

Stability of river bed forms

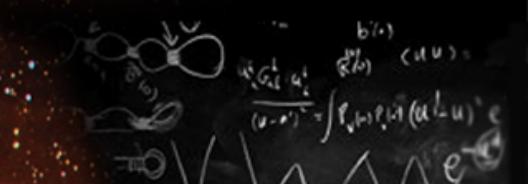
cont'd

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Ripples and Dunes

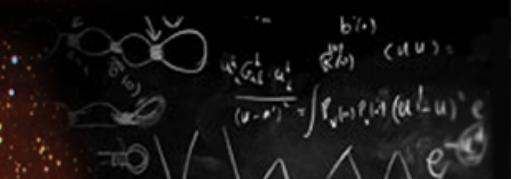


Ripples in the Hunter River, New South Wales, Australia

Image courtesy of M.C. Rygel



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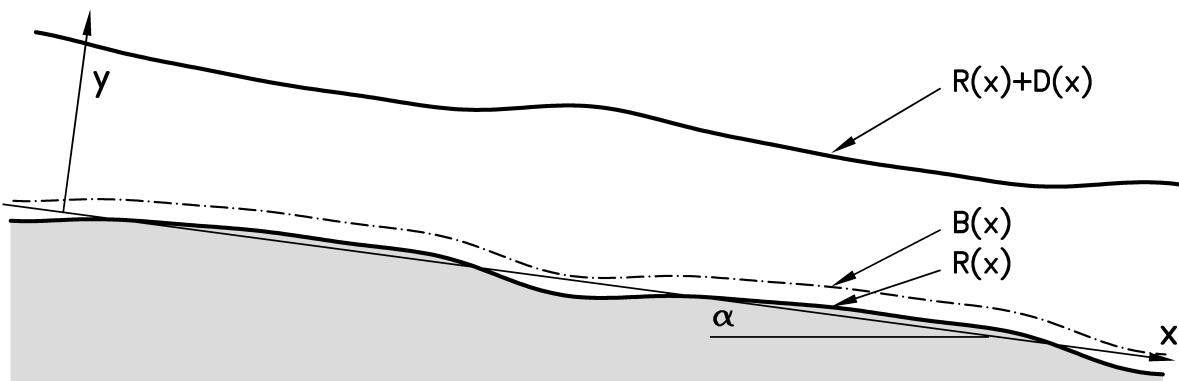
Previous Studies



Linear Stability Analyses

Authors	Flow Regime	Flow model	Results
Richards 1980 <i>J. Fluid Mech.</i>	rough	Rotational Effect of the flow depth included	Dunes Ripples
Engelund & Fredsoe, 1982, <i>Ann. Rev. Fluid Mech.</i>		Ripples associated to smooth flow regime	
Sumer & Bakioglu, 1984 <i>J. Fluid Mech.</i>	smooth	Rotational Boundary layer Infinite flow depth	Ripples
Fourriere et al, 2010 <i>J. Fluid Mech.</i>	rough	Rotational Effect of flow depth included	Ripples
Colombini & Stocchino, 2011, <i>J. Fluid Mech.</i>	rough transitional smooth	Rotational Effect of flow depth included	Dunes Ripples

All models consider bed load only



$$\eta = \frac{y - R}{D}$$

2D REYNOLDS EQUATIONS + CONTINUITY

$$u_{,t} + uu_{,x} + vu_{,y} + p_{,x} - S/Fr^2 - T_{xx,x} - T_{xy,y} = 0$$

$$v_{,t} + uv_{,x} + vv_{,y} + p_{,y} + 1/Fr^2 - T_{xy,x} - T_{yy,y} = 0$$

$$u_{,x} + v_{,y} = 0$$

BOUSSINESQ's TURBULENCE CLOSURE

$$T_{xx} = 2\nu_T u_{,x} \quad T_{yy} = 2\nu_T v_{,y} \quad T_{xy} = \nu_T (u_{,y} + v_{,x})$$

EDDY VISCOSITY

$$\nu_T = \ell^2 \sqrt{(u_{,x} - v_{,y})^2 + (u_{,y} + v_{,x})^2}$$

Rough vs. smooth regime



CONDUCTANCE COEFFICIENT (CHEZY)



$$C = \sqrt{\frac{8}{f}} = \frac{1}{\kappa} \ln \left(\frac{11.09}{2.5d} \right)$$

$$C = \sqrt{\frac{8}{f}} = \frac{1}{\kappa} \ln \left(\frac{\text{Re} \sqrt{f}}{3.41} \right)$$

Rough regime

(ASCE Task Committee, 1963)

Smooth regime

Rough vs. smooth regime



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Smooth regime

$$C = \sqrt{\frac{8}{f}} = \frac{1}{\kappa} \ln \left[\left(\frac{11.09}{2.5d} \right)^{1-\beta} \left(\frac{Re \sqrt{f}}{3.41} \right)^\beta \right]$$

Transitional regime
(Chang, 2008)

$$\beta = \exp \left[-0.11 \left(\ln Re_k \right)^{\frac{5}{2}} \right]$$

$$Re_k = \frac{u_*^* k_s^*}{\nu}$$

**Roughness
Reynolds number**

$$Re = \frac{4U^* D^*}{\nu}$$

**Flow
Reynolds number**



$$Re = \frac{4C}{2.5d} Re_k$$



MIXING LENGTH – Rough regime

$$l = \kappa D(\eta + \eta_0)(1 - \eta)^{1/2}$$

Rule of the thumb: the roughness height is about one thirtieth of the actual nondimensional roughness.

This result can be formally obtained by integrating the log-law along depth, relating the roughness height to the nondimensional Chezy coefficient)

$$1 = \int_0^1 u d\eta = \frac{1}{\kappa C} \int_0^1 \ln\left(\frac{\eta + \eta_0}{\eta_0}\right) d\eta \approx -\frac{1}{\kappa C} [1 + \ln(\eta_0)]$$

and then substituting the empirical relationship

$$C = \sqrt{\frac{8}{f}} = \frac{1}{\kappa} \ln\left(\frac{11.09}{2.5d}\right)$$

to obtain

$$\eta_0 = \exp(-\kappa C - 1) = \frac{d}{12} = \frac{2.5d}{30} = \frac{y_R}{30} = R_0$$

Rough vs. smooth regime



MIXING LENGTH – Smooth & Transitional regimes

$$l = \kappa D(\eta + \Delta\eta) \left[1 - \exp\left(-\frac{\eta + \Delta\eta}{A} D \frac{\text{Re}_k}{2.5d} \right) \right] (1 - \eta)^{1/2}$$



MIXING LENGTH – Smooth & Transitional regimes

$$l = \kappa D(\eta + \Delta\eta) \left[1 - \exp\left(-\frac{\eta + \Delta\eta}{A} D \frac{Re_k}{2.5d}\right) \right] (1 - \eta)^{1/2}$$

Van Driest (1956) – The exponential correction accounts for the existence of a viscous and buffer layer in the smooth and transitional regimes.



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Van Driest (1956) – The exponential correction accounts for the existence of a viscous and buffer layer in the smooth and transitional regimes.

Rotta (1962) – The velocity profiles for rough and smooth regimes can be similar provided a suitable displacement is introduced.



MIXING LENGTH – Smooth & Transitional regimes

$$\lambda = \kappa D(\eta + \Delta\eta) \left[1 - \exp\left(-\frac{\eta + \Delta\eta}{A} D \frac{Re_k}{2.5d}\right) \right] (1 - \eta)^{1/2}$$

Van Driest (1956) – The exponential correction accounts for the existence of a viscous and buffer layer in the smooth and transitional regimes.

