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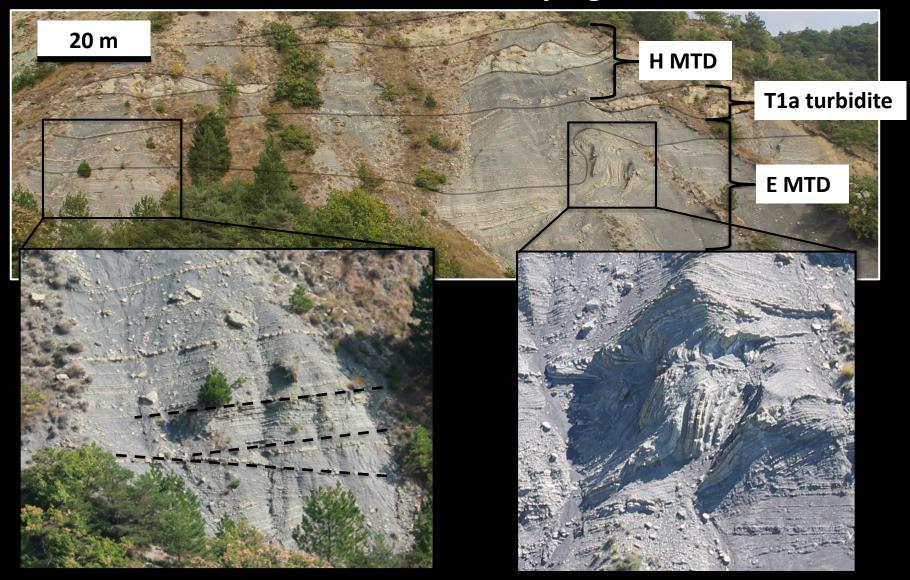
Single grains, dense suspensions (and disperse grainsize distributions) in turbidity currents; a geologist's perspective.

## Modern environments



## Rock Record (Aptian ~120 Ma)

## Relation turbidite sand with underlying MTD; Les Cosmes

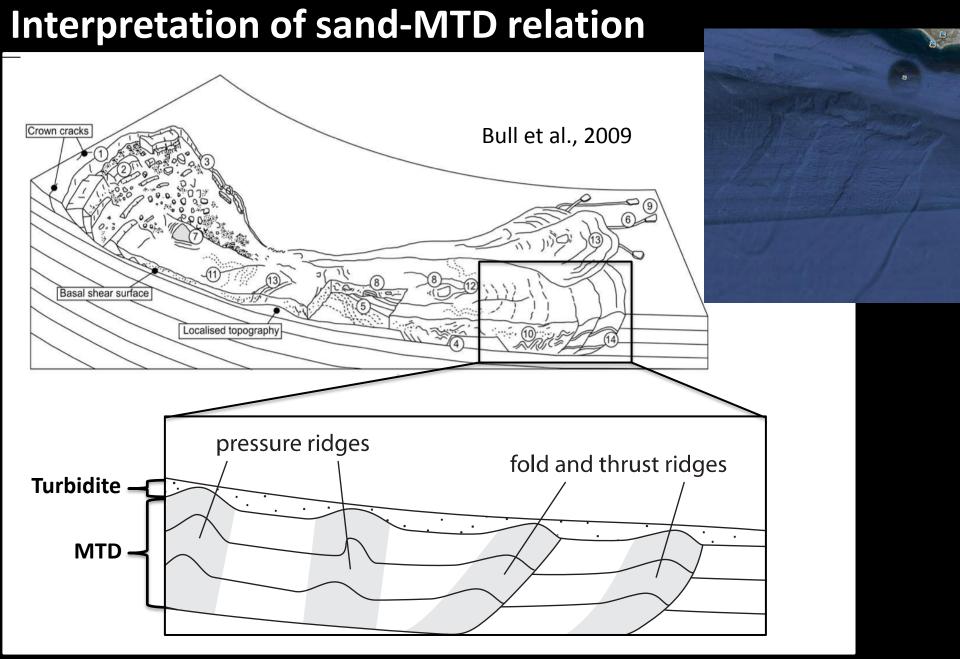


## Rock Record (Aptian ~120 Ma)

## Relation turbidite sand with underlying MTD; Les Cosmes



Rock Record (Aptian ~120 Ma)









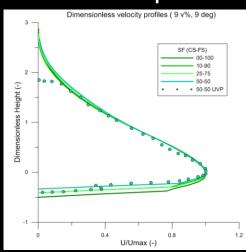
## **Rock Record**

### **Modern Environments**





## Numerical Experiment

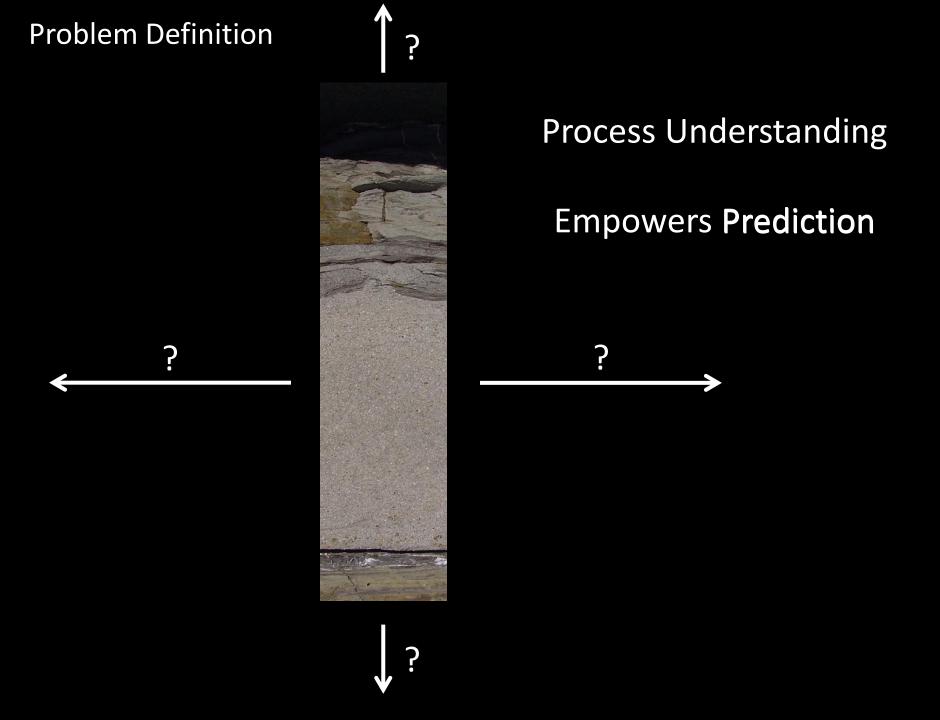


## **Physical Experiment**



## **Problem Definition**

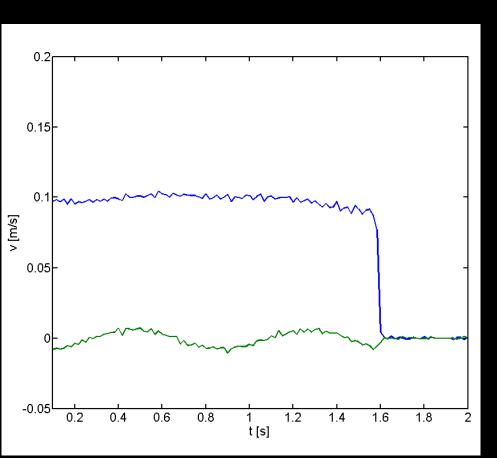




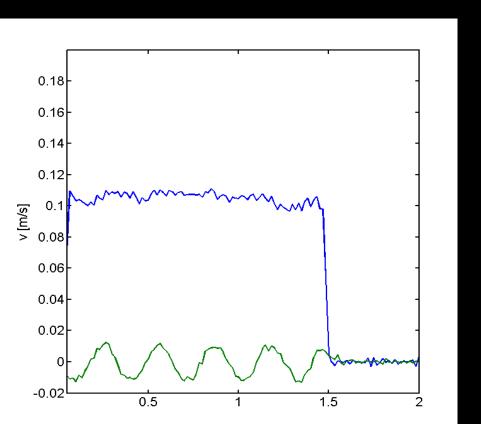
Parameter ranges: Grainsize Distribution [mud-8000 micron]; Grain shape: number of spheres = 0



