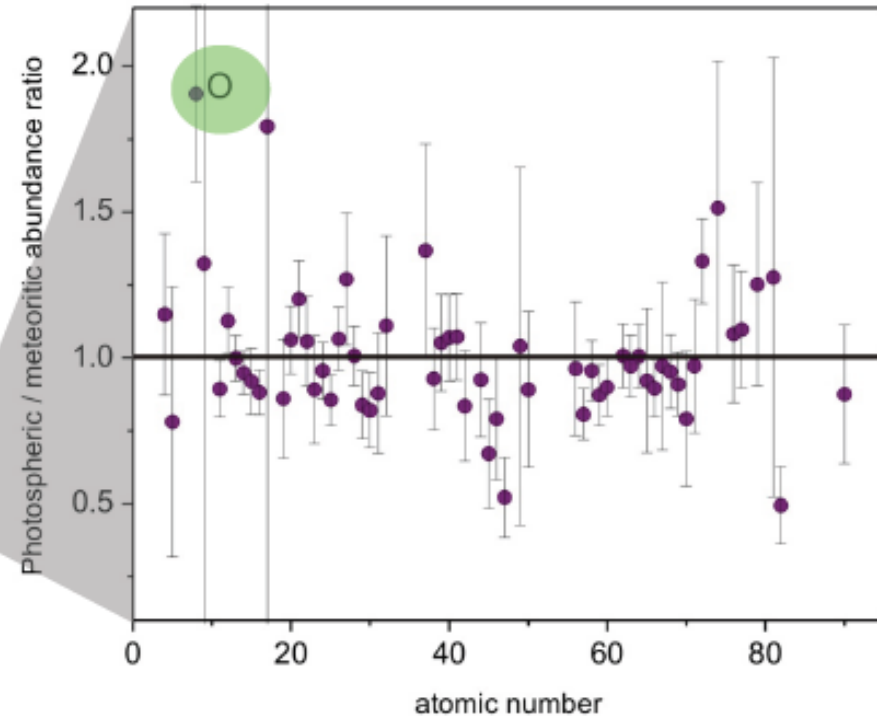
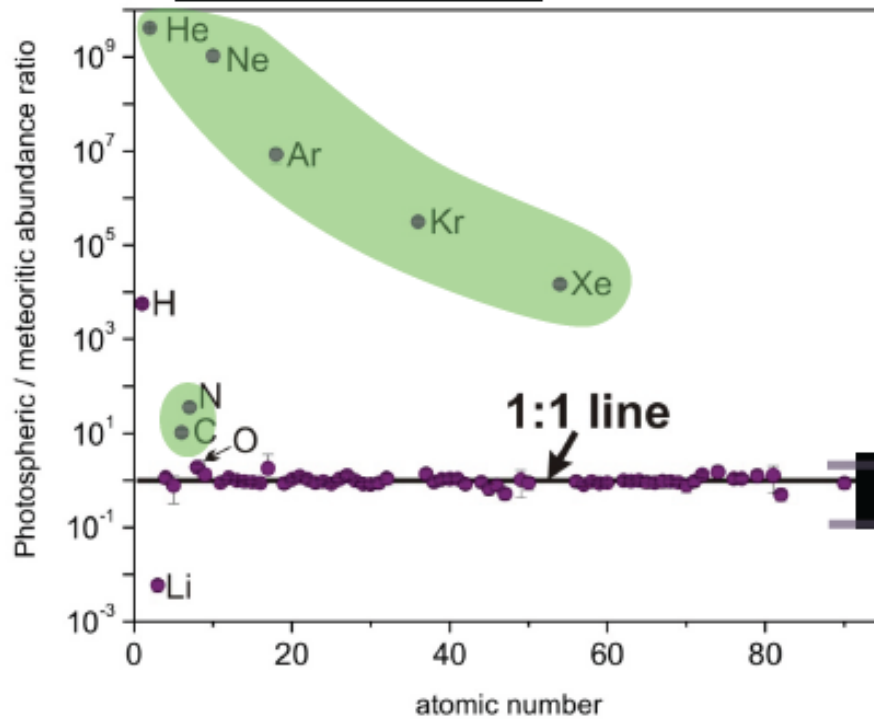
Chondrite (CHUR)=Bulk Solar System

2E30 Kg





Sun's Photosphere (today)
 CI chondrite (4.56 years old)



Asplund et al. 2009
 Lodders et al. 2009

All subsequent processing is recorded in the deviation from this correlation

Courtesy of K. McKeegan

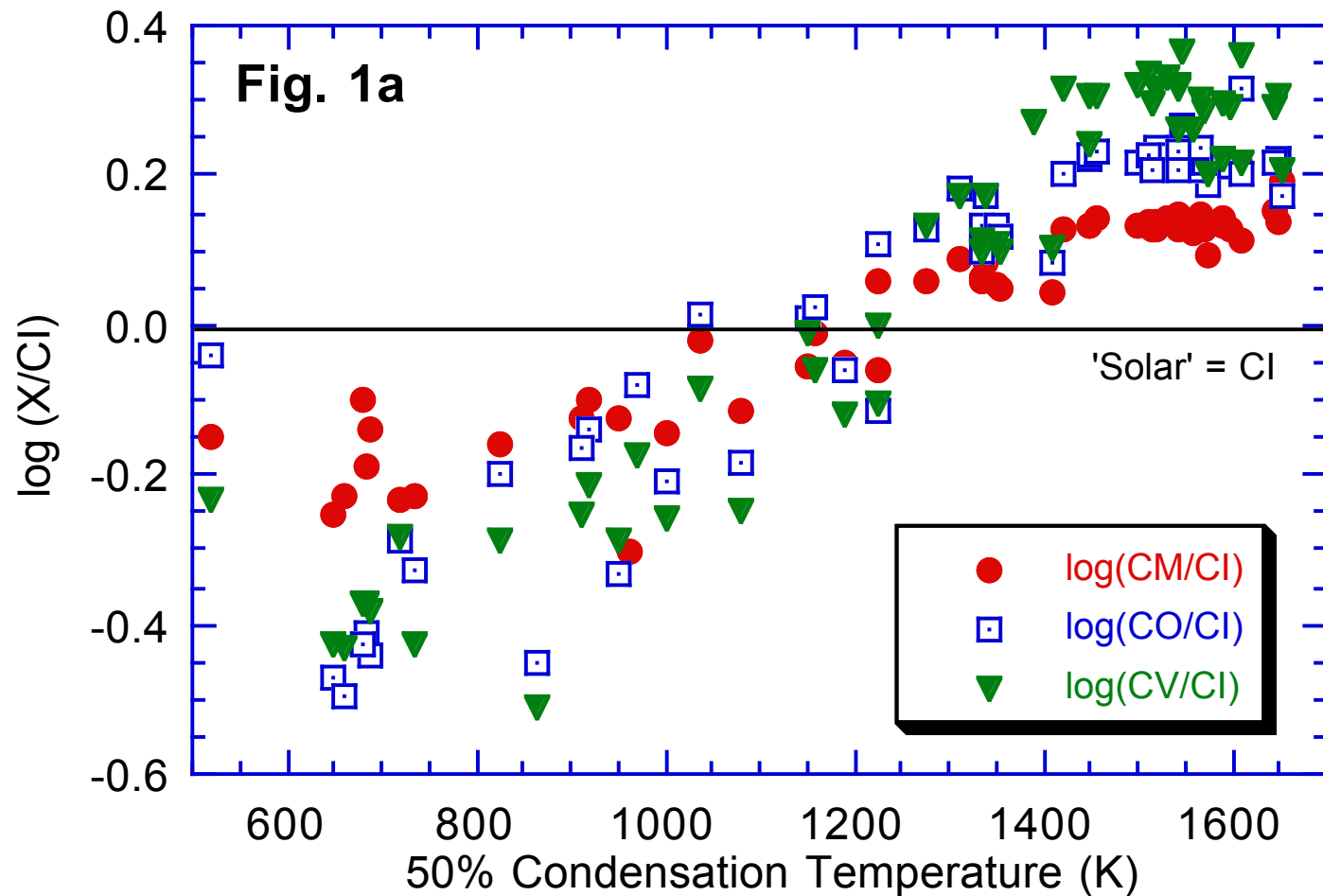
Where have all the volatiles gone?