Rotta (1962) – The velocity profiles for rough and smooth regimes can be similar provided a suitable displacement is introduced.

$$\Delta\eta^+ = \frac{Re_k}{30}$$

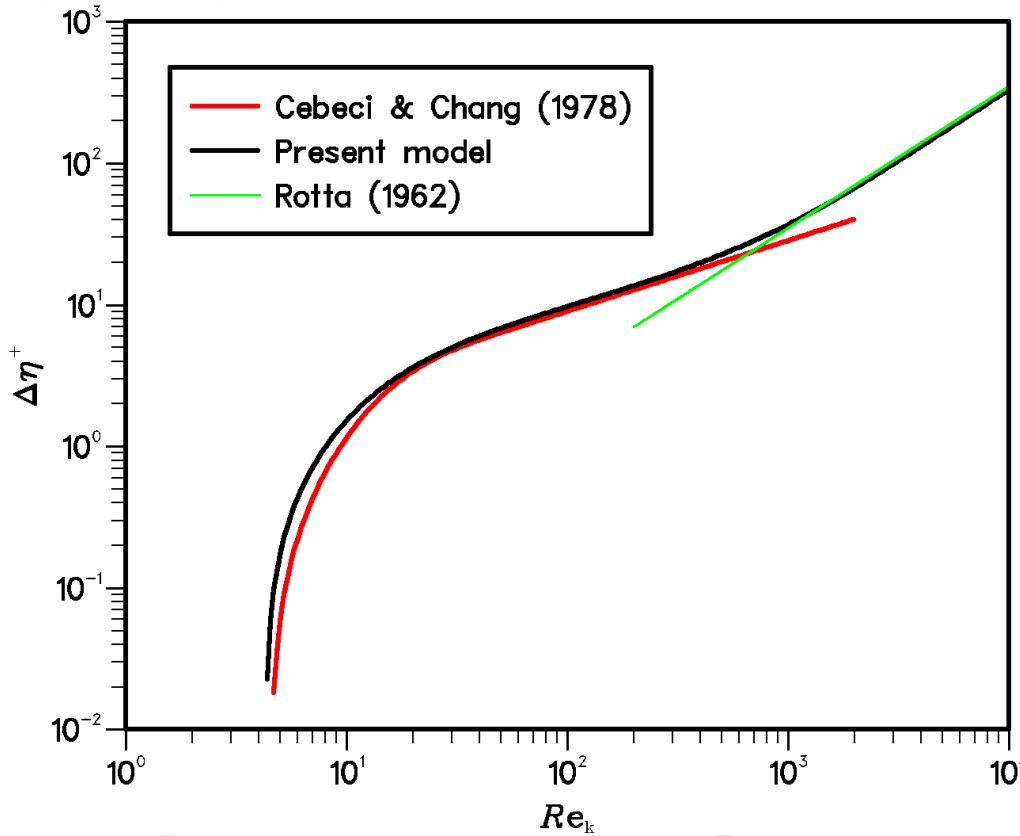
Rotta (1962)
Rough regime

$$\Delta\eta^+ = 0.9 \left[\sqrt{Re_k} - Re_k \exp\left(-\frac{Re_k}{6}\right) \right]$$

Cebeci & Chang (1978)
Smooth & Transitional regime

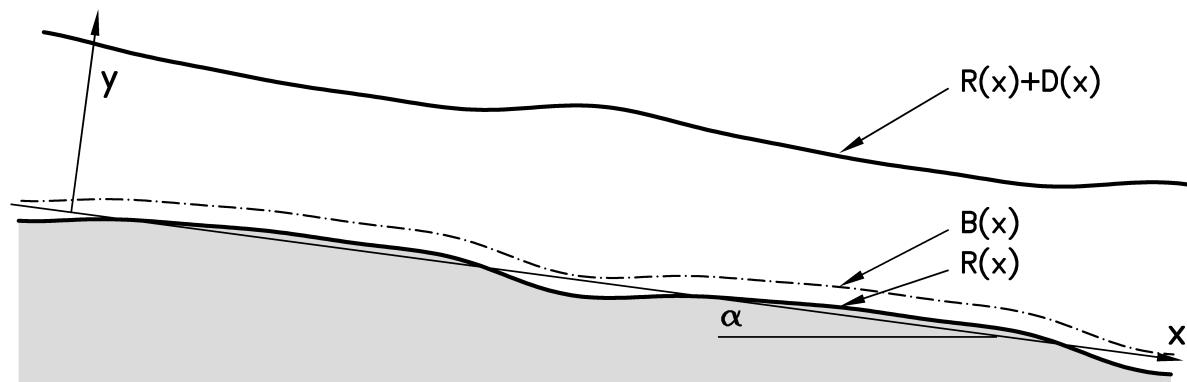
Rough vs. smooth regime

DISPLACEMENT



$$C = \frac{1}{\kappa} \ln \left[\left(\frac{11.09}{2.5d} \right)^{1-\beta} \left(\frac{Re \sqrt{f}}{3.41} \right)^\beta \right] \quad \rightarrow \quad \Delta\eta^+ = f(Re_k)$$

$$\beta = \exp \left[-0.11 (\ln Re_k)^{\frac{5}{2}} \right]$$



SEDIMENT CONTINUITY EQUATION (EXNER)

$$R_t + \mathcal{Q}_s \Phi_x = 0$$

$$\mathcal{Q}_s = \frac{1}{F} \frac{\sqrt{(s-1)d^3}}{1-p} \ll 1$$

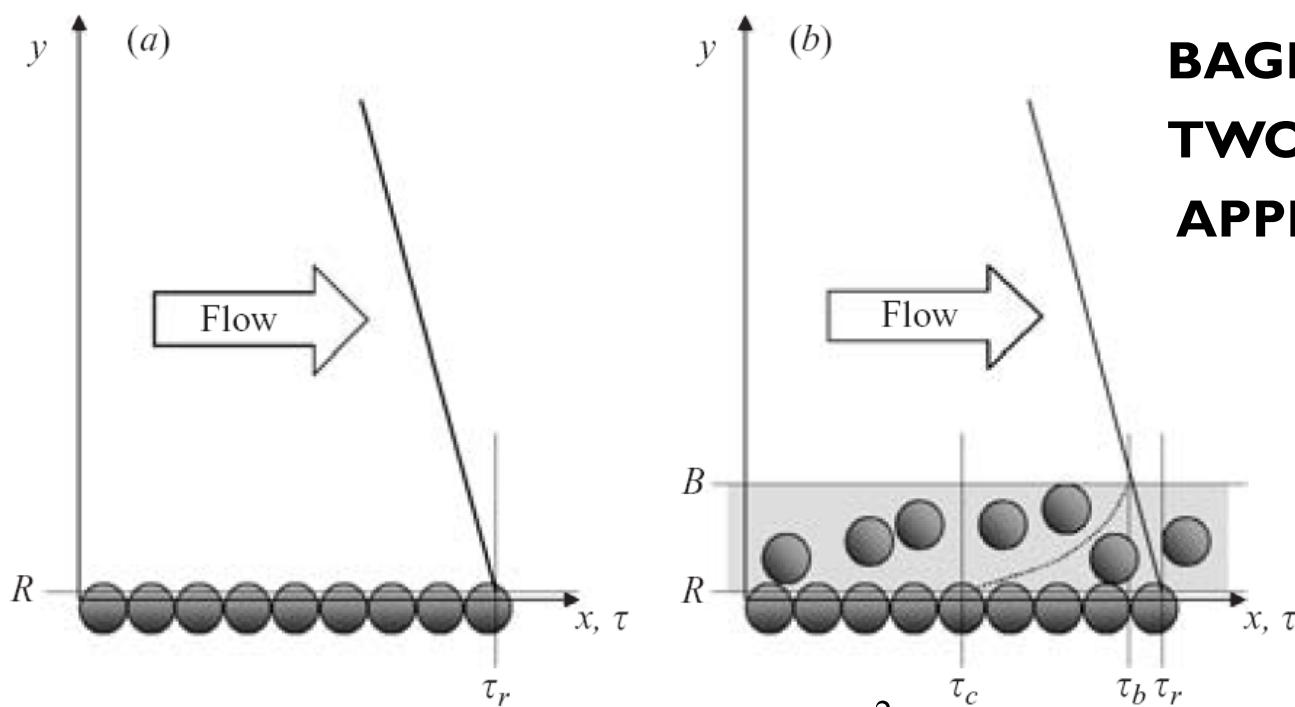
BEDLOAD FUNCTION (MPM after Wong & Parker (2006))

