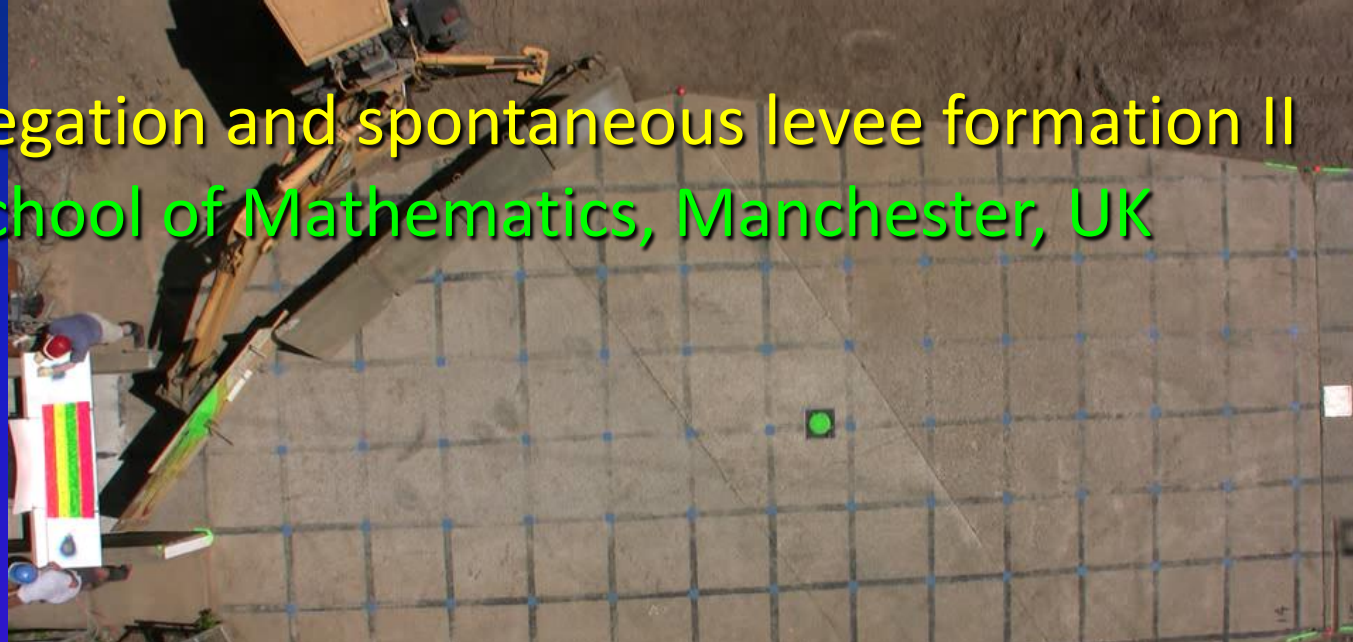


Particle-size segregation and spontaneous levee formation II

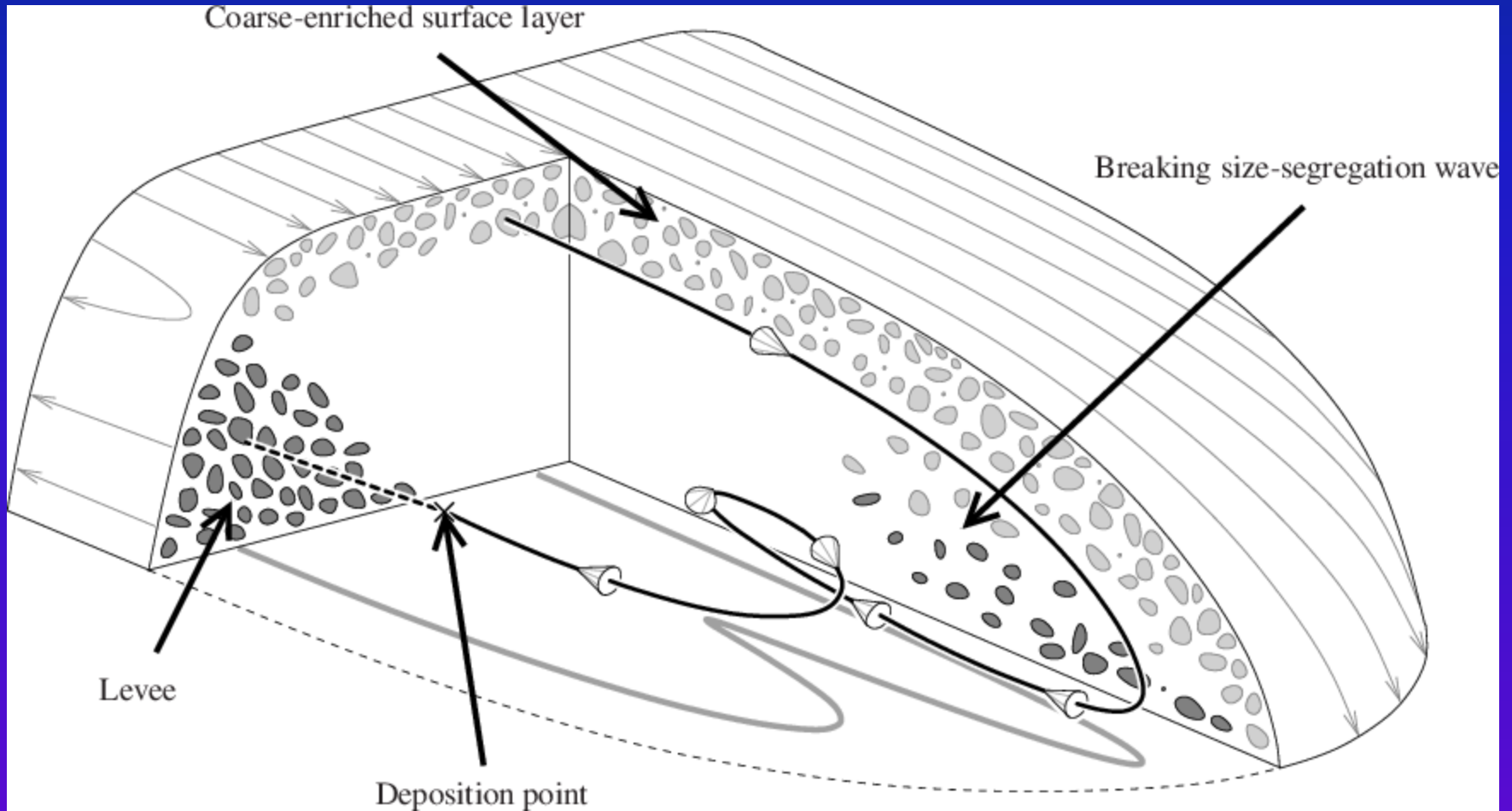
Nico Gray, School of Mathematics, Manchester, UK