d= 1.1 mm  $v_s = 0.098 \text{ m/s}$   $v_{spred} \text{ (Ferguson\&Church, 2004)= 0.13 m/s}$   $St = m_p \, v_s \, / \, 6\pi \, \mu \, r^2 = 32$   $Re_p = r \, v_s \, / \, v = 54$  Time to rest = 0.05 s



```
d= 1.1 \text{ mm}
v_s = 0.10 \text{ m/s}
v_{spred} = 0.13 \text{ m/s}
St = 36
Re_p 60
Time to rest = 0.1 \text{ s}
```



```
d= 2.0 mm

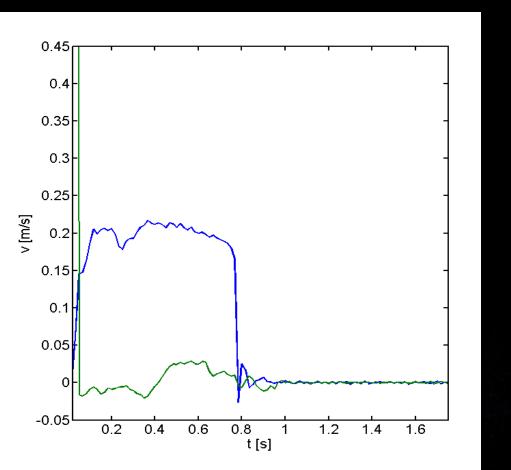
v_s= 0.21 m/s

v_{spred} = 0.19 m/s

St = 118

Re_p 201

Time to rest = 0.22 s
```



```
d= 3.1 mm

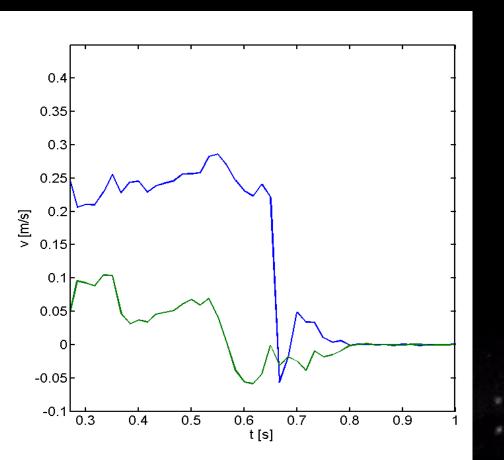
v_s= 0.24 m/s

v_{spred} = 0.25 m/s

St = 222

Re_p 377

Time to rest = 0.15 s
```



d= 4.4 mm

 $v_s = 0.29 \text{ m/s}$ 

 $v_{spred} = 0.30 \text{ m/s}$ 

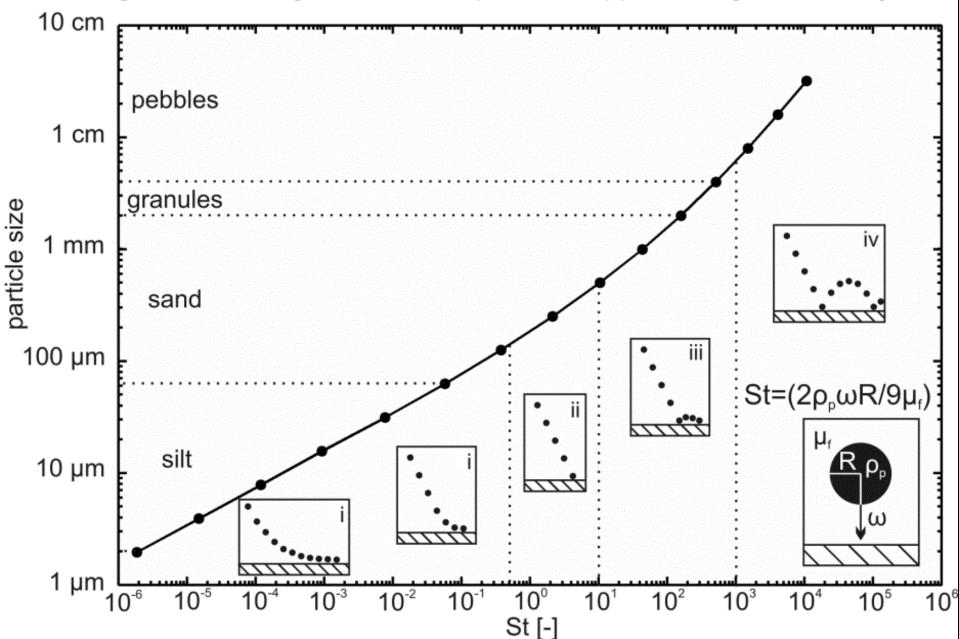
St = 375

 $Re_p$  640

Time to rest = 0.23 s

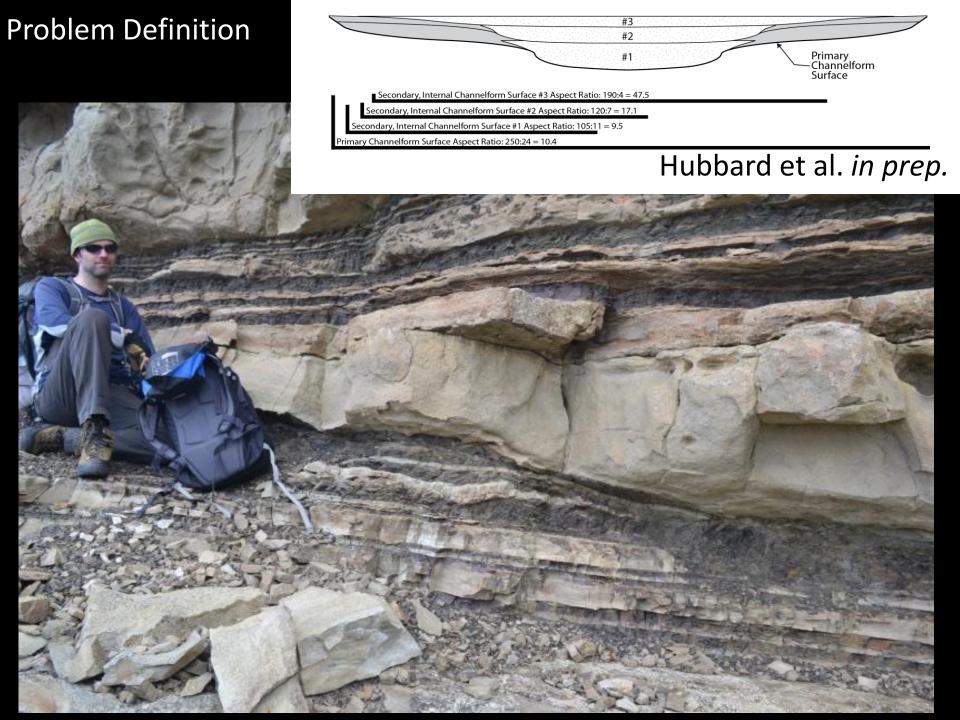


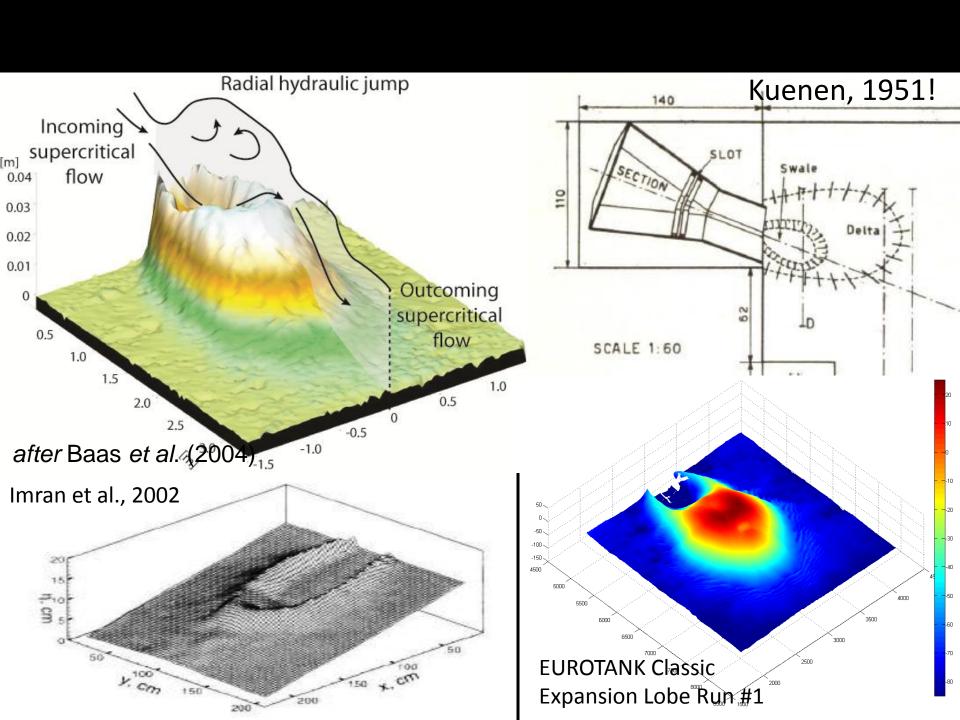
Figure 4. Settling behaviour of particles approaching a boundary



