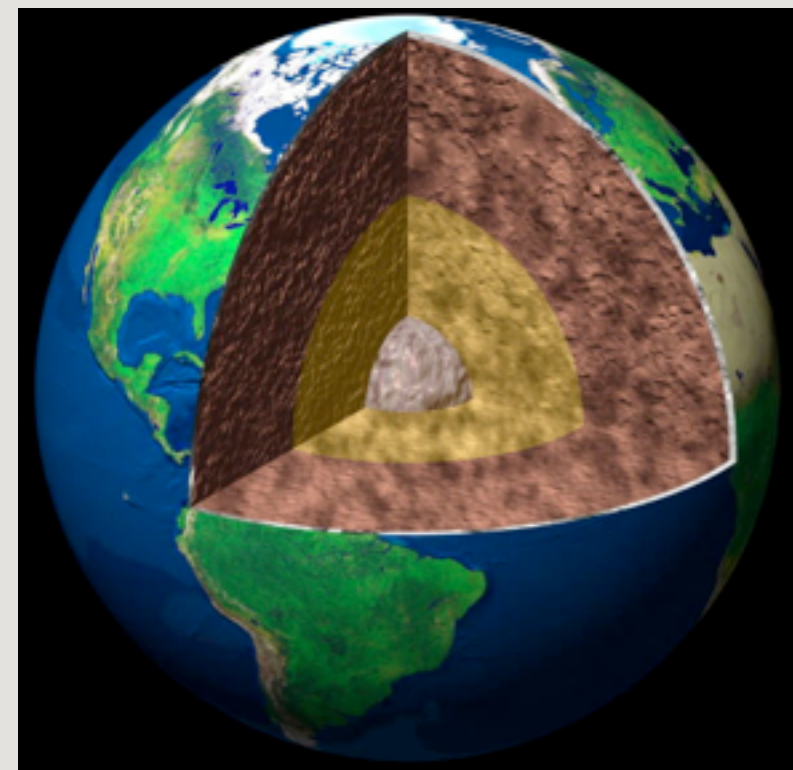
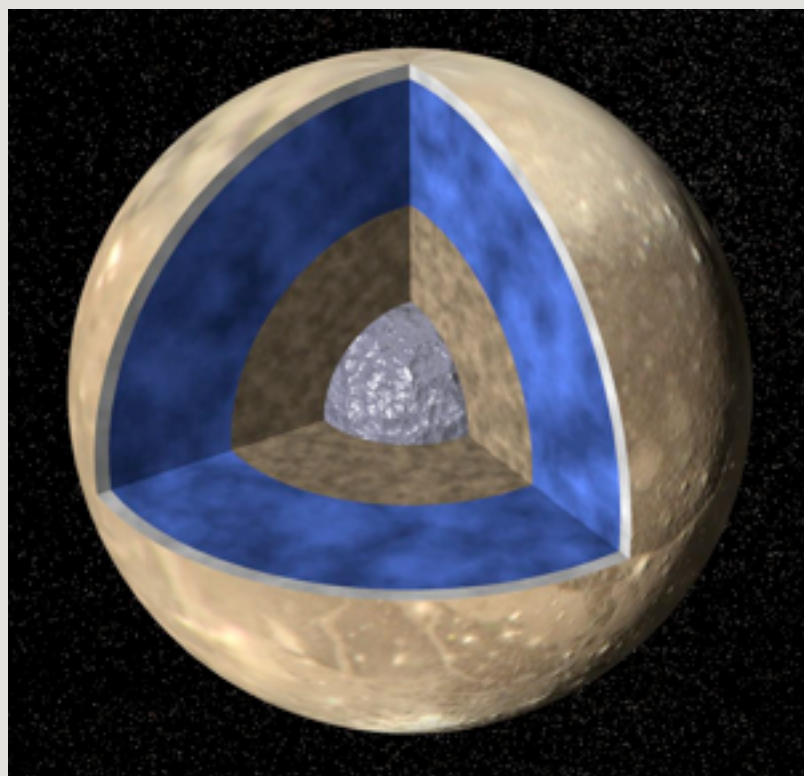
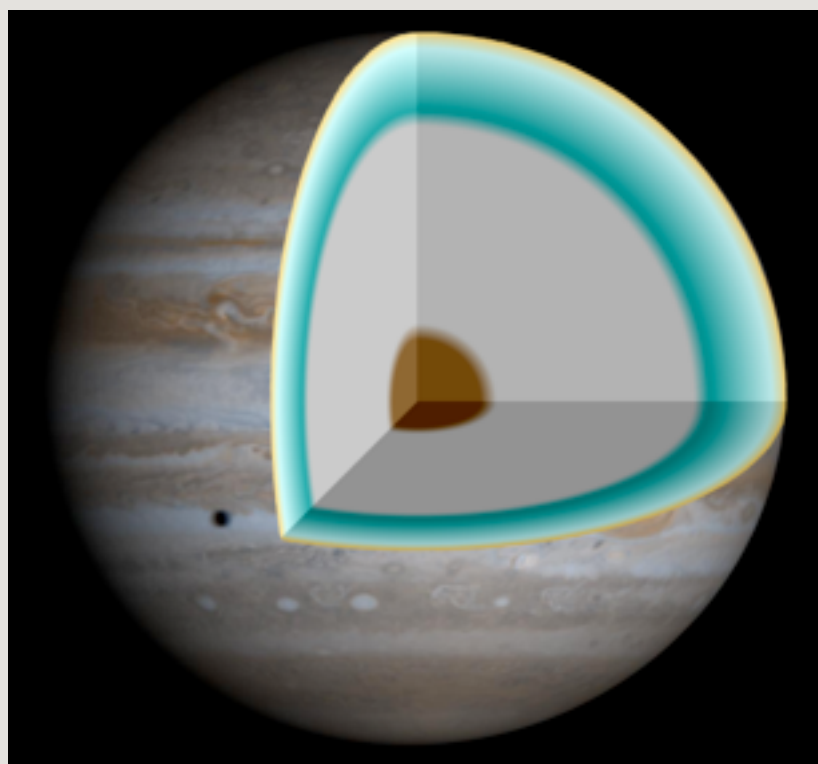




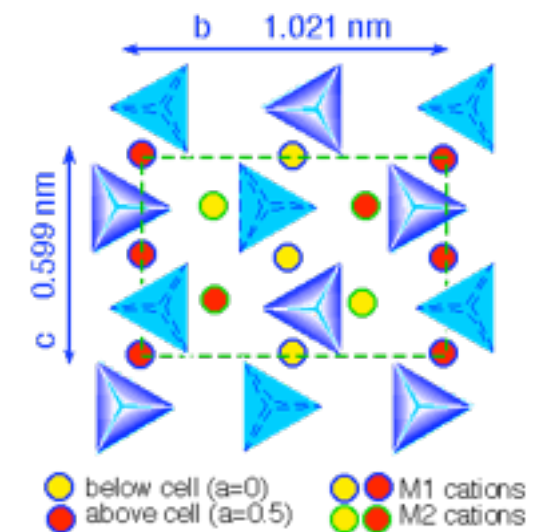
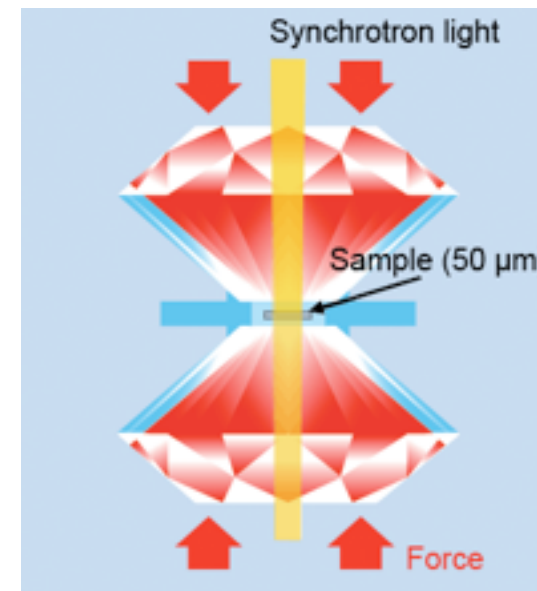
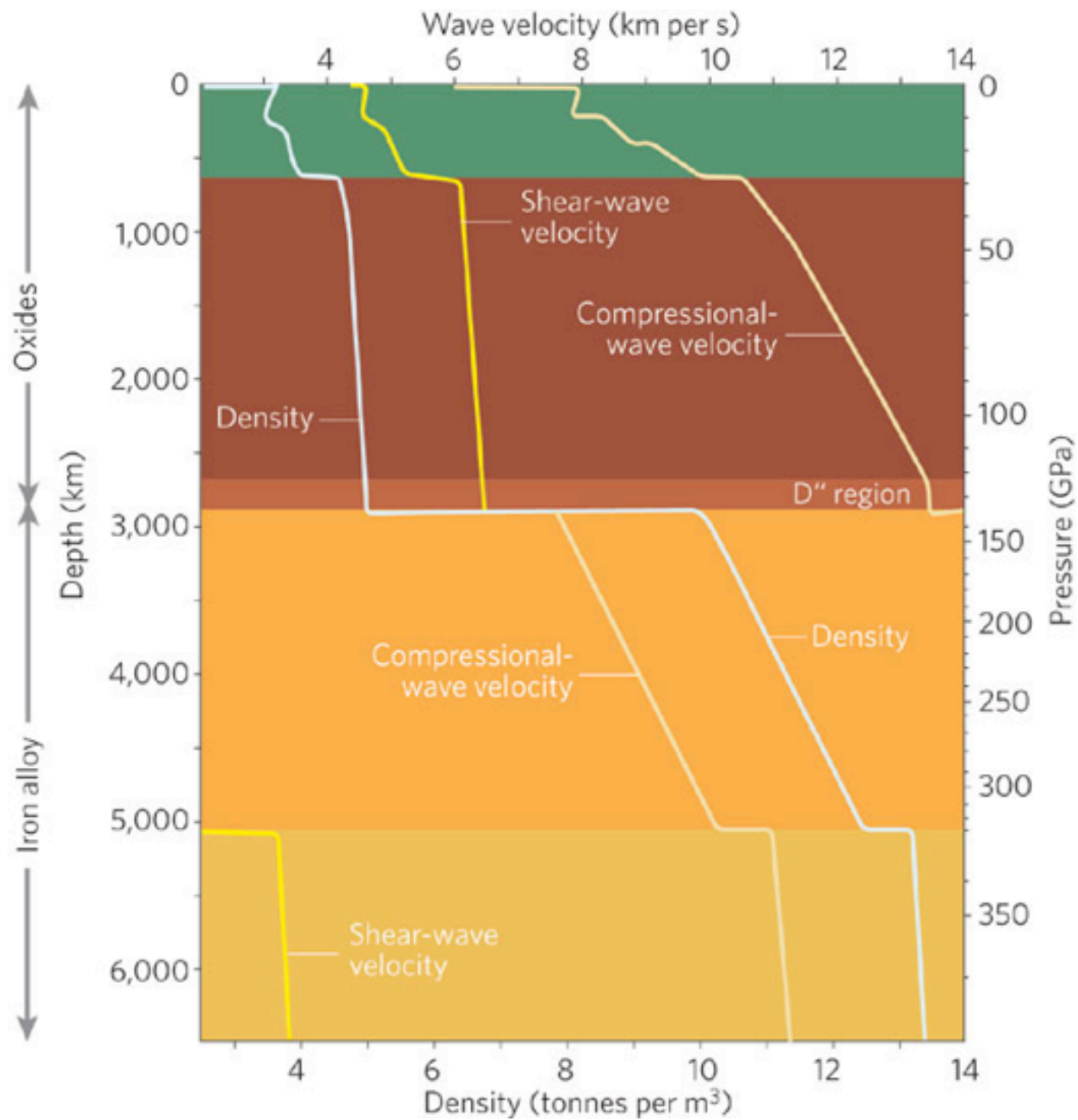
Observatoire  
de la CÔTE d'AZUR

# Terrestrial Exoplanet Radii, Structure and Tectonics

Diana Valencia (OCA), KITP, 31 March 2010



# Earth's Structure

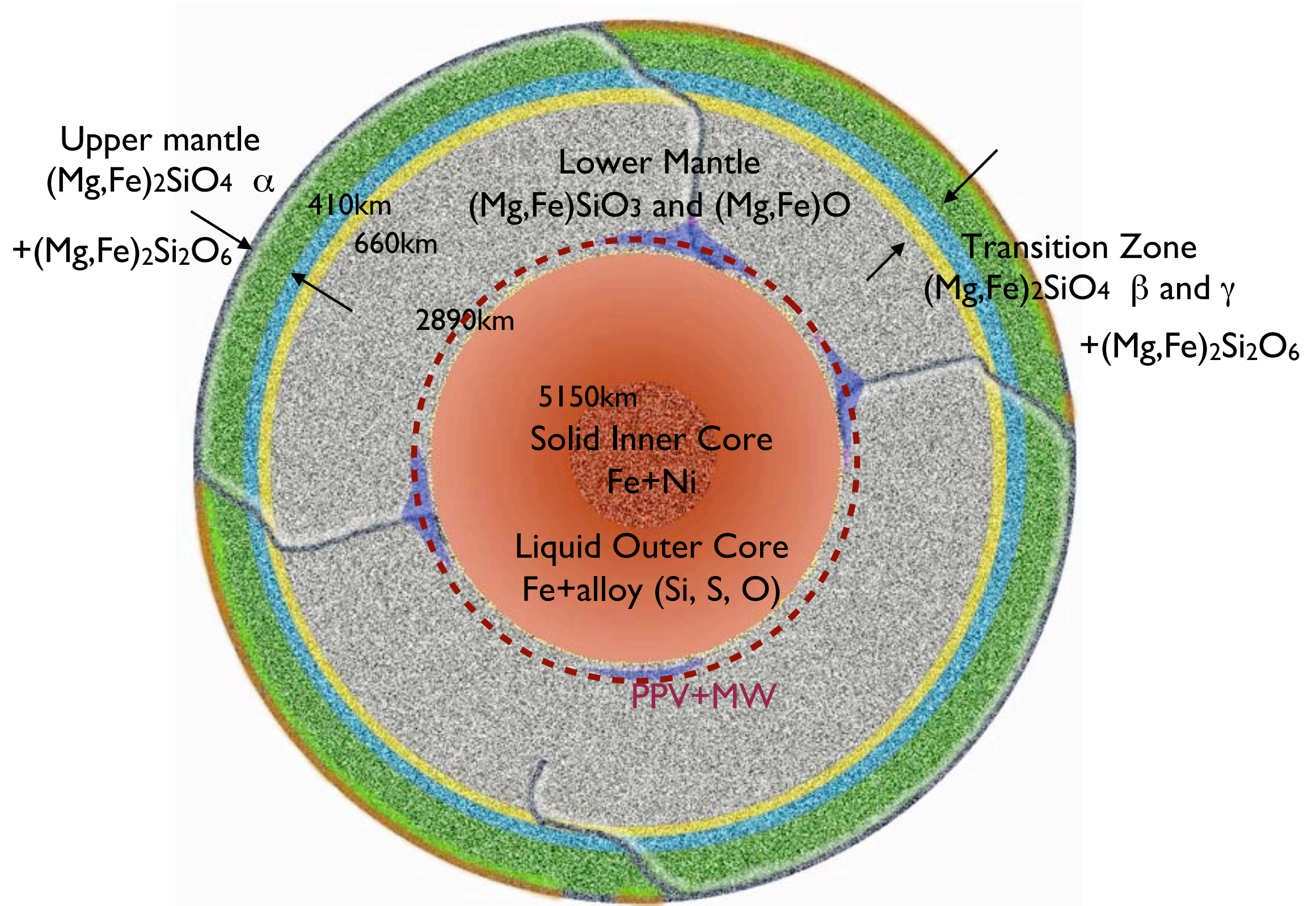


Olivine Structure

Romanowicz, 2008



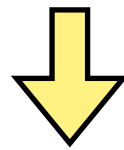
# Earth's Structure



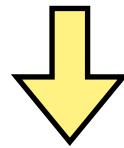
# Internal Structure Model

Planets made of Mg, Si, O, Fe, H<sub>2</sub>O, H/He

Input:  $M, P_{\text{surf}}, T_{\text{surf}}, \text{guess } R, g_{\text{surf}}, \chi$



$$\begin{aligned} \frac{d\rho}{dr} &= -\frac{\rho(r)g(r)}{\phi(r)} & \frac{dP}{dr} &= -\rho(r)g(r) \\ \frac{dg}{dr} &= 4\pi G\rho(r) - 2G\frac{m(r)}{r^3} & \frac{dT}{dr} &= -\frac{q}{k} \quad \text{and} \quad \frac{dT}{dr} = -\frac{g(r)\gamma(r)}{\phi(r)}T \\ \frac{dm}{dr} &= 4\pi r^2\rho(r) & & + \text{EOS}(\chi) \end{aligned}$$