Inner Solar System is fractionated
relative to bulk Solar System

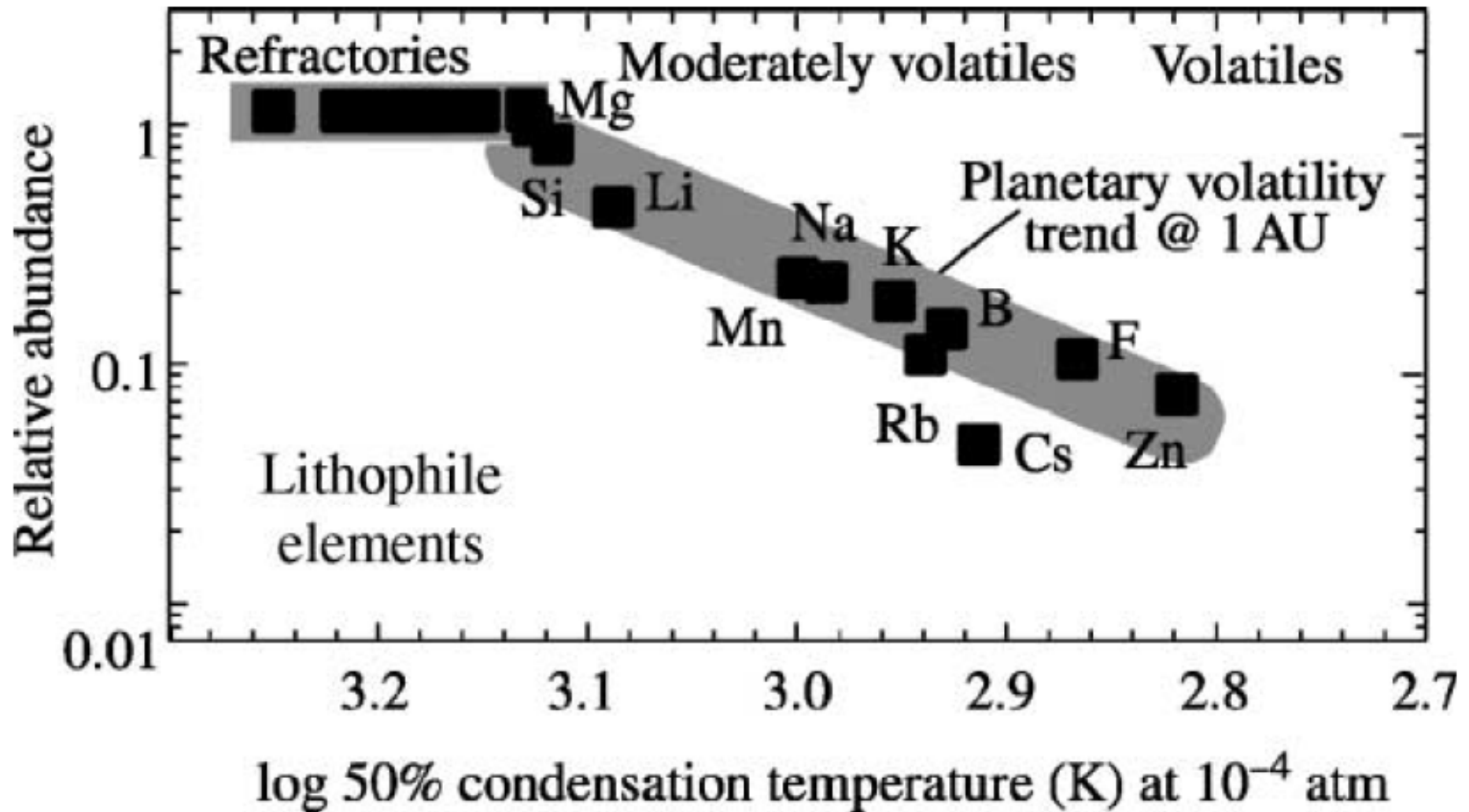
Refractory
enriched

Volatiles
depleted

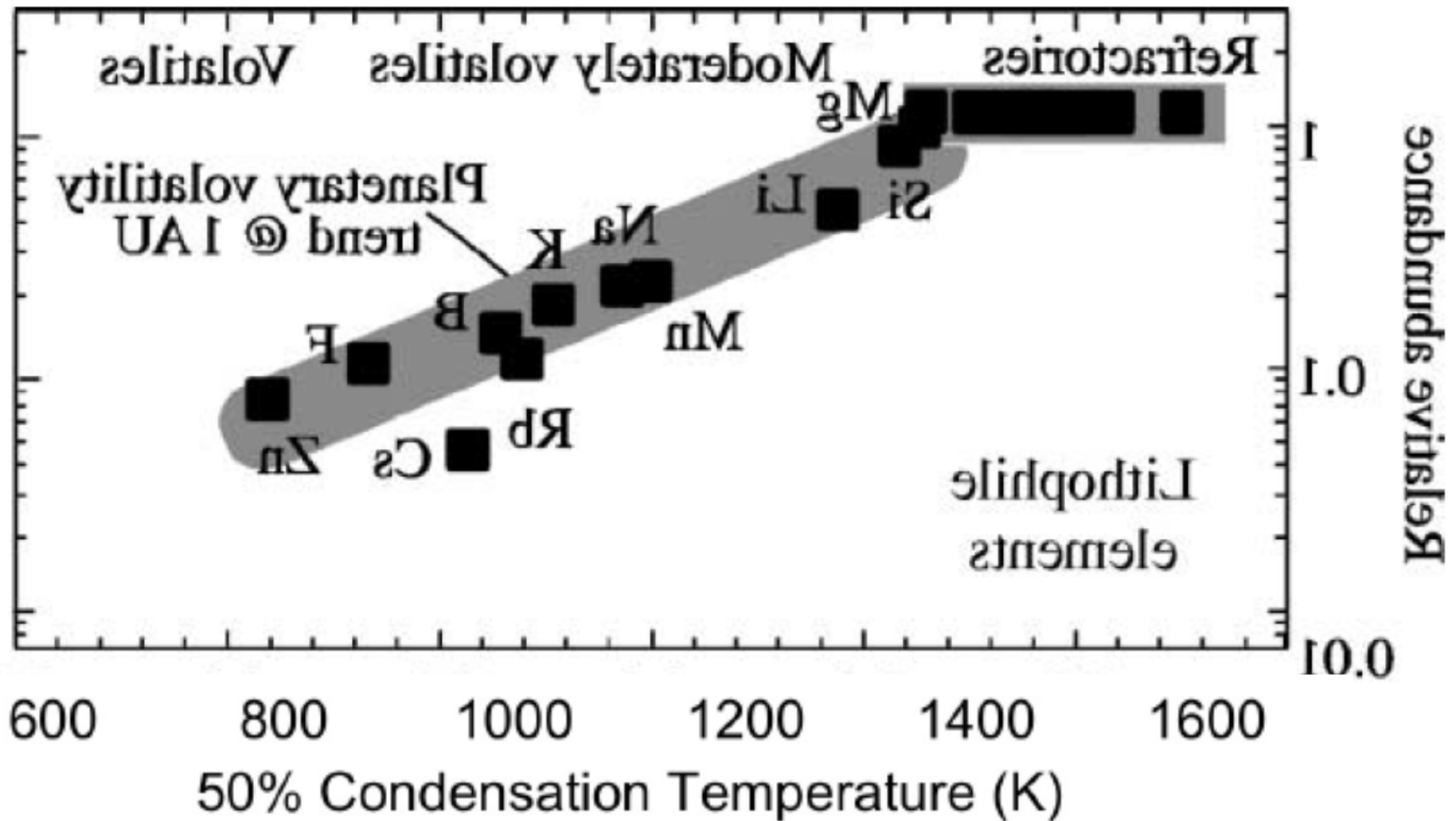
Yin (2005)