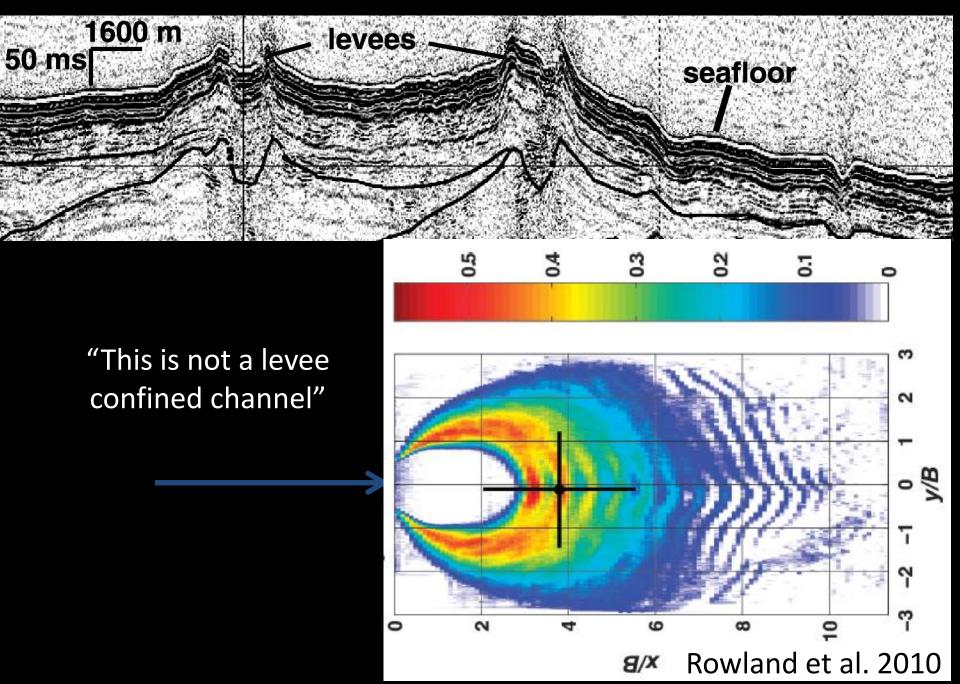
#### Some observations and remarks

- Cd is periodic for intermediate grains (St 30)
- Restitution has rotational & lateral components when grains are irregular; [all natural grains are irregular]; therefore all natural grains have rotational and lateral components to restitution.
- Splash mobilisation only occurs for large grains (>2 mm) in water
- For large grains, the wake-jet has resuspension potential
- The dampened and floating regime are also relevant for tutbidity currents [1-8000 micron]

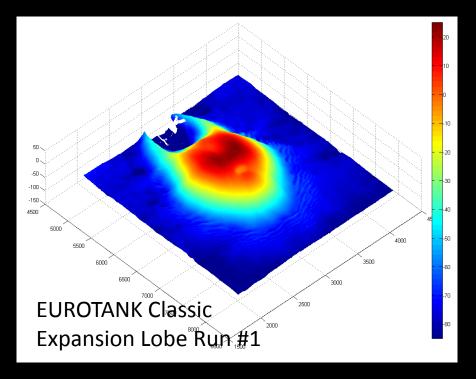




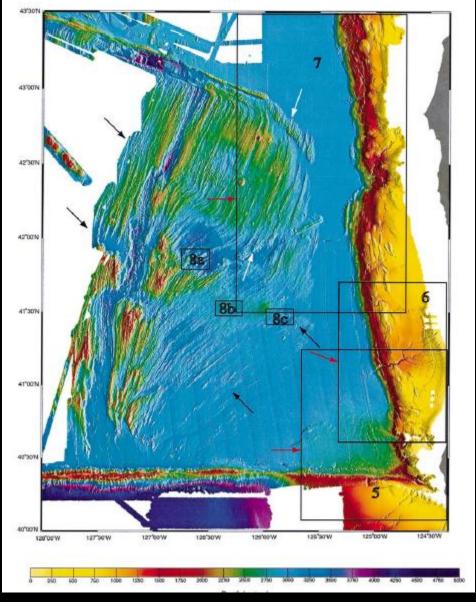
Rowland et al. 2010: "This is a levee confined channel"



#### This architecture is an analogue to some natural systems:

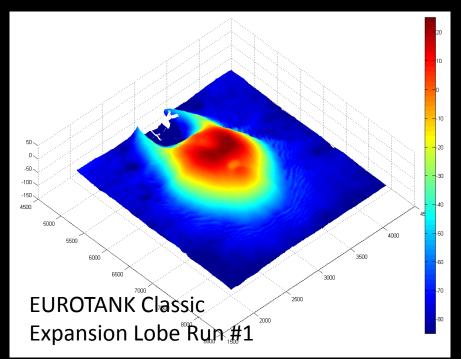


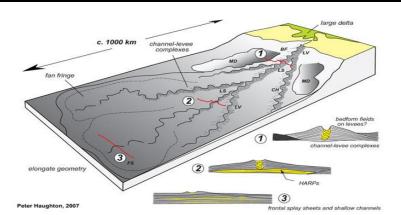
Plunge pools at canyon-fan transitions at the toes of steep active margins (Pacific; Indian Ocean)

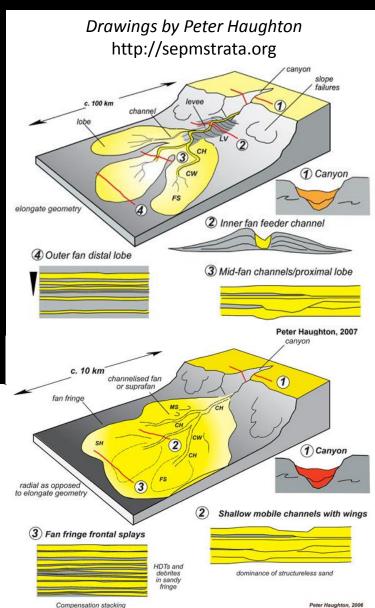


Dziak et al., 2001

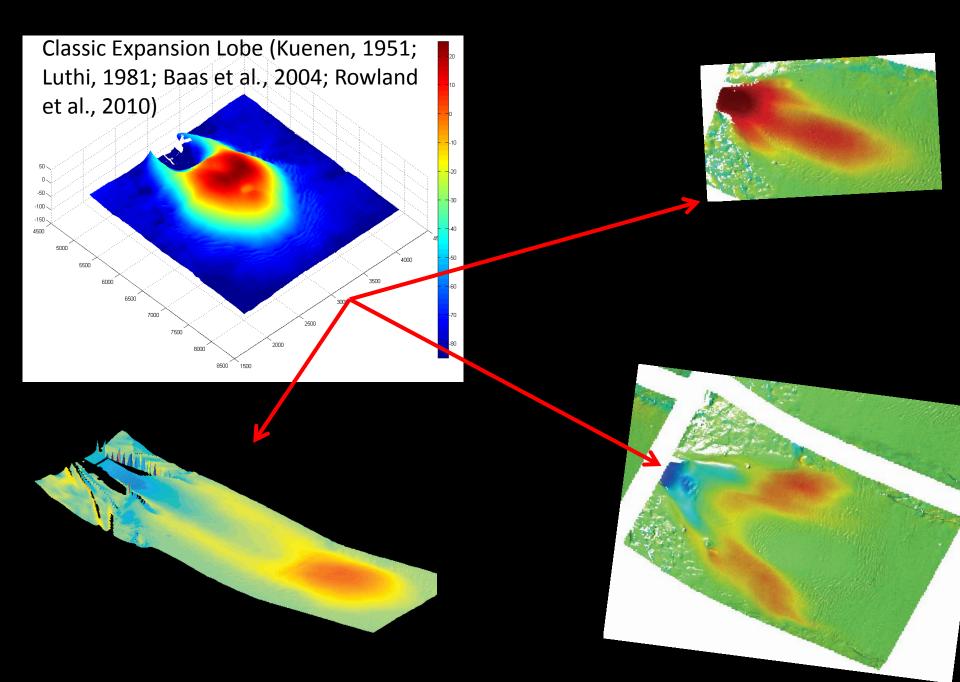
# But this experimental architecture is not analogue to these types of systems