$$\Phi = 3.97 (\vartheta_B - \vartheta_C)^{3/2}$$

BEDLOAD LAYER THICKNESS (scaled with grain size)

$$I_B = 1 + 1.3 \left(\frac{\vartheta_0 - \vartheta_{C0}}{\vartheta_{C0}} \right)^{0.55}$$

Sediment transport model



$$\vartheta_B = \vartheta_0 T_B = T_{xy} \Big|_{y=B}$$

$$B = R + \eta_B D$$

$$\vartheta_0 = \frac{F^2}{C^2(s-1)d}$$

$$\eta_B = l_B d$$

CORRECTION FOR SEDIMENT WEIGHT

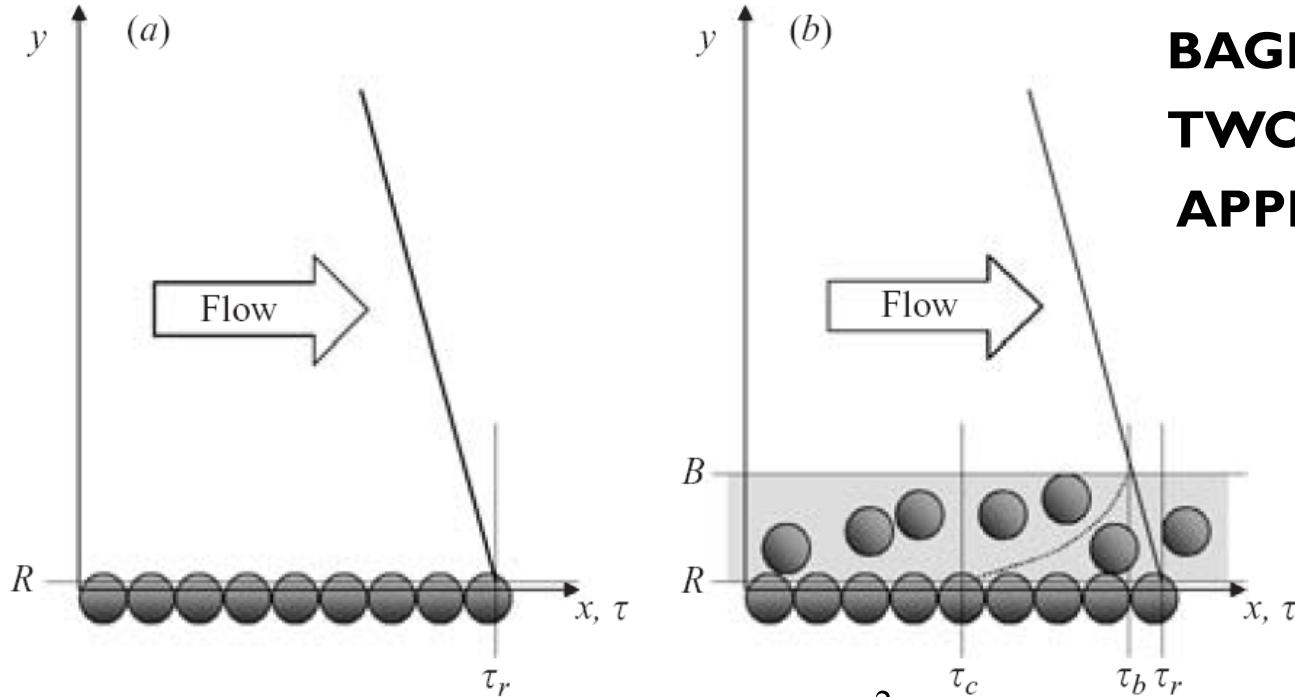
$$\vartheta_C = \vartheta_{CH} - \mu_x (S - R_{,x})$$

$$\mu_x = 0.1$$

Fredsøe (1974)

BAGNOLD'S TWO-LAYER APPROACH

Sediment transport model



$$\vartheta_B = \vartheta_0 T_B = T_{xy} \Big|_{y=B}$$

$$B = R + \eta_B D$$

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$$\eta_B = l_B d$$

CORRECTION FOR SEDIMENT WEIGHT

$$\vartheta_C = \vartheta_{CH} - \mu_x (S - R_{,x}) \quad \mu_x = 0.1 \quad \text{Fredsøe (1974)}$$

$$\vartheta_{CH} = 0.22 \text{Re}_P^{-0.6} + 0.06 \exp(-17.73 \text{Re}_P^{-0.6}) \quad \text{Brownlie (1981)}$$

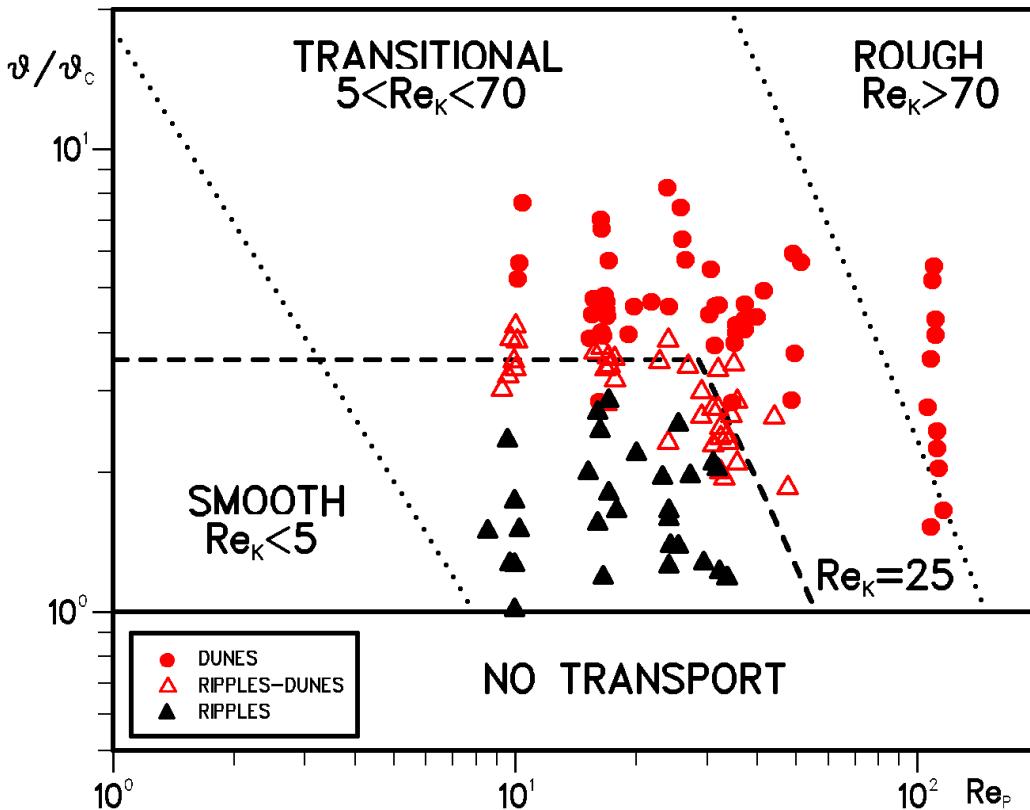
$$\vartheta_{CH} = 0.0495 \quad \text{Wong & Parker (2006)}$$