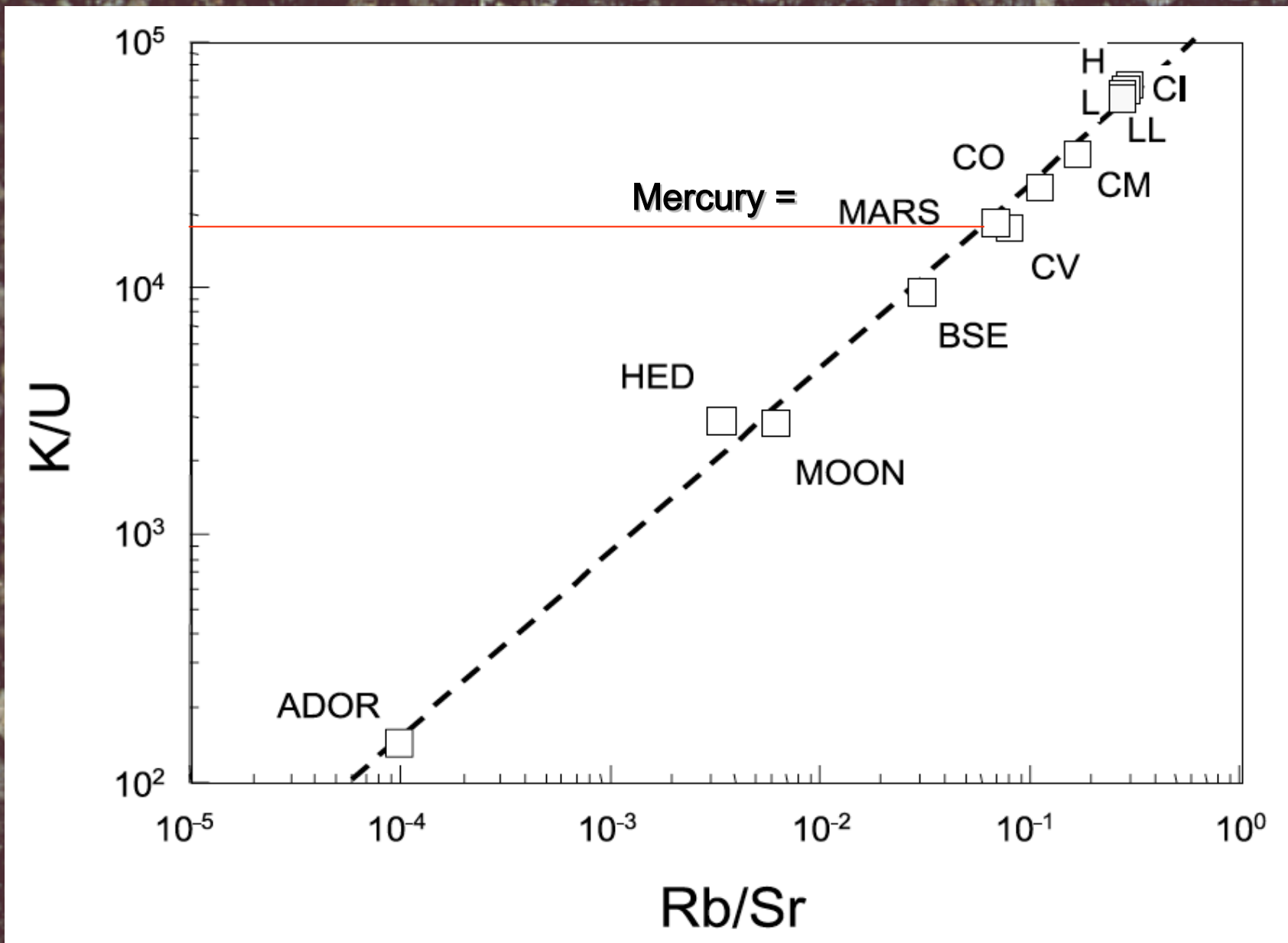


Silicate Earth (McDonough 2003)



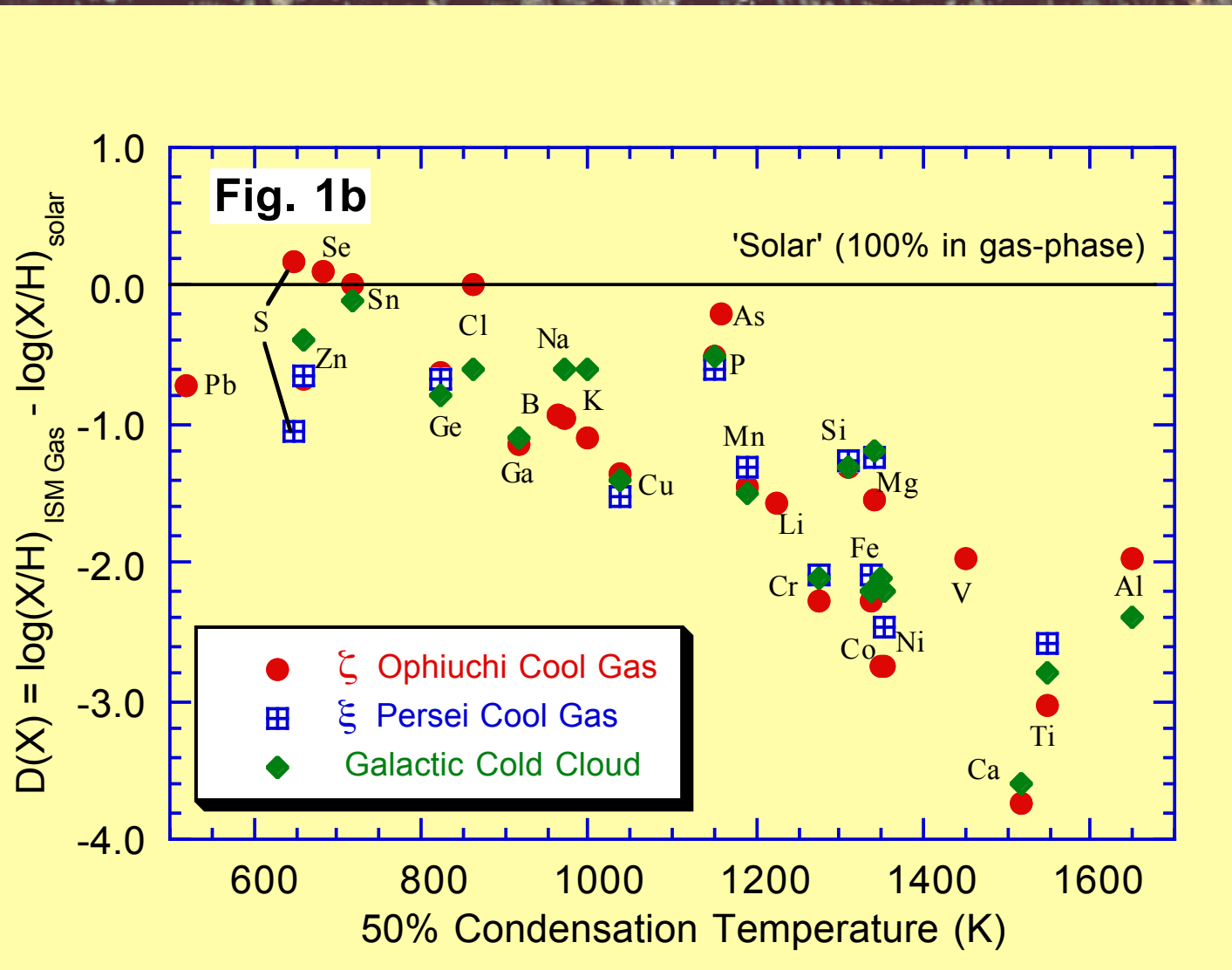
Silicate Earth (McDonough 2003)