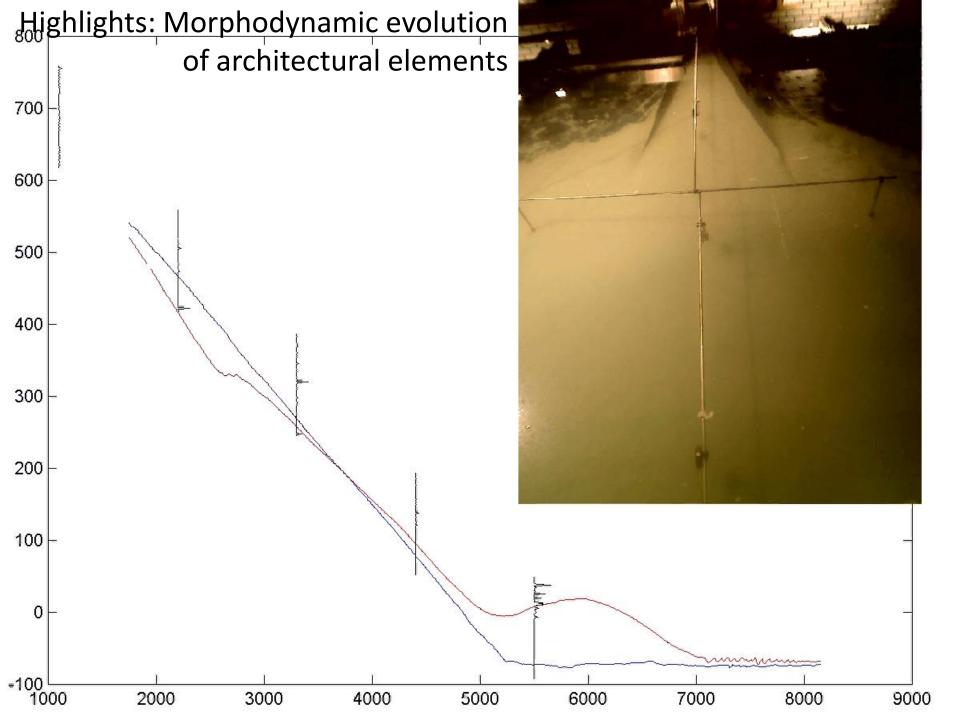


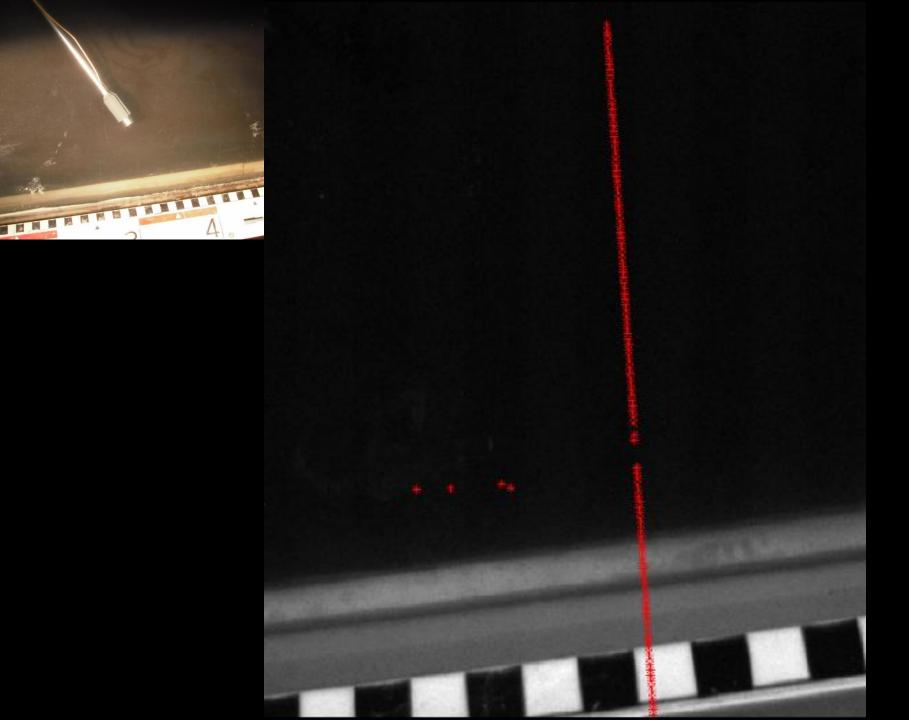


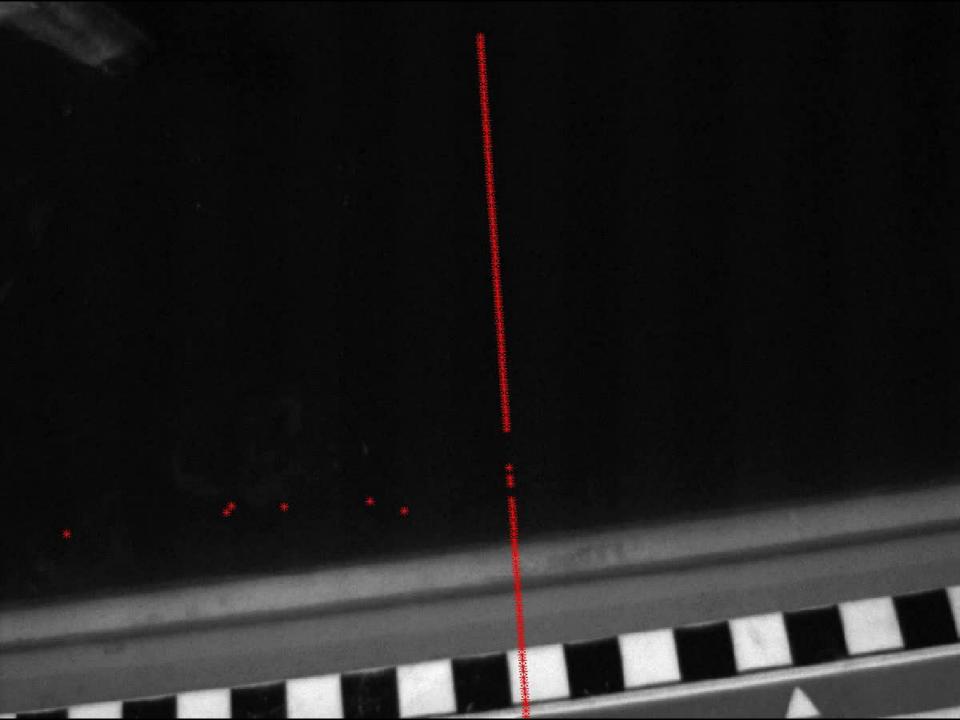


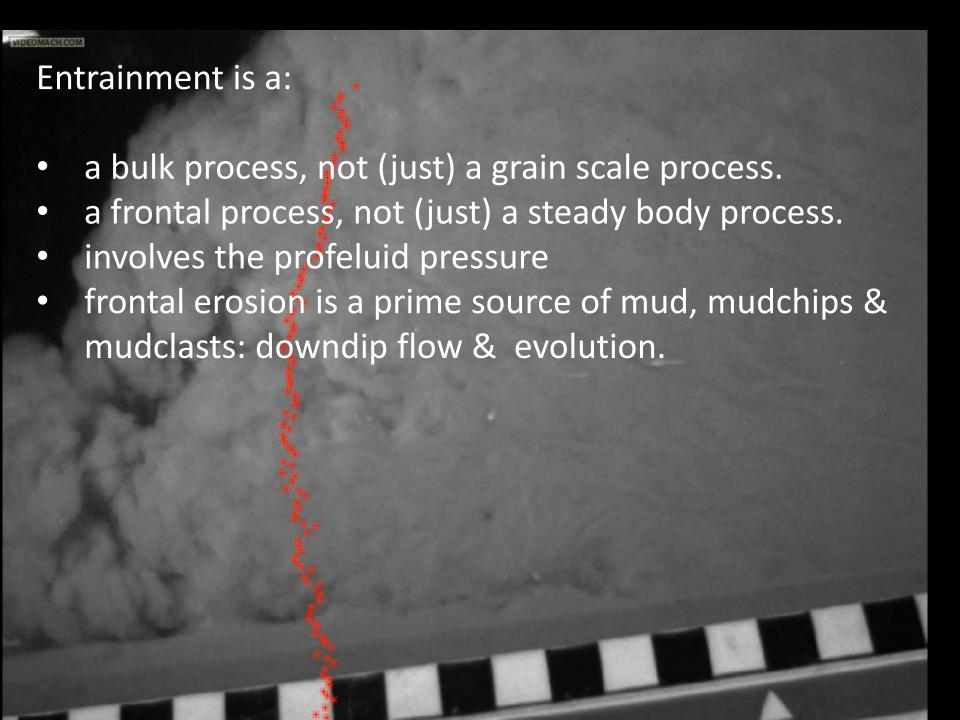
## Diverse morphodynamic evolution of architectural elements