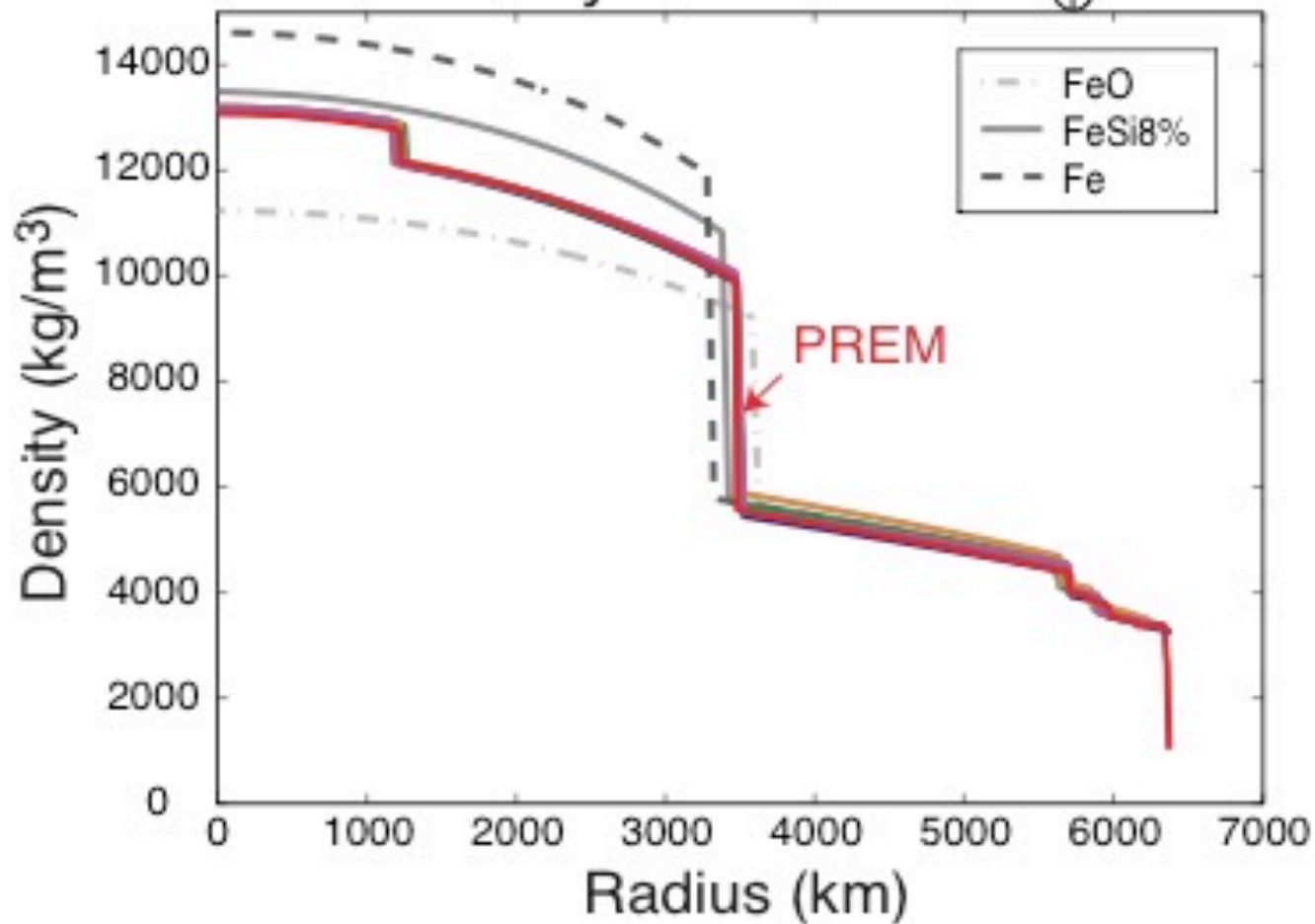
Output:  $R(M, \chi);$

$\rho(r), P(r), g(r), m(r), I/MR^2, D, \dots$

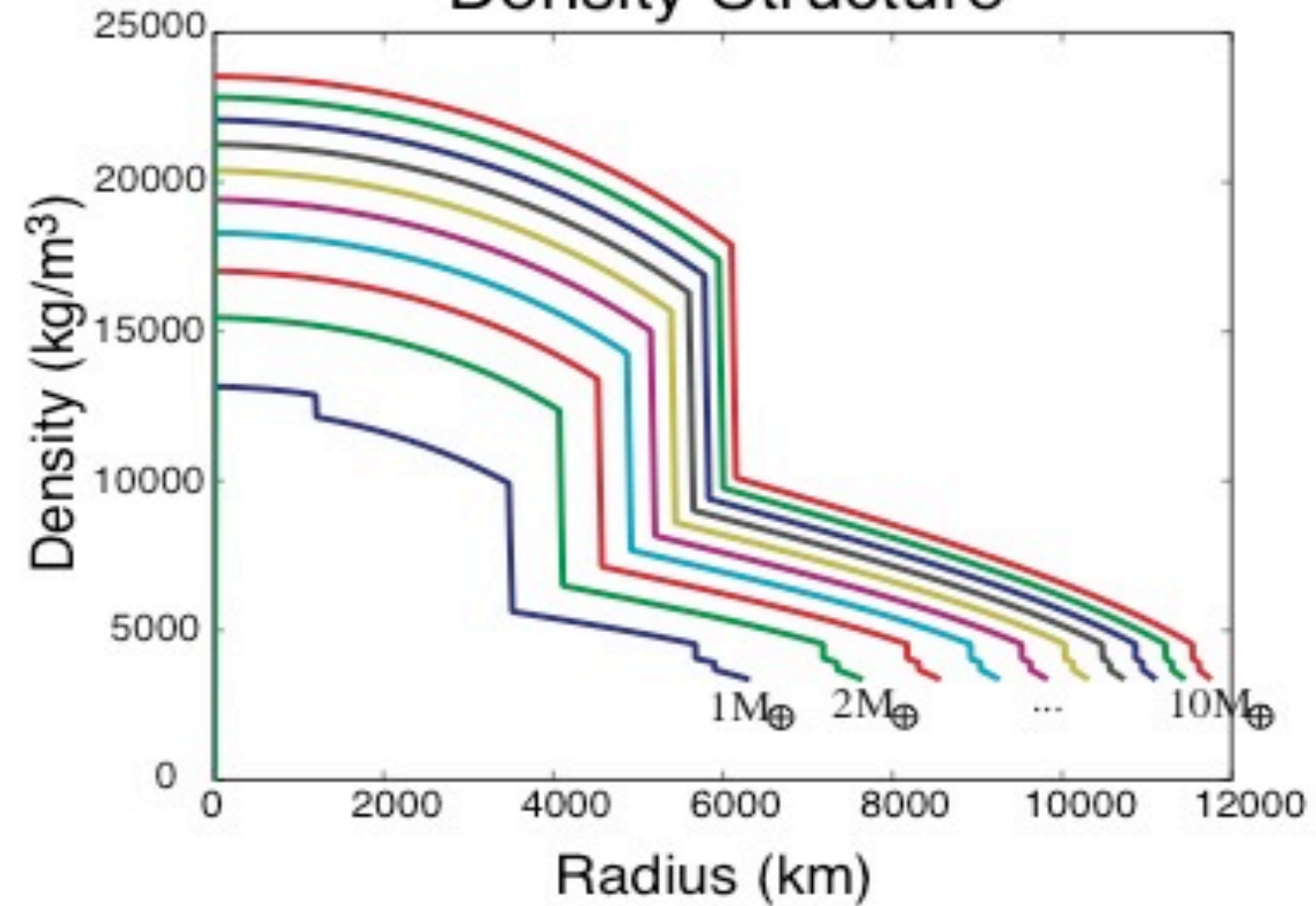


# Earth and rocky super-Earths

Density Profile for  $1M_{\oplus}$

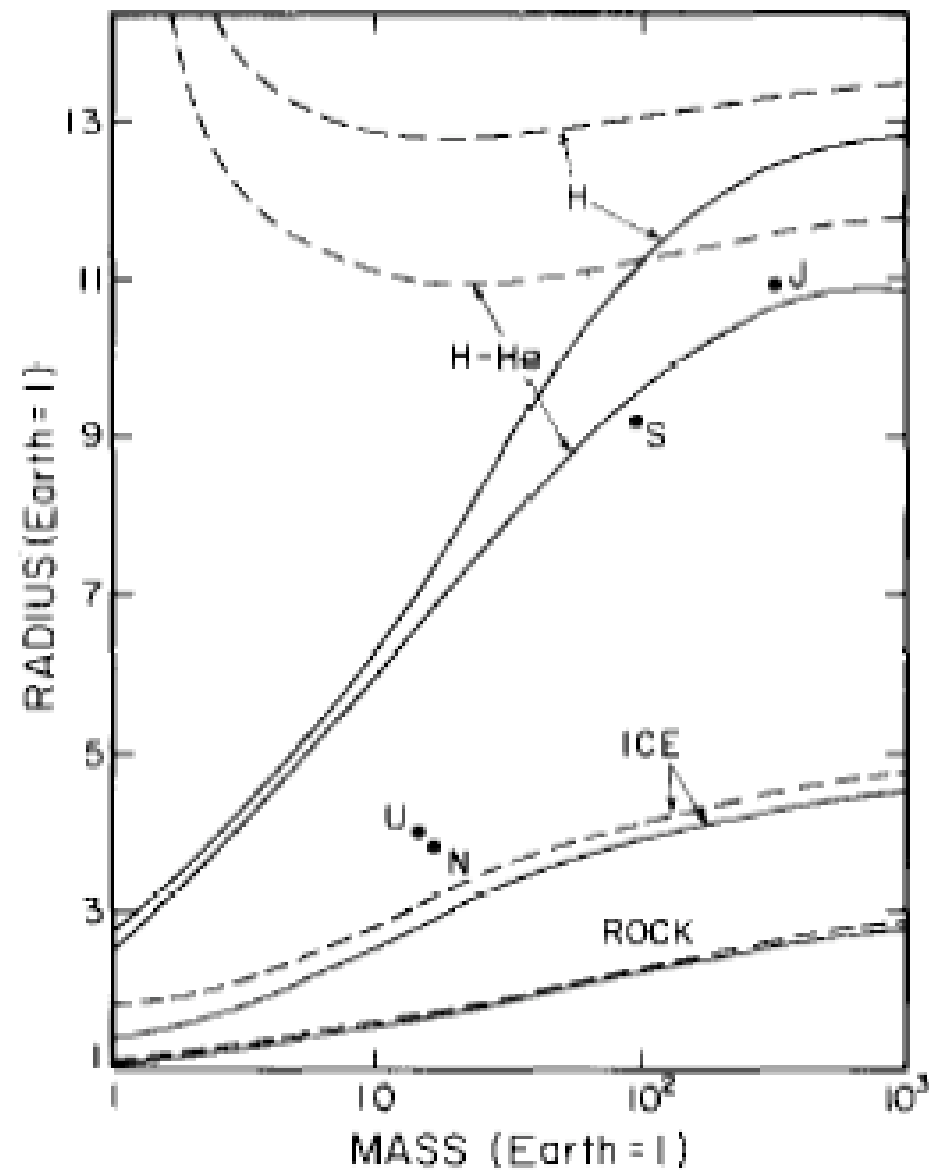


Density Structure

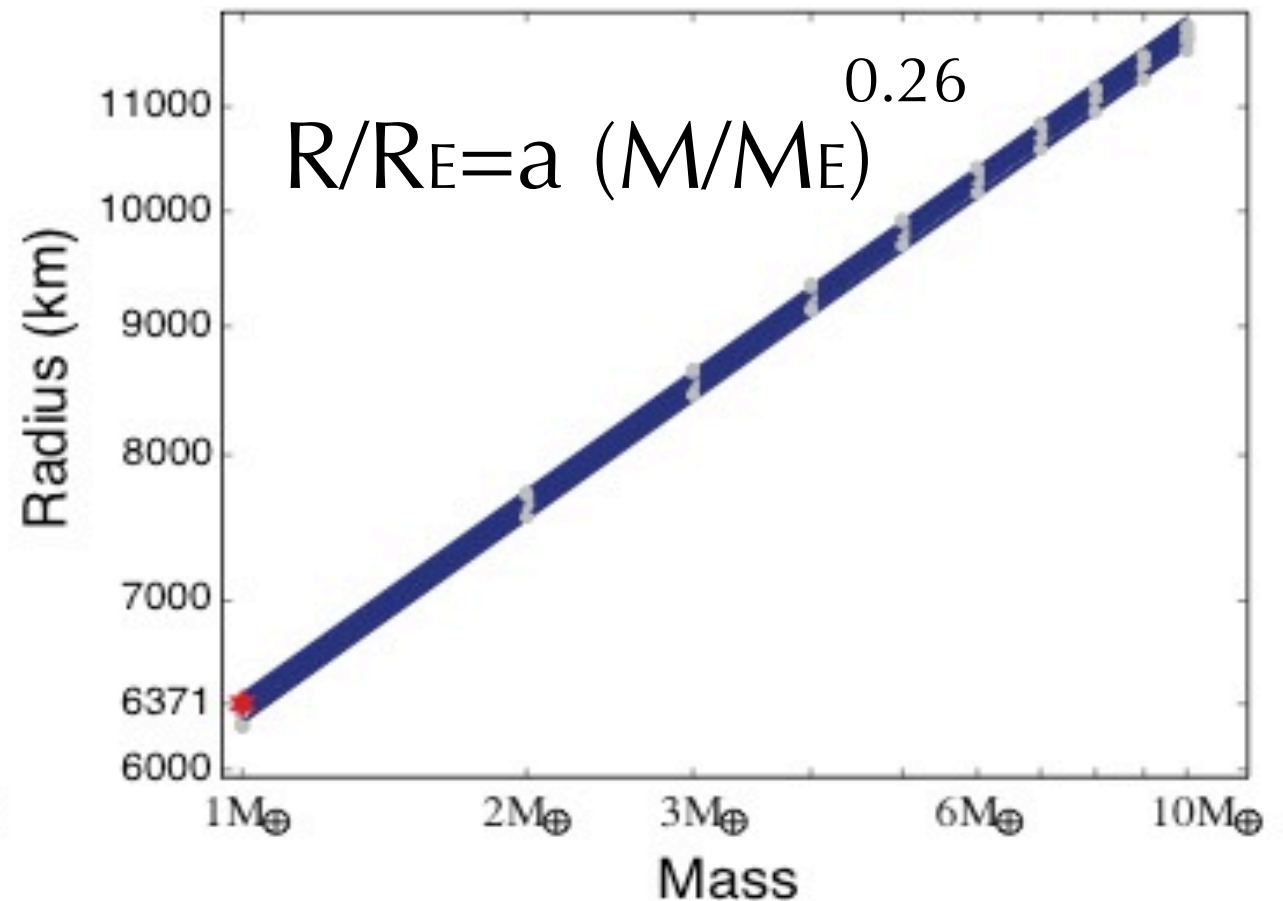


Valencia et. al, 2006

# Mass-Radius Relationships



Salpeter & Zapolsky 1967  
Stevenson 1982



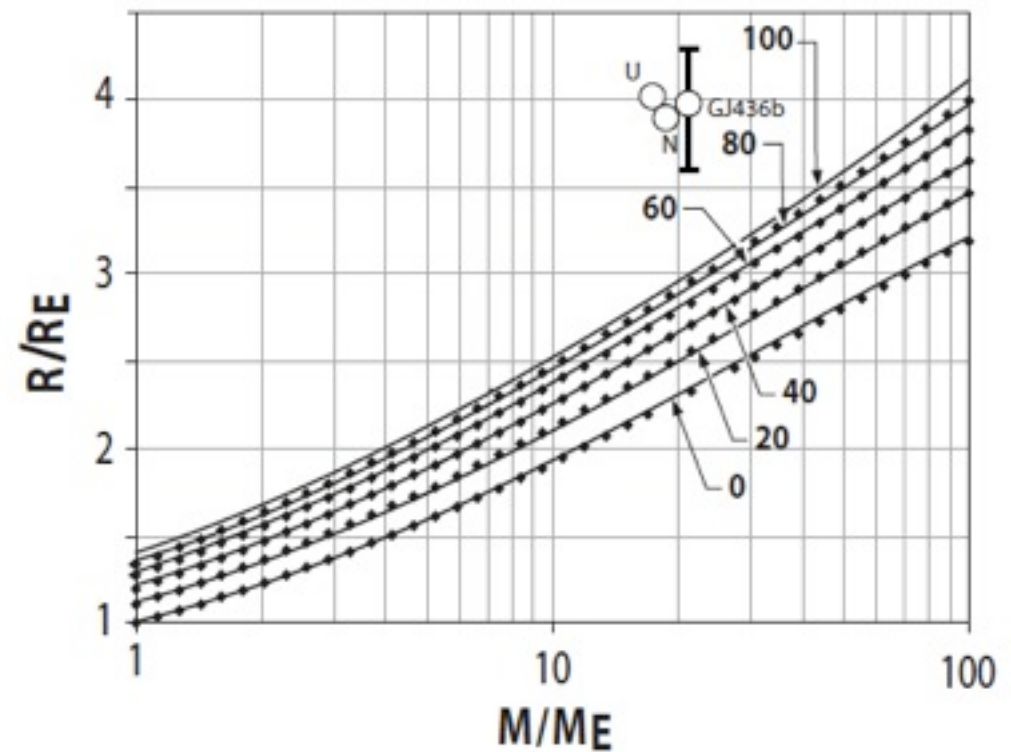
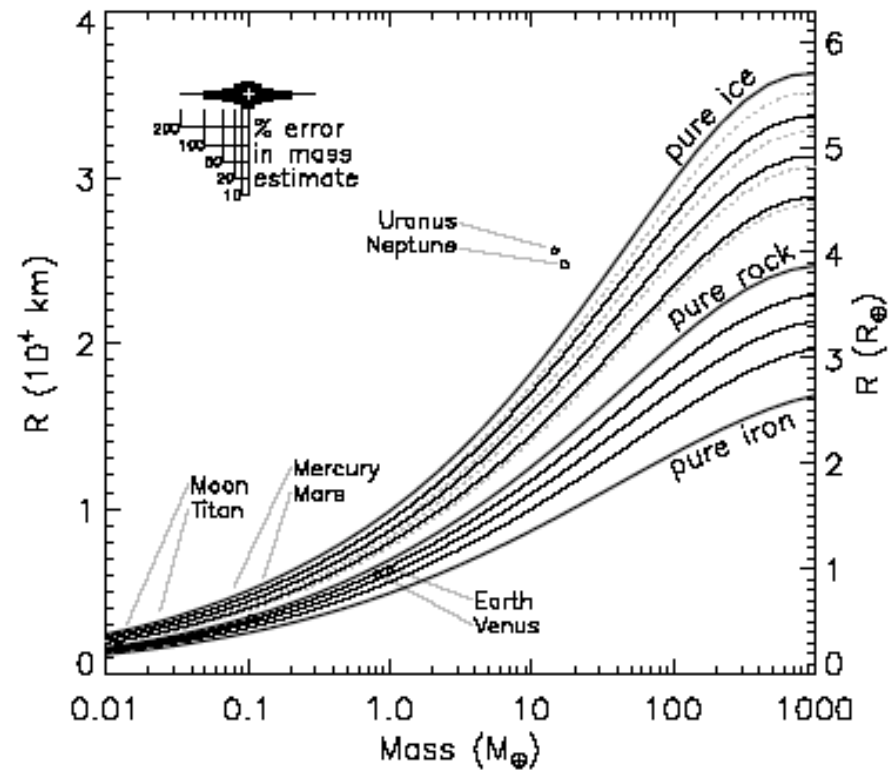
Valencia et. al, 2006

Exponent robust to EOS,  
Temperature Structure and CMF



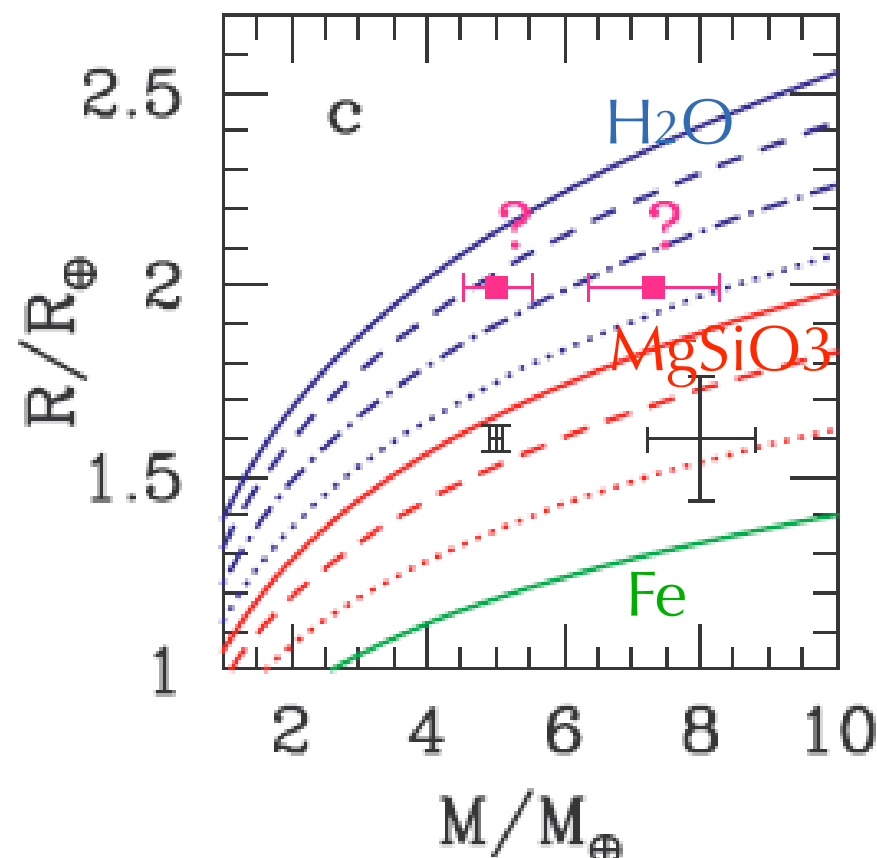
# Mass-Radius Relationships

Fortney et al 2007

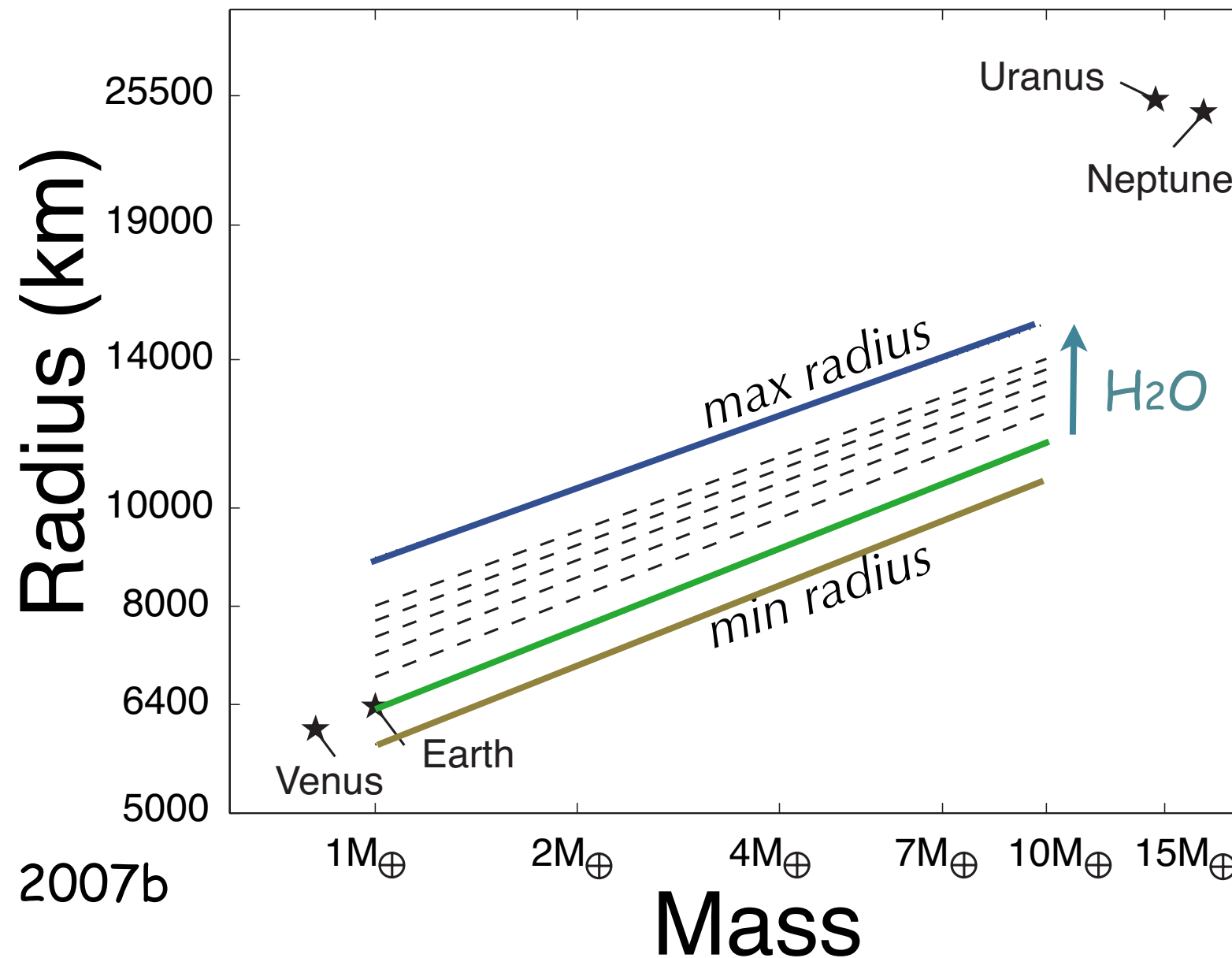


Sotin et al 2007,  
Grasset et al 2009

Seager et al 2007



# Generalized M-R relationship



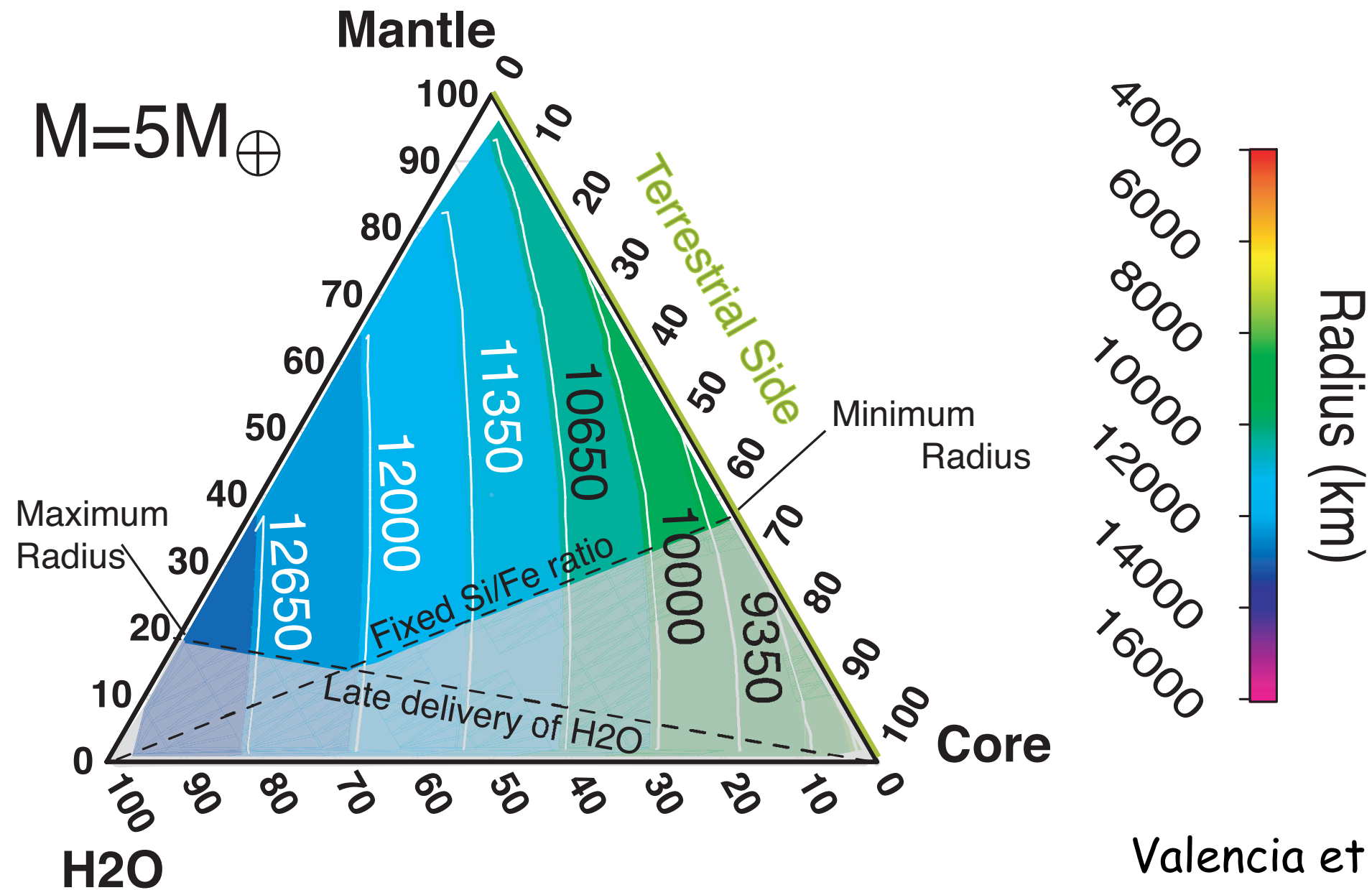
(Earth-like:  
Fe/Si)  
IMF: ice-mass  
fraction

Valencia et al. 2007b

$$R/R_{\text{Earth}} = (1 + 0.56 \text{ IMF}) (M/M_{\text{Earth}})^{0.262(1 - 0.138 \text{ IMF})}$$

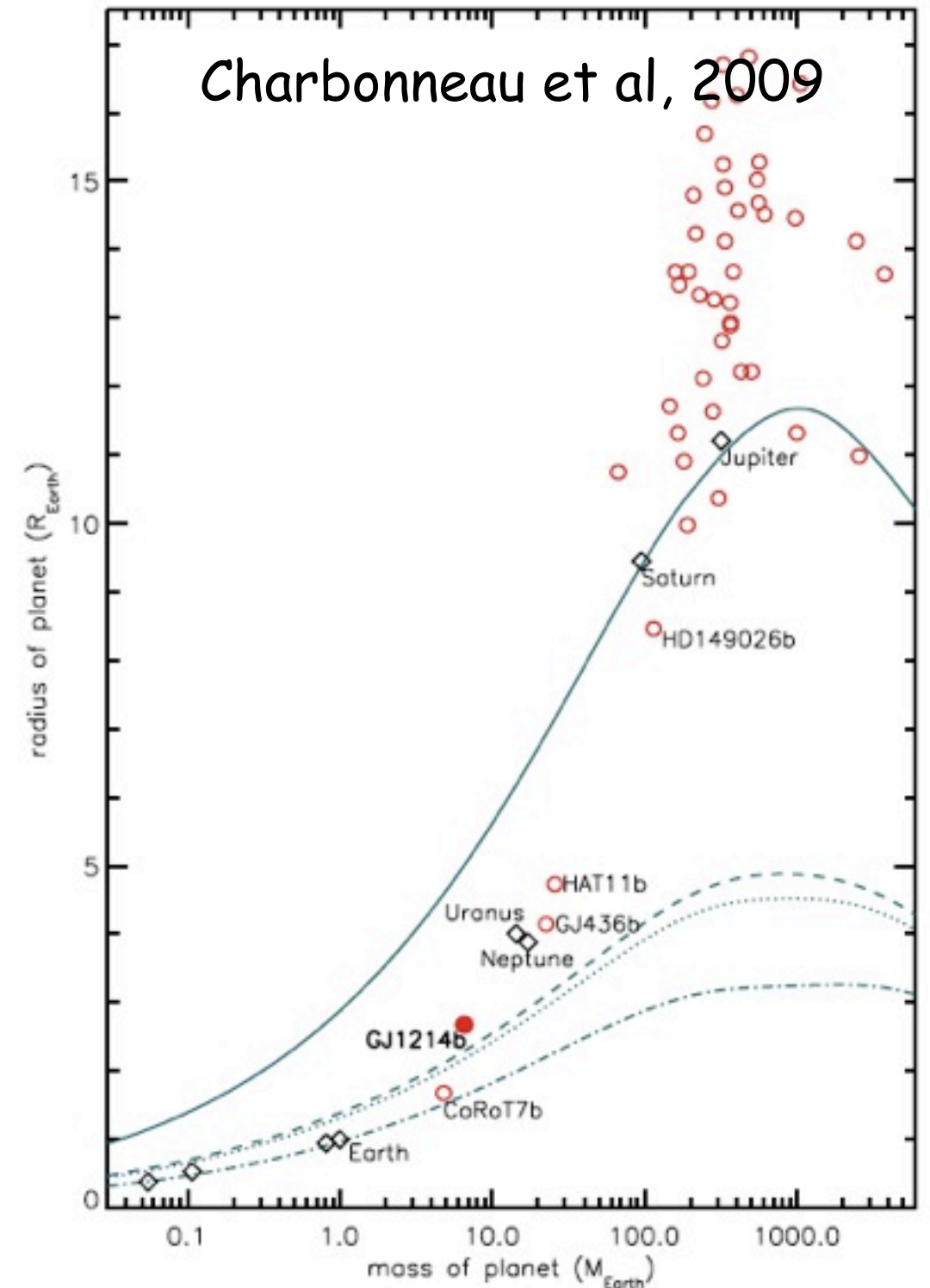
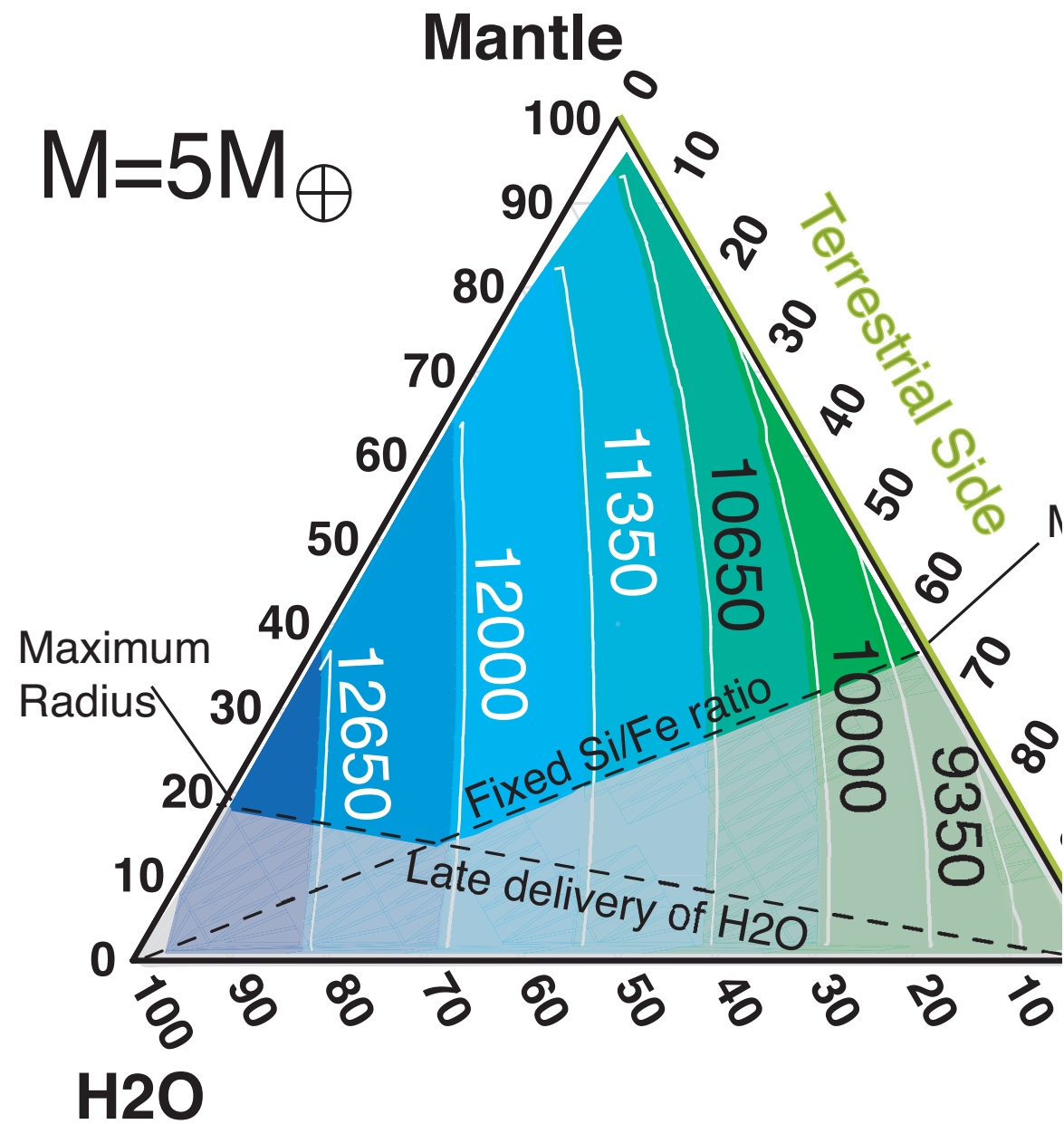


# Super-Earth Composition



There is degeneracy in composition.  
Some compositions are improbable.

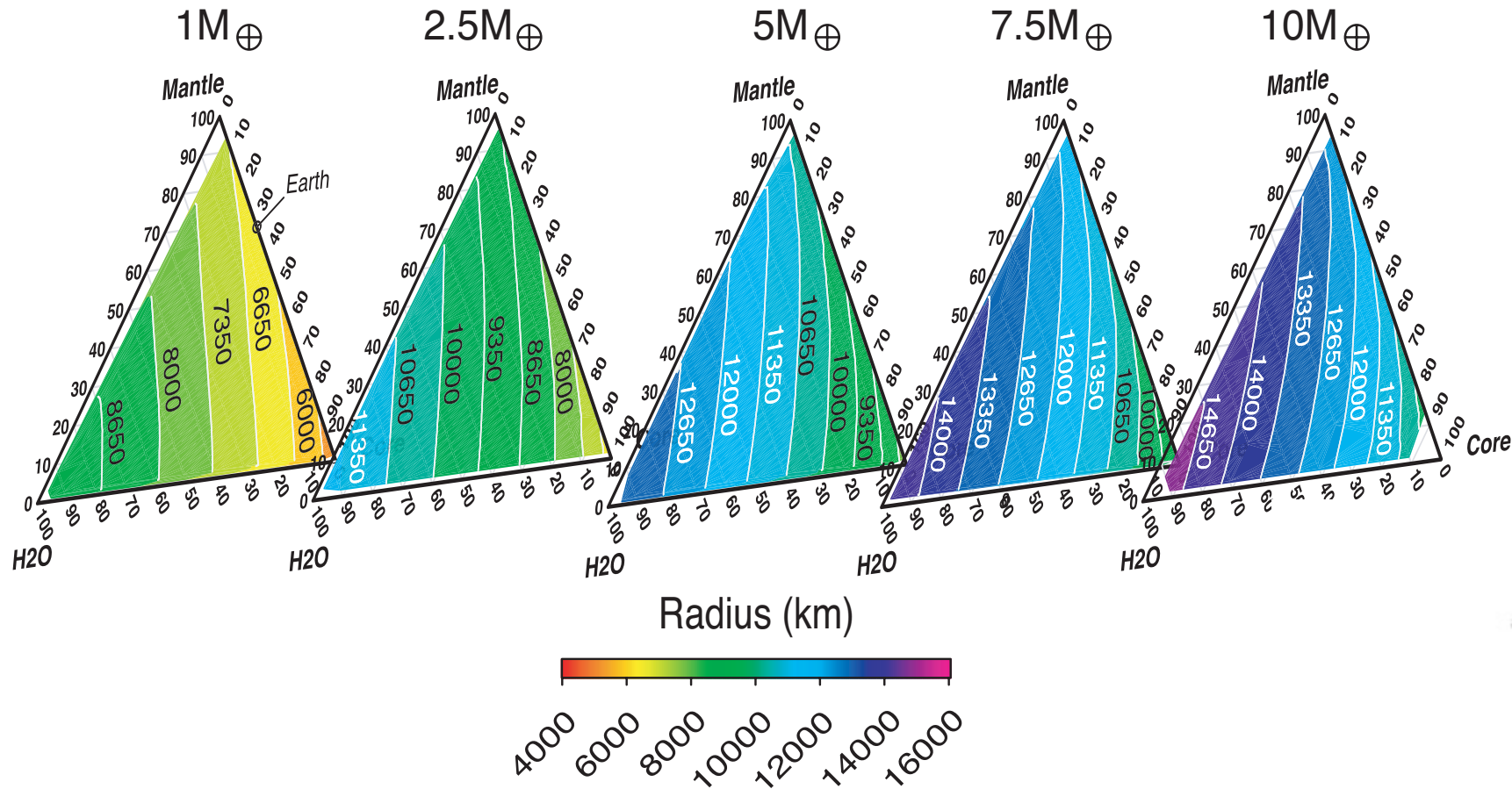
# Super-Earth Composition



There is degeneracy in  
Some compositions are improbable.