BAGNOLD'S TWO-LAYER APPROACH

Experimental Observations



Experiments by Guy, Simons & Richardson (1966)



$$Re_k = \frac{u_f^* k_s^*}{\nu}$$

**Roughness
Reynolds number**

$$Re_p = \frac{\sqrt{(s-1)gd^* d^*}}{\nu}$$

**Particle Reynolds
number**

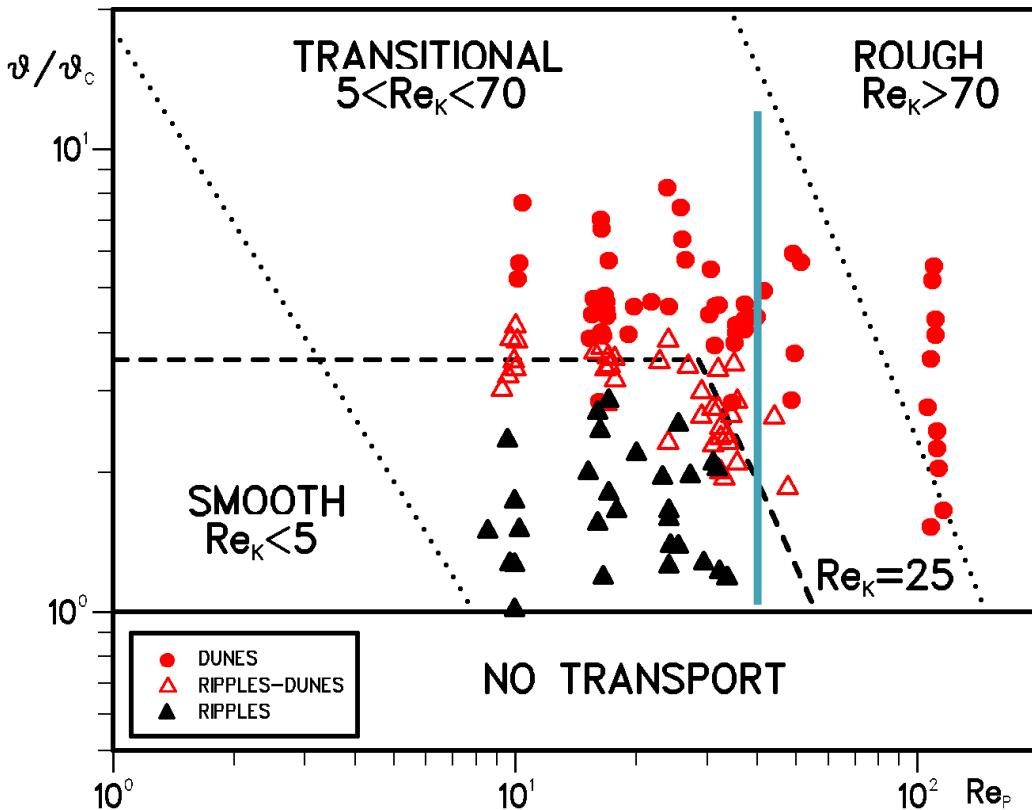


$$Re_k = \sqrt{\vartheta} Re_p 2.5$$

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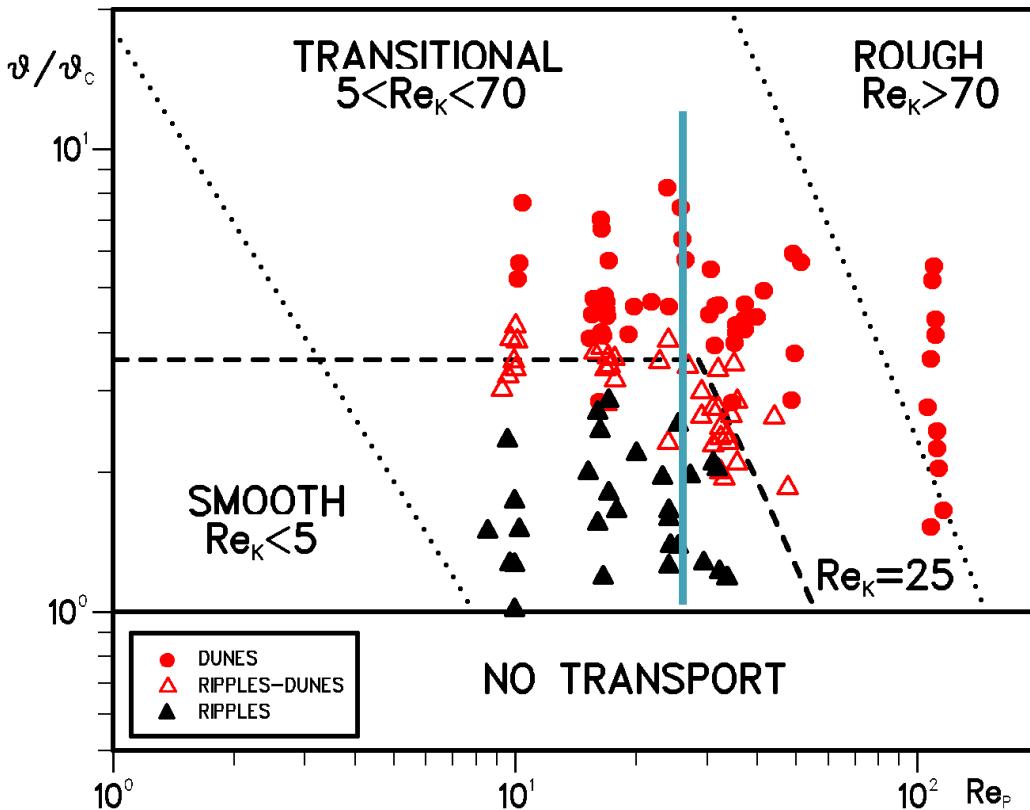