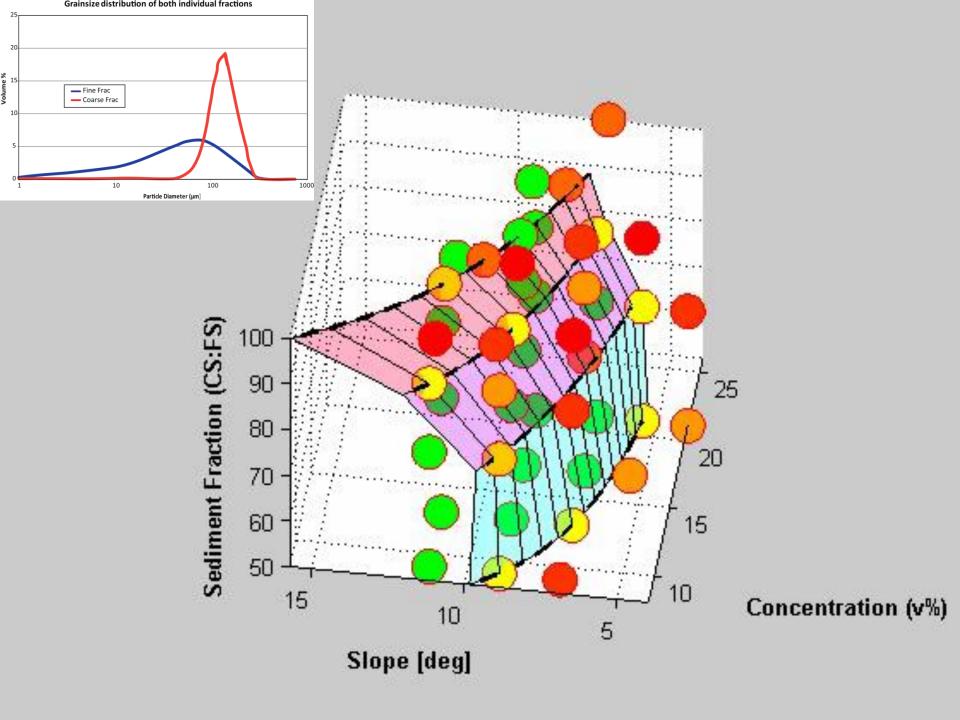






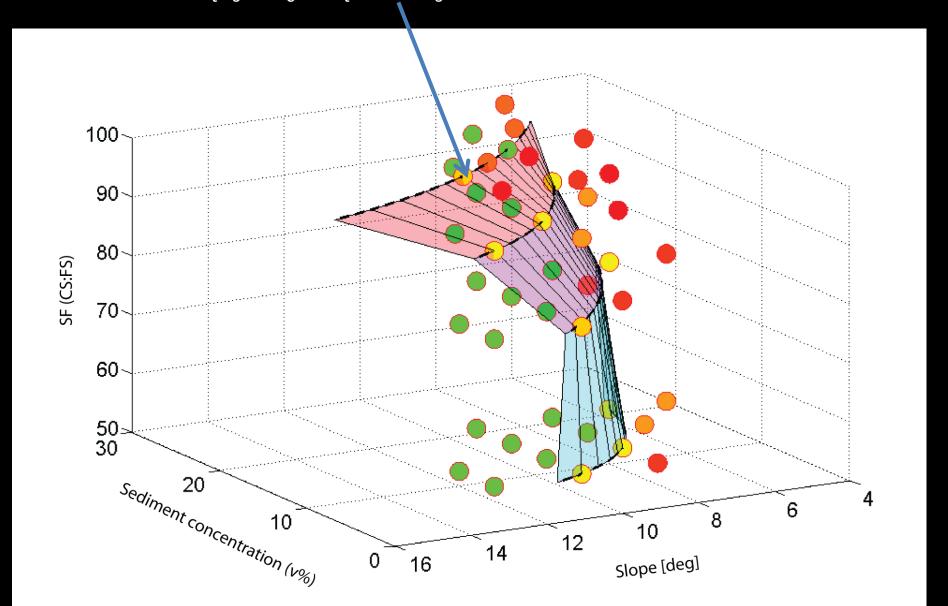




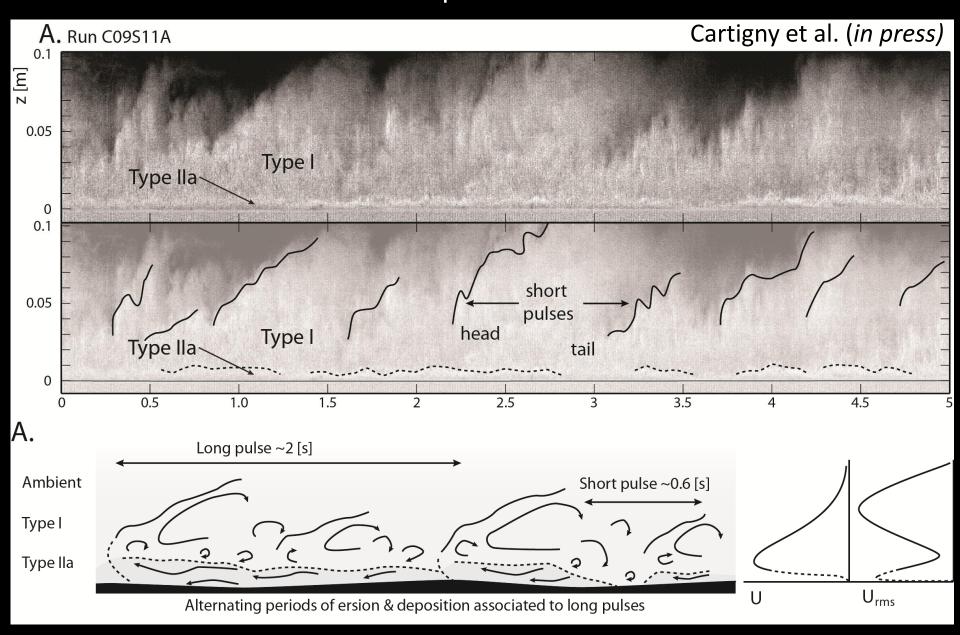


### What is the function for the bypass condition?

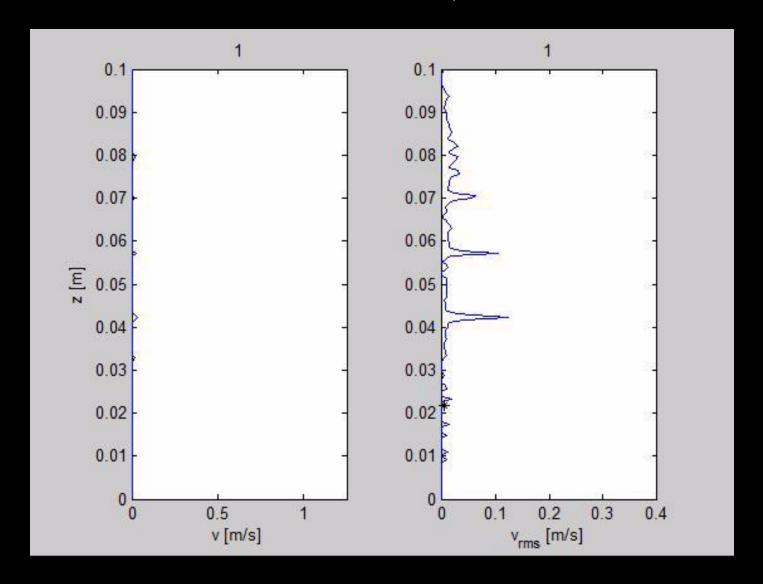
Bypass= F{ Ri<sub> $\tau$ </sub>v<sub>s</sub> =  $\kappa_c$  (Re<sub> $\tau$ </sub>); u\*/v<sub>s</sub>= 1; k\*; Fr'; S; Q; d; c; composition}



## What are the flow conditions and phenomena near the bed at deposition?



#### Flow conditions; u\*?



Location of the bed?

Few mm's of uncertainty...