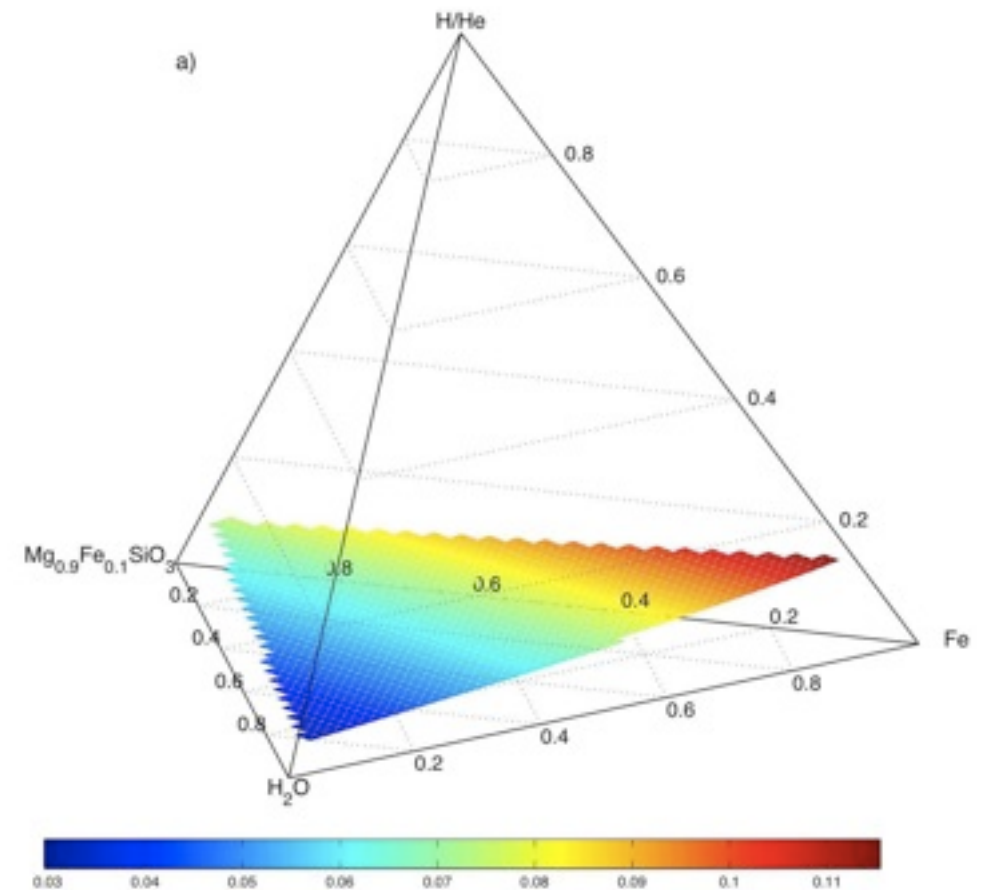
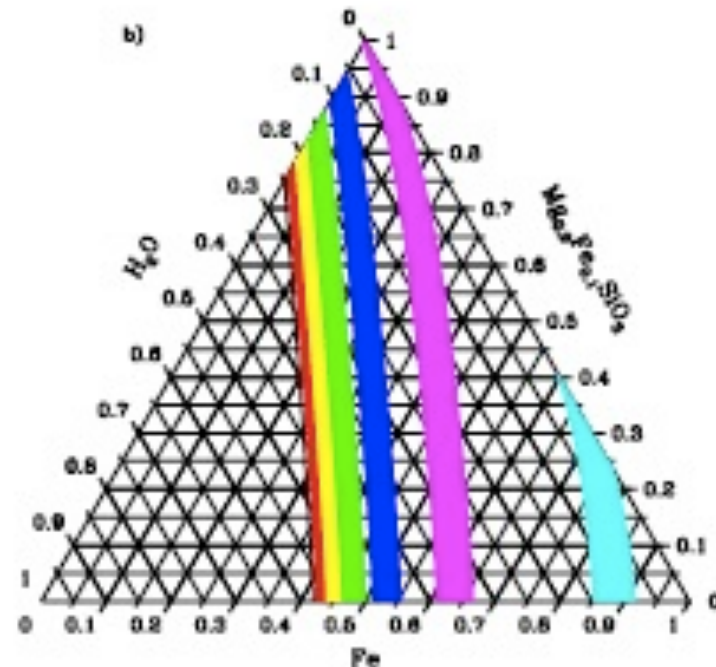


# 'Toblerone' Diagram

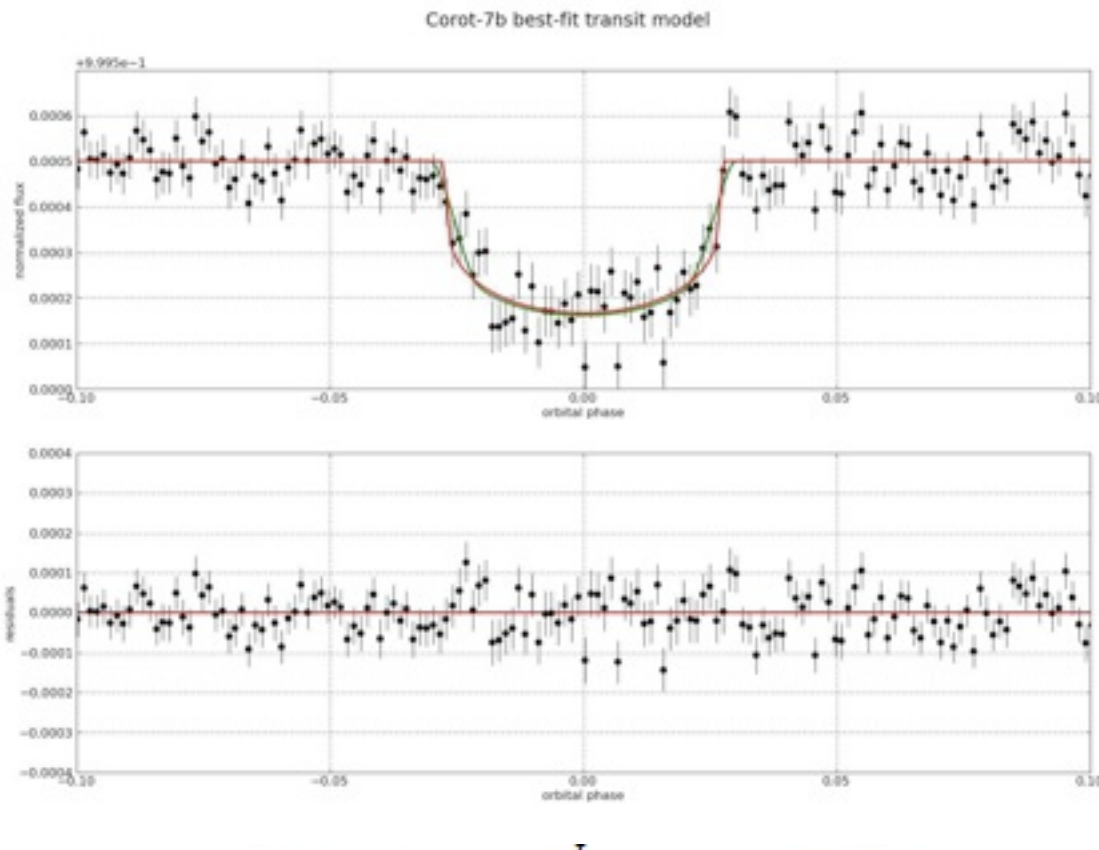


Valencia et al. 2007b

Rogers & Seager,  
2010



# CoRoT-7b: the first transiting super-Earth

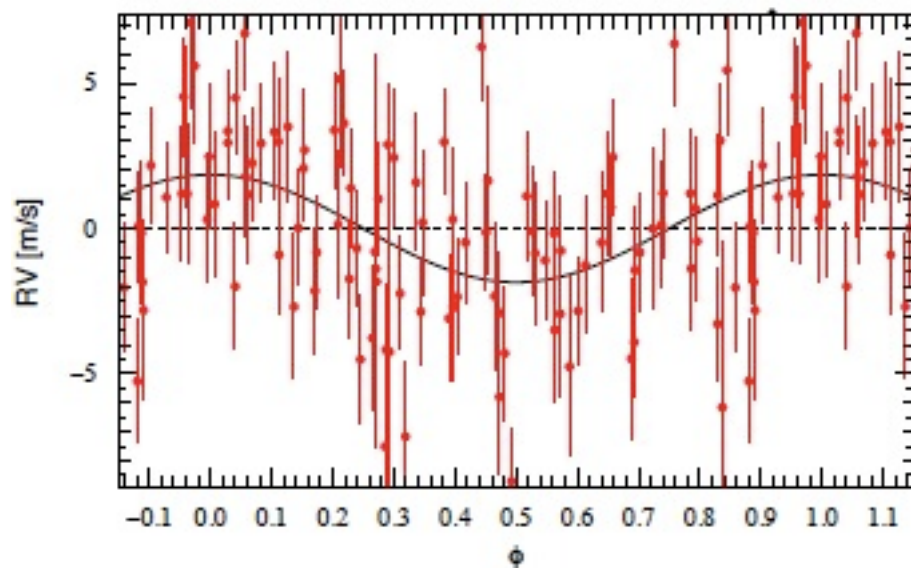


Radius =  $1.68 \pm 0.09 R_E$

Period = 0.854 days

Age = 1.2-2.3 Gy

Mass =  $4.8 \pm 0.8 M_E$



Leger et al, 2009

Queloz et al, 2009



# Can it retain an atmosphere?

No, unless it is constantly resupplied

Energy limited calculation based on UV flux

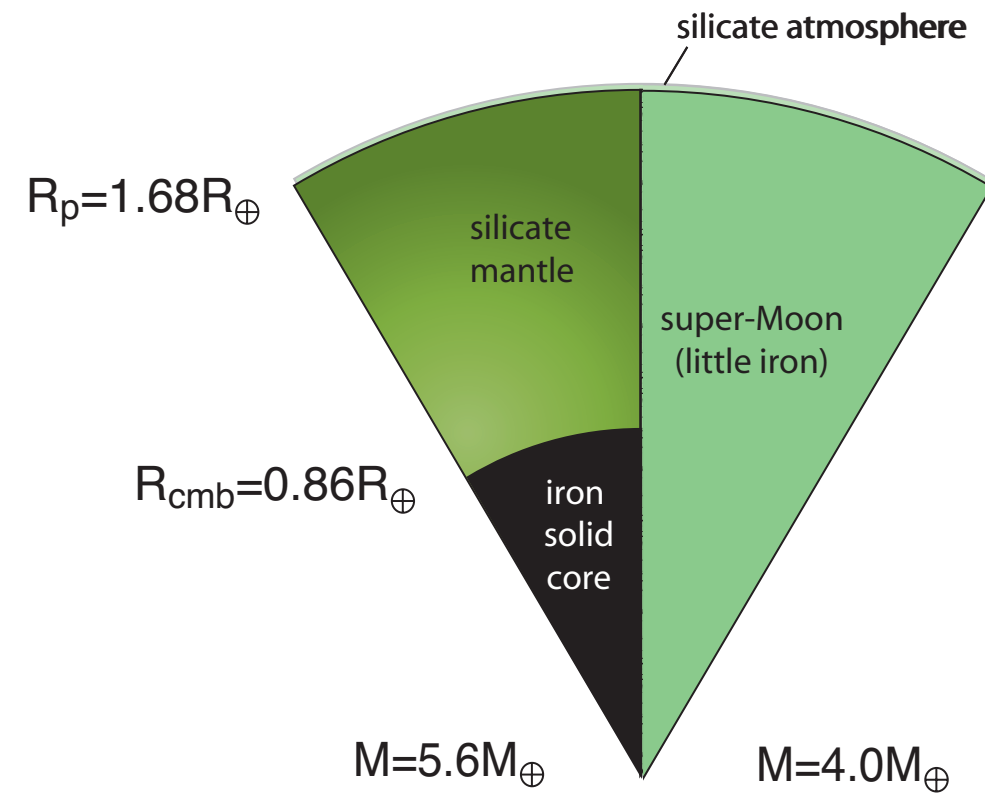
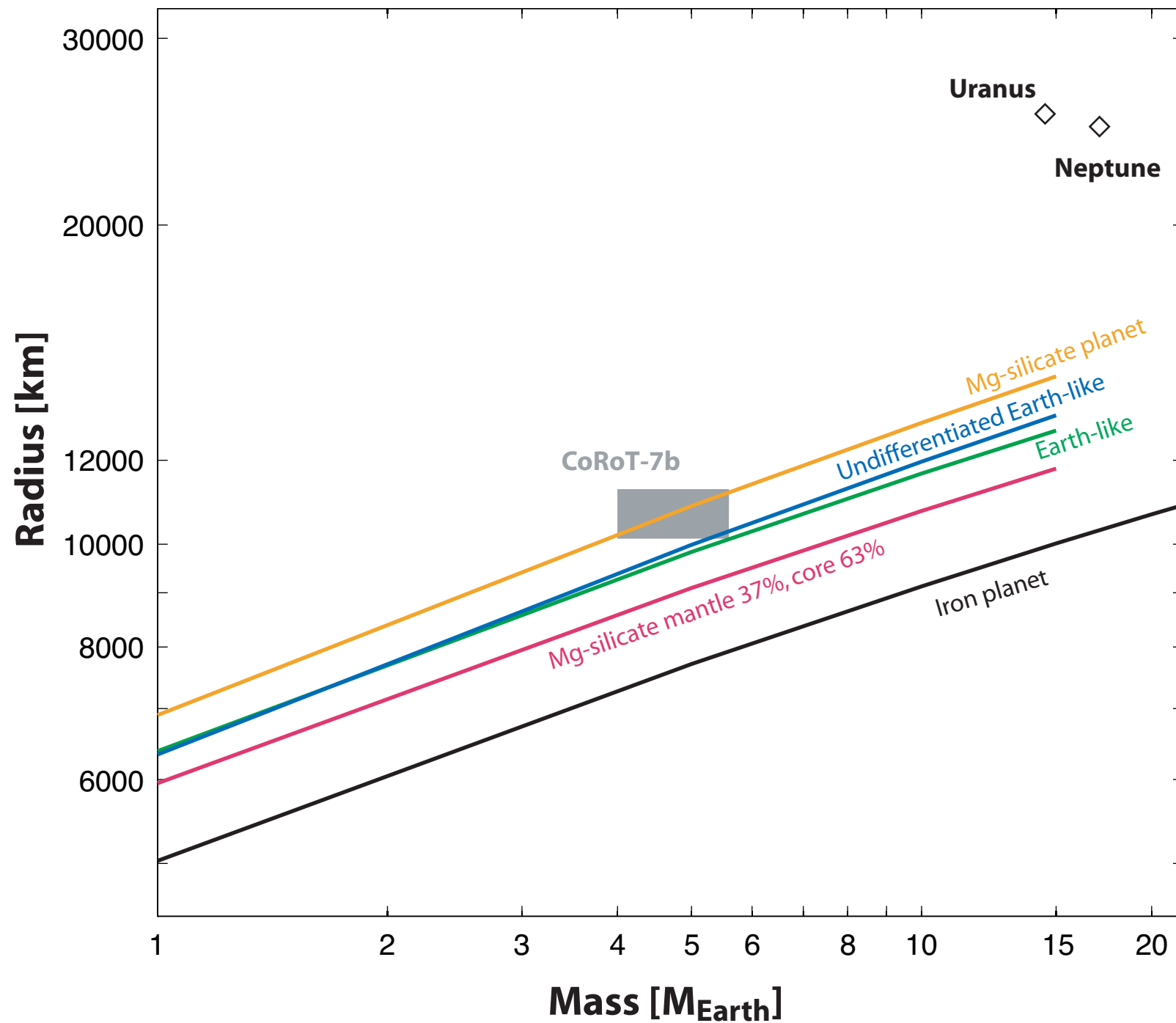
$$\frac{dM}{dt} = \frac{3 \epsilon F_{EUV}}{4 G \rho K_{\text{tide}}} = 10^{11} \text{ g/s}$$

For more details on  $\epsilon$   
see Lammer et al 09

within an order of magnitude of the escape rate  
of HD 209458b

Even if it has a silicate atmosphere, it is thick  
enough for UV absorption

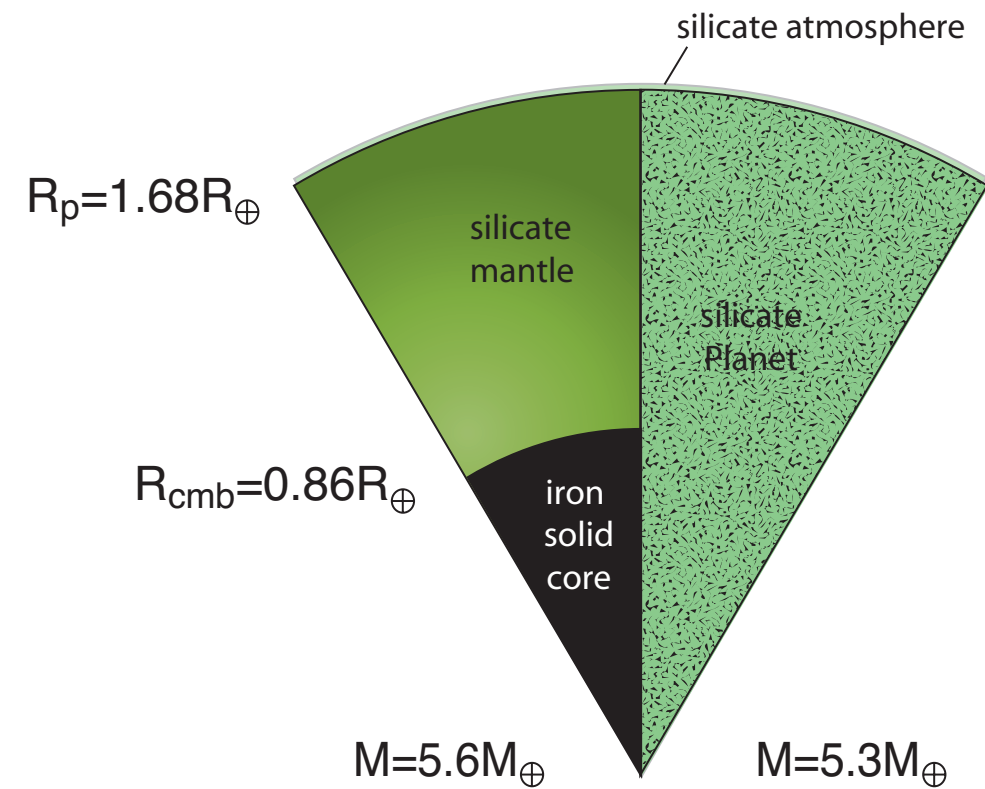
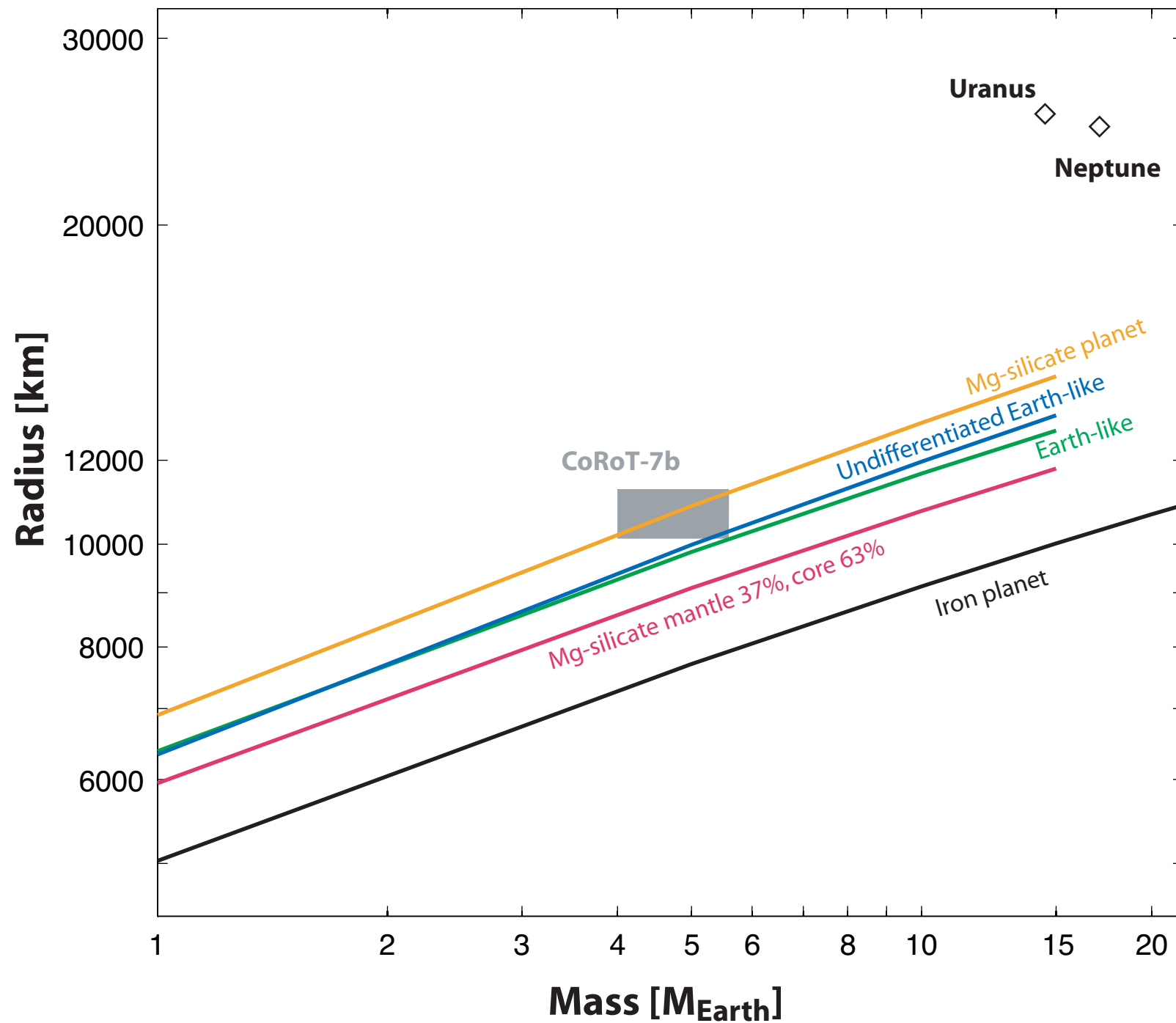
# Rocky CoRoT-7b



Valencia et al, 2010

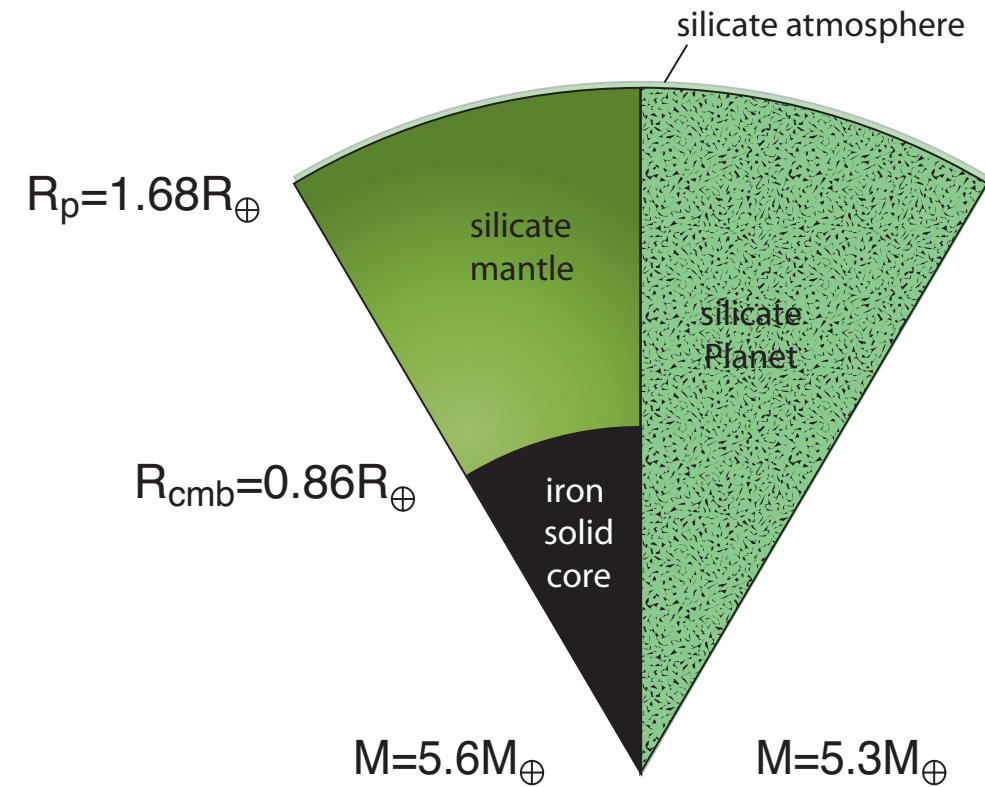
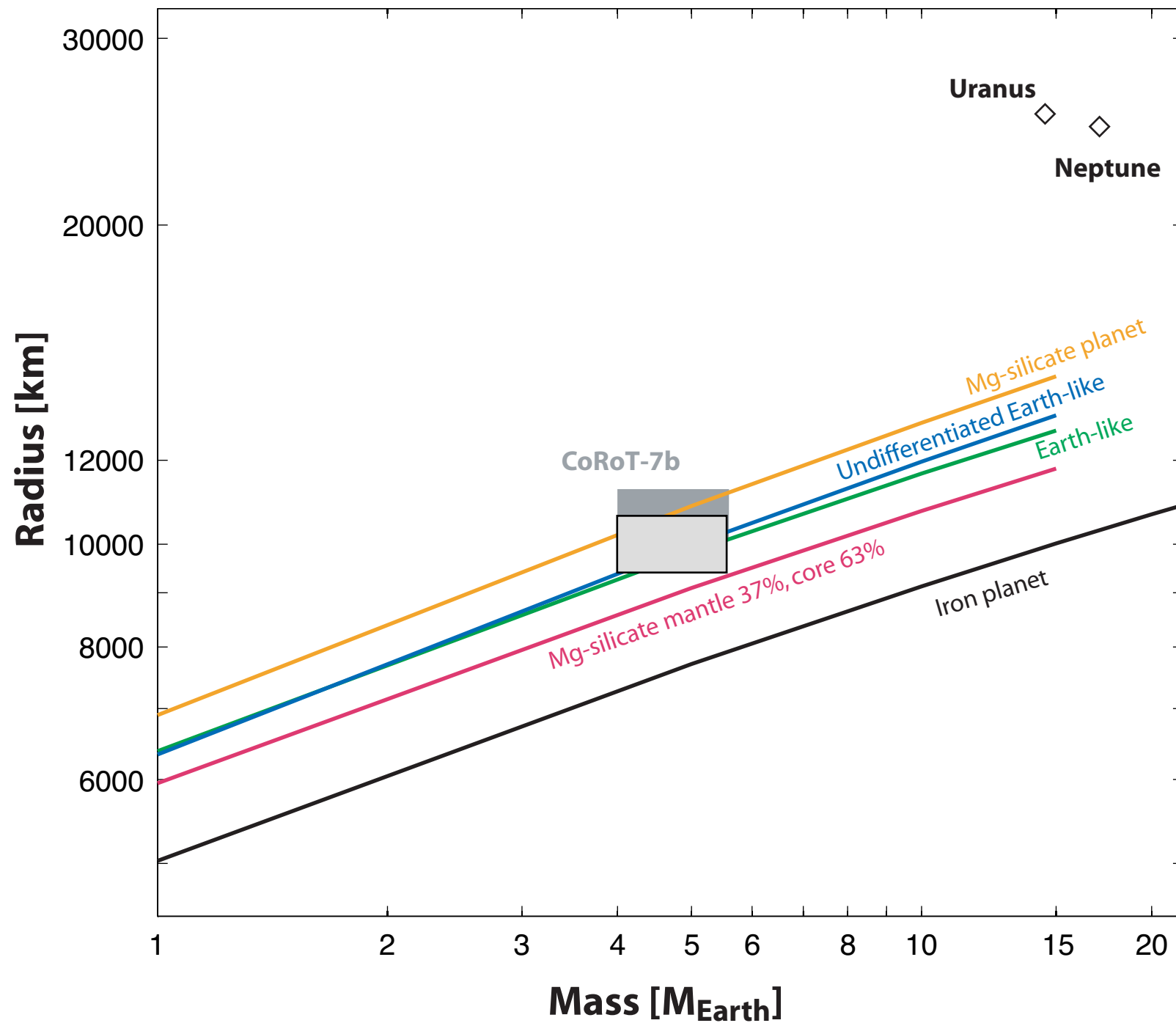