Halliday and Porcelli (2001)

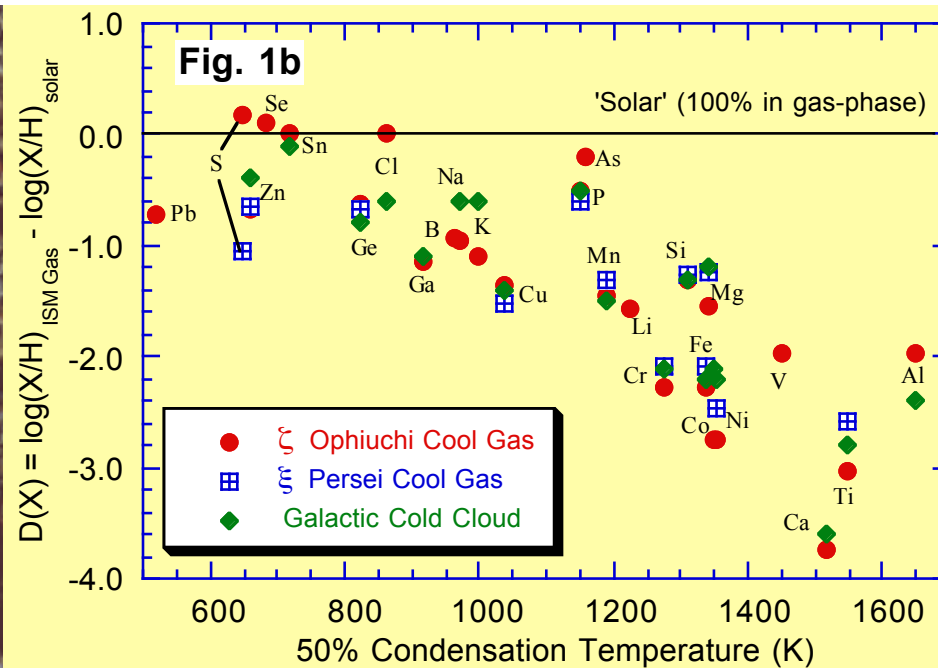
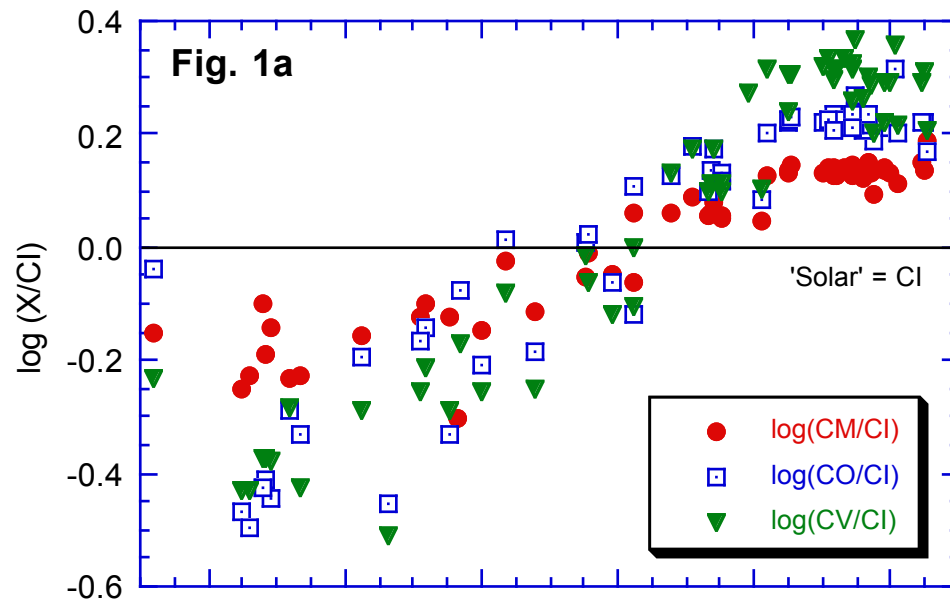
ISM Gas Phase Composition



Data Source: Savage and Sembach (1996); Welty et al.(1999)

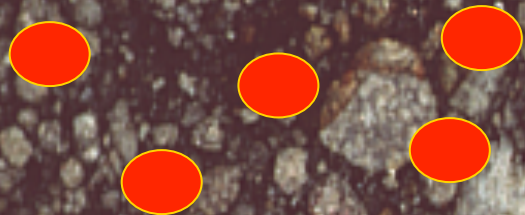
For moderately volatile elements, interstellar gas phase data is “mirror-imaged” by the meteorite data.

Yin (2005)

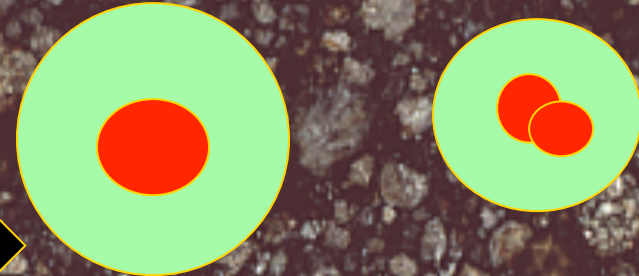


Inheritance Model

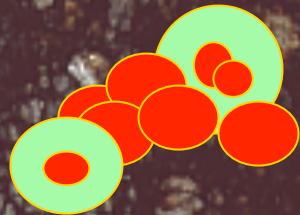
1. Interstellar Stage:



2. Molecular Cloud Stage:

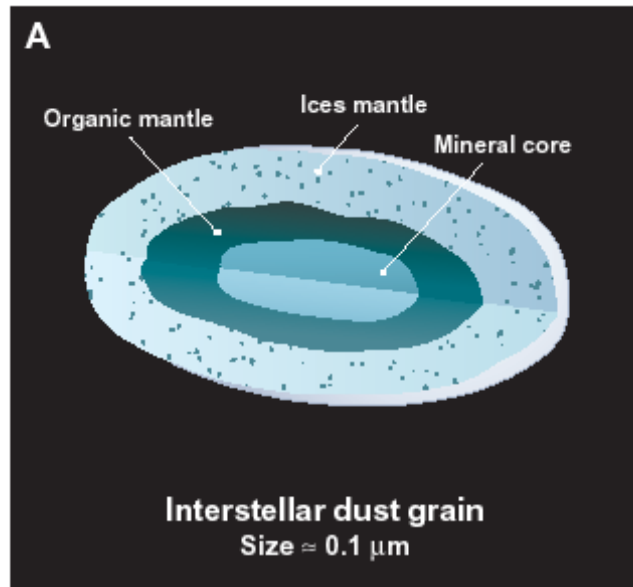


3. Solar Nebula Stage:

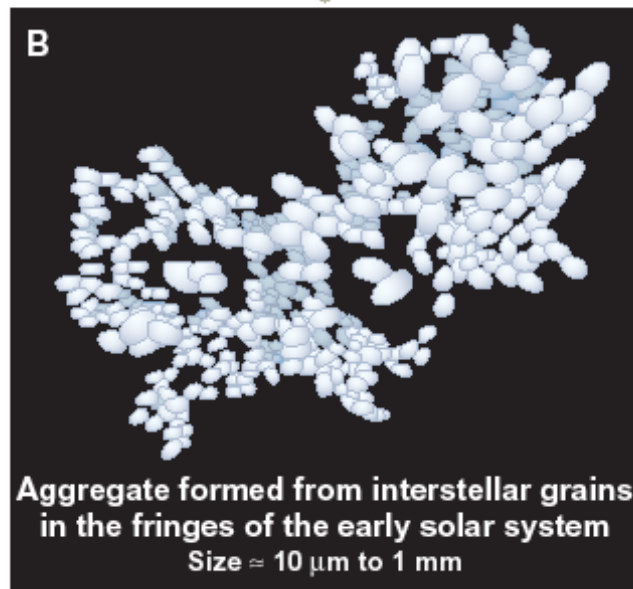


Red: Refractory Grains

Green: Ice mantle with volatiles



Molecular clouds formed



**Levasseur-Regourd
(2004)**

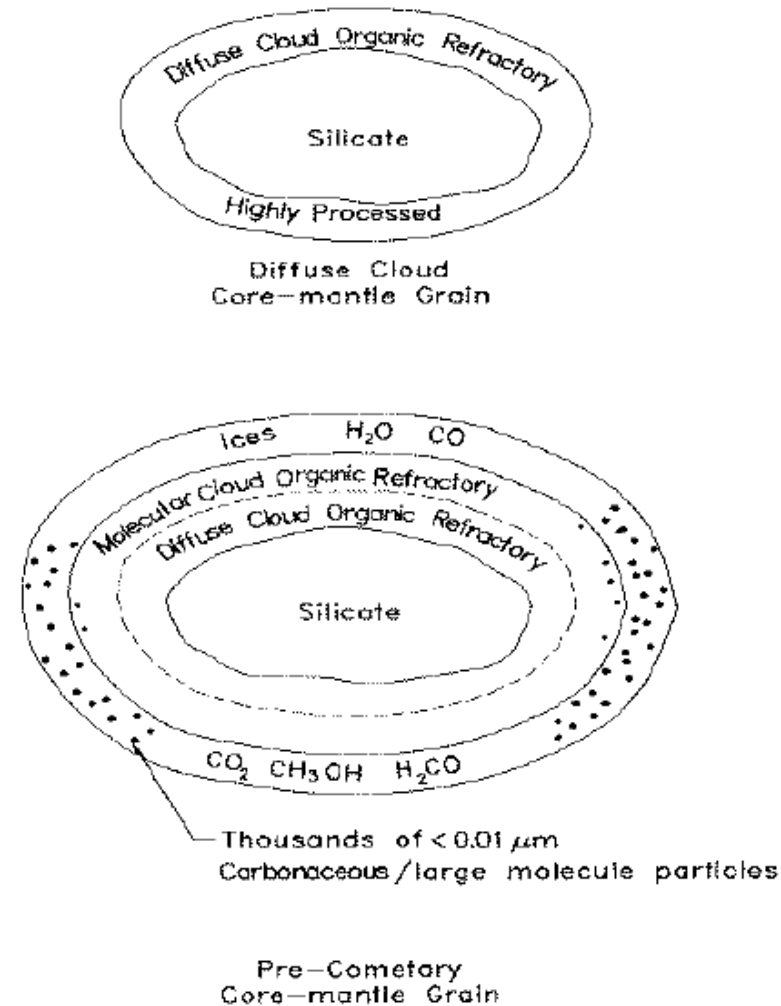
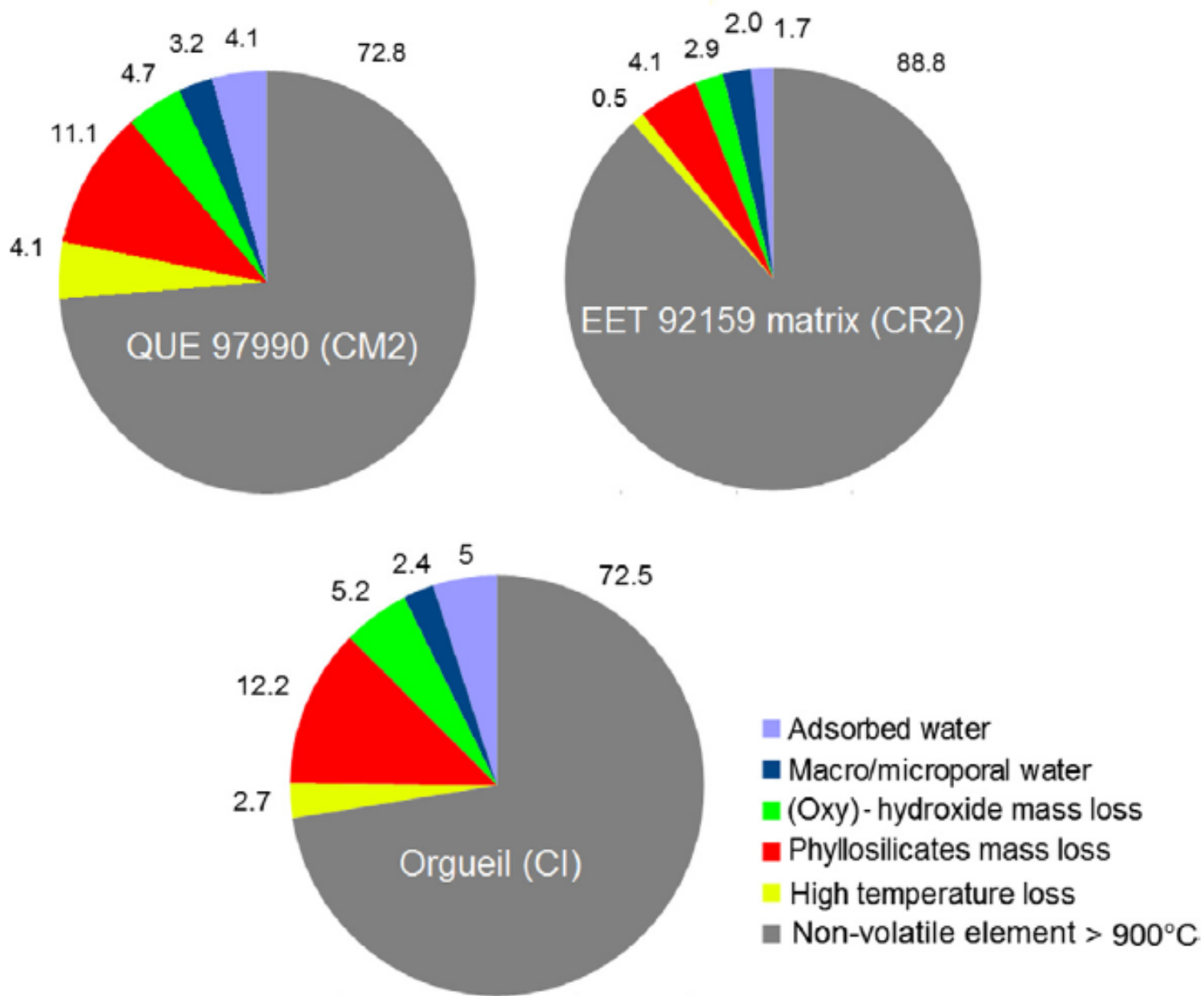