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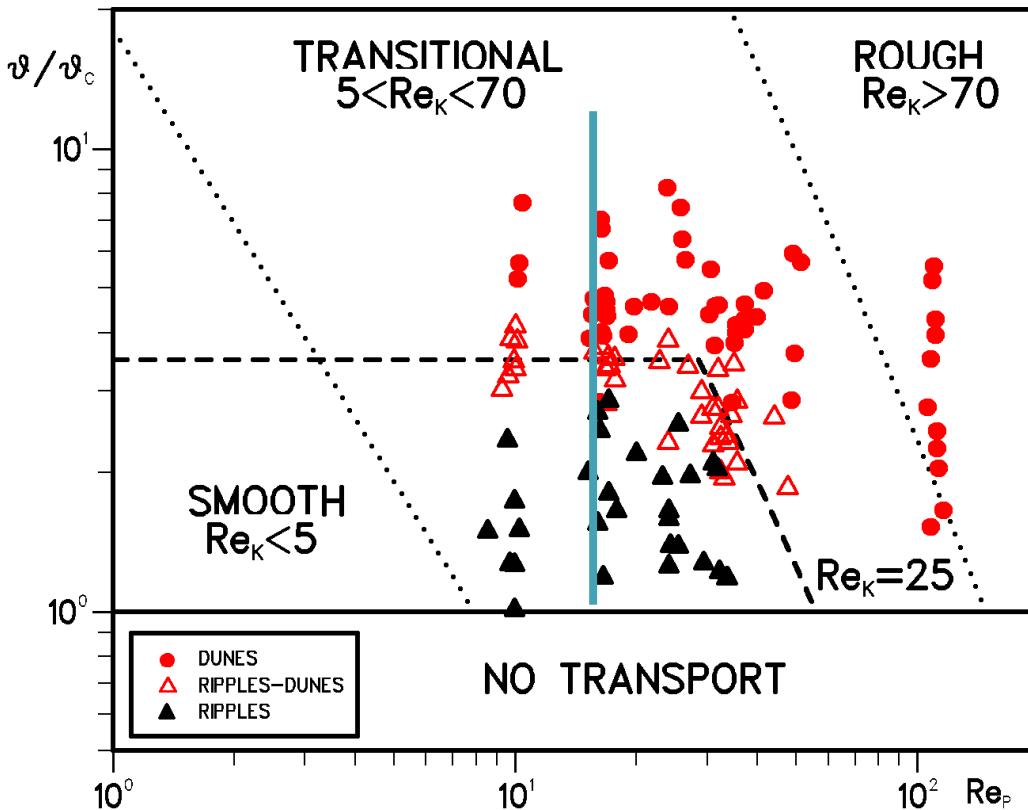


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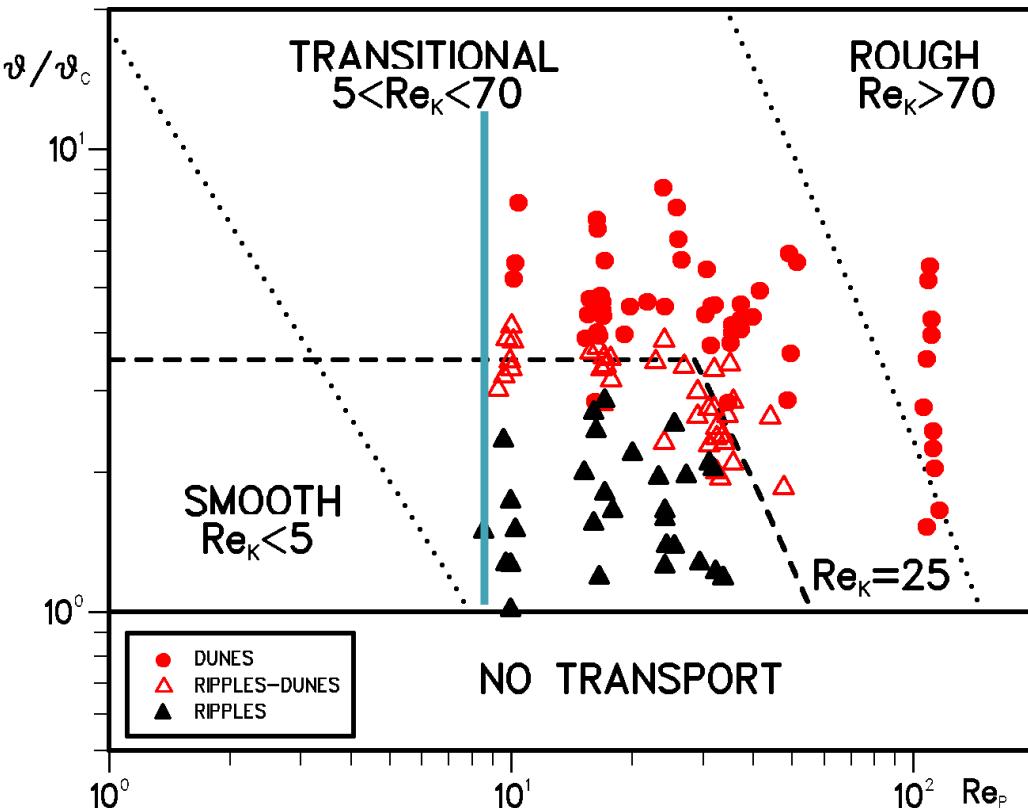
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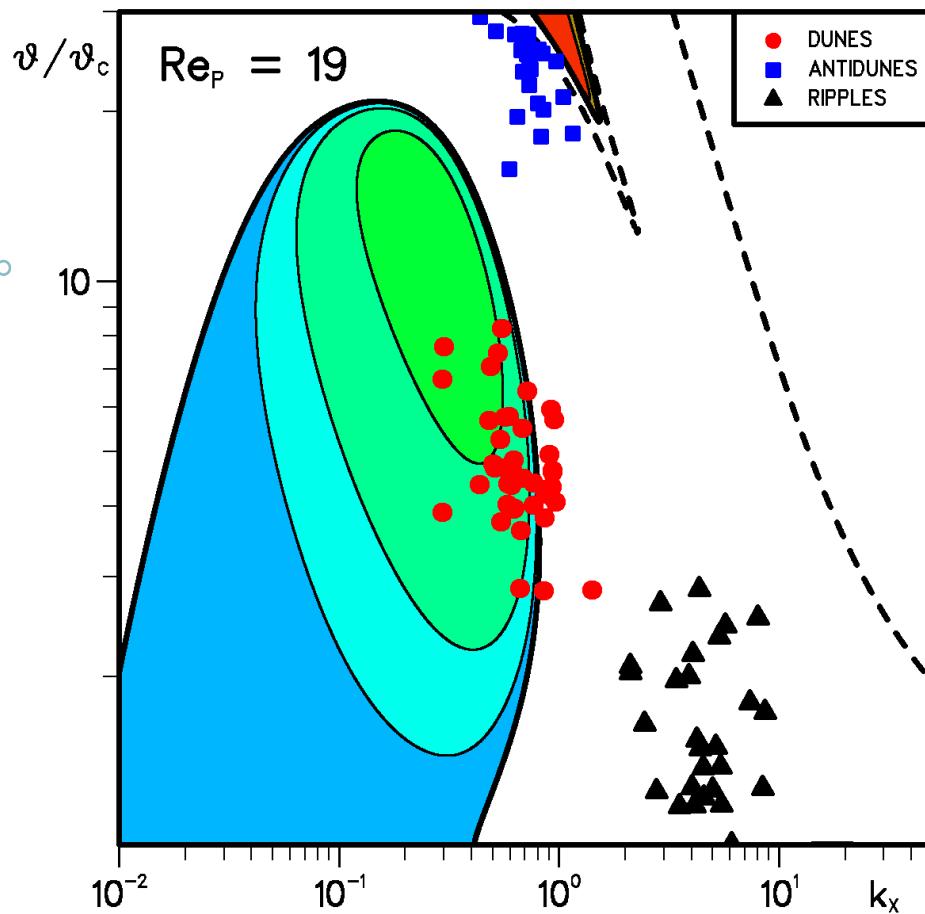
**Particle Reynolds
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$$Re_k = \sqrt{\vartheta} Re_p 2.5$$