# Rocky CoRoT-7b



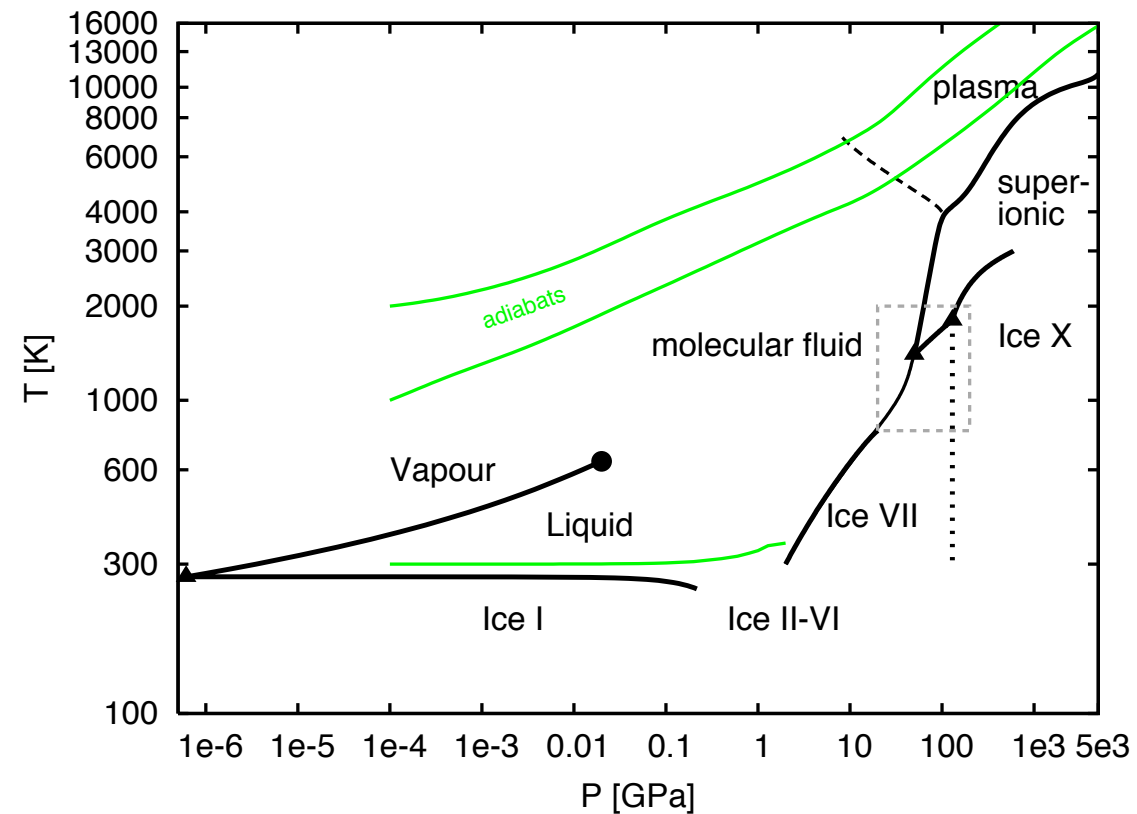
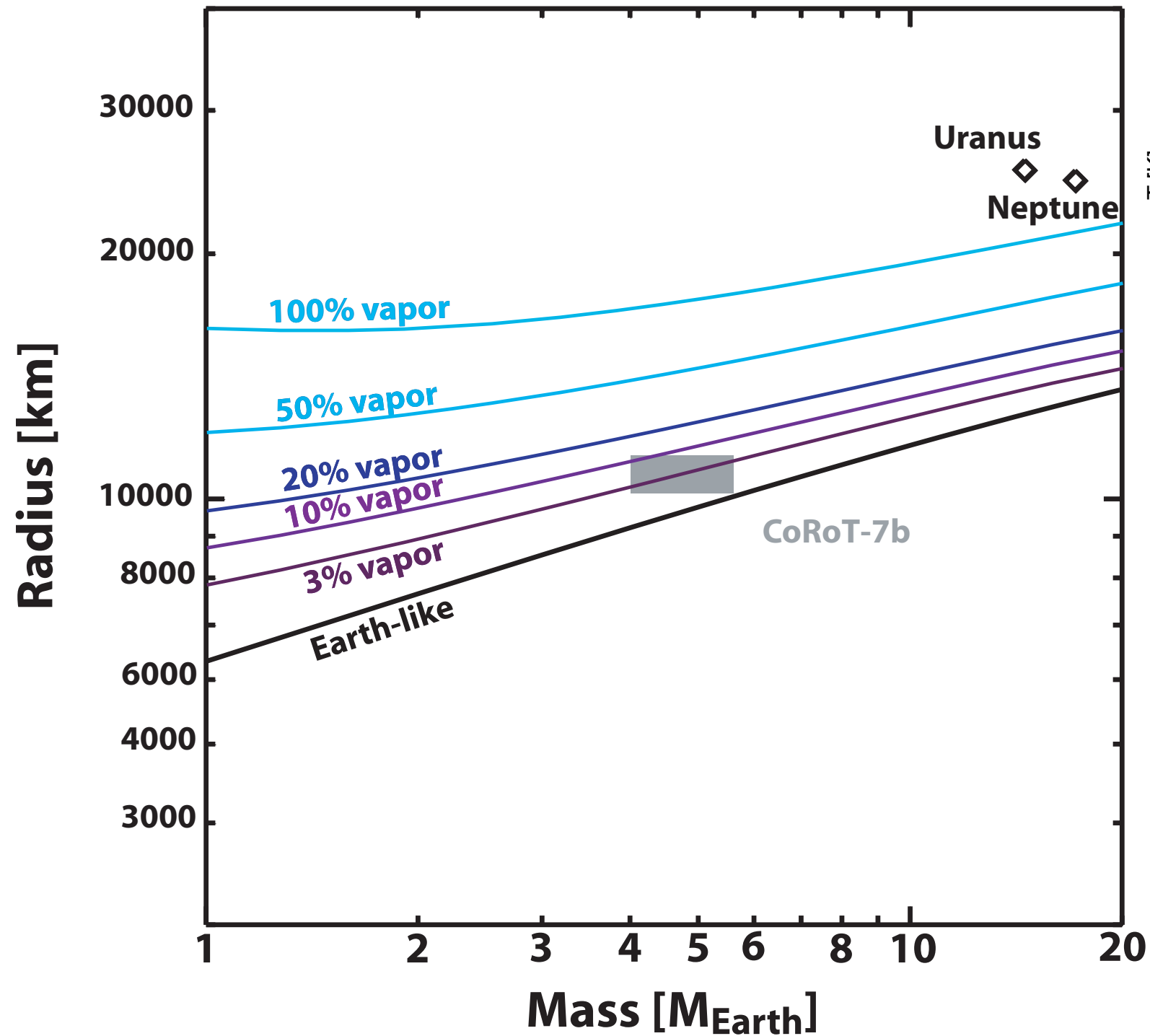
Valencia et al, 2010

# Rocky CoRoT-7b



Valencia et al, 2010

# H2O Vapor CoRoT-7b

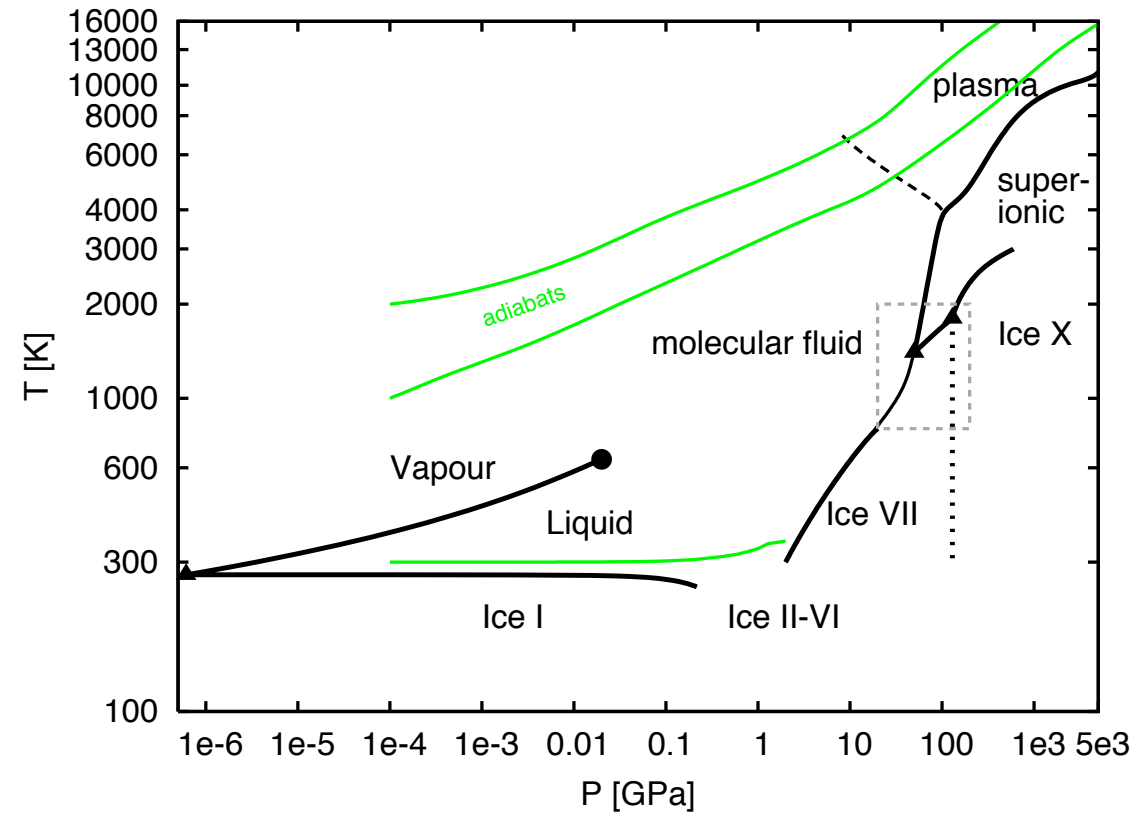
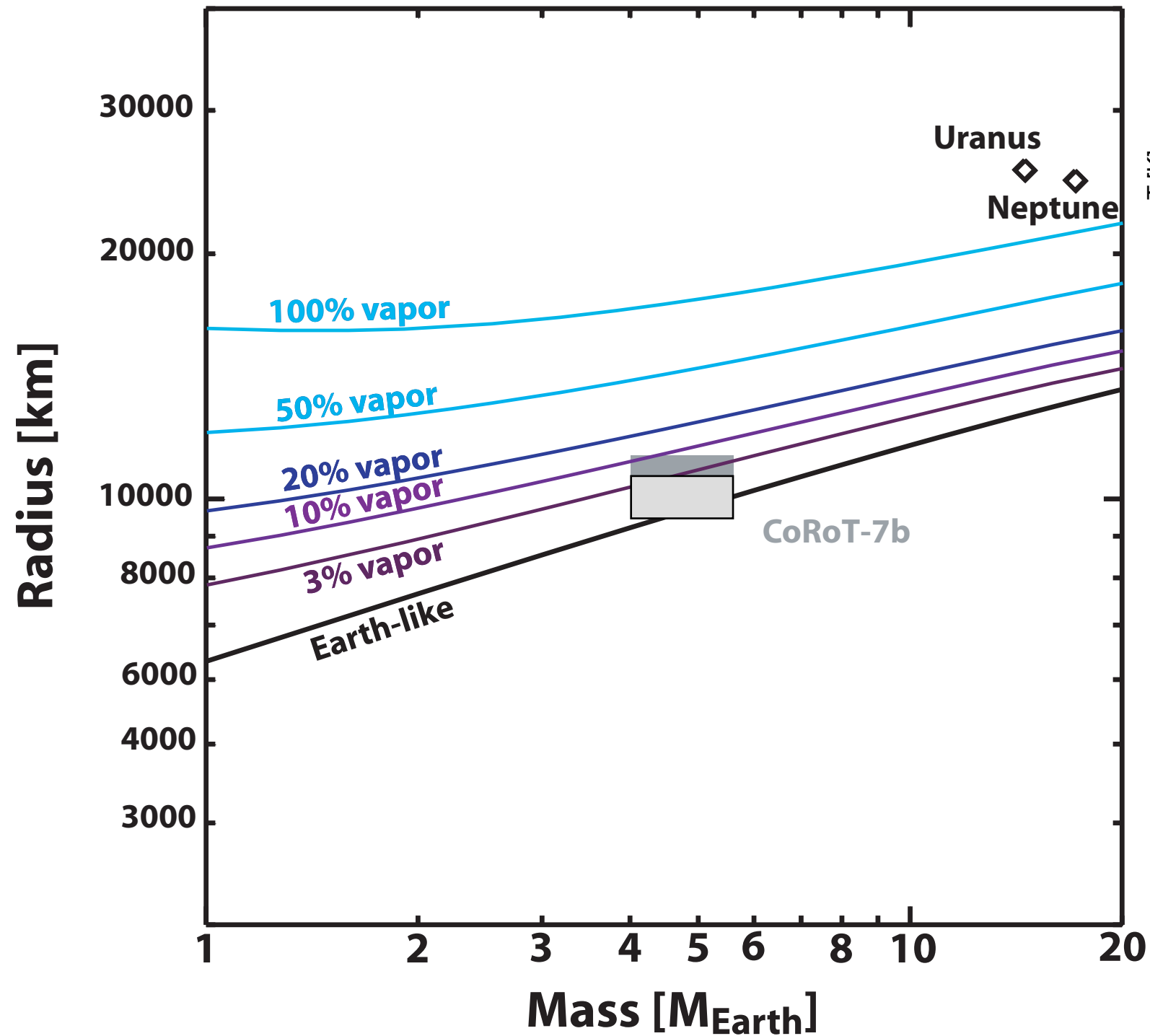


Evaporation  
Timescale  $\sim 1 \text{ Gy}$

Valencia et al, 2010



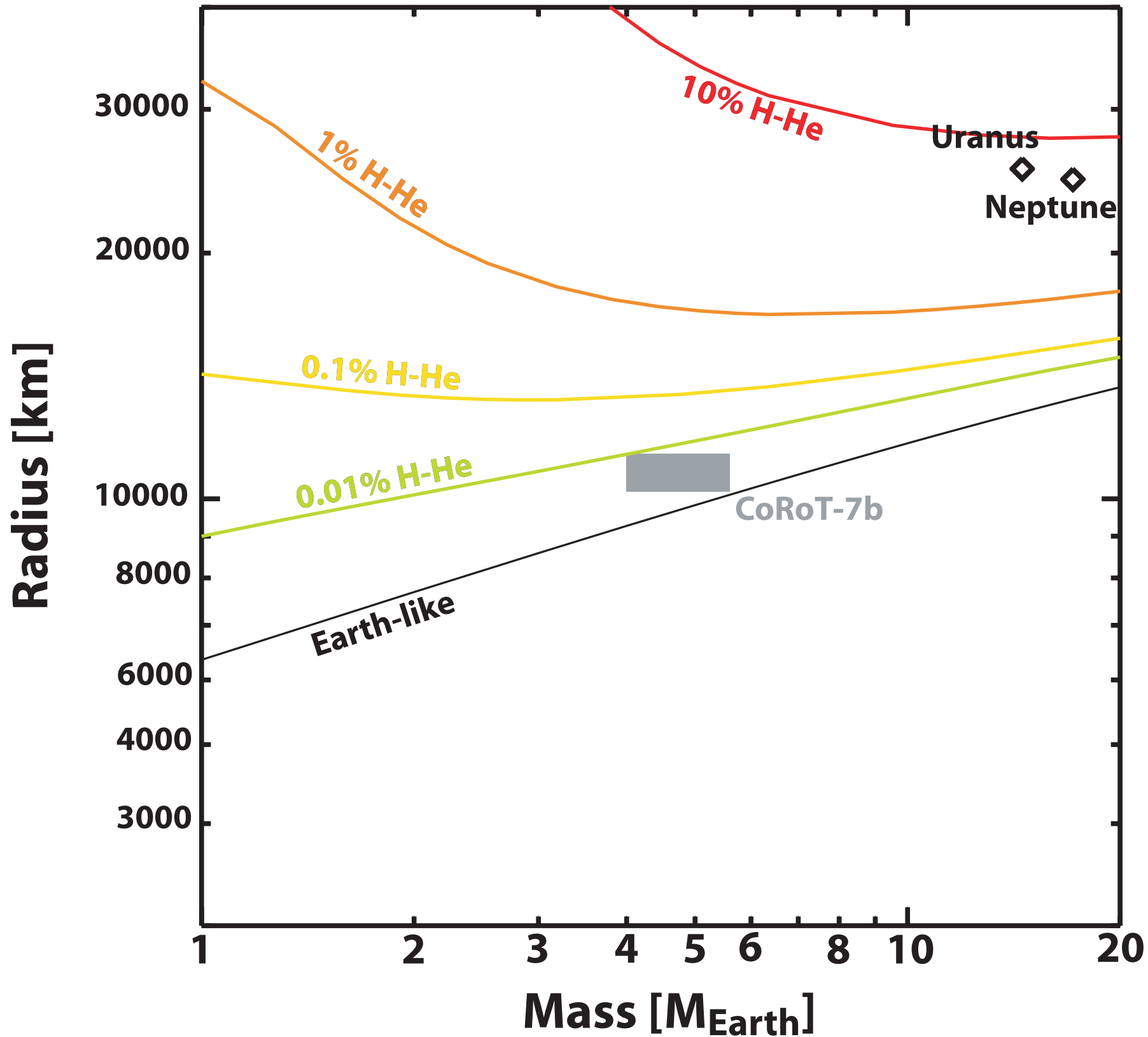
# H2O Vapor CoRoT-7b



Evaporation  
Timescale  $\sim 1$  Gy

Valencia et al, 2010

# Hydrogen or Helium?



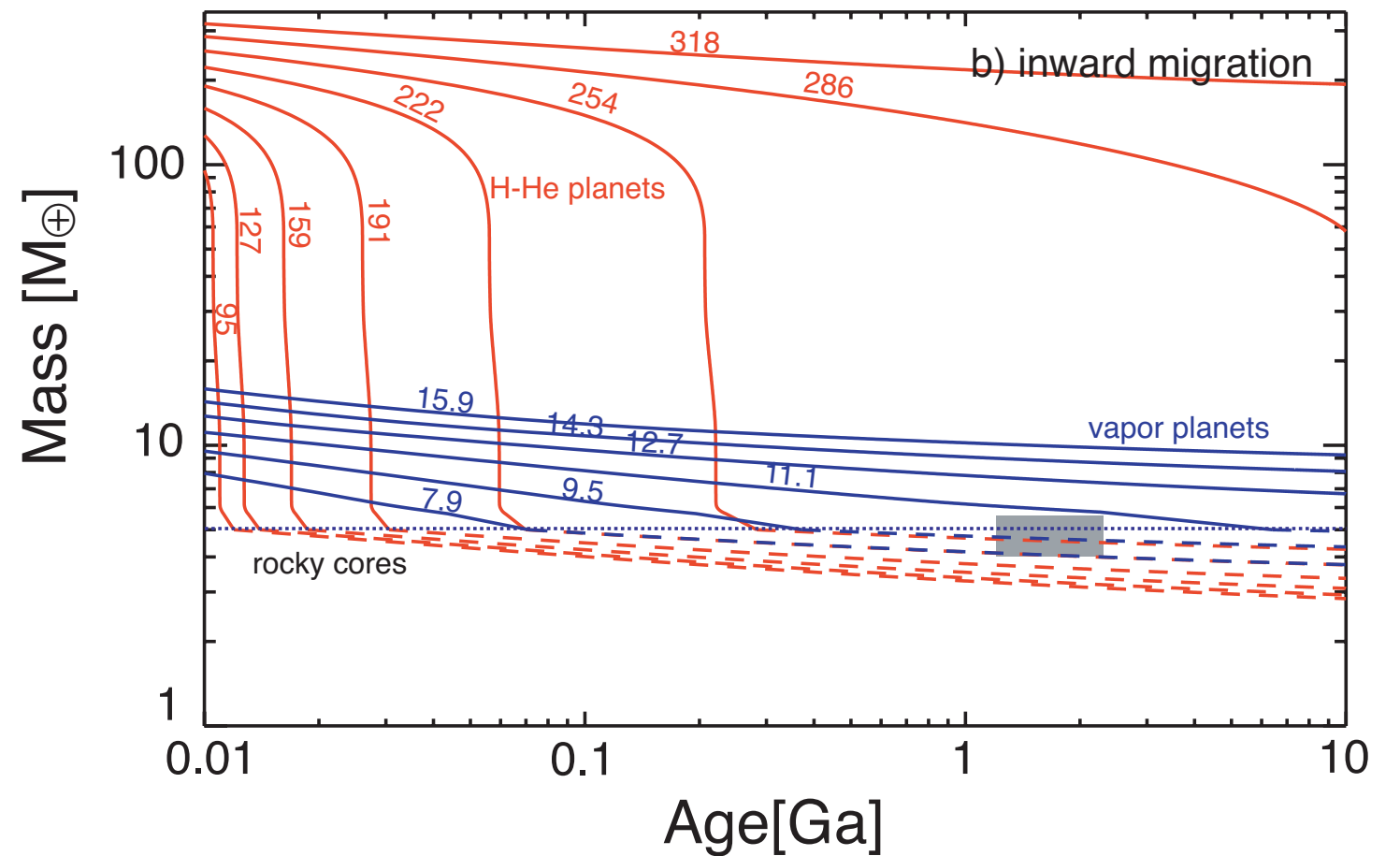
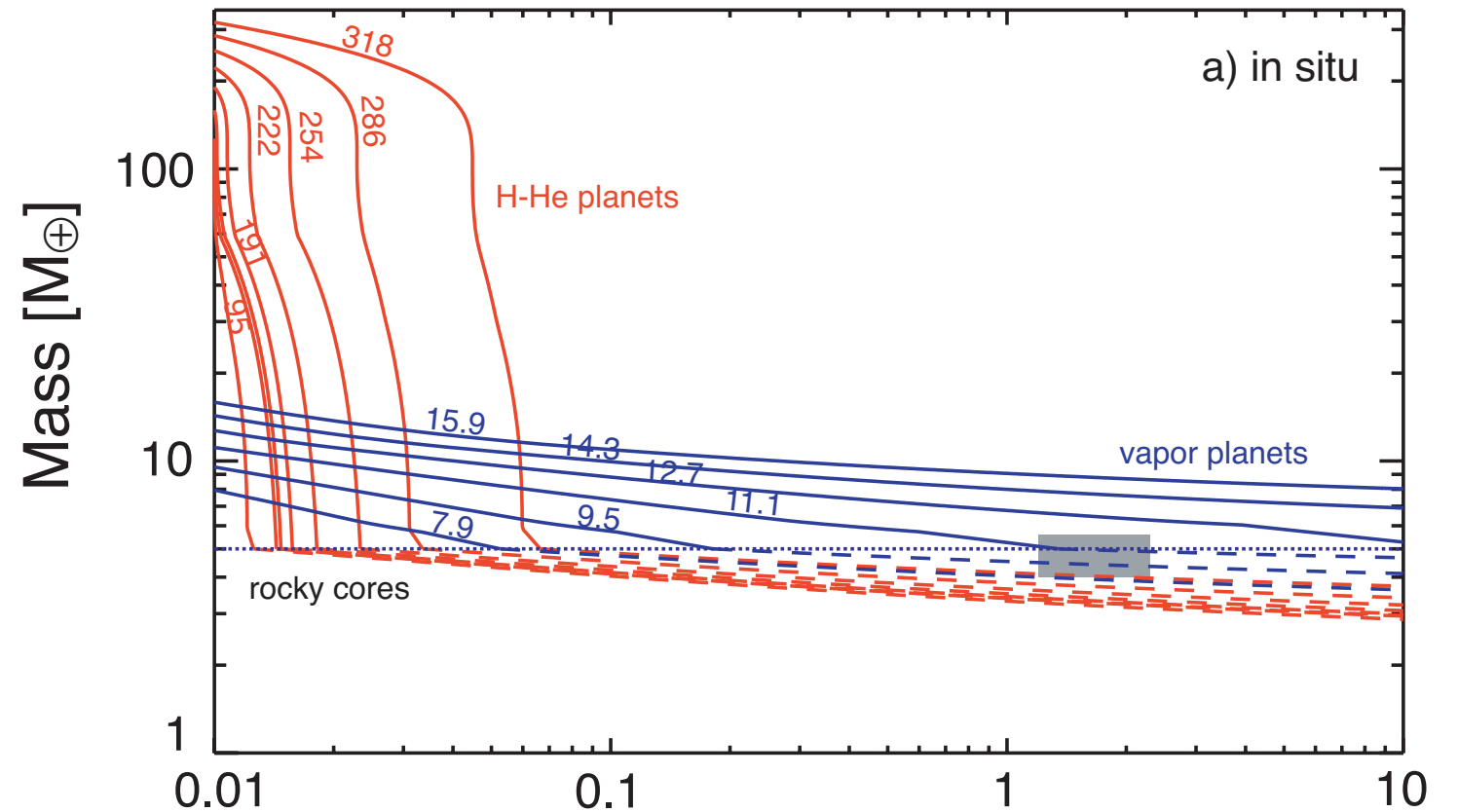
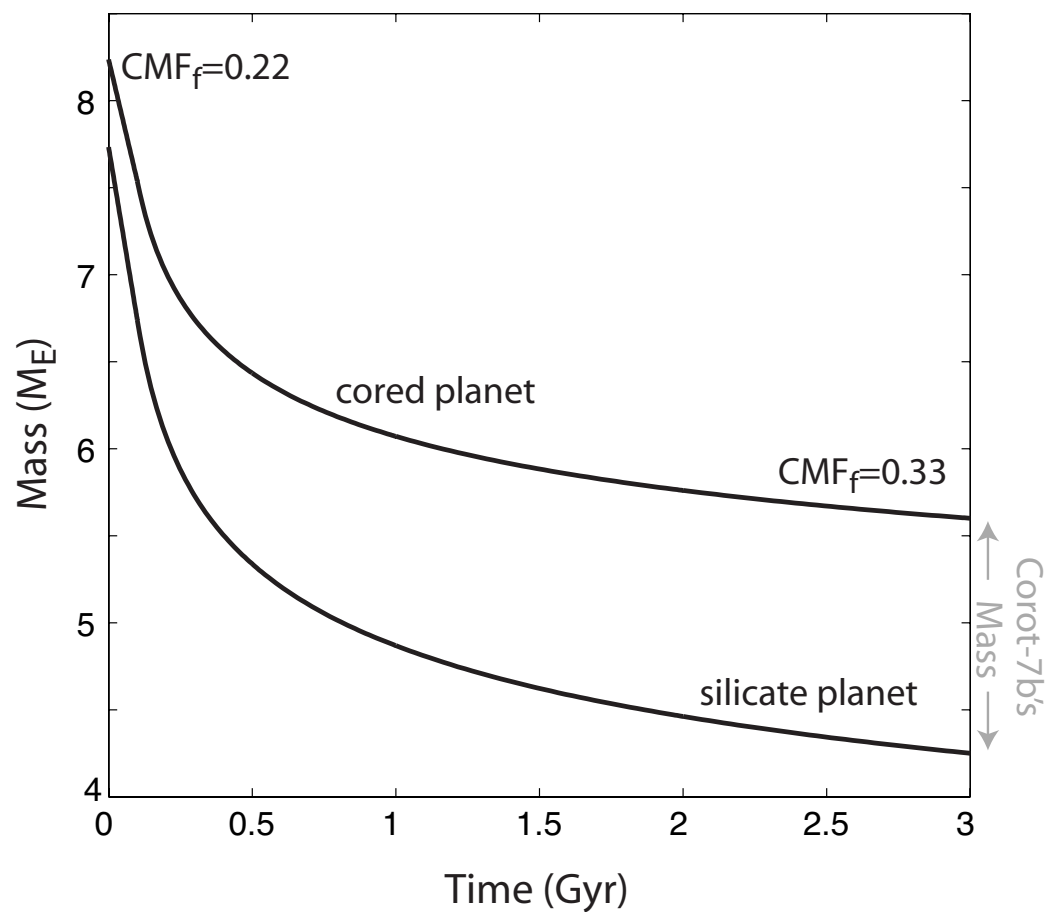
Evaporation  
Timescale  $\sim 1$  My!