Fig. 1. A schematic description of the morphological and chemical structure of core-mantle interstellar dust grains in diffuse cloud regions and in the latest stage of the collapse of an interstellar cloud. The ices are both accreted and created along with the molecular cloud organics. The very small particle/large molecule components of the interstellar dust accrete along with the ices in the dense cloud.

**J. M. Greenberg
(1998)**



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Altwegg et al (2015)

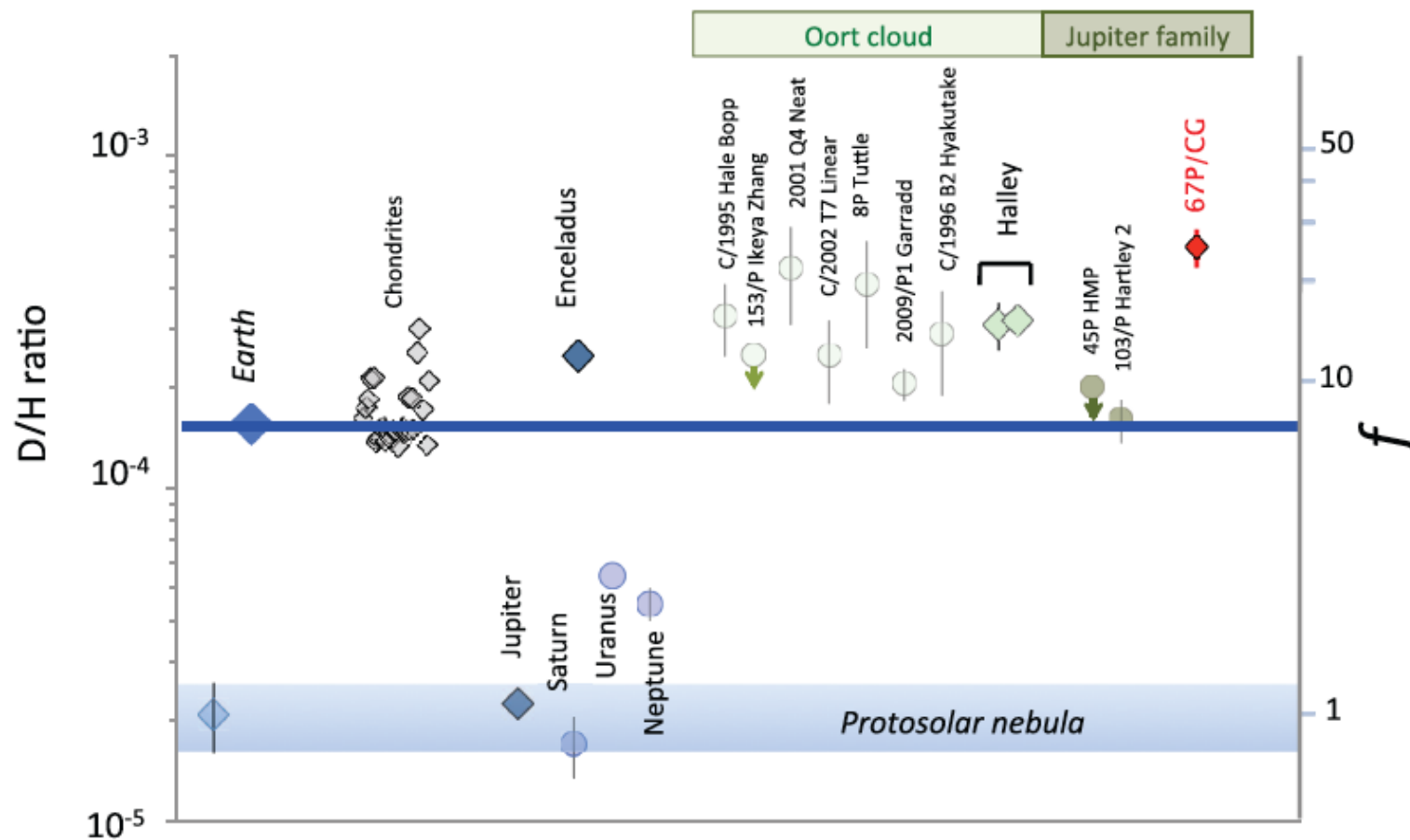
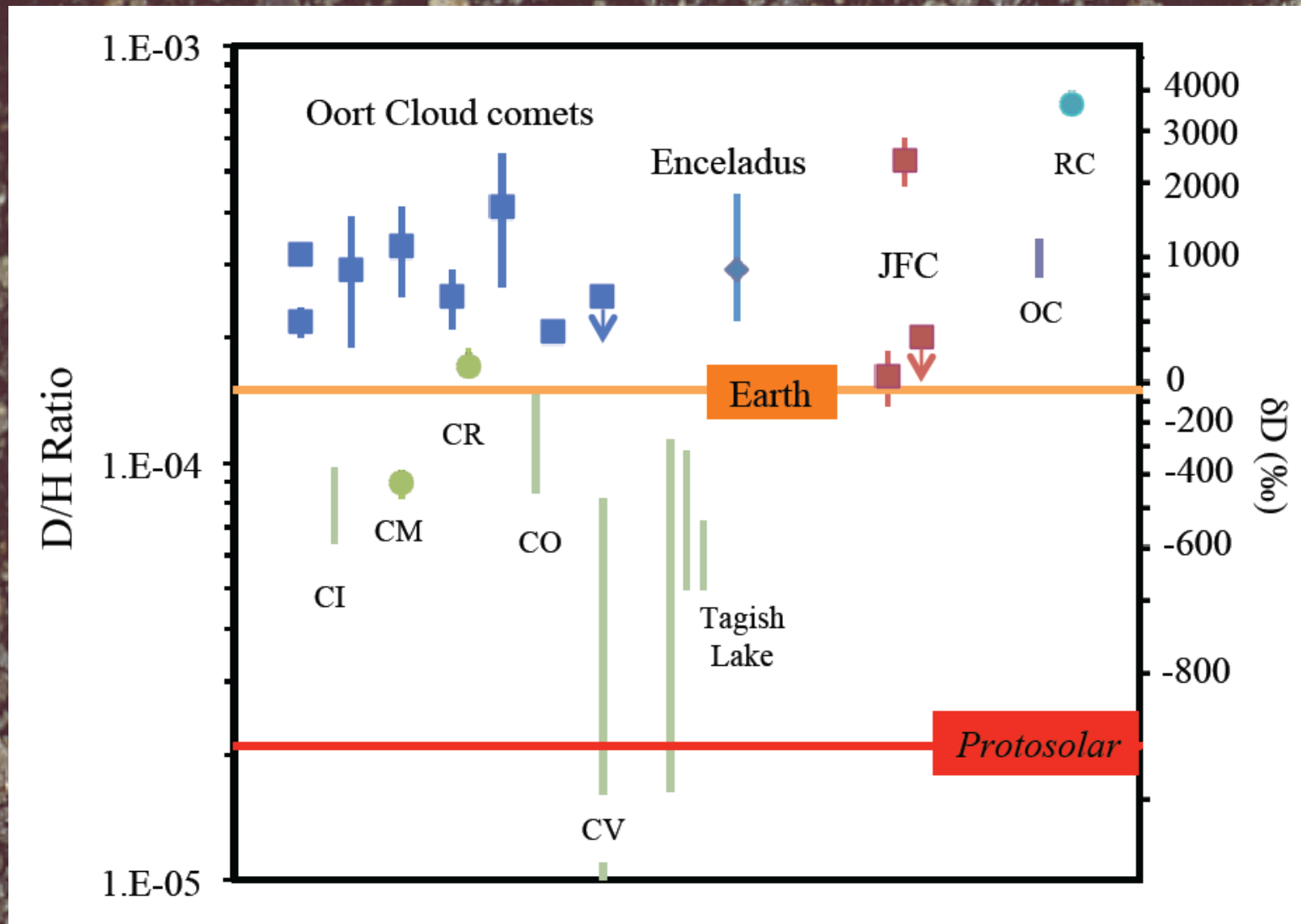