#### u\*, z0, uncertain bed position

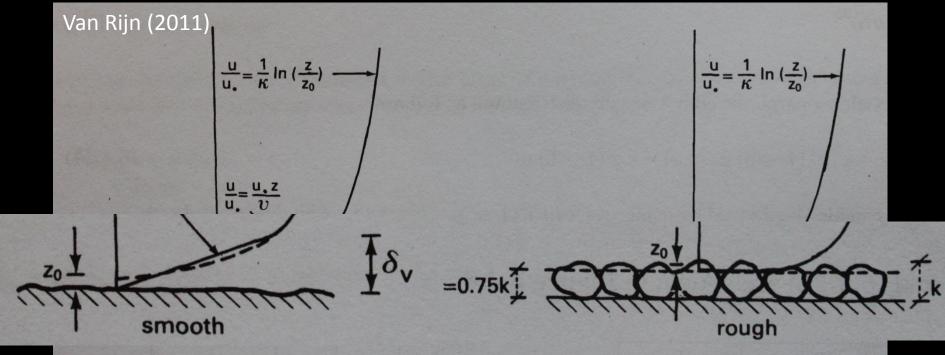
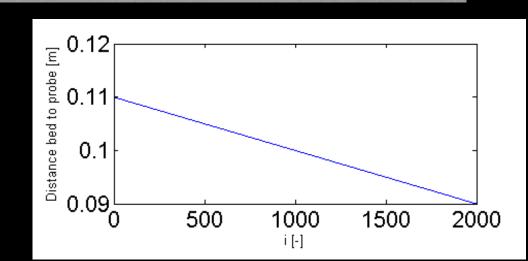
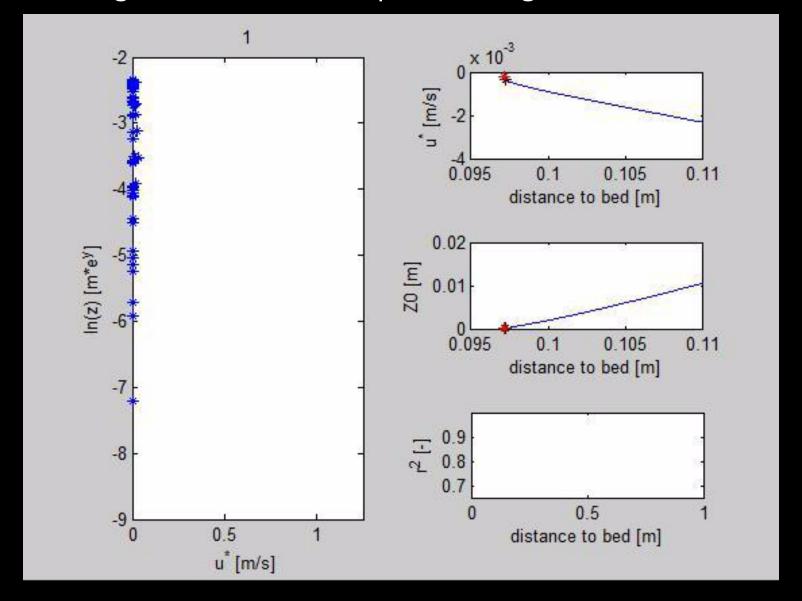


Figure 6.4.9 Velocity distributions in hydraulically smooth and rough flow

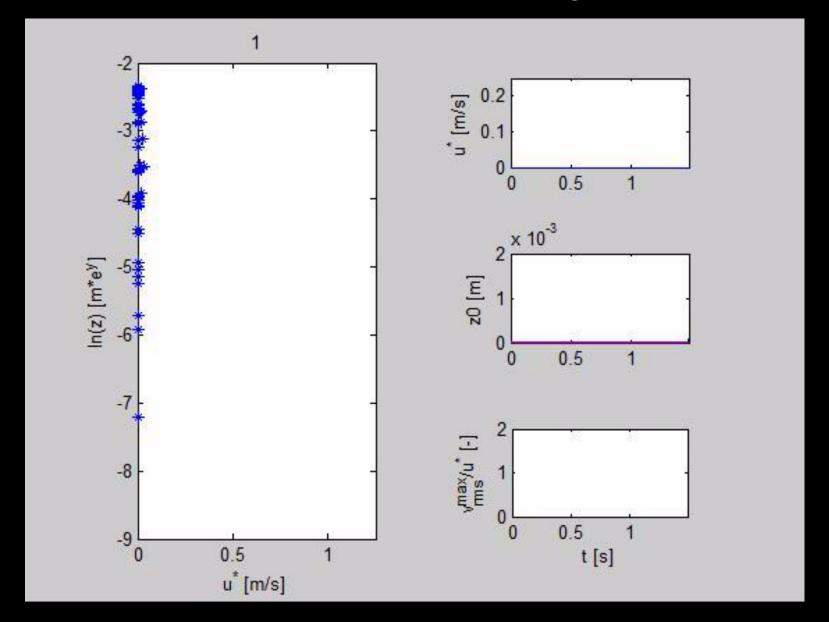
- Empeirical z0
- Best-fit for different bed locations