Valencia et al. 2010

# CoRoT-7b's origin?

## Volatile Origin

## Rocky Origin



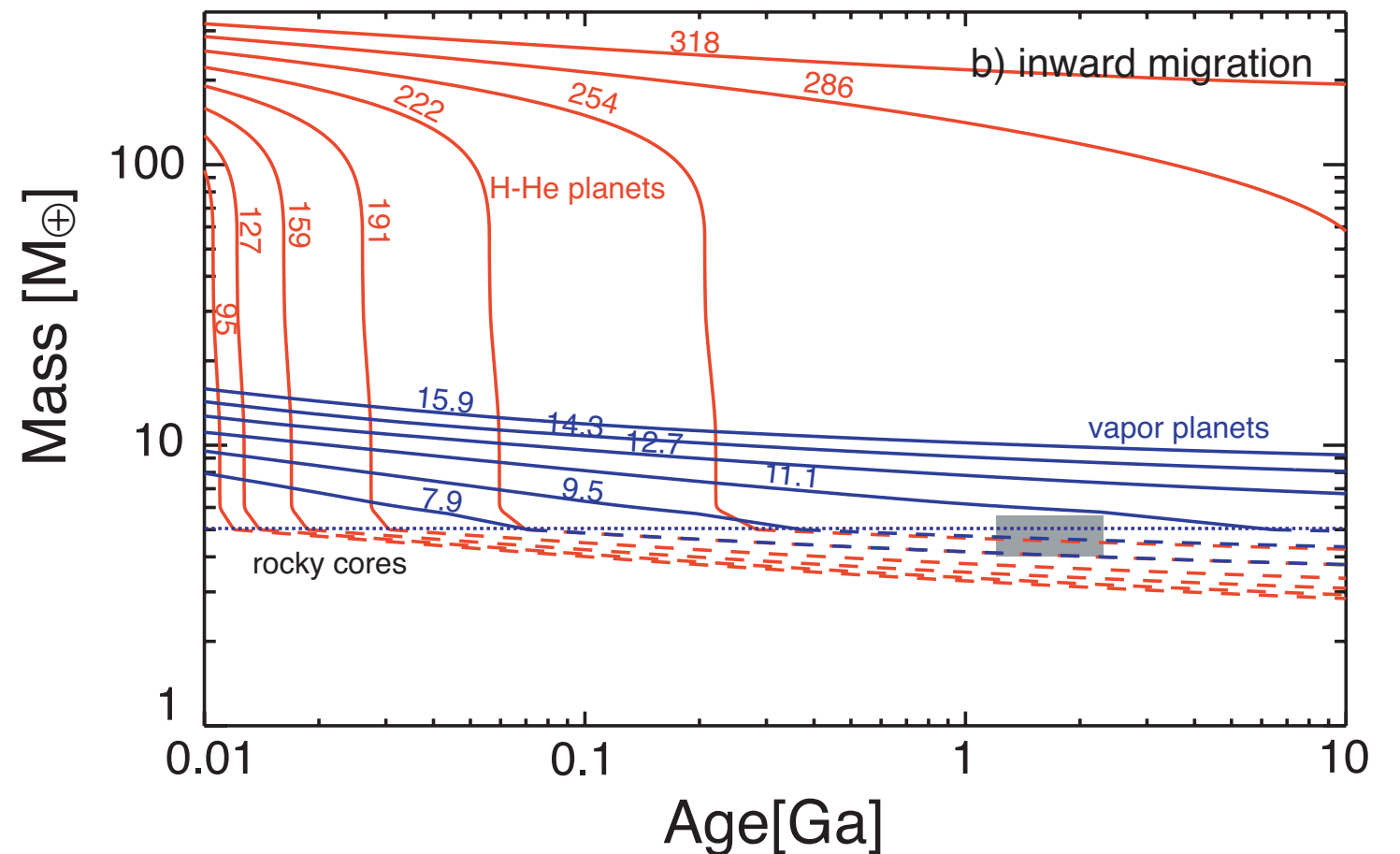
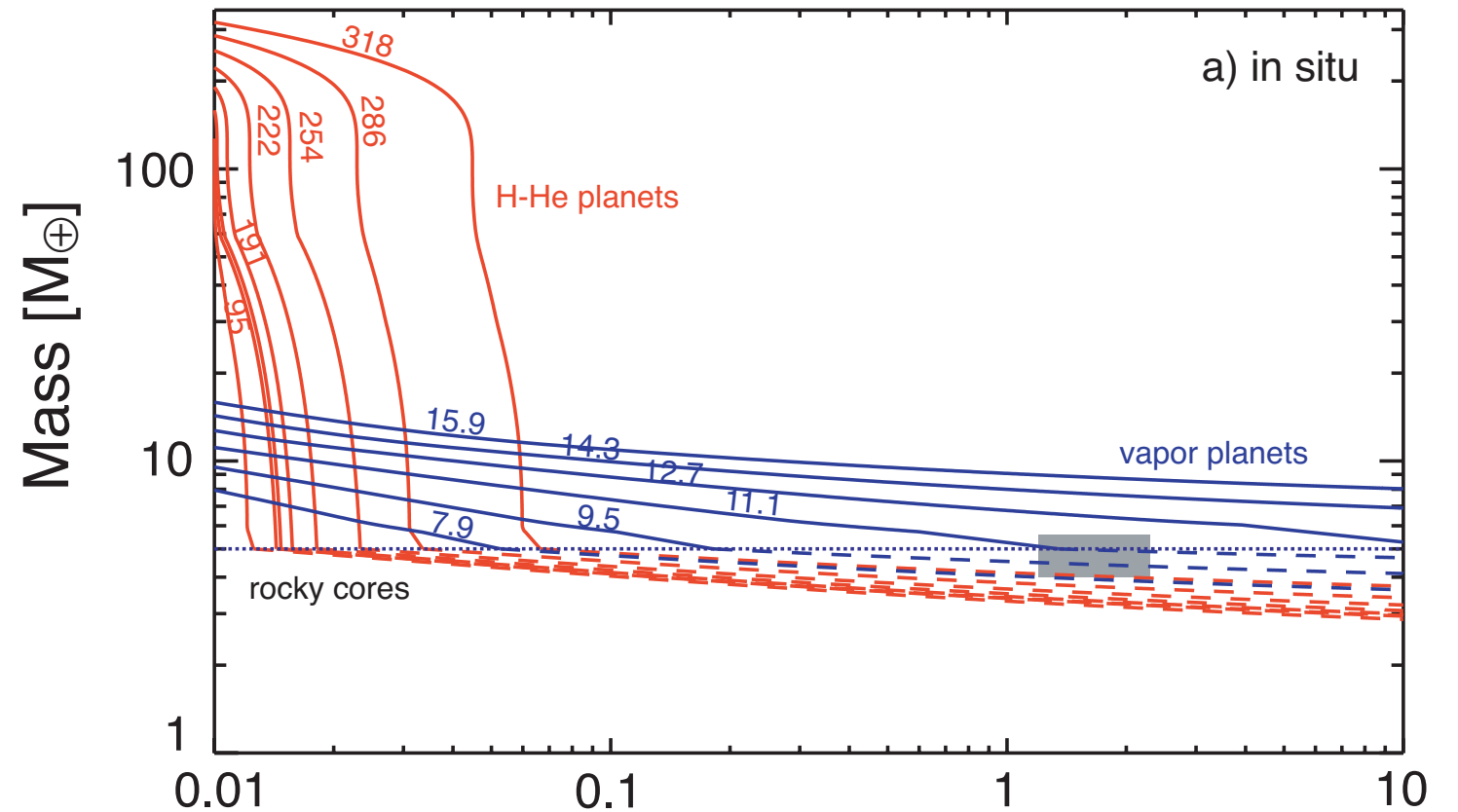
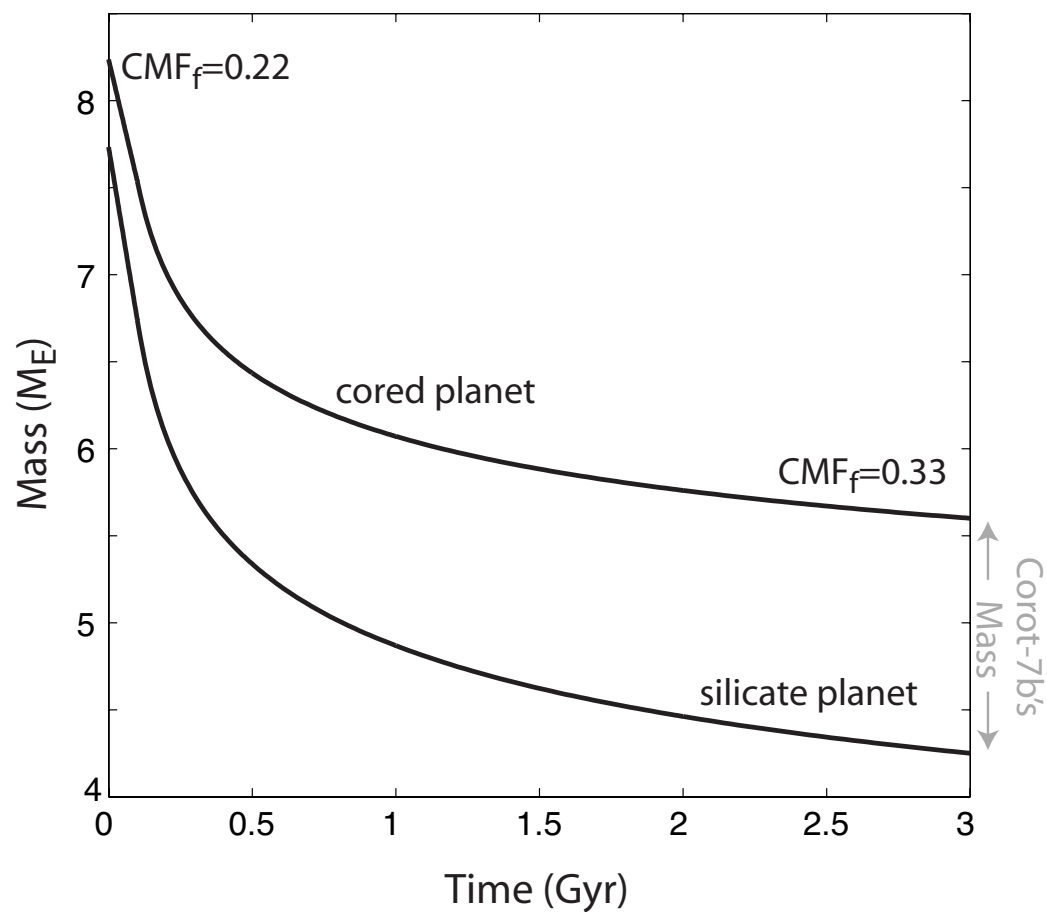


# CoRoT-7b's origin?

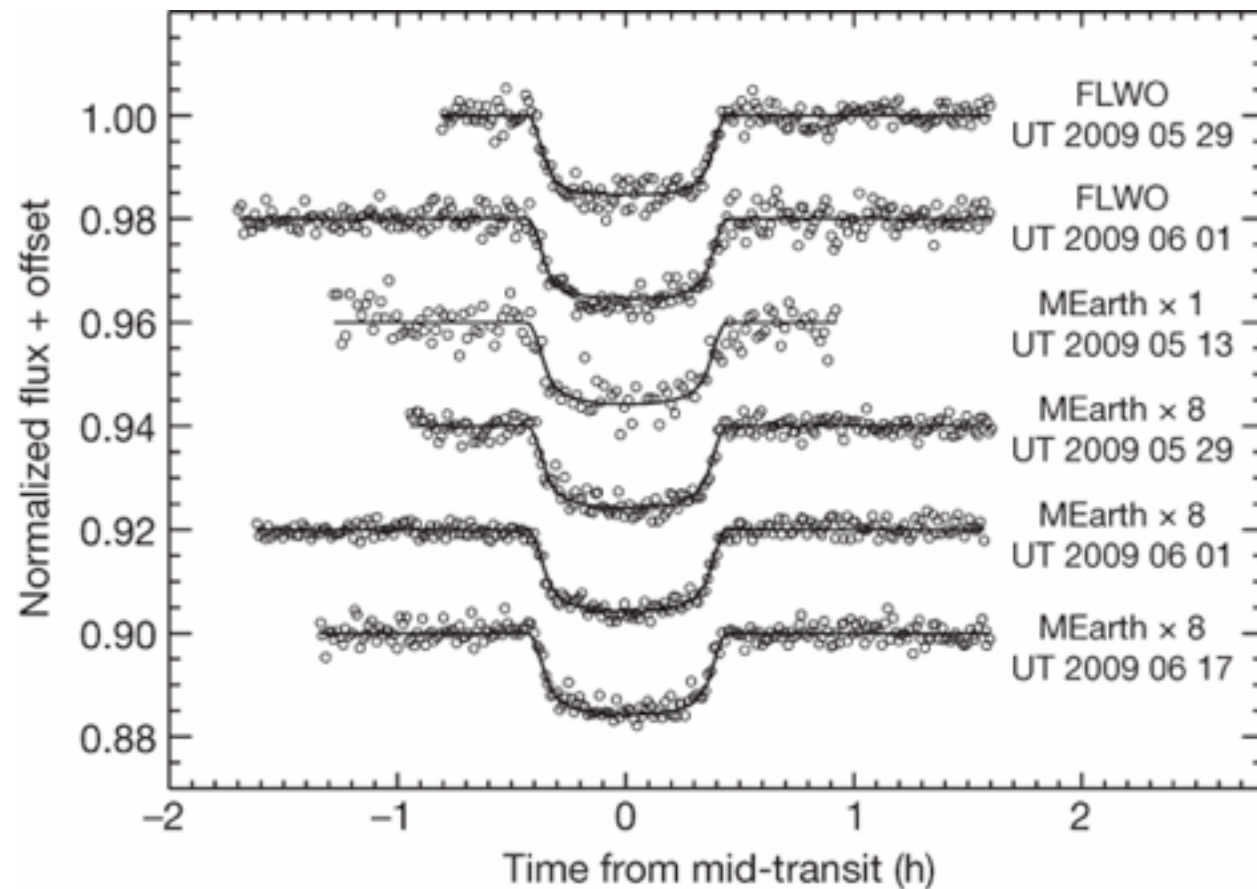
# Volatile Origin

Unconstrained

Rocky Origin



# GJ 1214b

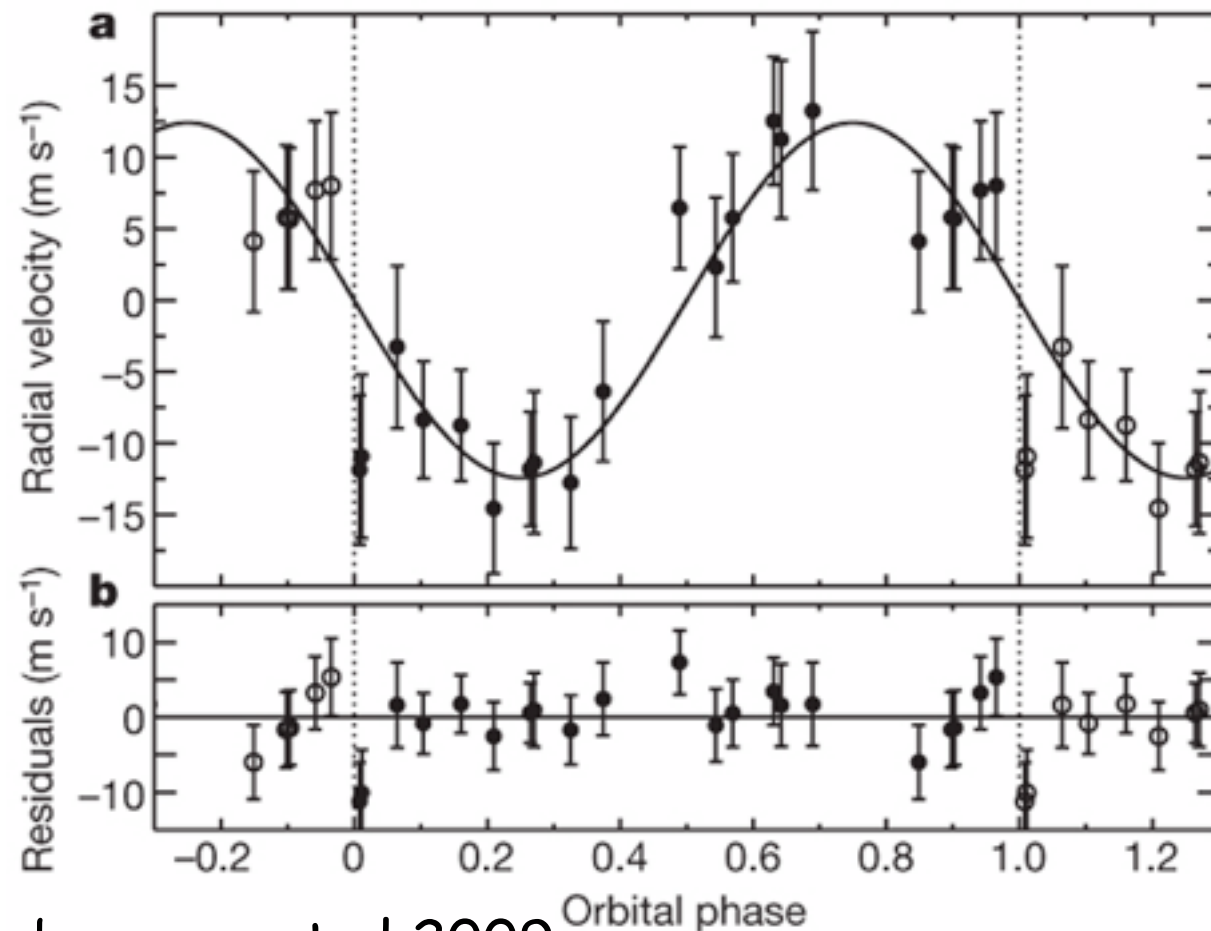


Radius =  $2.678 \pm 0.13 R_E$

Mass =  $6.55 \pm 0.98 M_E$

Period = 1.58 days

Temp = 393-555 K

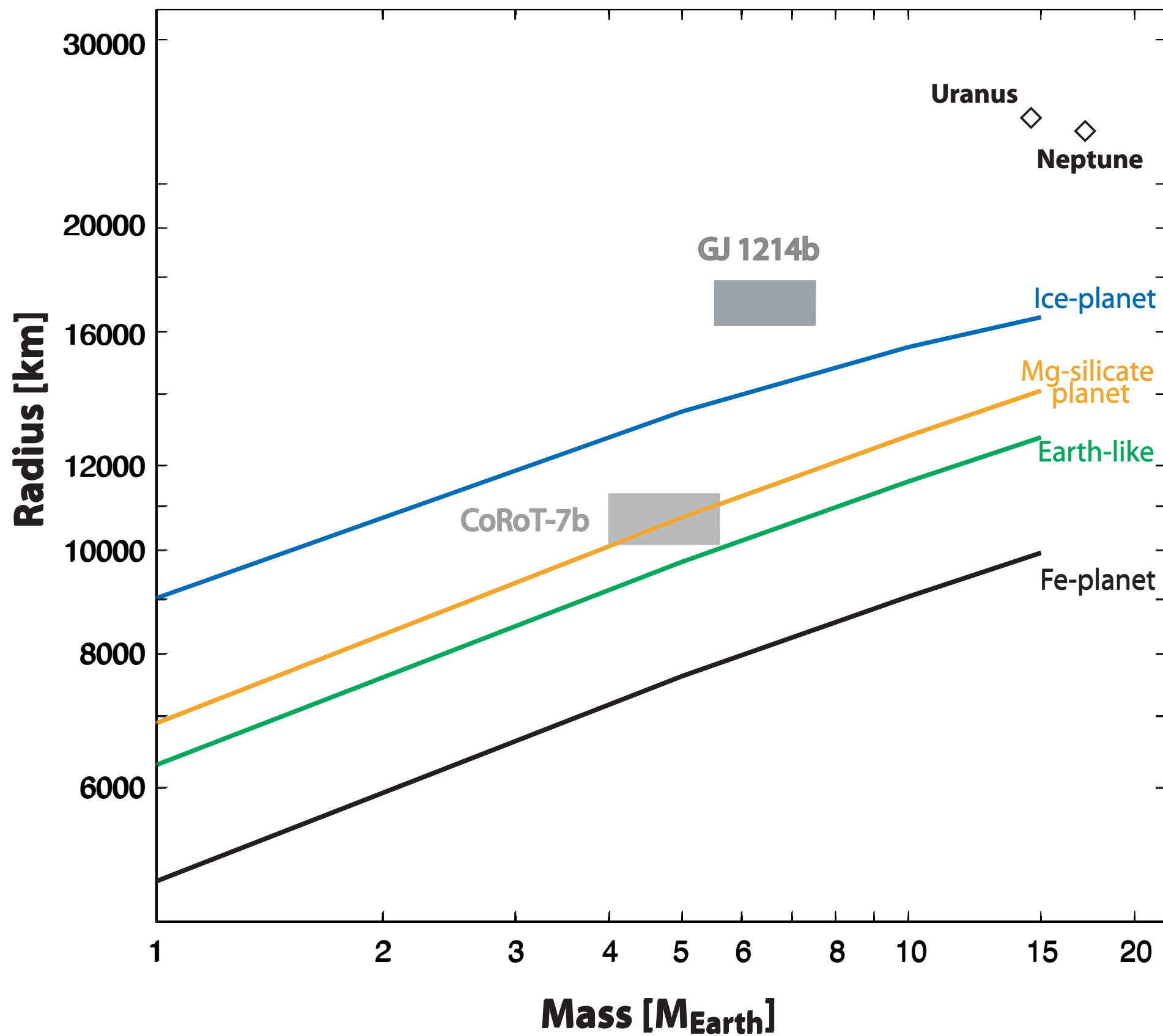


See also: Rogers & Seager, 2010

Miller-Ricci & Fortney 2010

Charbonneau et al 2009

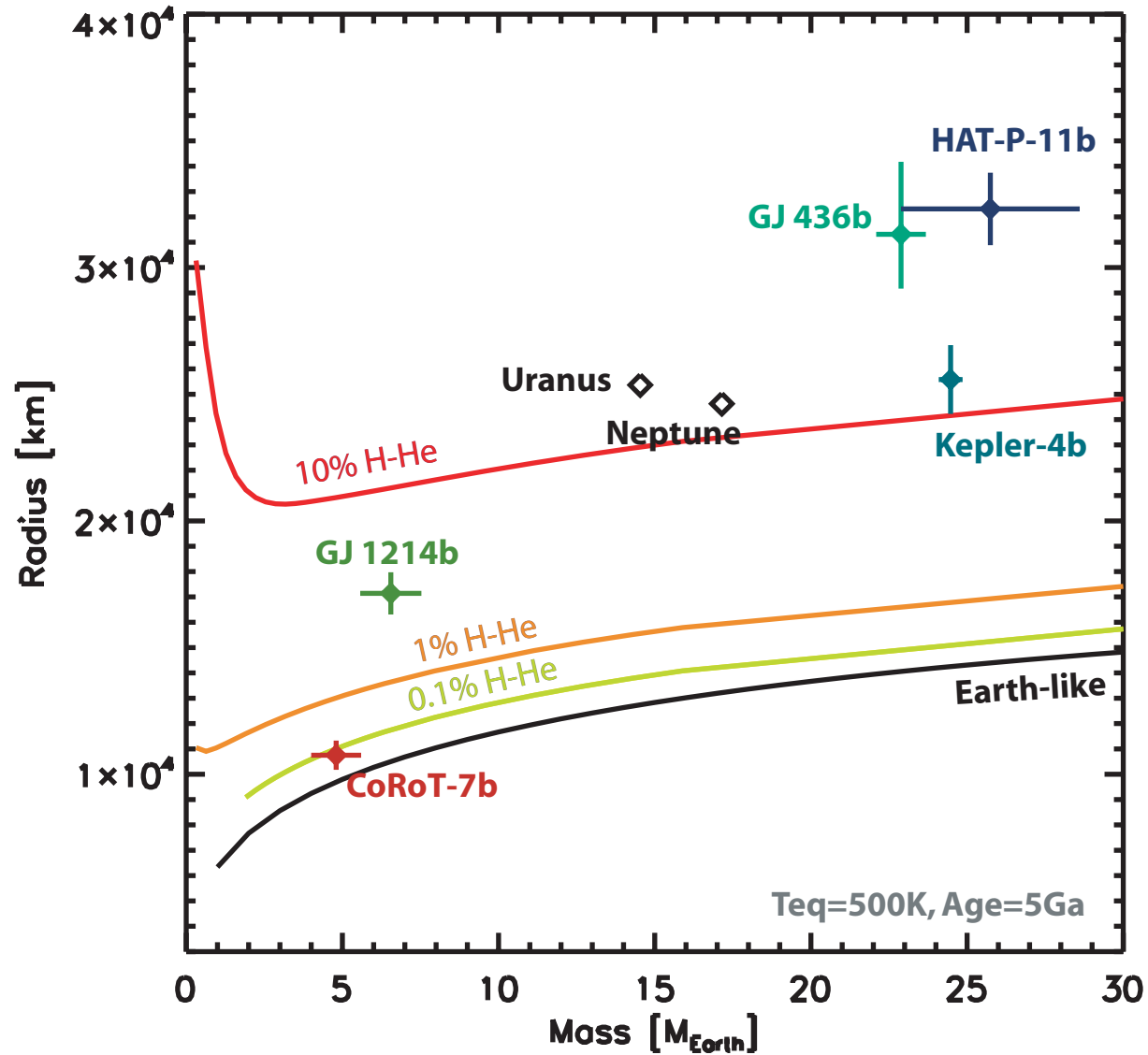
# Solid GJ 1214b ?