Fig. 3. D/H ratios in different objects of the solar system. Data are from (1, 2, 5–7, 26–28) and references therein. Diamonds represent data obtained by means of in situ mass spectrometry measurements, and circles refer to data obtained with astronomical methods.

Alexander et al (2012)



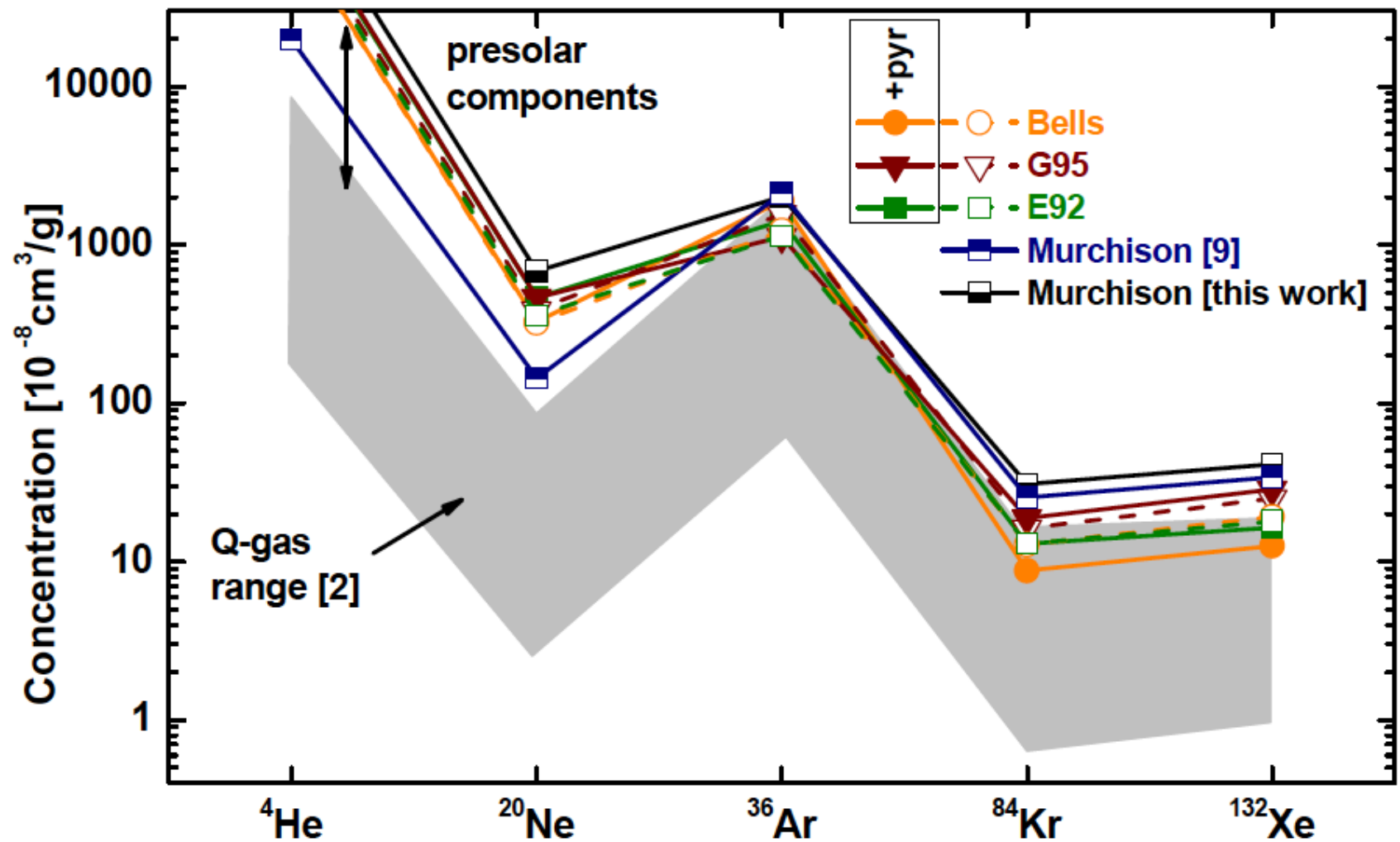


Fig. 1: Gas concentrations in E92, G95, Bells and Murchison IOM before and after pyridine treatment.