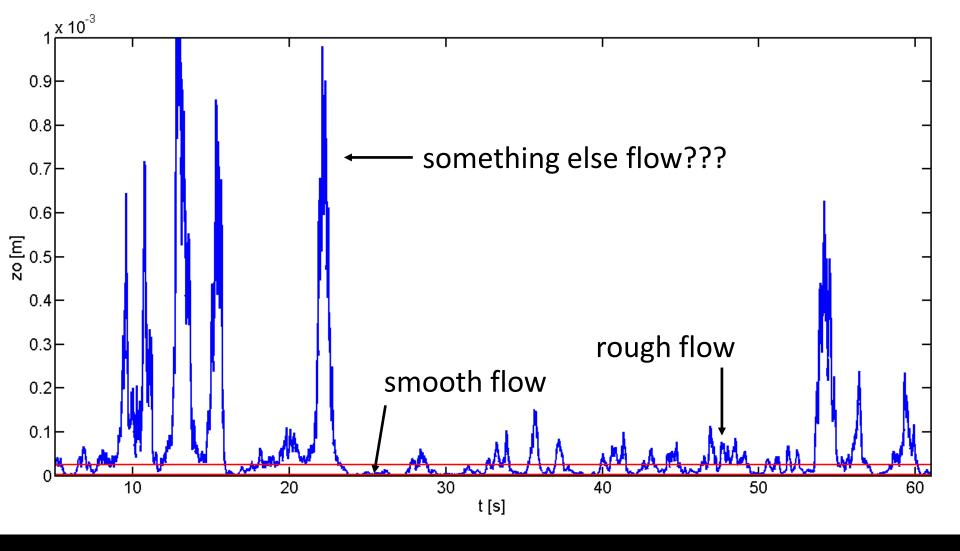


#### Fitting u\* and z0 without pre-defining the bed location.

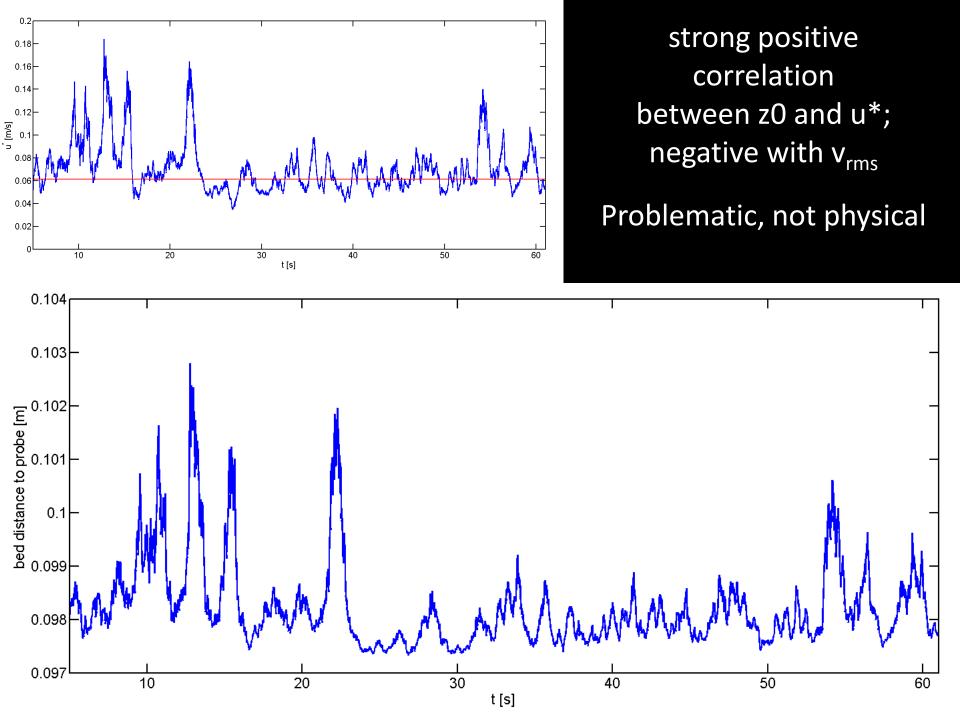


#### Evolution of $u^*$ , z0, and $r^2$ through time.

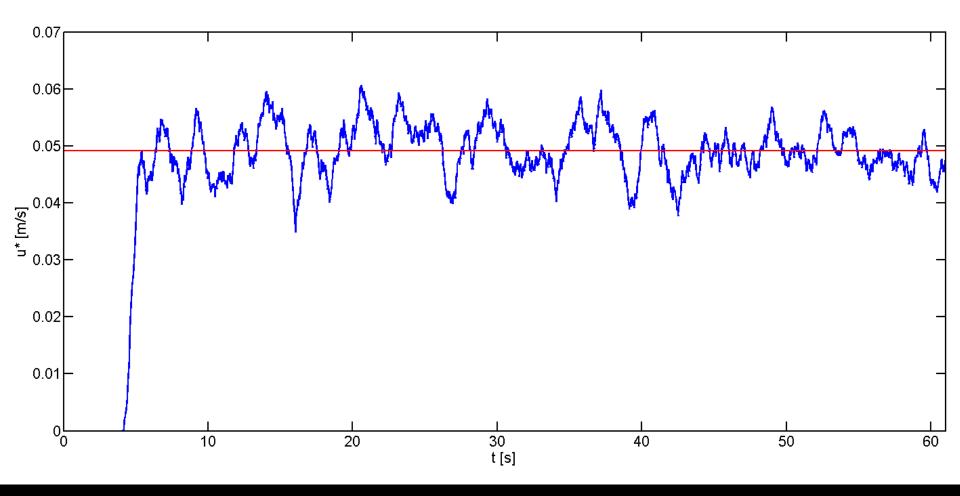




$$k^* = u^*k_s/v = 46$$
 transitionally rough  $k_s = 3 d_{90}$  (van Rijn)

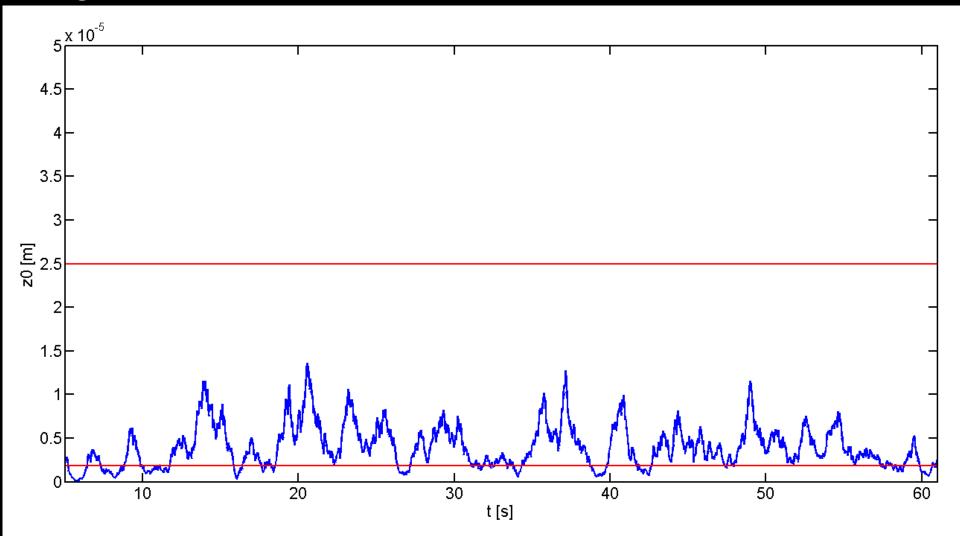


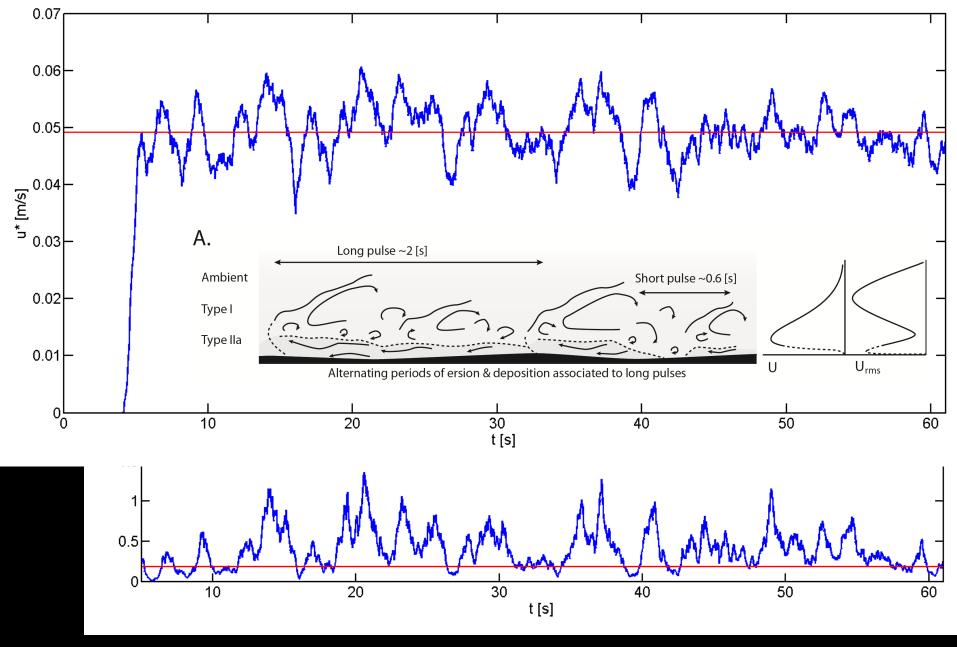
#### Fixing the bed at 0.0975 m



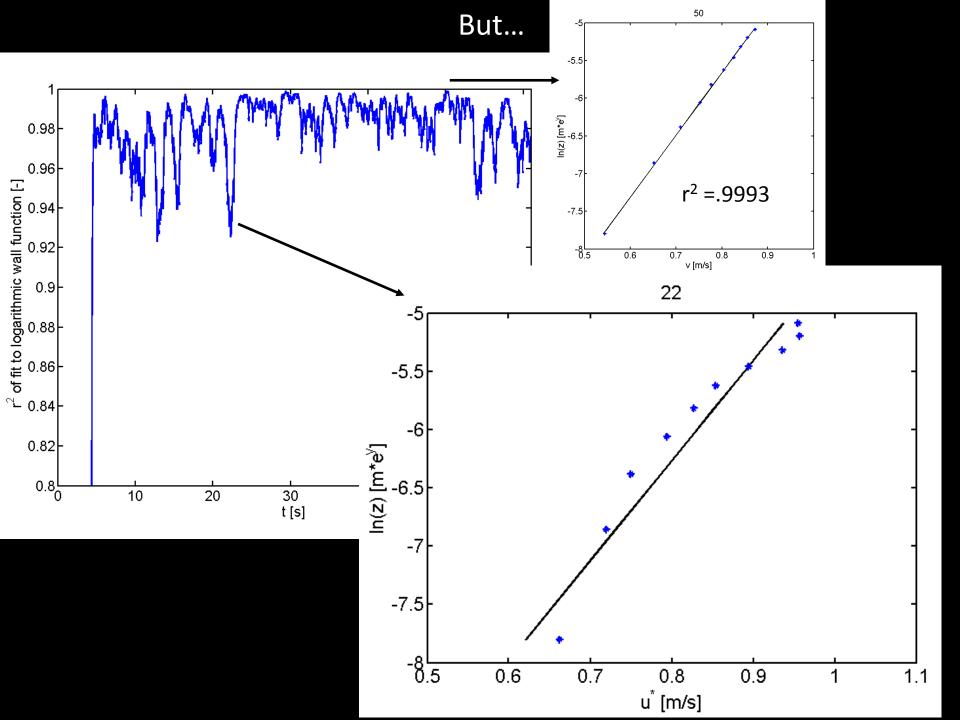
 $u^* = 0.049 \text{ m/s}; u^*_{rms} = 0.004 \text{ m/s}$ 

#### Fixing the bed at 0.0975 m

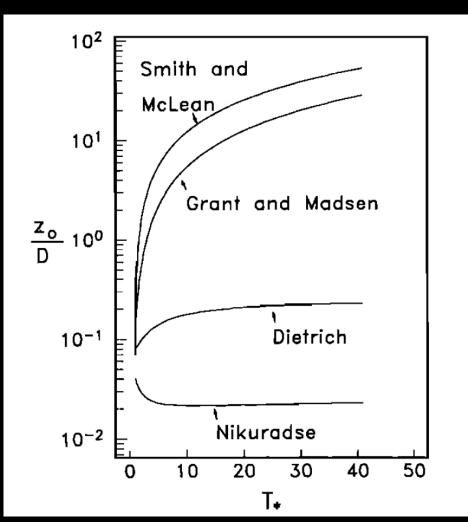




Positive correlation:  $k^* = 30-45$ ;  $u^* = 0.04-0.06$ ;  $z^0 = small-rough$ 



Lesson learned (methodological):  $u^*$ ,  $\tau_b$ , z0? When bed location & transport stage uncertainty: freedom, freedom (with misleading high  $r^2$ ).