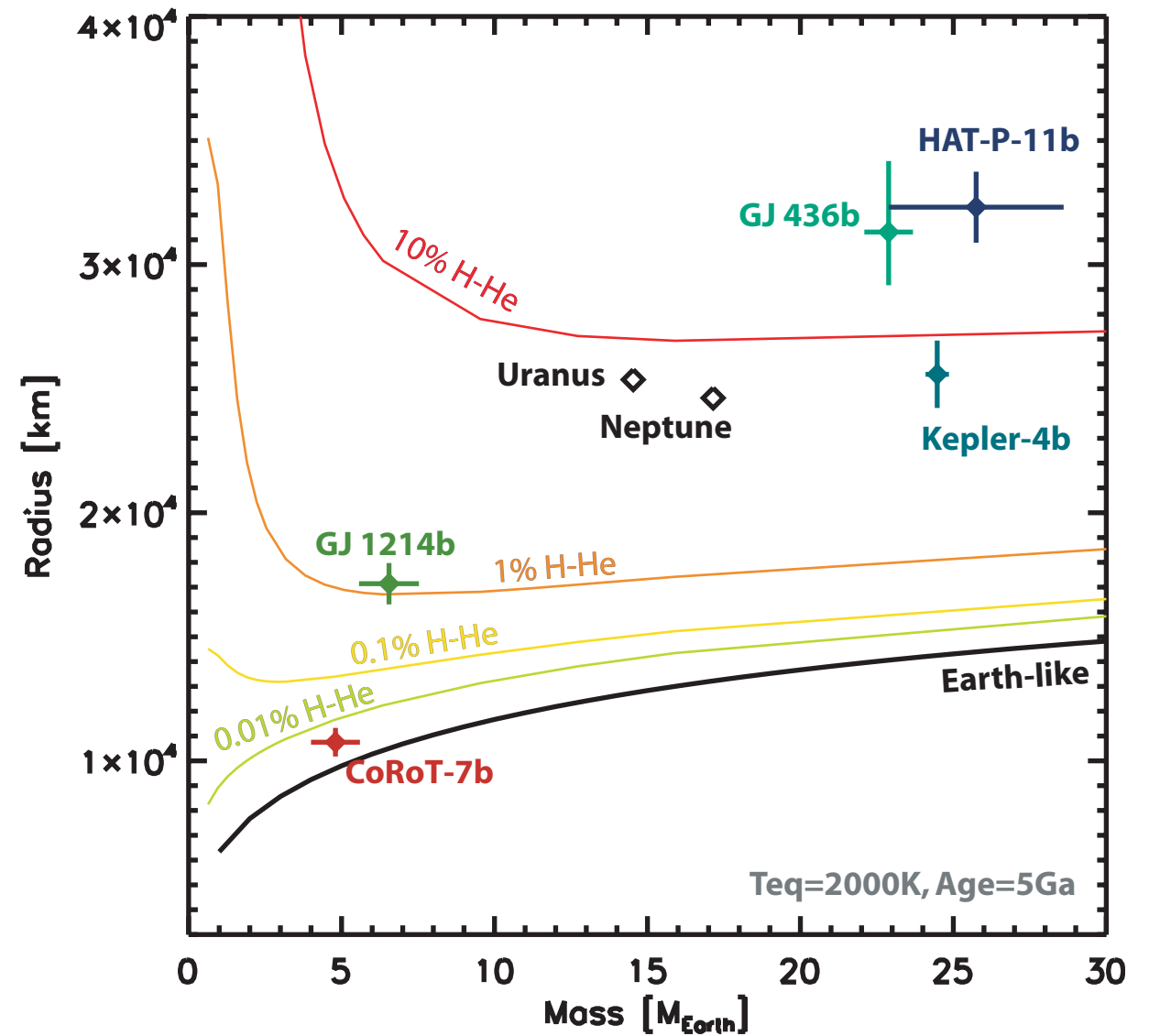


# H/He in GJ 1214b ?

Teq=500K



Teq=2000K



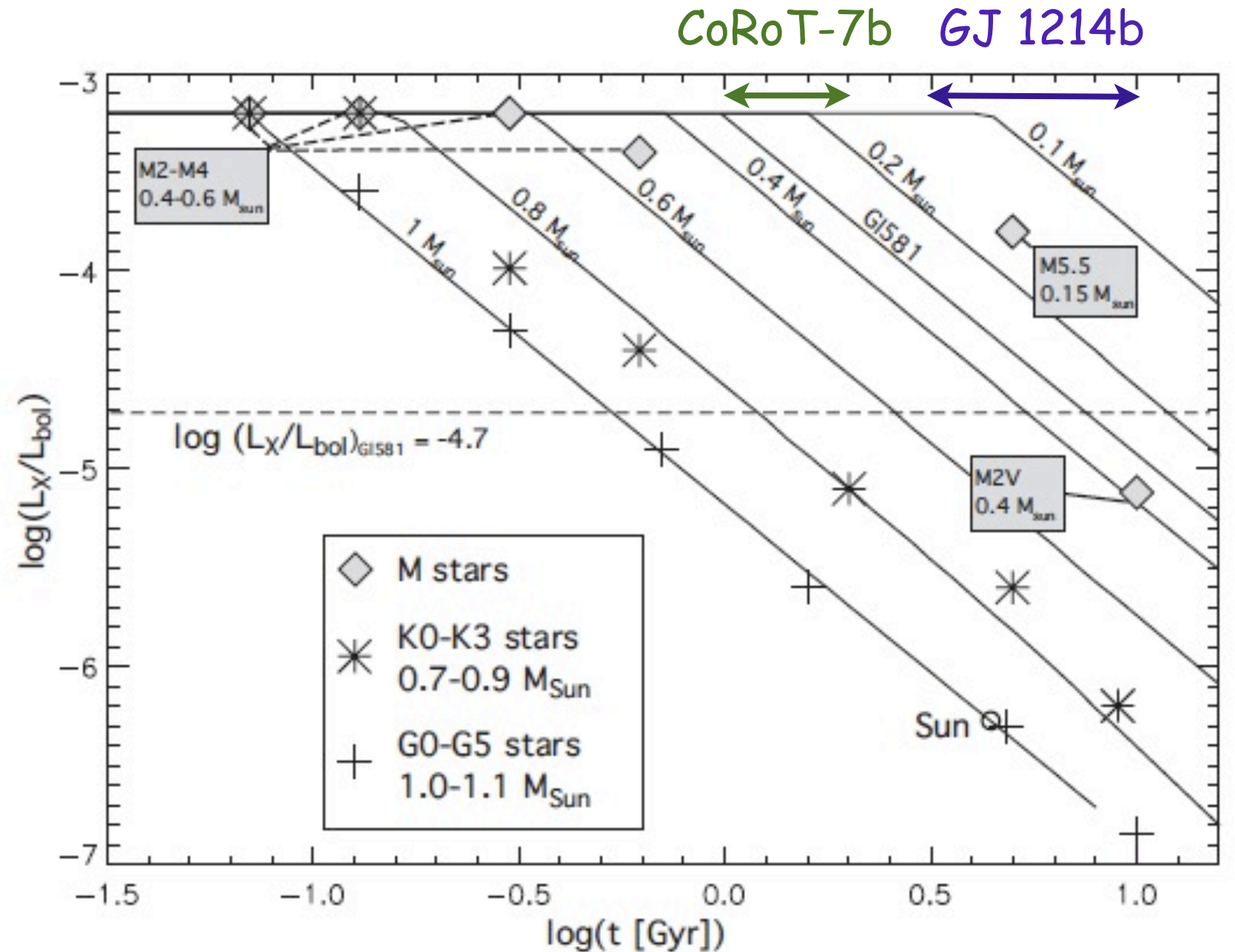
In preparation

# GJ 1214b: escape rate

$$\rho = 0.34 \rho_E$$

$$\text{Age} = 3\text{-}10 \text{ Gy}$$

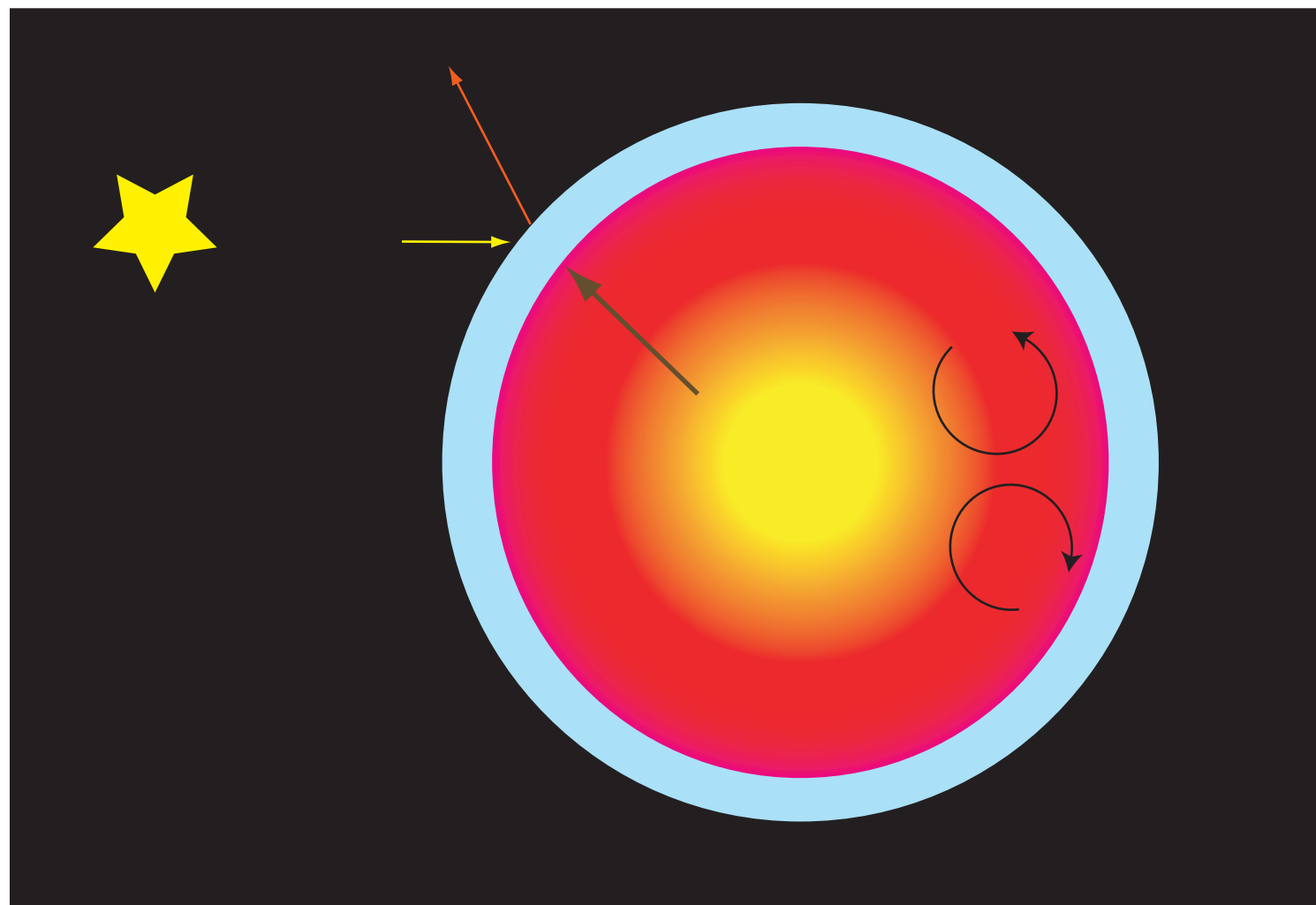
$$\frac{dM}{dt} = \frac{3 \epsilon F_{\text{EUV}}}{4 G \rho K_{\text{tide}}}$$



Selsis et al, 2007

# Habitability from a Planetary Perspective: Follow the water ...

Surface temperature depends on insolation, interior heat flux and atmospheric response

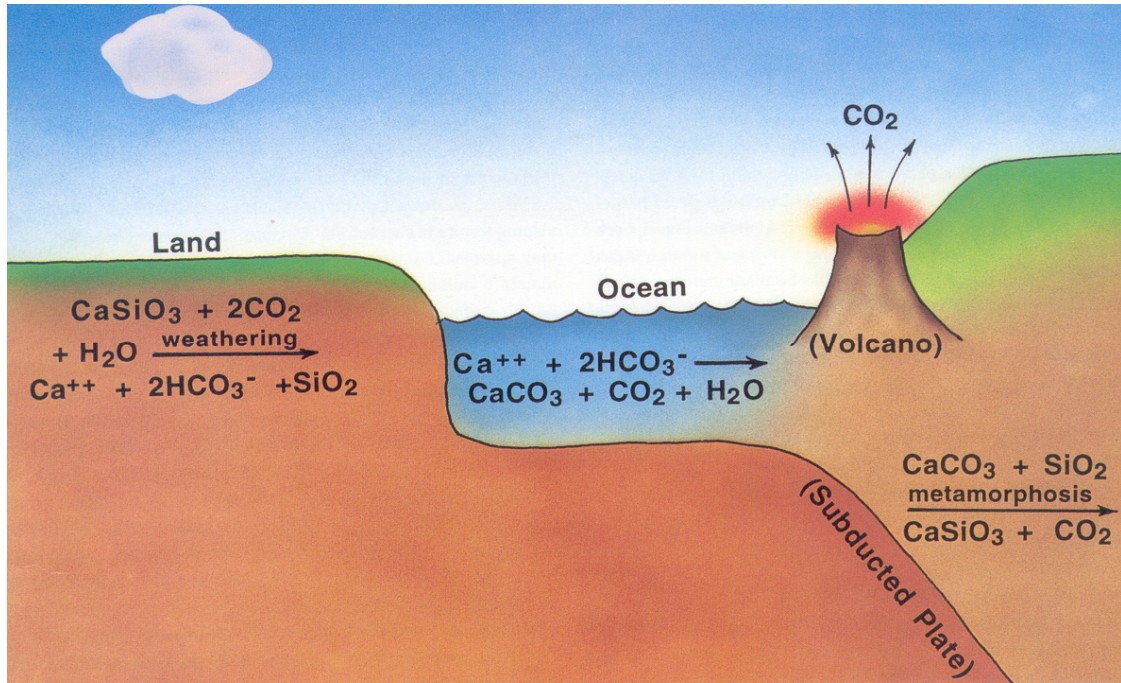
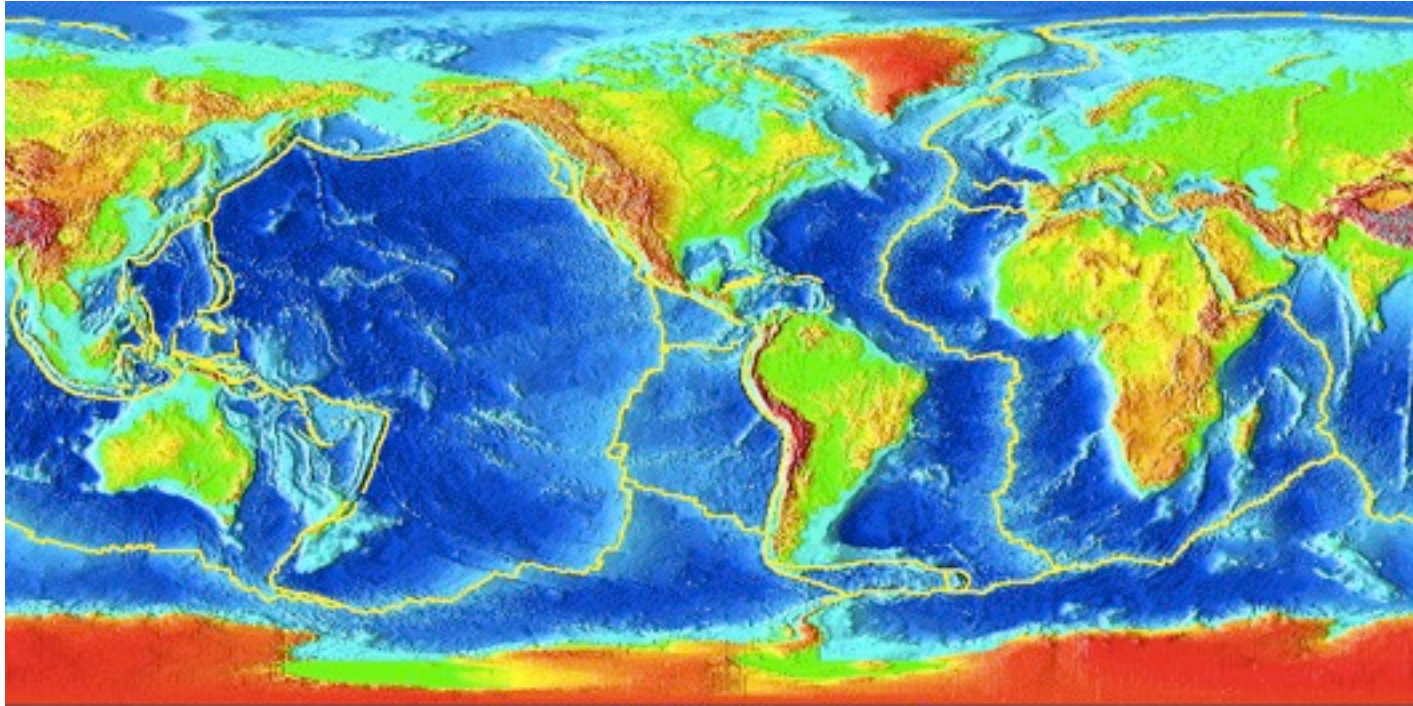


atmospheric composition

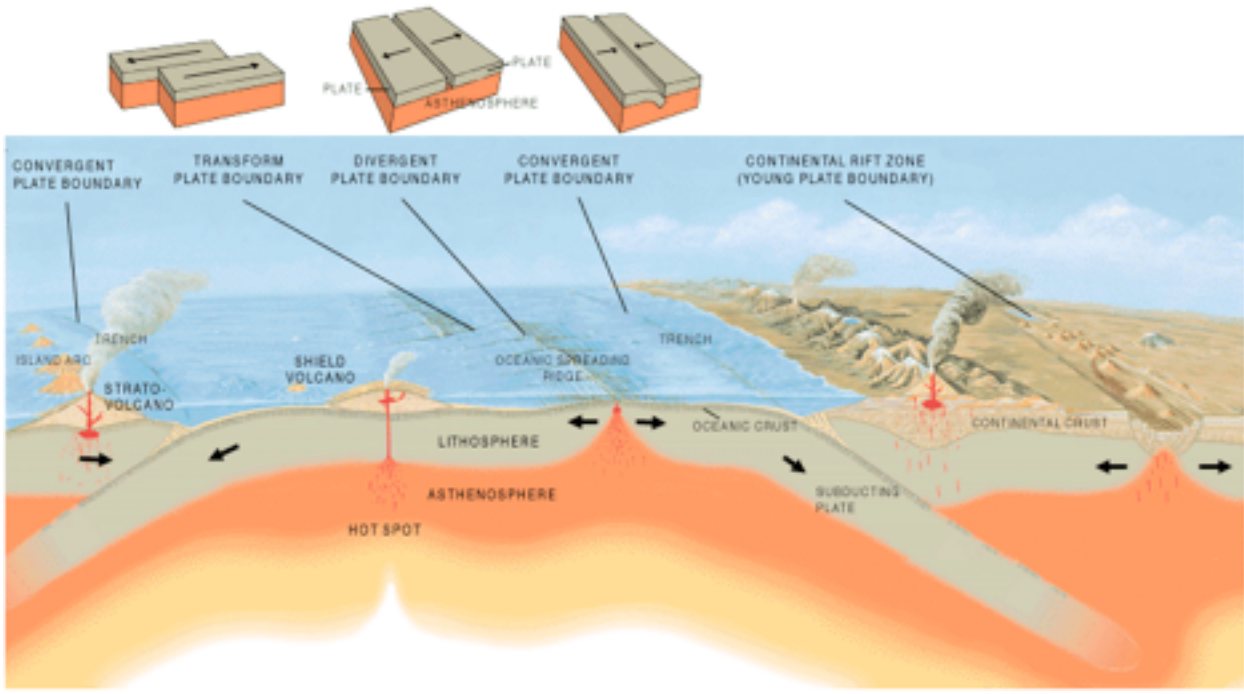
Interior processes:  
tectonics, volcanism,  
magnetic field



# Can a massive rocky planet have plate tectonics?



Walker 1981, Kasting 1996

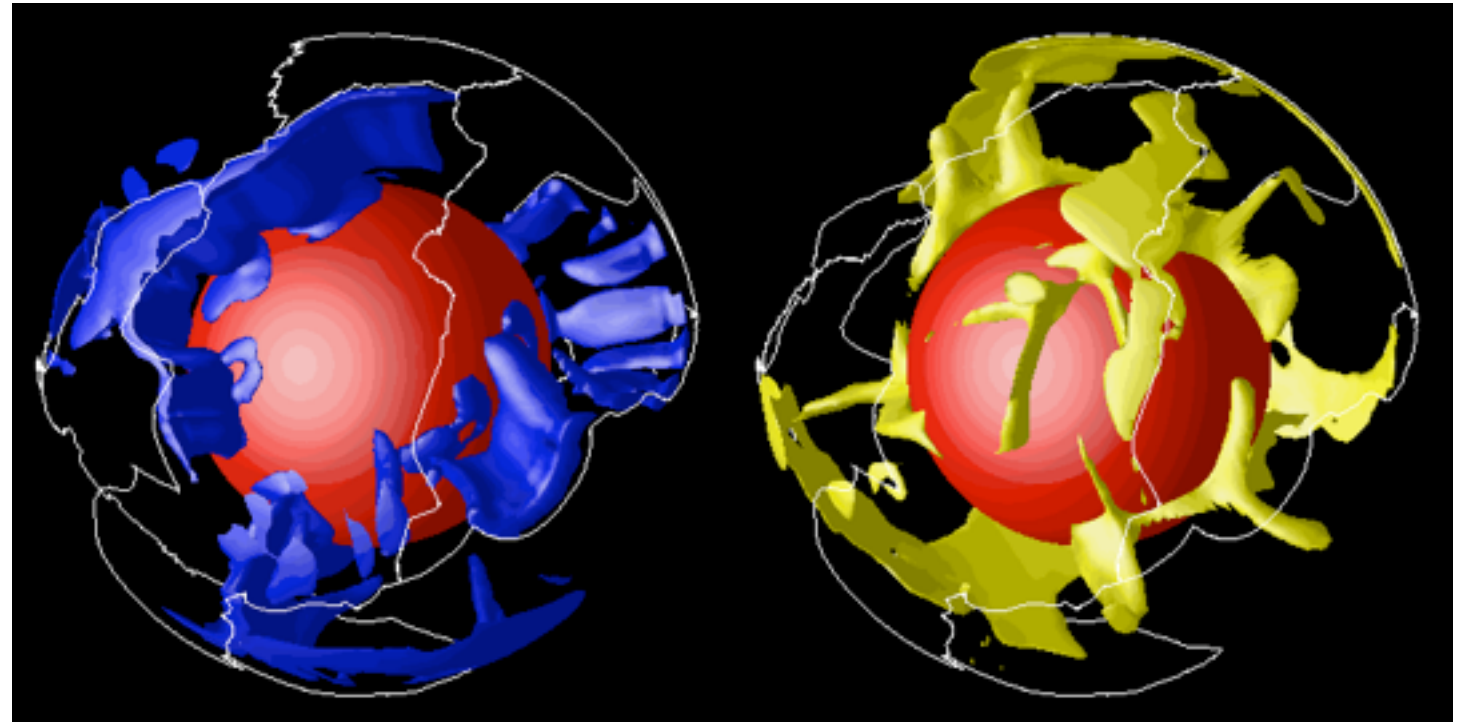


Strong, coherent plate;  
deformation on boundaries;  
PT is the surface expression  
of mantle convection

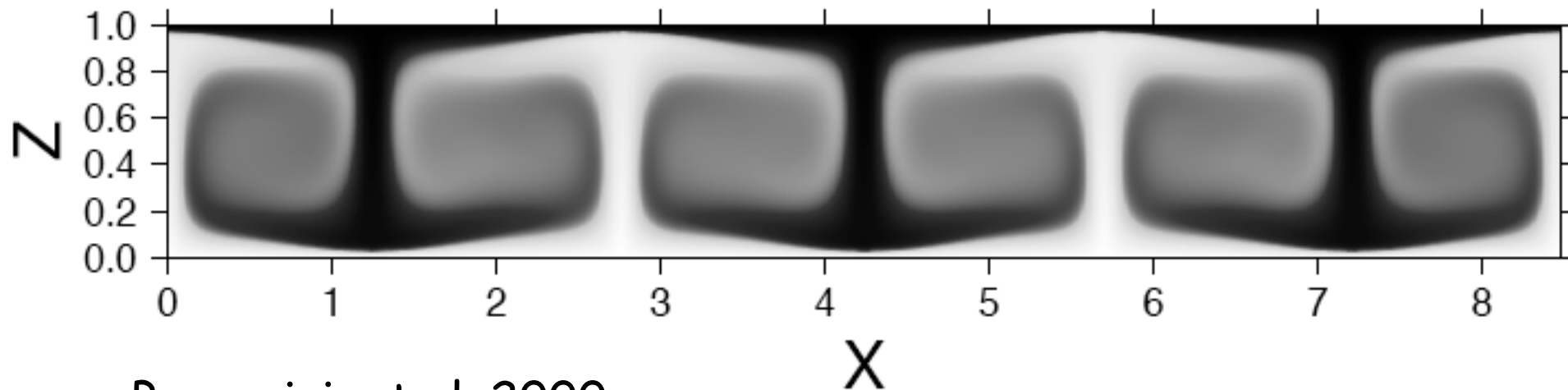


# Plate Tectonics and Mantle Convection

- Navier Stokes equations
- Classical Boundary layer theory
- Numerical modeling
- Laboratory Experiments



Bernard Convection



Bercovici, et al. 2000

# Conditions for PT

1. Deformation of the plate

Valencia et al, 2007, 2009  
O'Neill & Lenardic 2007  
Van Heken & Tackley, 2009  
Sotin & Jackson, 2009

2. Negative buoyancy

Valencia, 2009  
Kite et al 2009

3. Energy dissipation at subduction zones is not enough to halt PT

Valencia et al, 2009

Can terrestrial super-Earths  
have plate tectonics?



# Can terrestrial super-Earths have plate tectonics?

under debate

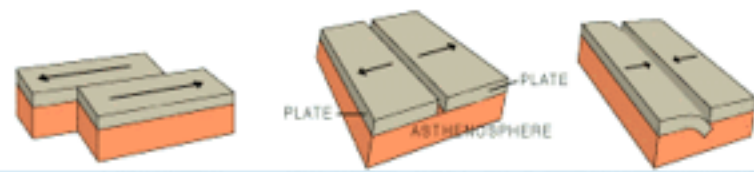
# Can terrestrial super-Earths have plate tectonics?

under debate

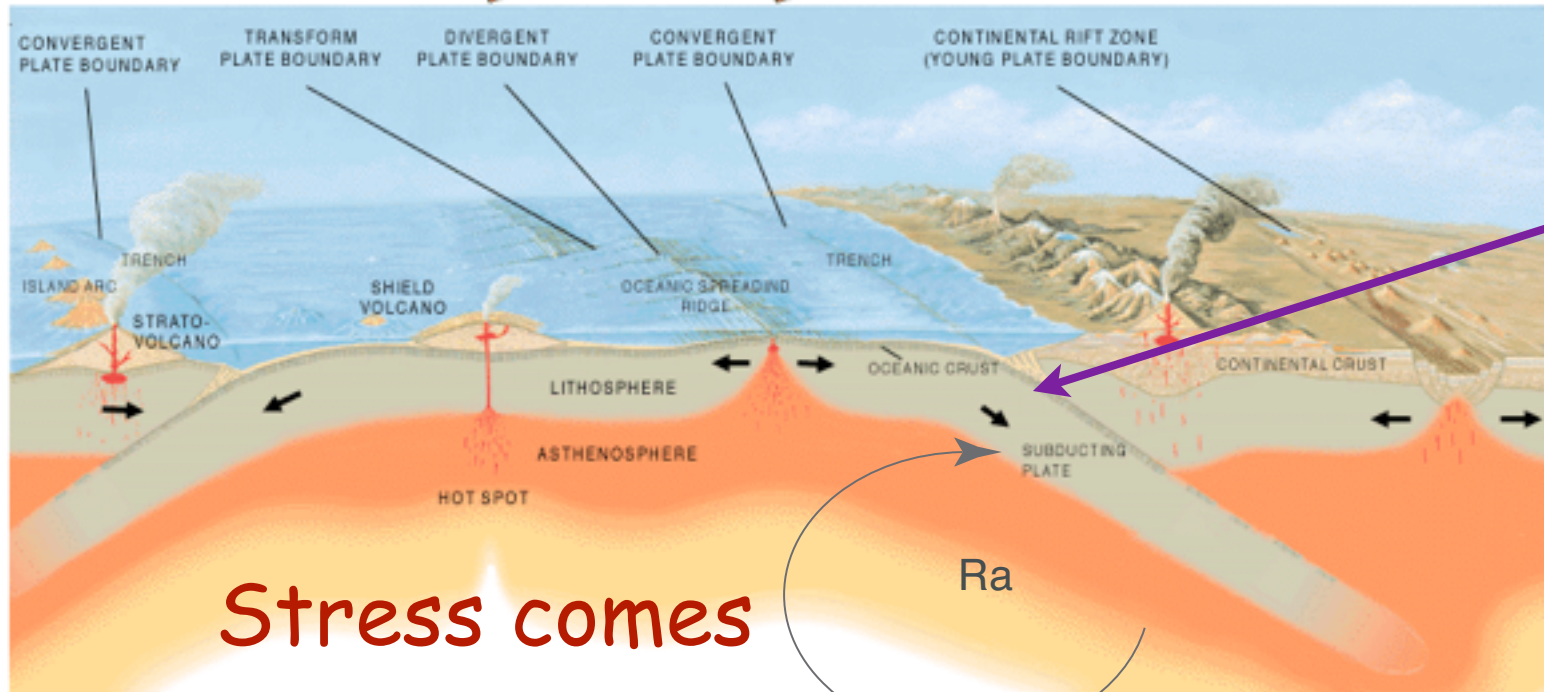
Valencia et al 2007: "... as mass increases, the process of subduction, and hence plate tectonics, becomes easier. Therefore, massive super-Earths will very likely exhibit plate tectonics"

O'Neill and Lenardic 2007: "...these results suggest super-Earths may in fact be in an episodic or stagnant-lid regime, rather than a mobile lid regime similar to Earth's plate tectonics."