Dune-ripple transition



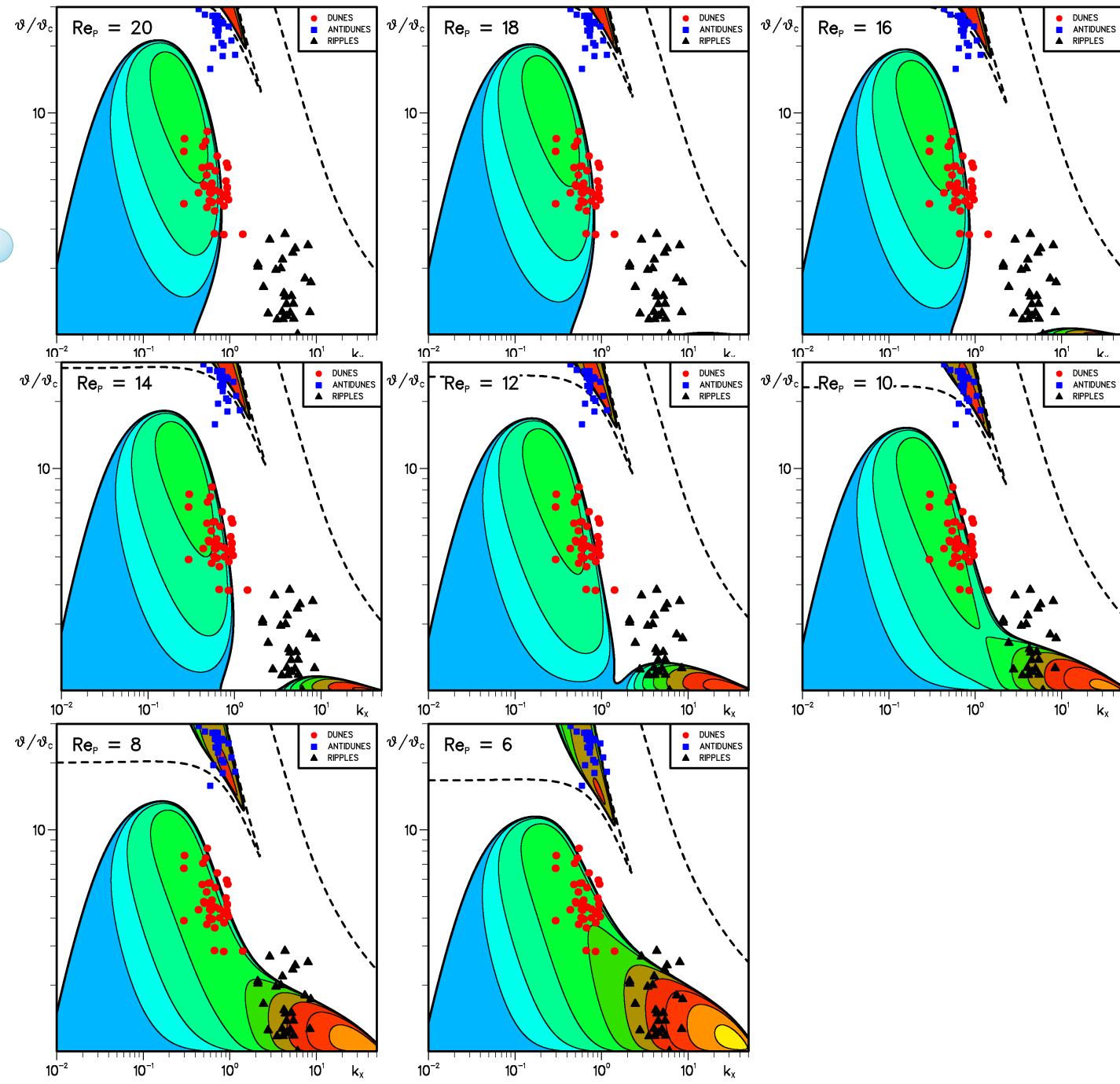
$$Re_p = \frac{\sqrt{(s-1)gd^*} d^*}{\nu}$$

Re_p is the parameter that controls the transition between ripples and dunes

Re_p decreasing $\Rightarrow d^*$ decreasing \Rightarrow finer sediment

d constant, d^* decreasing \Rightarrow shallower flow

Dune-ripple transition



GEOFLOWS 13 - Fluid-Mediated Particle Transport in Geophysical Flows, Sept. 23~Dec 20, 2013, Santa Barbara, California

The neverending story



- **Kennedy (1963,1969) - Irrotational** D A R
- **Engelund (1970), Fredsøe (1974) - Slip velocity** D A R
- **Parker (1975) - Irrotational** D A R

- **Richards (1980) - Rotational** D A R
- **Sumer & Backioglu (1984) - Rotational** D A R
- **Coleman & Fenton (2000) - Irrotational** D A R
- **Colombini (2004) - Rotational** D A R
- **Fourriére, Claudin & Andreotti (2010) - Rotational** D A R
- **Colombini & Stocchino (2011) - Rotational** D A R

The neverending story



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- **Fourrière, Claudin & Andreotti (2010) - Rotational** D A R
- **Colombini & Stocchino (2011) - Rotational** D A R

- **Charru, Andreotti & Claudin (2013) - Sand ripples and dunes. Ann. Rev. Fluid Mech.** D A R
- **Andreotti & Claudin (2013) - Aeolian and subaqueous bedforms in shear flows. Phil. Trans. R. Soc. A** D A R

Linear Coarsening

Marco Colombini

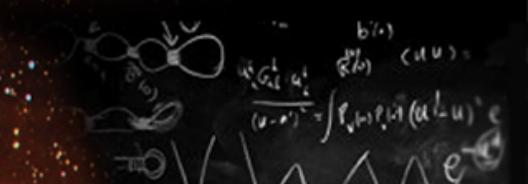
on behalf of

L. Ridolfi, C. Camporeale, R. Vesipa

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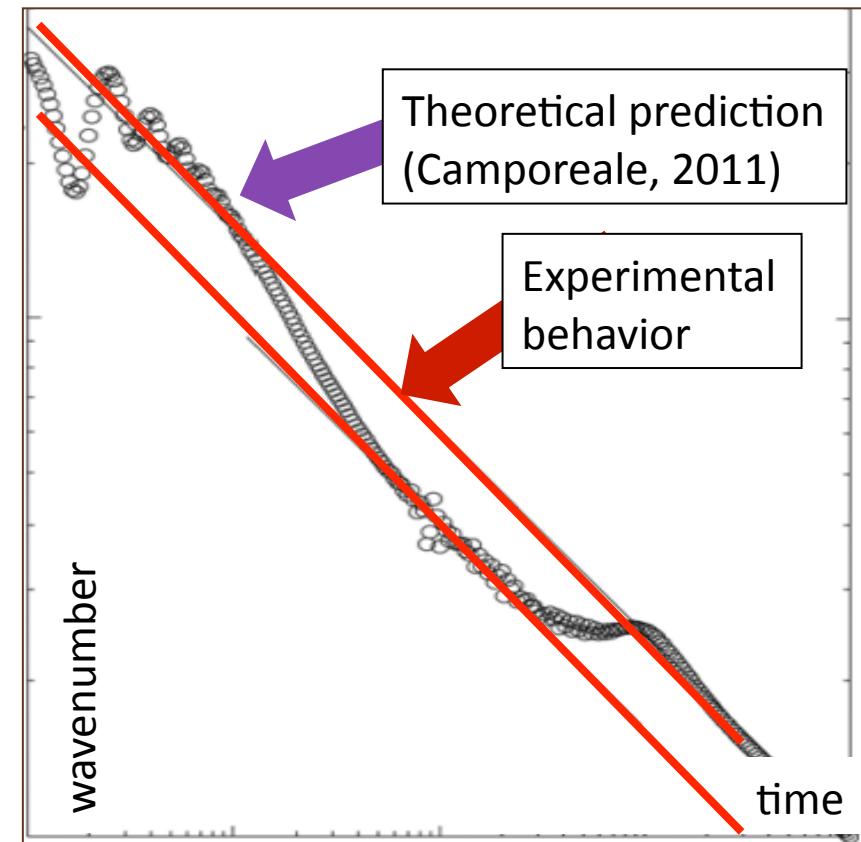
Nonmodal analysis and river morphodynamics