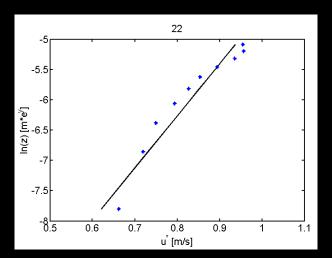


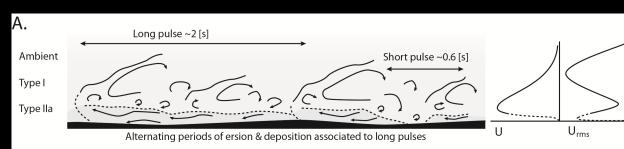
Wiberg & Rubin 1989 "Bed Roughness produced by saltating sediment"

Implication Hypothesis:
standard methods of determining
u\* and z0 break down because
the logarithmic law of the wall
does not apply, applying it
anyway results in
"ghost roughness"

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Nikuradse [1933]
        Owen [1964]
Smith & McLean [1977]
Grant & Madsen [1982]
        Dietrich [1982]
        ???
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Lessons learned (phenomenological): The velocity structure periodically fluctuates away from logarithmic

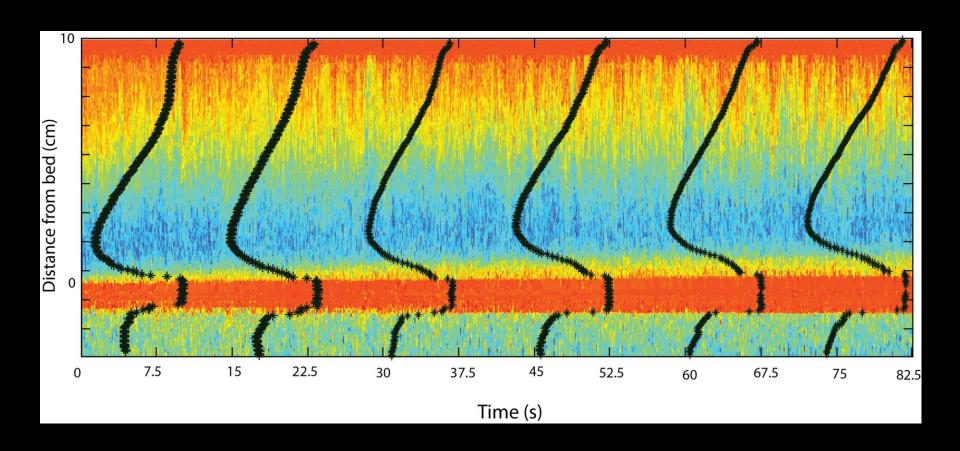




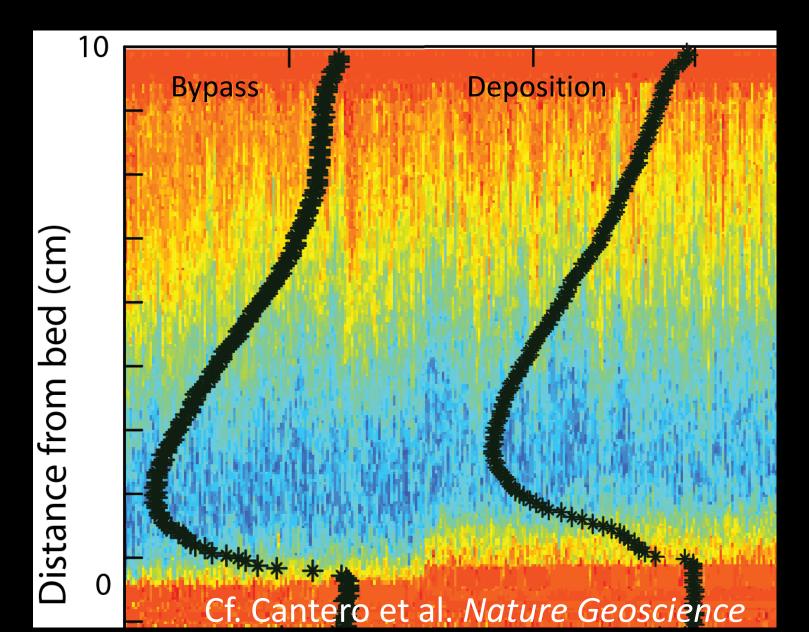
#### Hypothetic interpretations:

- Density stratification?  $\tau(z)/\rho(z) = u^*(z)$
- Turbulence dampening sensu Bala?  $Ri_{\tau}v_{s} > \kappa_{c}(Re_{\tau})$
- Formation of a bed extracts energy from the flow?
- Transport of momentum by settling of high velocity grains?
- Extra energy sink from collisions of particles in high density region?
- ...?
- Experimental methodological explanations? (Doppler; space-time averaging; density dependence of ultrasound velocity)

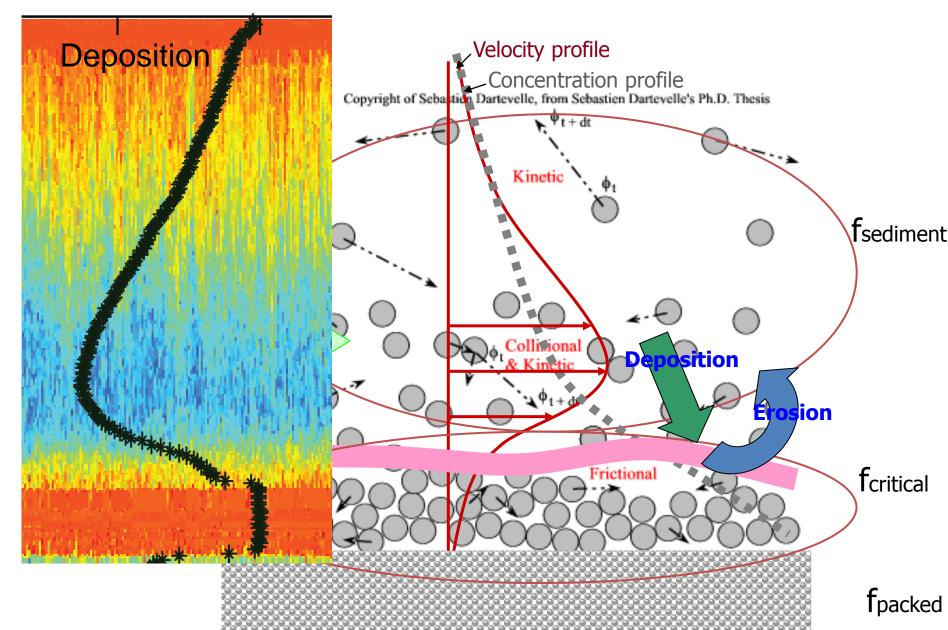
### Application to flow over aggrading bed Bypassing flow structure ≠ depositing flow structure



# The rheology of this basal zone is different from Newtonian Turbulent Flow?



## The deposit is formed in a modified basal rheological regime.



# Deposition from turbidity currents is the transformation from turbulent suspension to fixed bed.

This "Depositional Transformation" is a difficult subject:

- What is the role of porefluid pressure gradients?
- Collisional& frictional vs. viscous lubrication?
- How does the evolution of this basal layer affect the over-riding flow dynamics?
- How to model?
  - DNS resolving grains & porefluid (sensu Eckart's talk)?
  - Can a non-Newtonian continuum rheology be applied?