# Similarities



Fixed Tsurf

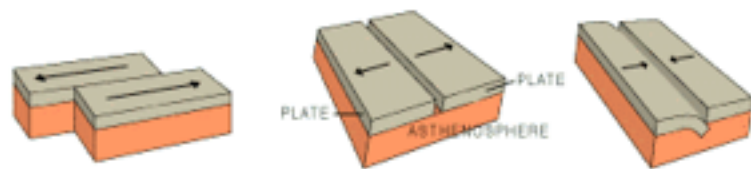


Coulomb failure criterion for plate deformation

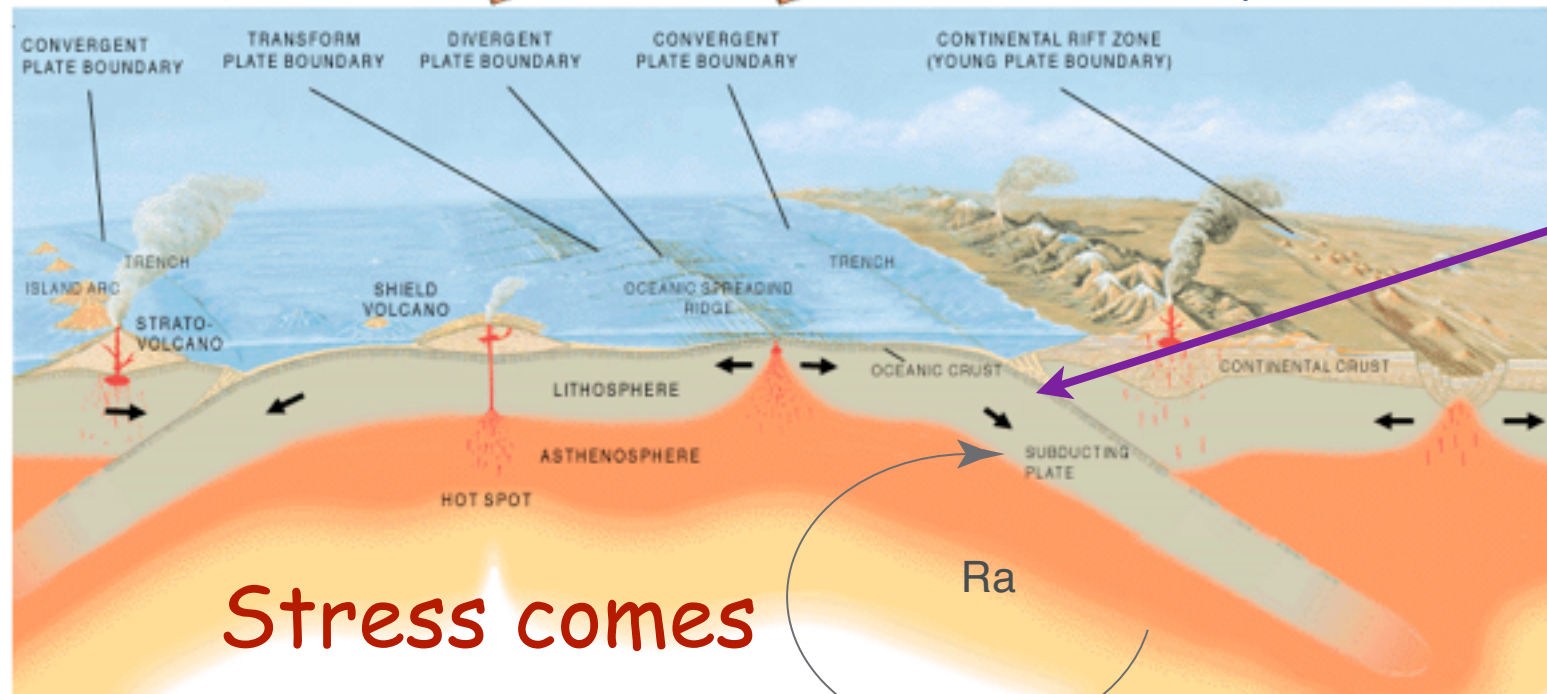
Stress comes from convection

Ra

# Similarities



Fixed Tsurf



Coulomb failure criterion for plate deformation

Stress comes from convection

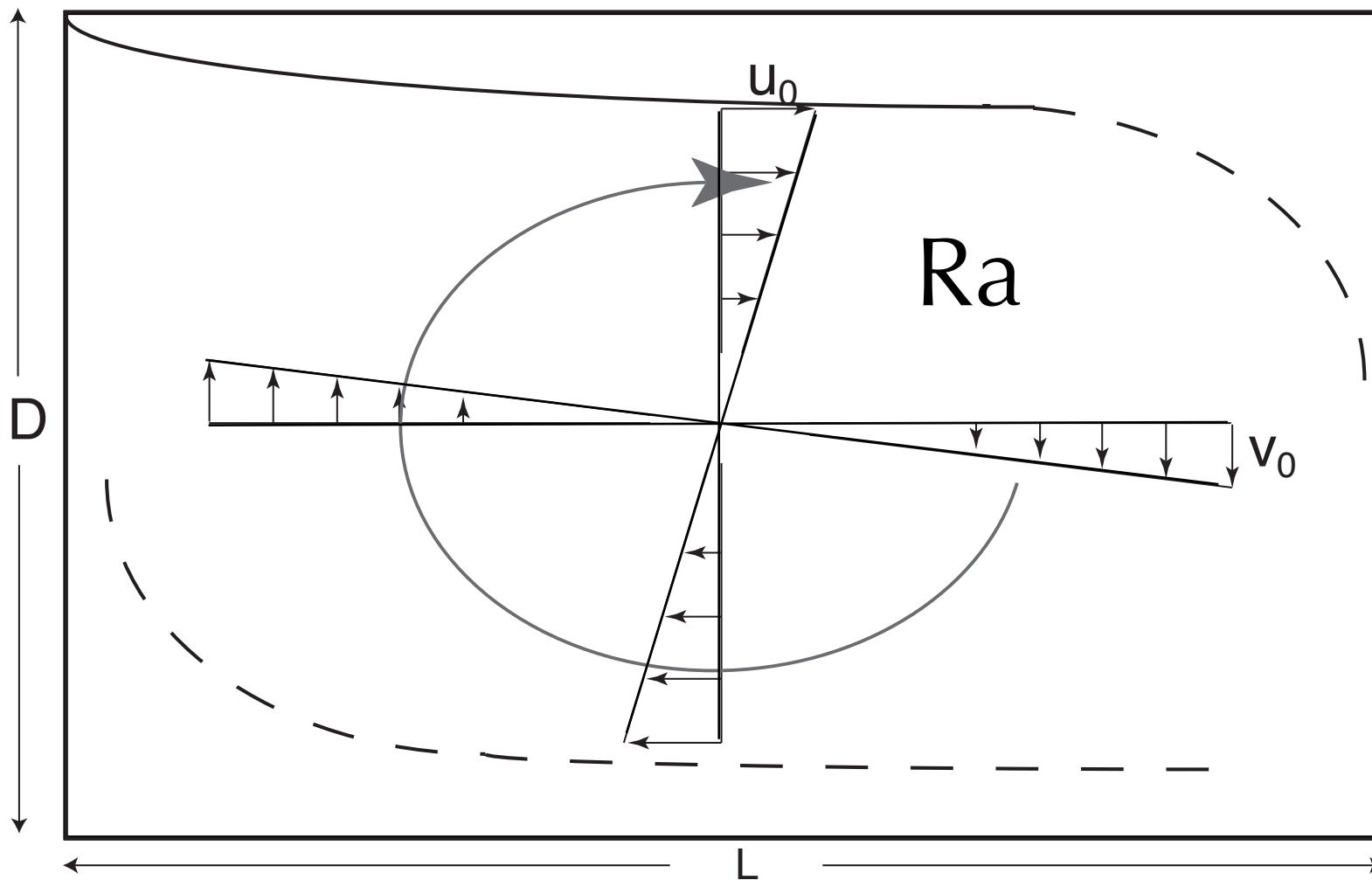
# Differences

Valencia et al:  
Classical Boundary Layer  
Structure model to scale  
parameters  
Dominated by radioactive heating

OL 2007  
Scaled a numerical model  
(Moresi and Solomatov, 1998)  
Constant density scaling  
Mixed heating



# Classical Boundary Layer Theory



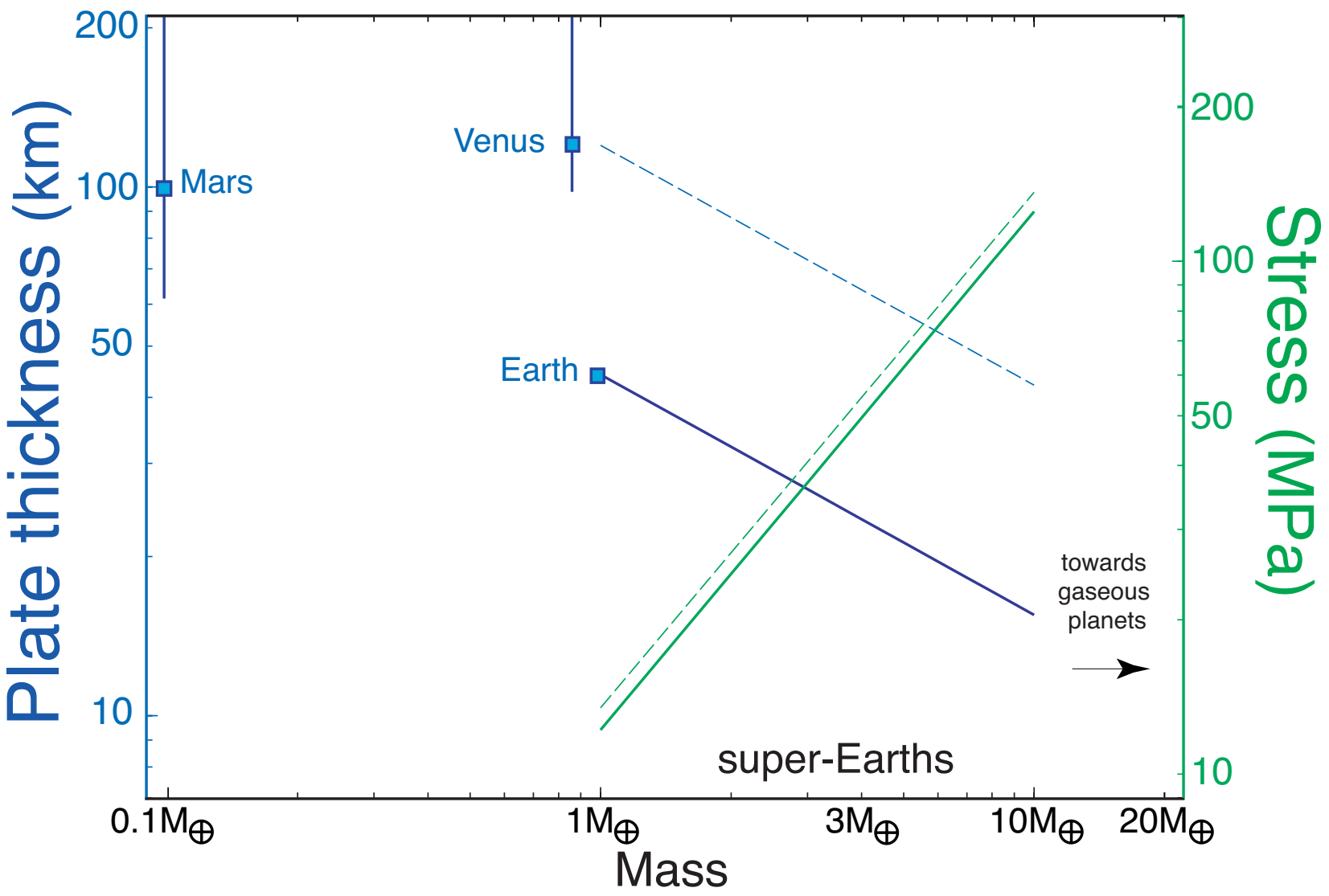
velocity  $u$   
 stress  $\Delta \tau_{xy}$   
 plate length  $L$   
 timescale

Turcotte & Schubert, 2002

**Internally heated:**  $Ra = \rho g \alpha D^4 q / \kappa \eta (T) k$

$$D/2\delta \sim Ra^{1/4}$$

# Convective Parameters on super-Earths



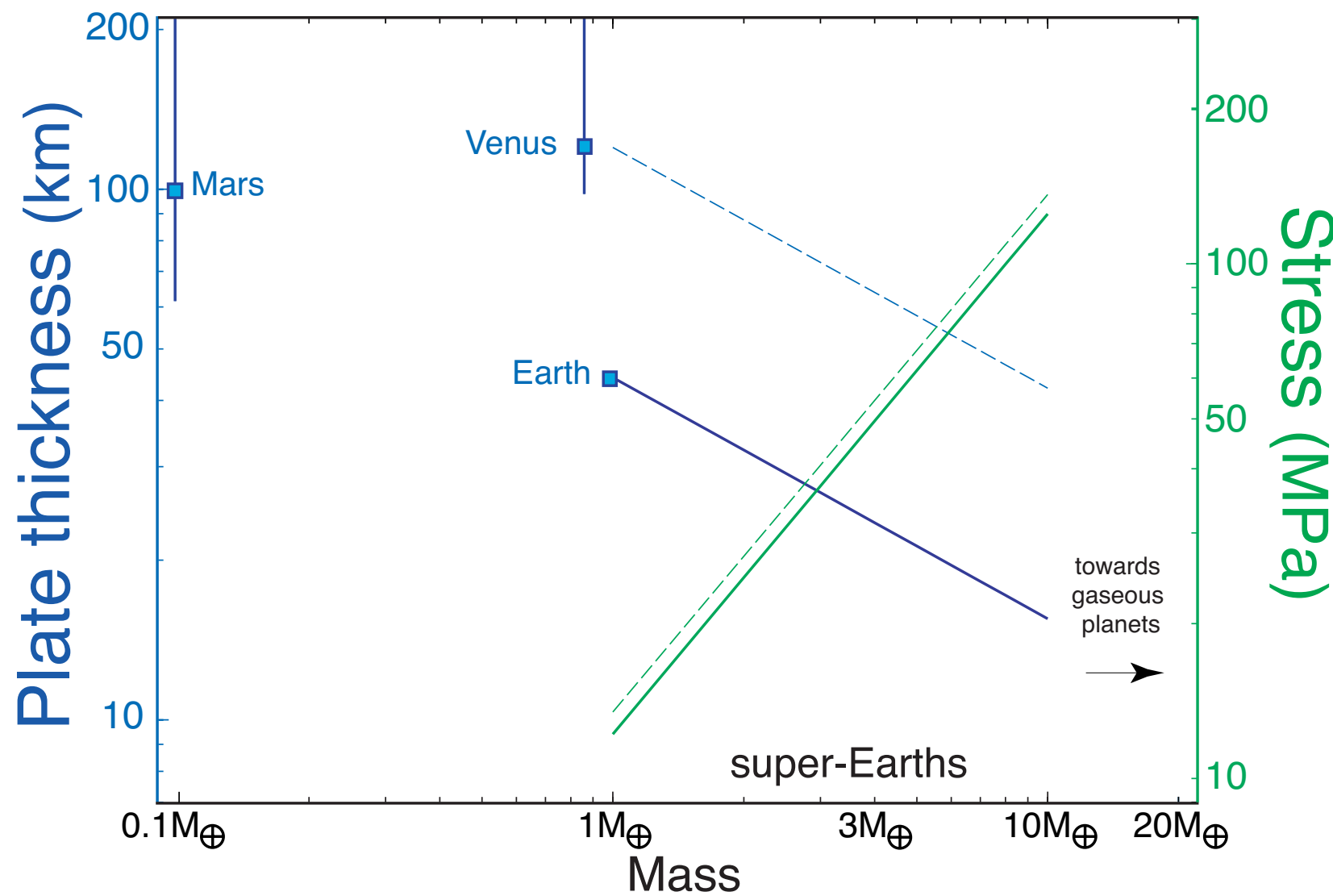
$$u \sim \kappa/D \text{ Ra}^{1/2}$$

$$\Delta \tau_{xy} \sim \eta(T) u/D$$

Valencia et al., 2007c

# Convective Parameters on super-Earths

Plate P-T structure is nearly invariant with mass



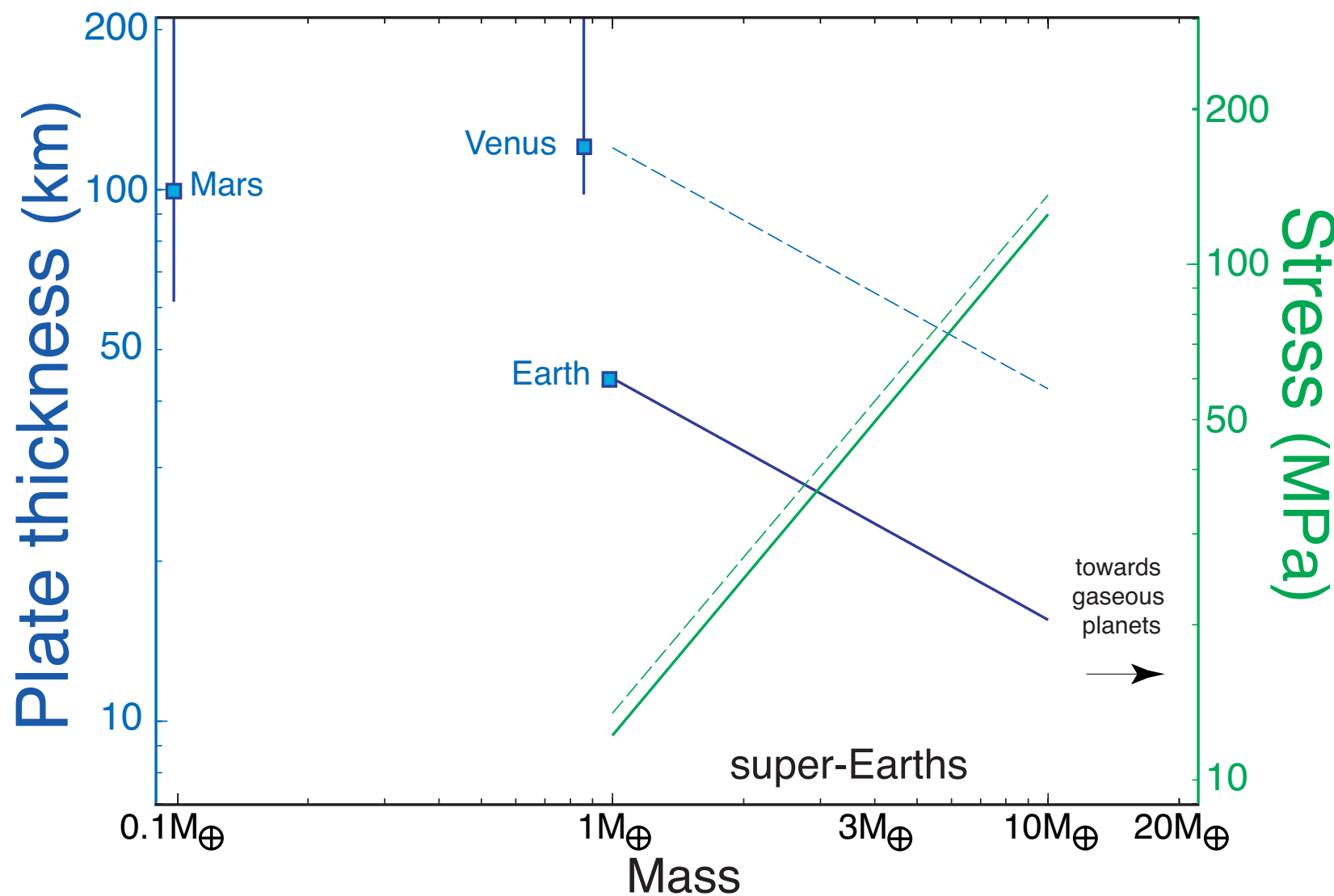
$$u \sim \kappa/D \text{ Ra}^{1/2}$$

$$\Delta \tau_{xy} \sim \eta(T) u/D$$

Valencia et al., 2007c

# Convective Parameters on super-Earths

Plate P-T structure is nearly invariant with mass



$$u \sim \kappa/D \text{ Ra}^{1/2}$$

$$\Delta \tau_{xy} \sim \eta(T) u/D$$

Valencia et al., 2007c

Super-Earth's: thinner plates and larger driving forces



# Plate deformation: Coulomb Failure Criterion

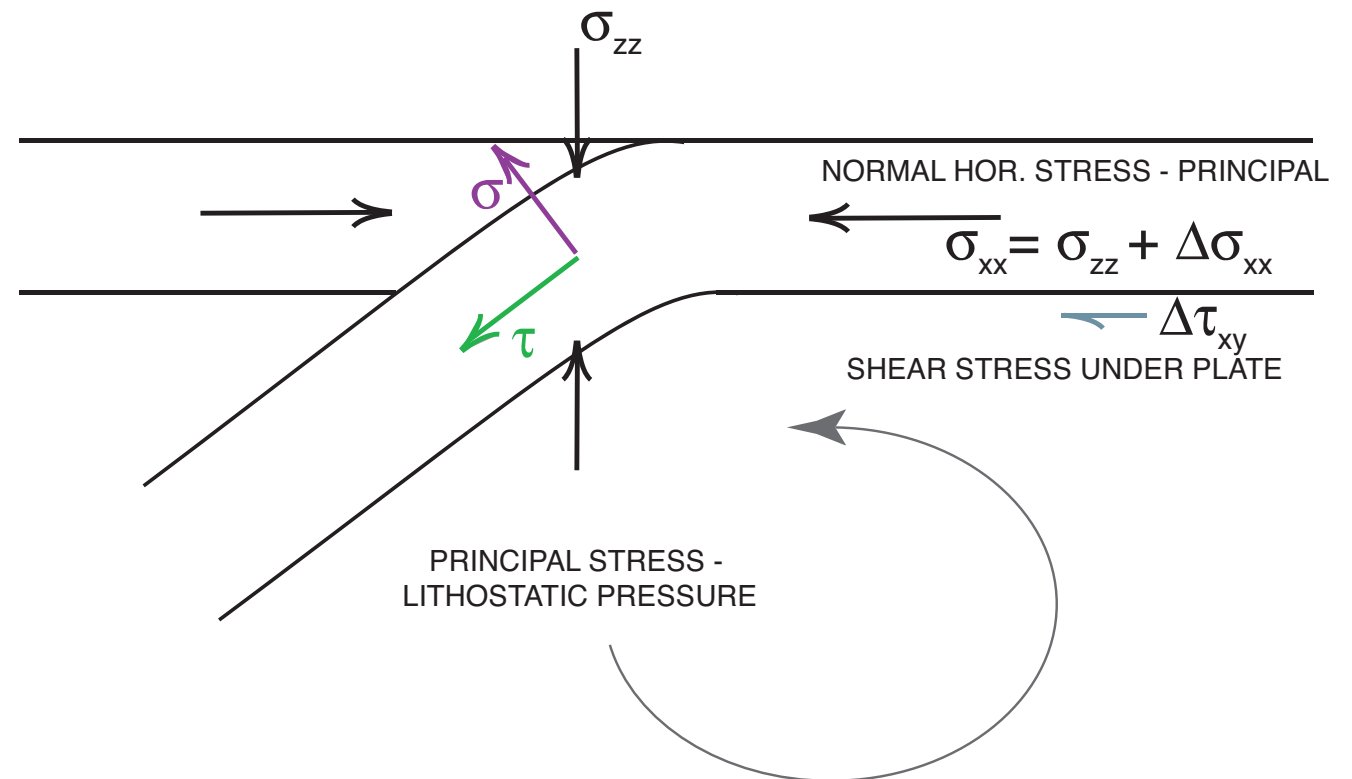
For deformation:

$$\frac{\tau}{\text{yield stress}} > 1$$

$$\tau = 0.5 \Delta\sigma_{xx} \sin 2\theta$$

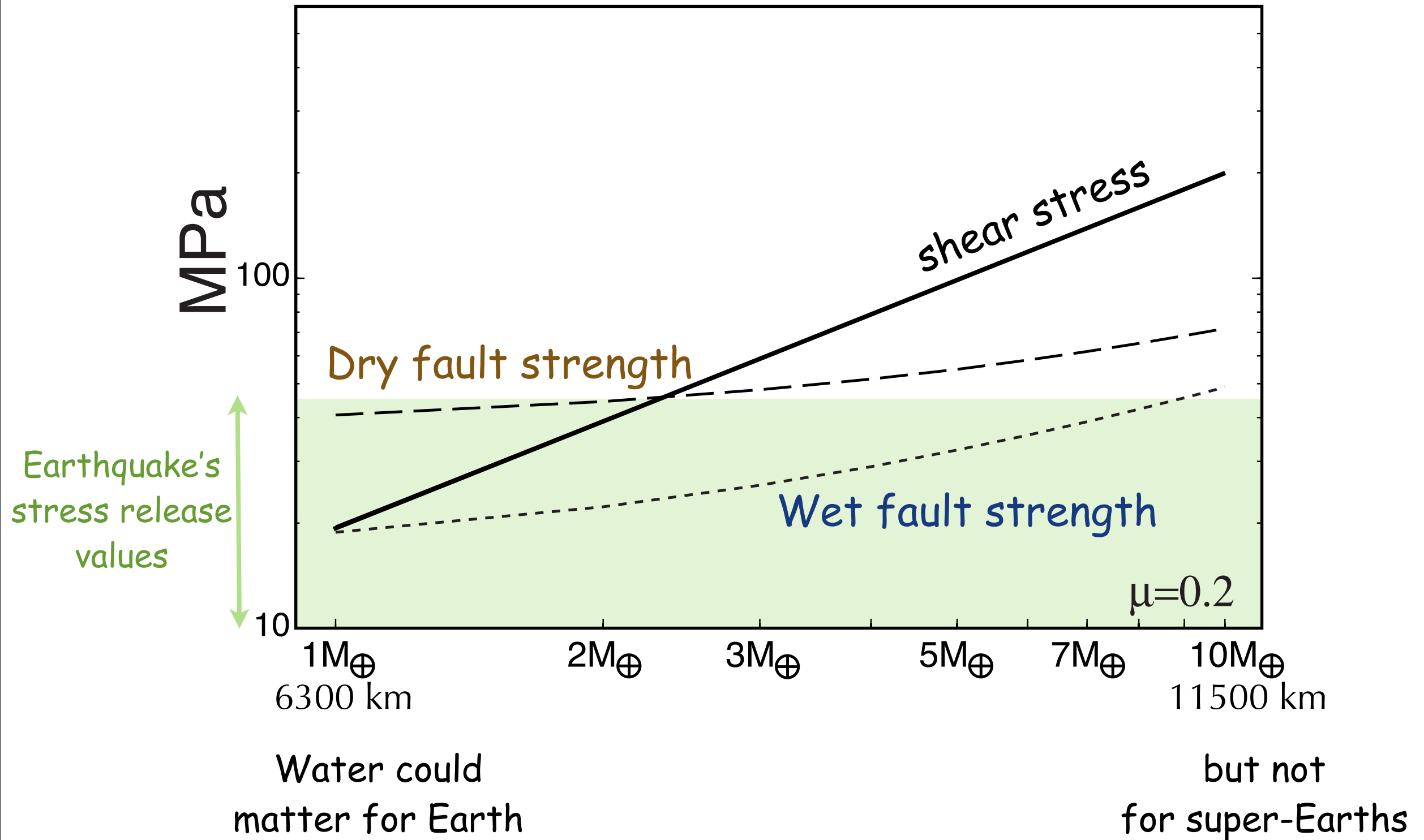
$$\sigma = \sigma_{zz} + 0.5 \Delta\sigma_{xx} (1 + \cos 2\theta)$$

$$\text{yield stress} = S_0 + \mu(\sigma - \lambda\sigma_{zz})$$



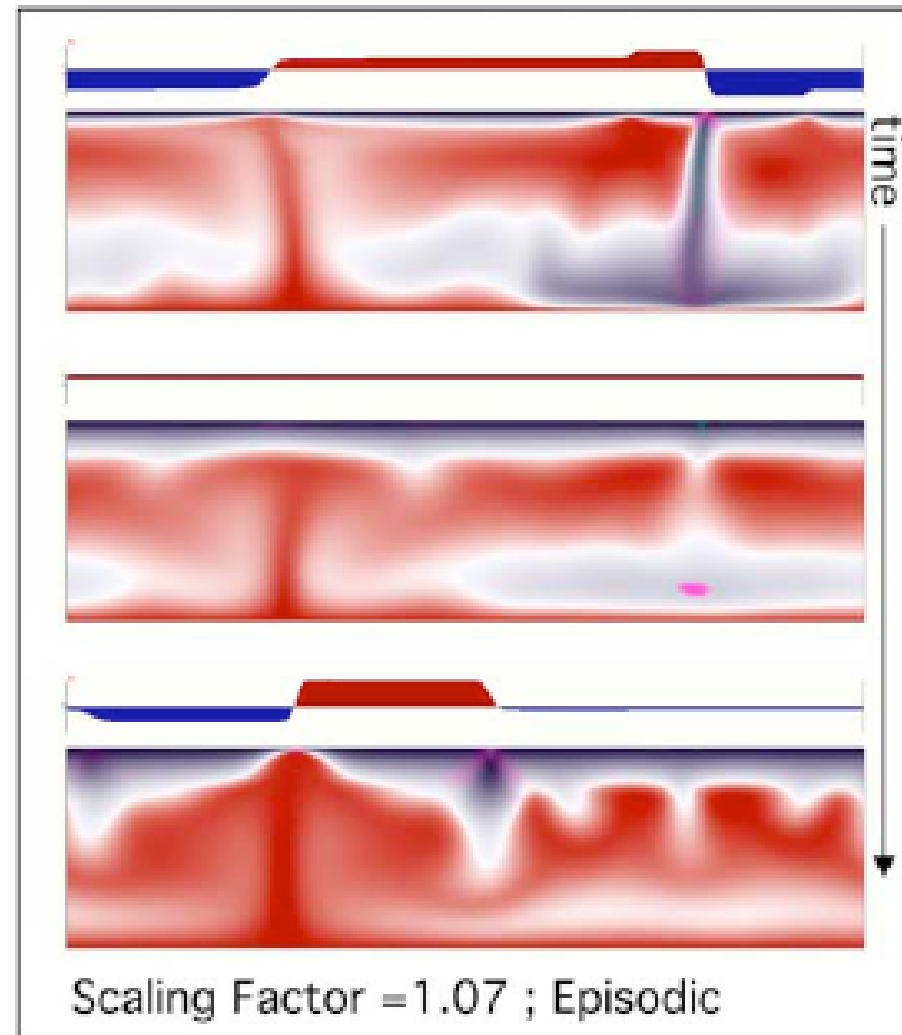
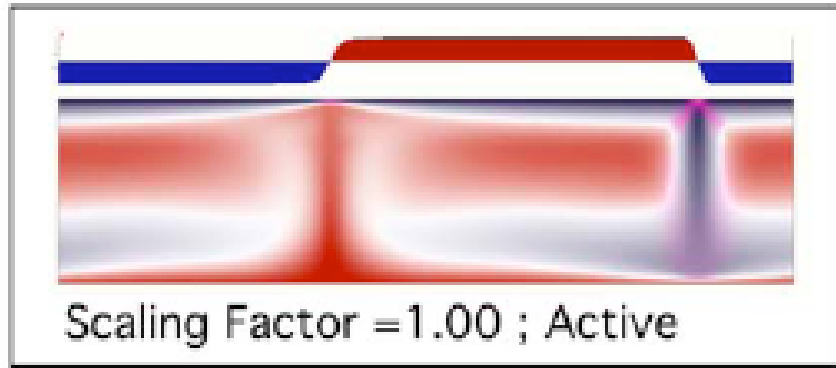
Water reduces strength

# Stress vs. Strength on Faults



Valencia et al., 2009b

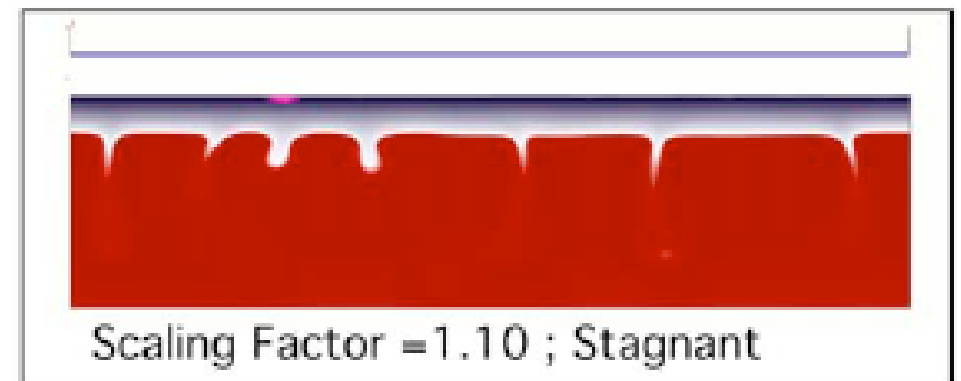
# PT on super-Earths ?