Nonmodal analysis has been proposed for different morphological systems:

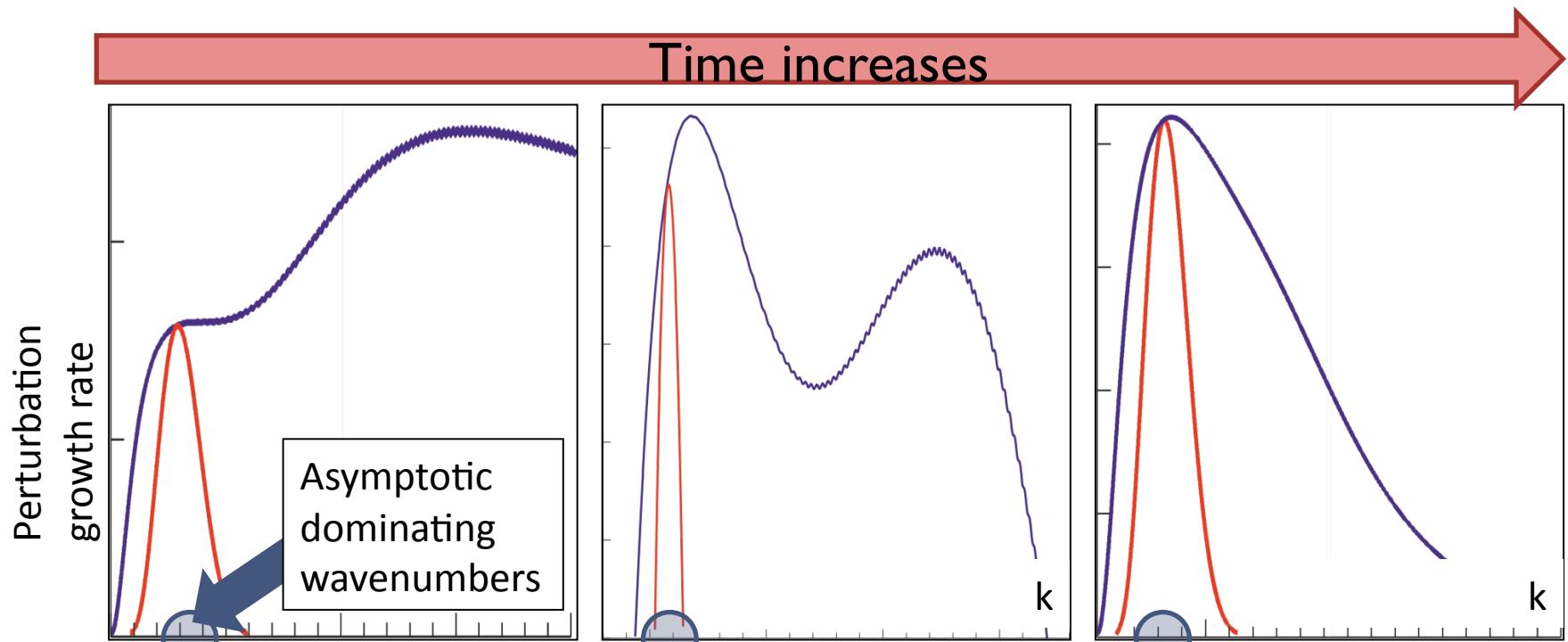
- Long sediment waves and bars (Camporeale, 2011; Vesipa 2012)
- Dunes and antidunes (Camporeale, 2011)

Nonmodal analysis can possibly explain a number of issues:

- Wavelength at initial times are shorter than wavelength at asymptotic times (linear transient excitation of stable modes)
- Sediment wave amalgamation is induced by linear mechanisms (excitation of modes with different phase velocity)



Non modal analysis and bed form wavelengths



At the very beginning

- short waves dominates
- the higher k , the higher the phase velocity

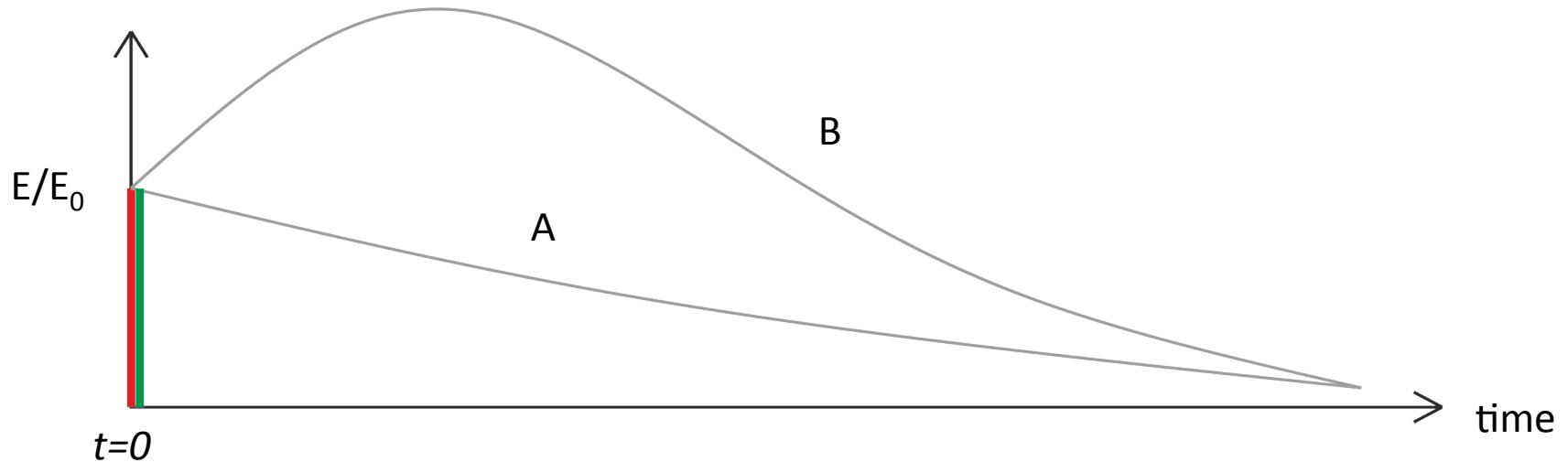
At intermediate times

- short and asymptotic waves compete
- waves with different phase velocity can be seen
- amalgamation among waves