I-Pu-Xe Systematics and Missing Xe

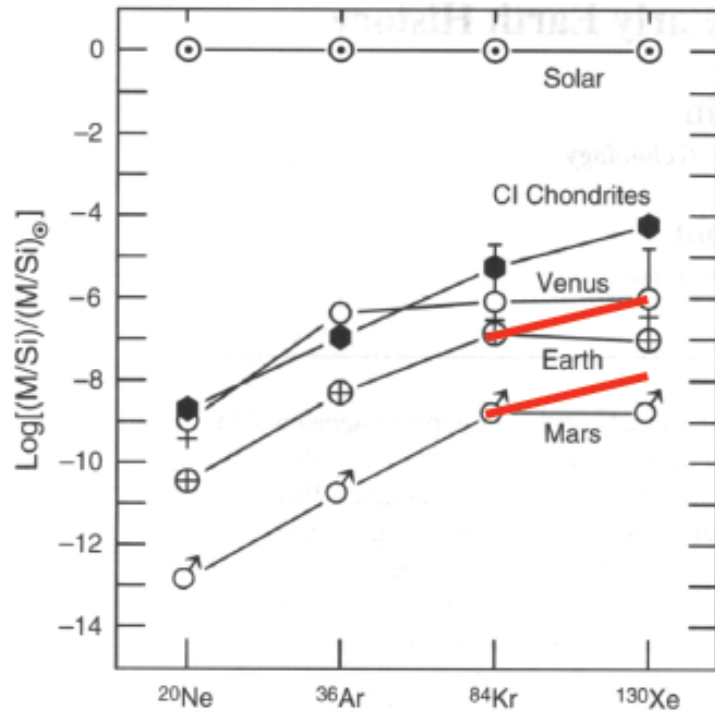
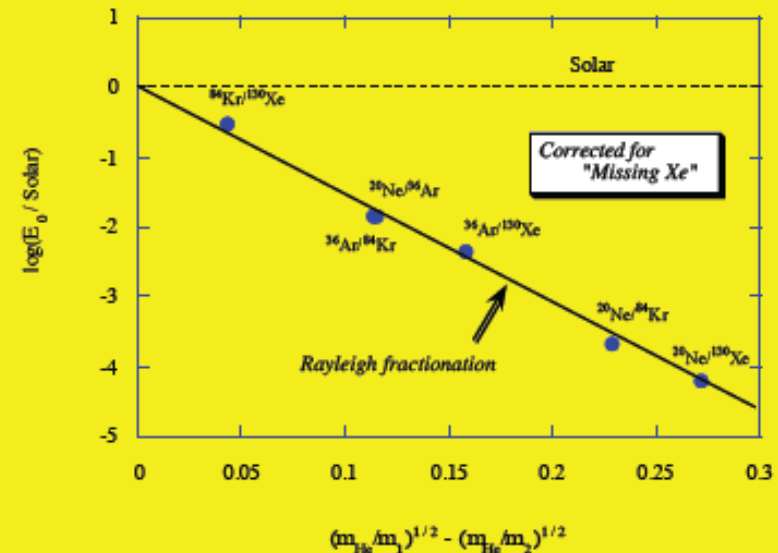
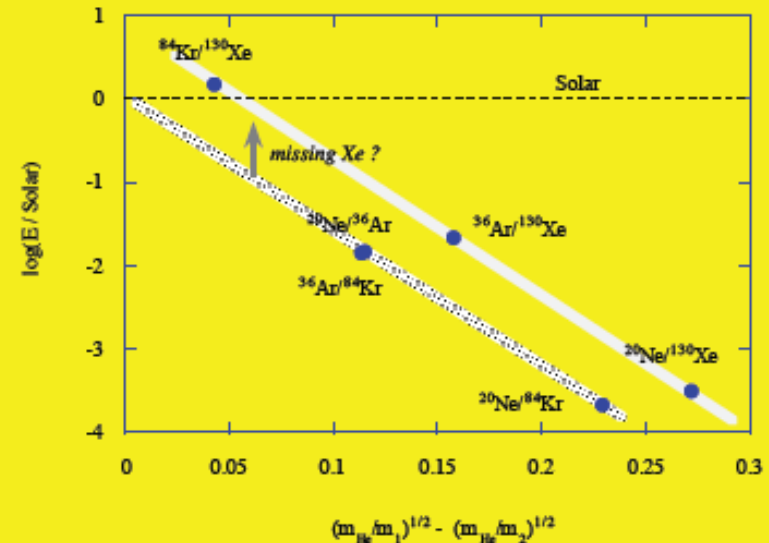


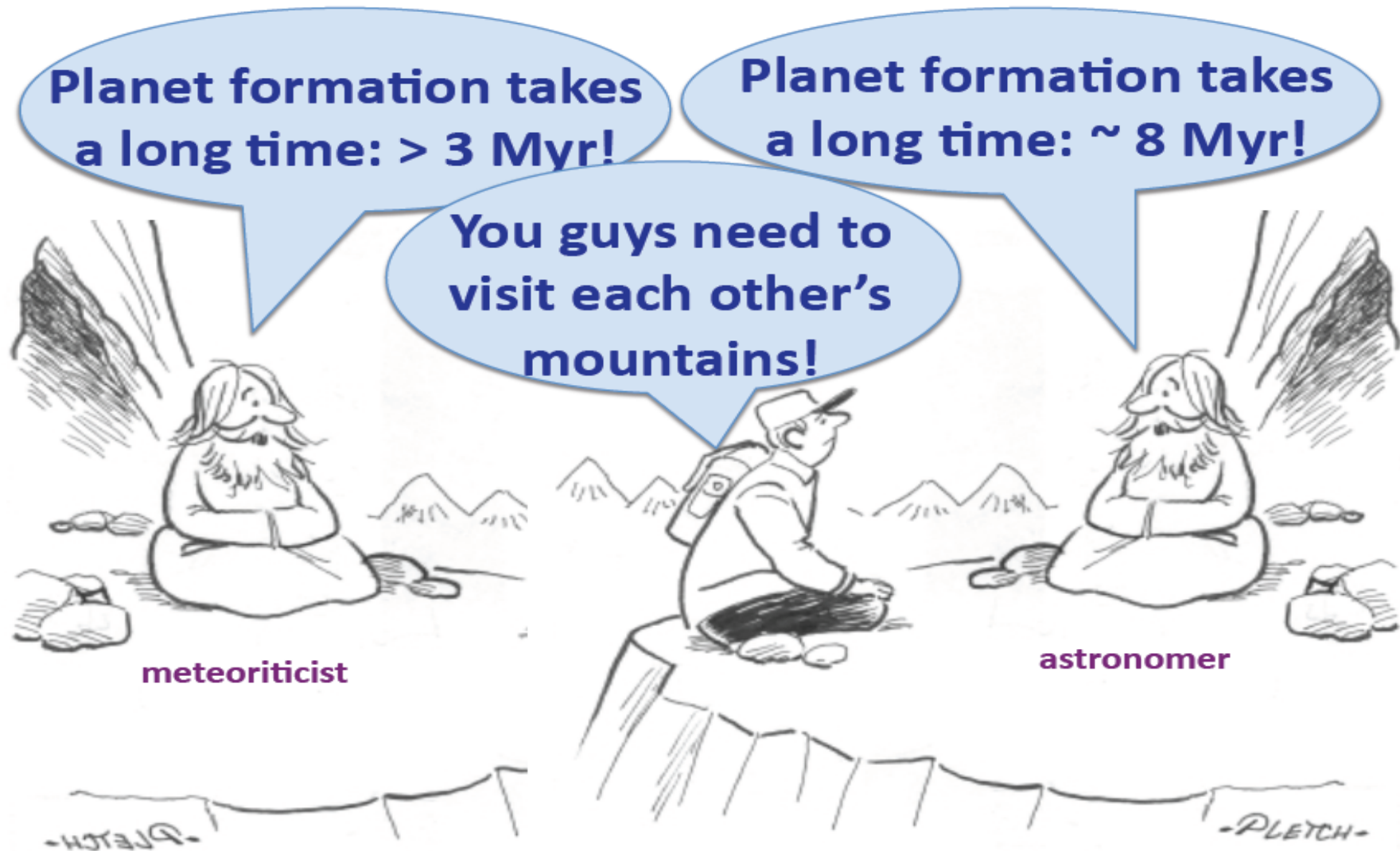
Fig. 1. Noble gas abundances in planetary atmospheres and CI chondrites, plotted as the atom concentration relative to Si divided by the corresponding solar ratio. Data are from a compilation by *Pepin (1991)*. Note that ranges of Kr and Xe values are shown for Venus.

Qualitatively, missing Xe of about one log unit is discernible. ca. 90%



For mass-dependent Rayleigh distillation, all the points should fall on a straight line through the origin.
 Ozima and Podosek (1999): First quantitatively derived the missing Xe amount to be 86% based on the least square fit.

Conclusions



Courtesy of S. Desch via S. Stewart