Numerical Model by:  
Moresi and Solomatov (1998)

Coulomb Failure Criterion

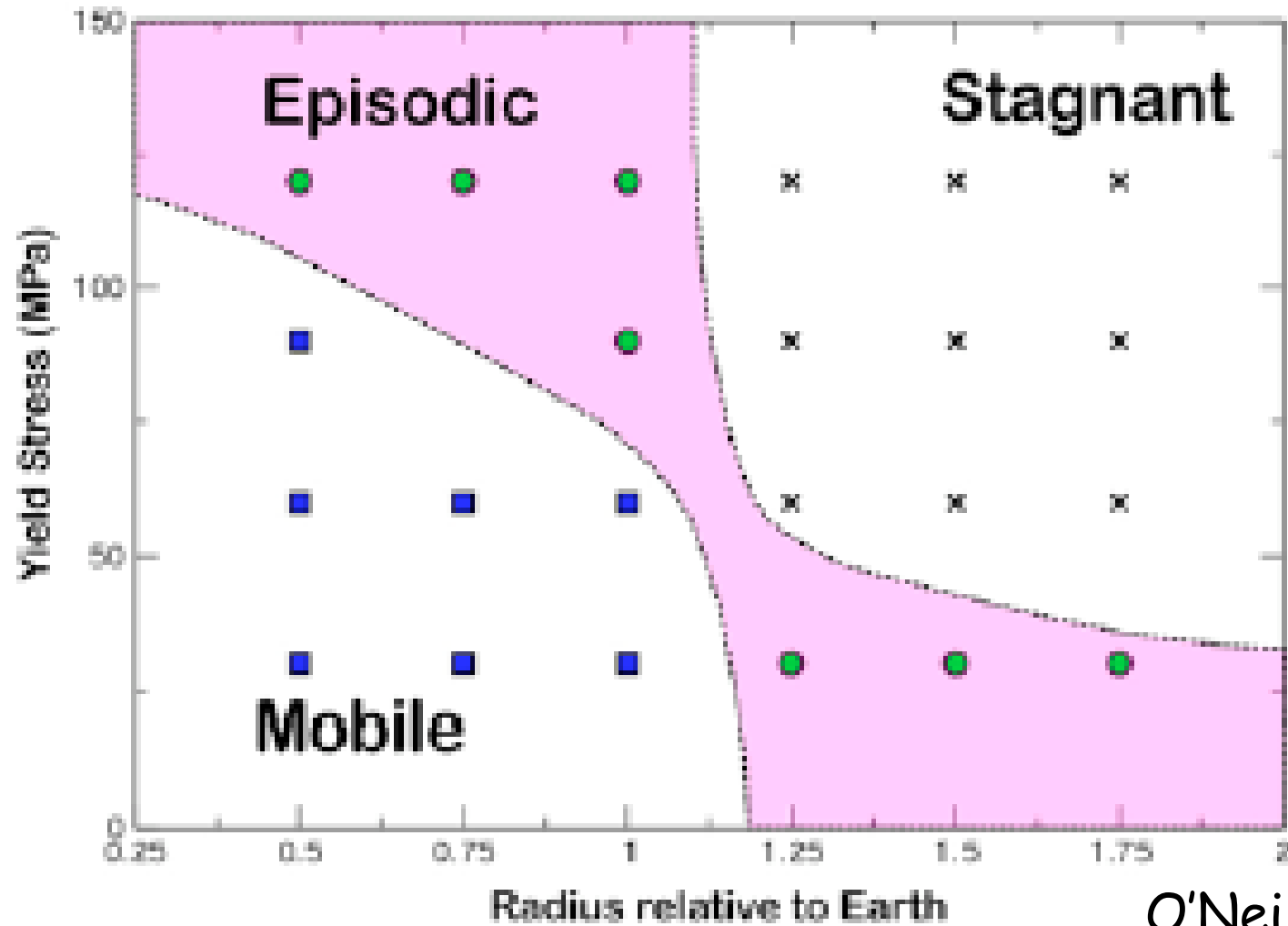
Non-newtonian rheology (plateness, zones of weakness)



Scaling Factor =  $R/R_E$

O'Neill & Lenardic, 2007

# Plate Tectonics on super-Earths



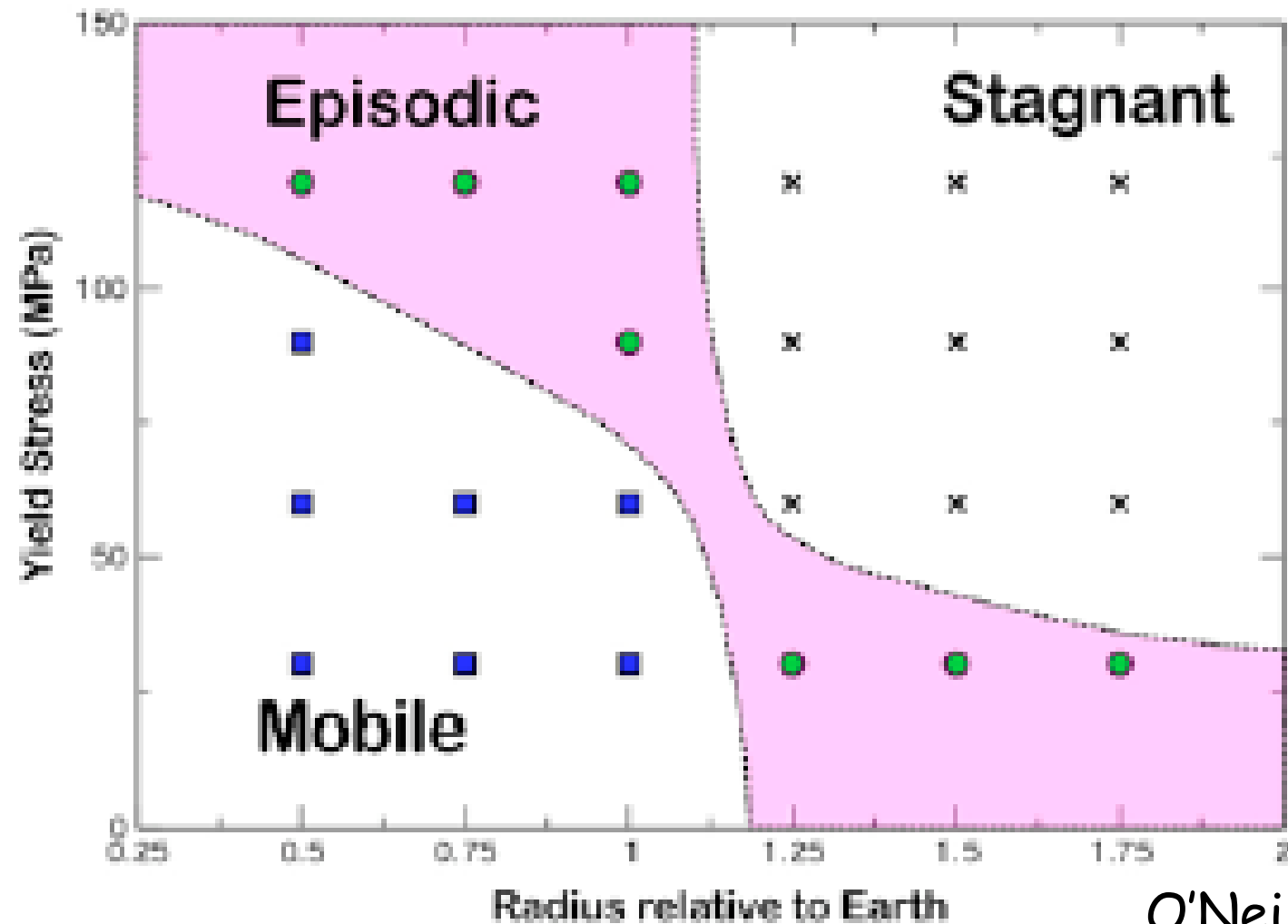
O'Neill & Lenardic, 2007

$$u \sim \kappa/D \text{ Ra}^{1/2}$$

$$\Delta \tau_{xy} \sim \eta(T) u/D$$



# Plate Tectonics on super-Earths



O'Neill & Lenardic, 2007

$$u \sim \kappa/D \text{ Ra}^{1/2}$$

$$\Delta \tau_{xy} \sim \eta(T) u/D$$

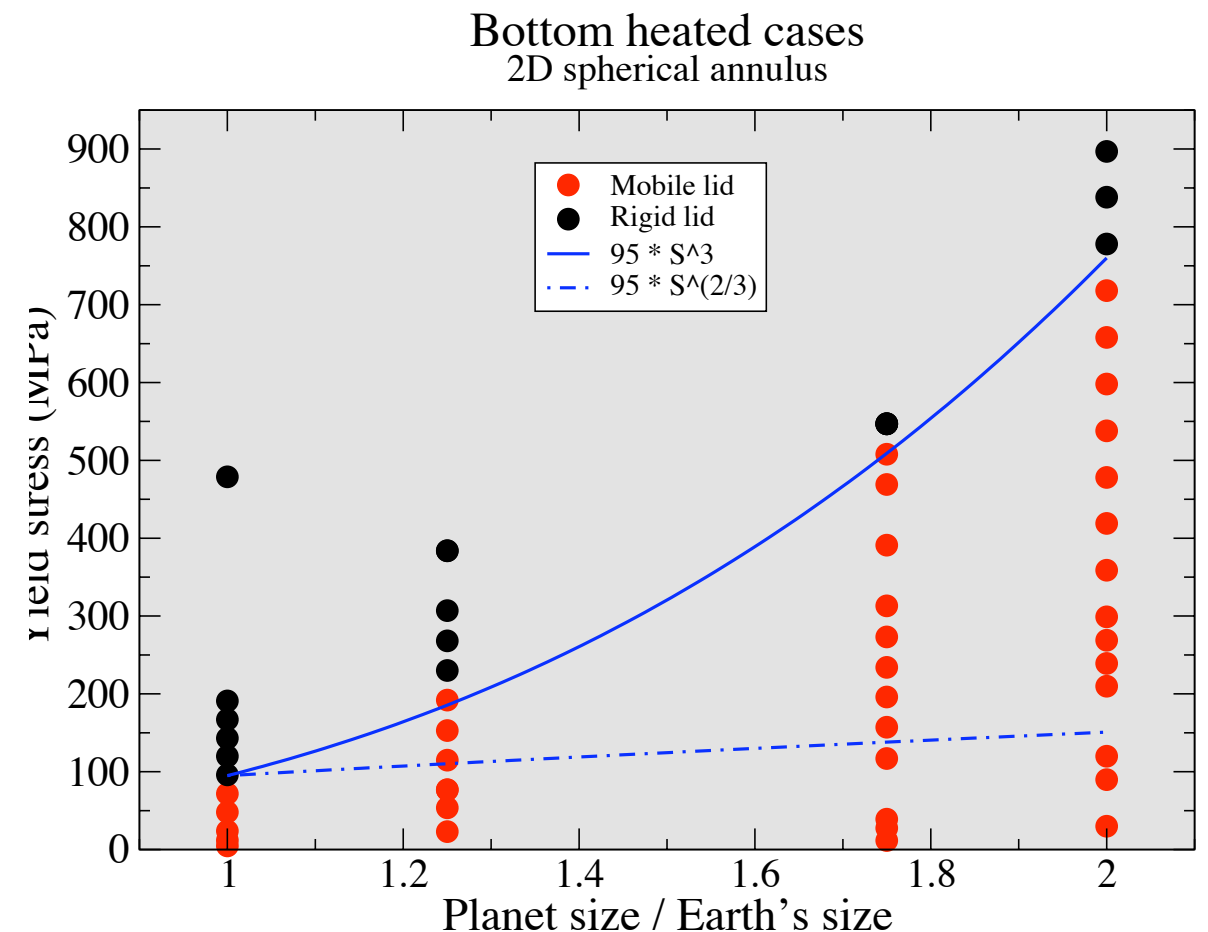
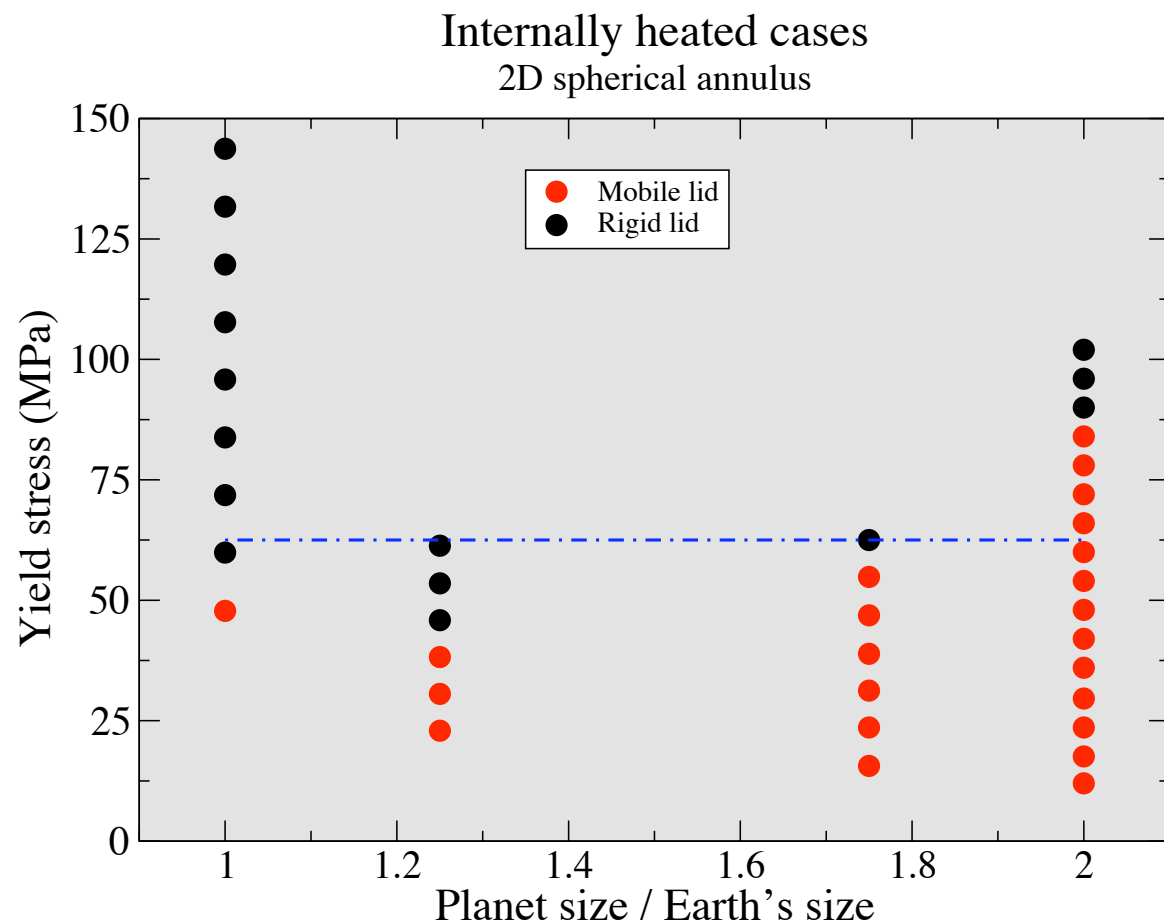
Small planets are cold and large planets are hot.

Convective stress is dominated by internal viscosity, so hotter means lower convective stress. If the yield of the lithosphere is constant, hot planets can not have plate tectonics

# Other Studies

Sotin & Jackson 2009 predict a high stress/strength ratio for internally heated and mixed heated systems

Van Heck & Tackley 2009 show that for planets that are heated internally the stress/strength ratio is constant with mass, and for planets that are heated from below the ratio increases

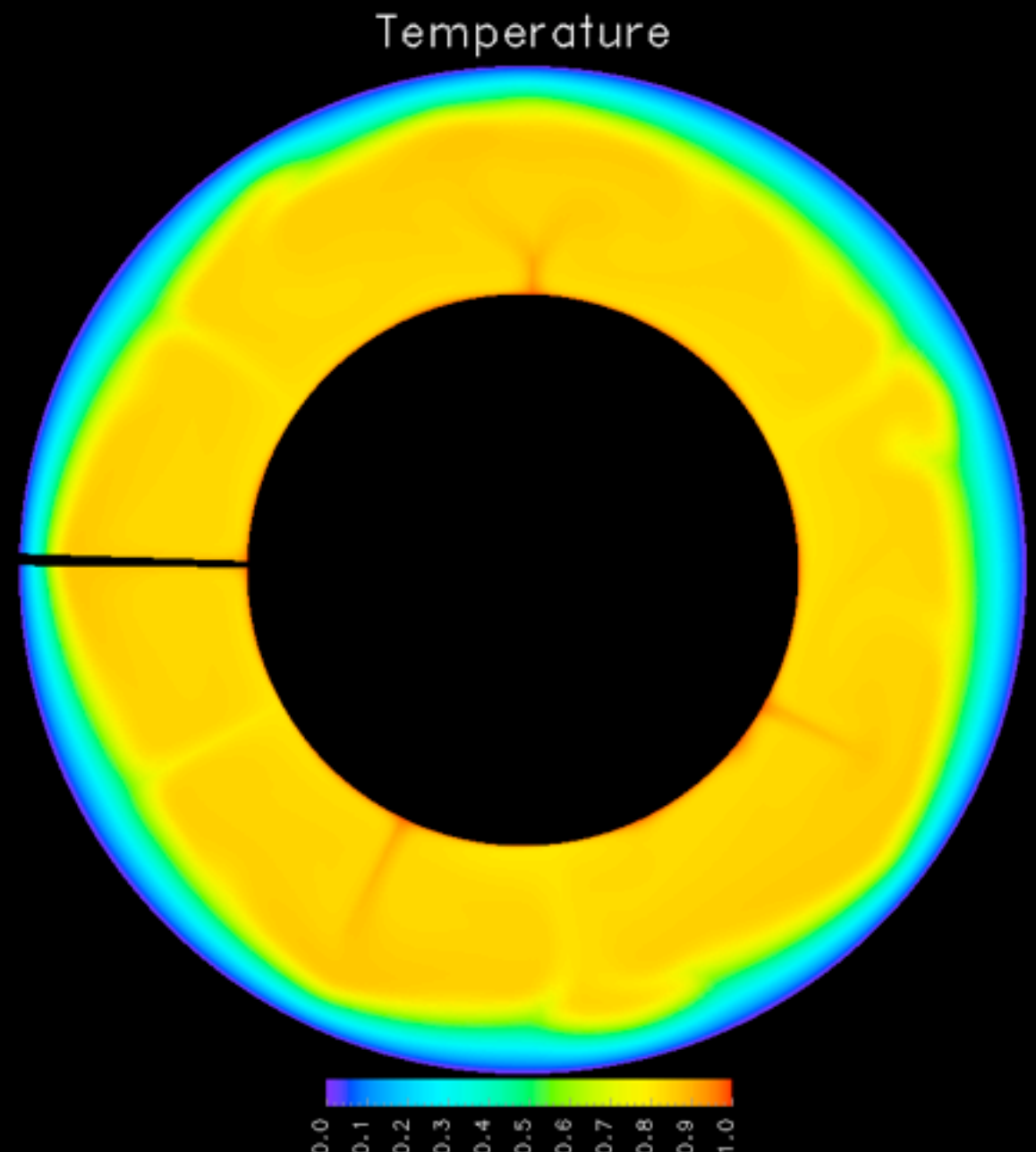
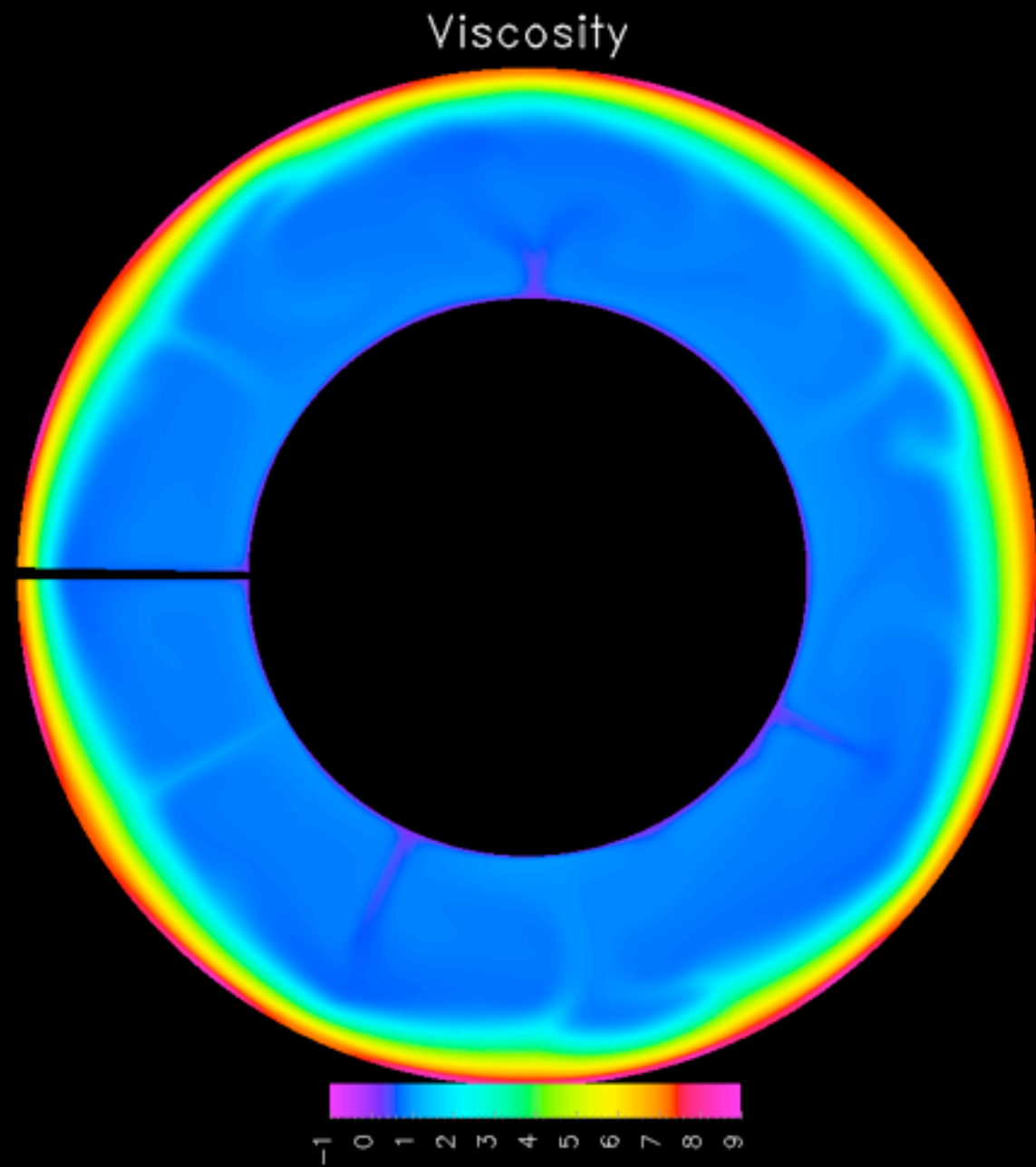


# Plate Tectonics on Exo-Earths?

$$R/R_E=2$$

Van Heck & Tackley, 2009

# Plate Tectonics on Exo-Earths?



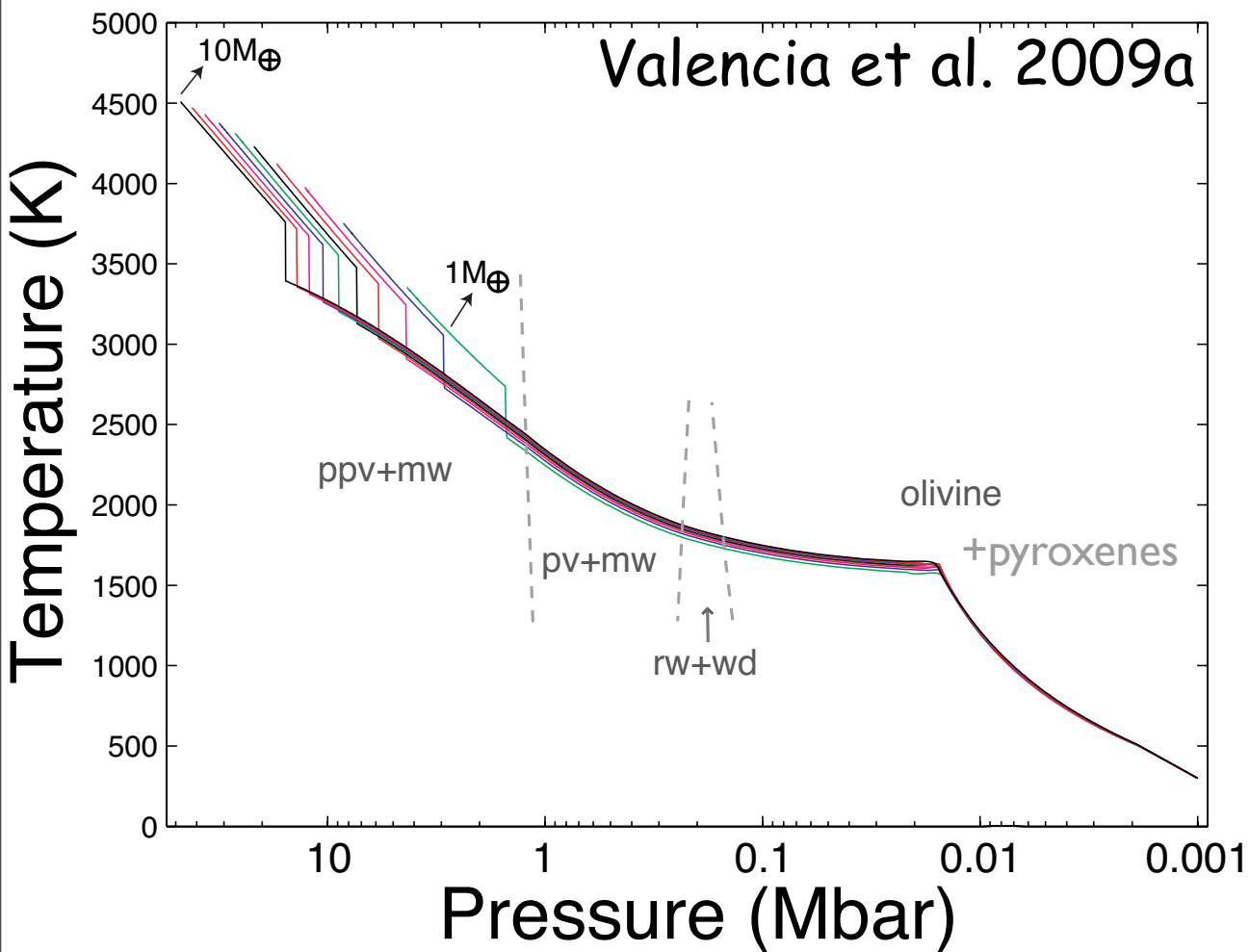
$R/R_E = 2$

Van Heck & Tackley, 2009

# Additional slides



# Uncertainties in the Interior



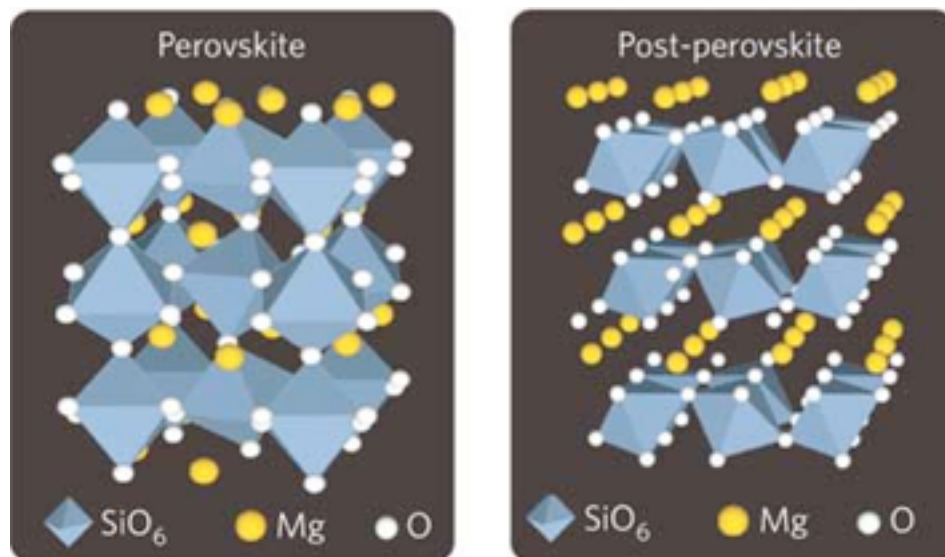
Super-Earths' mantles are mostly composed of PPV

What is PPV's stability region?  
Are there any other phase transitions?

Virtually incompressible oxide:  
 $Gd_3Ga_5O_{12}$  (Mashimo et al '06)

Dissociation of silicates at high P?  
 $MgSiO_3 \rightarrow MgO + SiO_2$

(Umemoto et al '06)  
(Grocholski et al '10 doesn't agree)



Murakami et al 2004