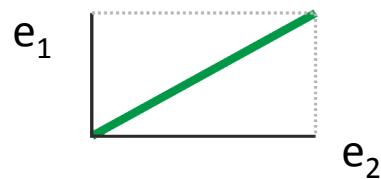
At asymptotic times

- long waves dominates

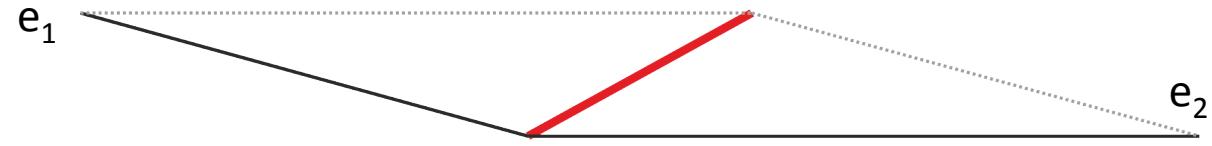
Energy exchanges among (stable) modes



$t=0$

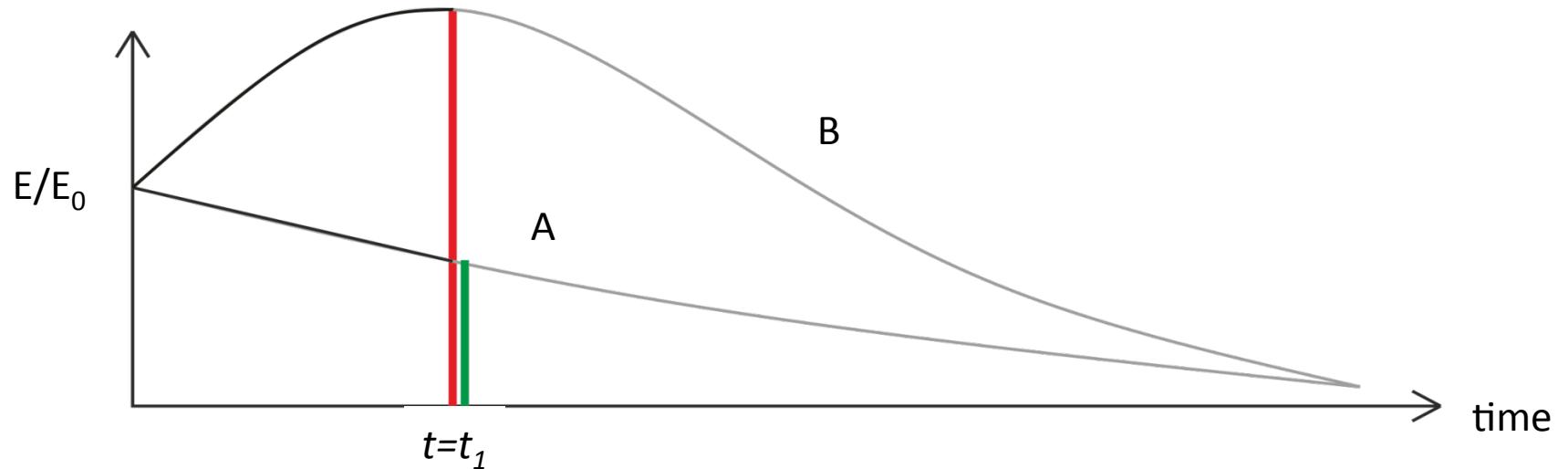


A- Eigenvectors (e_1, e_2) are orthogonal



B- Eigenvectors (e_1, e_2) are not orthogonal

Energy exchanges among (stable) modes



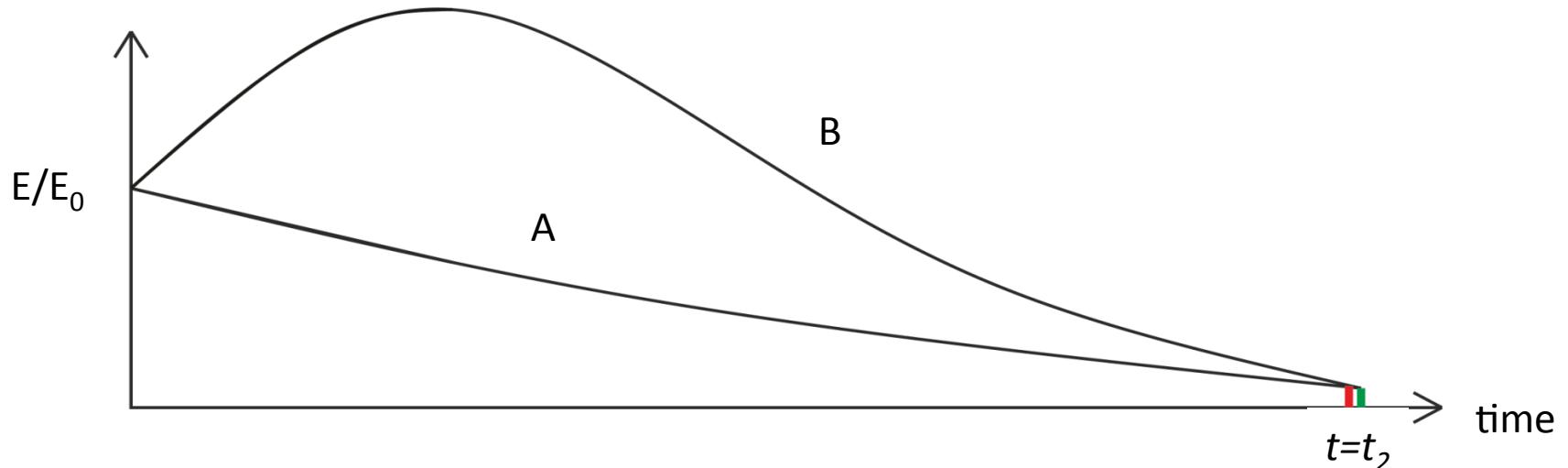
$t=t_1$



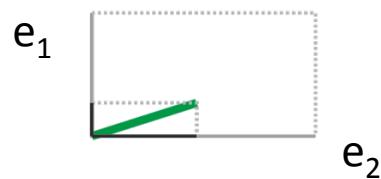
A- Each energetic components of the pert. reduces, the total energy reduces

B- Each energetic components of the pert. reduces, **but the total energy increases**

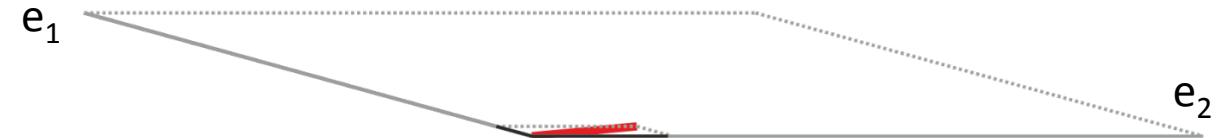
Energy exchanges among (stable) modes



$t=t_2 \approx \infty$



A- For **asymptotic time** the perturbation energy has decayed



B- After a **transient growth**, unperturbed conditions are restored

References

Camporeale, C. and Ridolfi, L. 2009 **Nonnormality and transient behavior of the de Saint-Venant-Exner equations** *Water Resources Research* 45 W08418

Camporeale, C. and Ridolfi, L. 2011 **Modal versus nonmodal linear stability analysis of river dunes** *Physics of Fluids* 23(10) 104102

Schmid P.J. and Henningson D.S. 2001 *Stability and Transition in Shear Flows*, Springer

Vesipa, R. and Camporeale, C. and Ridolfi, L. 2012 **Transient growths of stable modes in riverbed dynamics** *EPL* 100(6) 64002