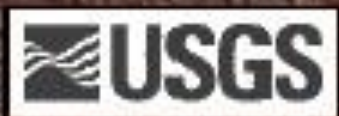
Johnson *et al* (2012) *J. Geophys. Res.* **117**, F01032



Schematic diagram of the levee formation process



- larger particles are shouldered to the sides to create levees
- this is an example of a segregation-mobility feedback effect



Mount St Helens May 18, 1980



Mt St Helens July 22, 1980

Mount St Helens, USA, 1980

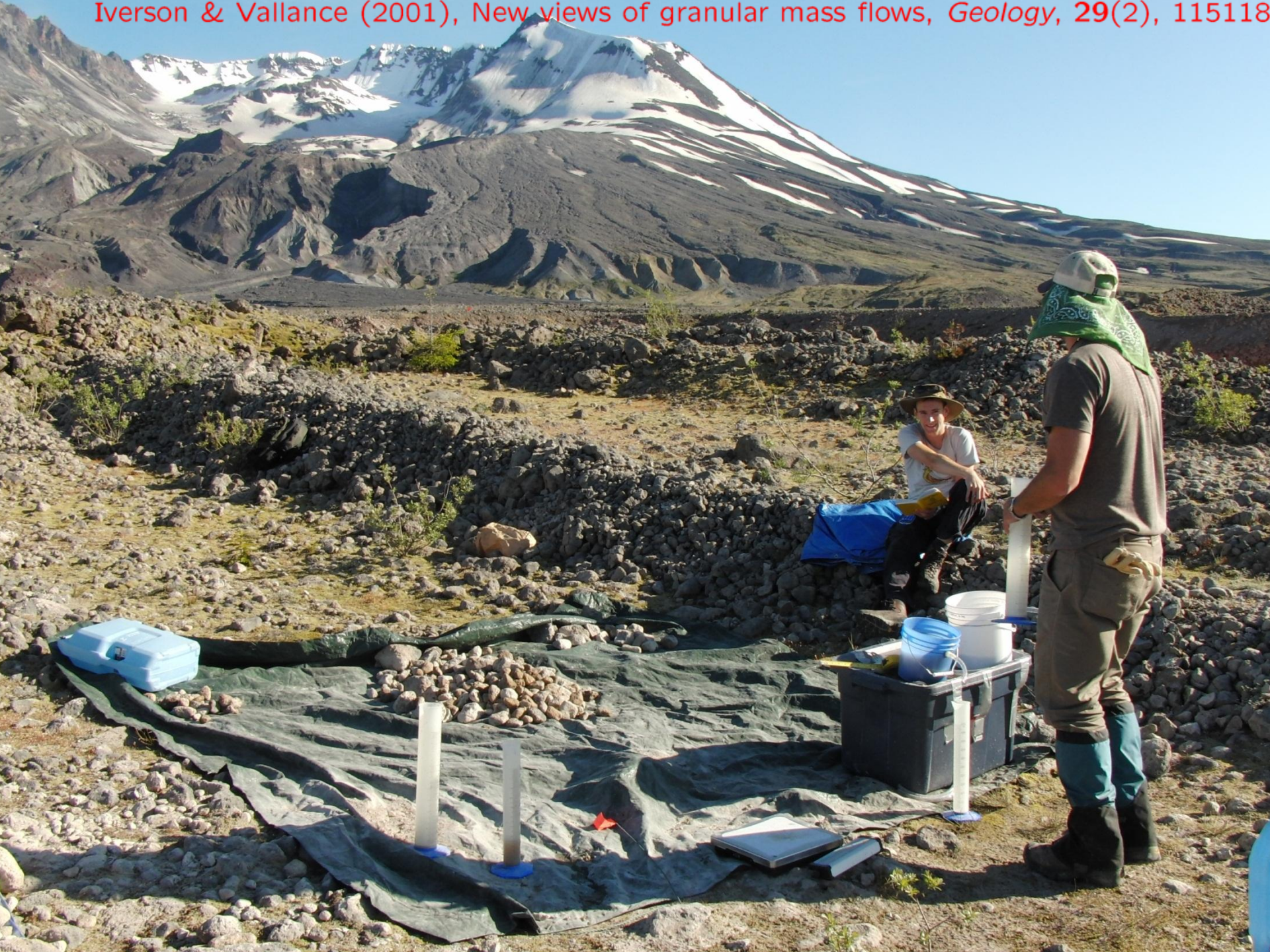
Finer grained interior



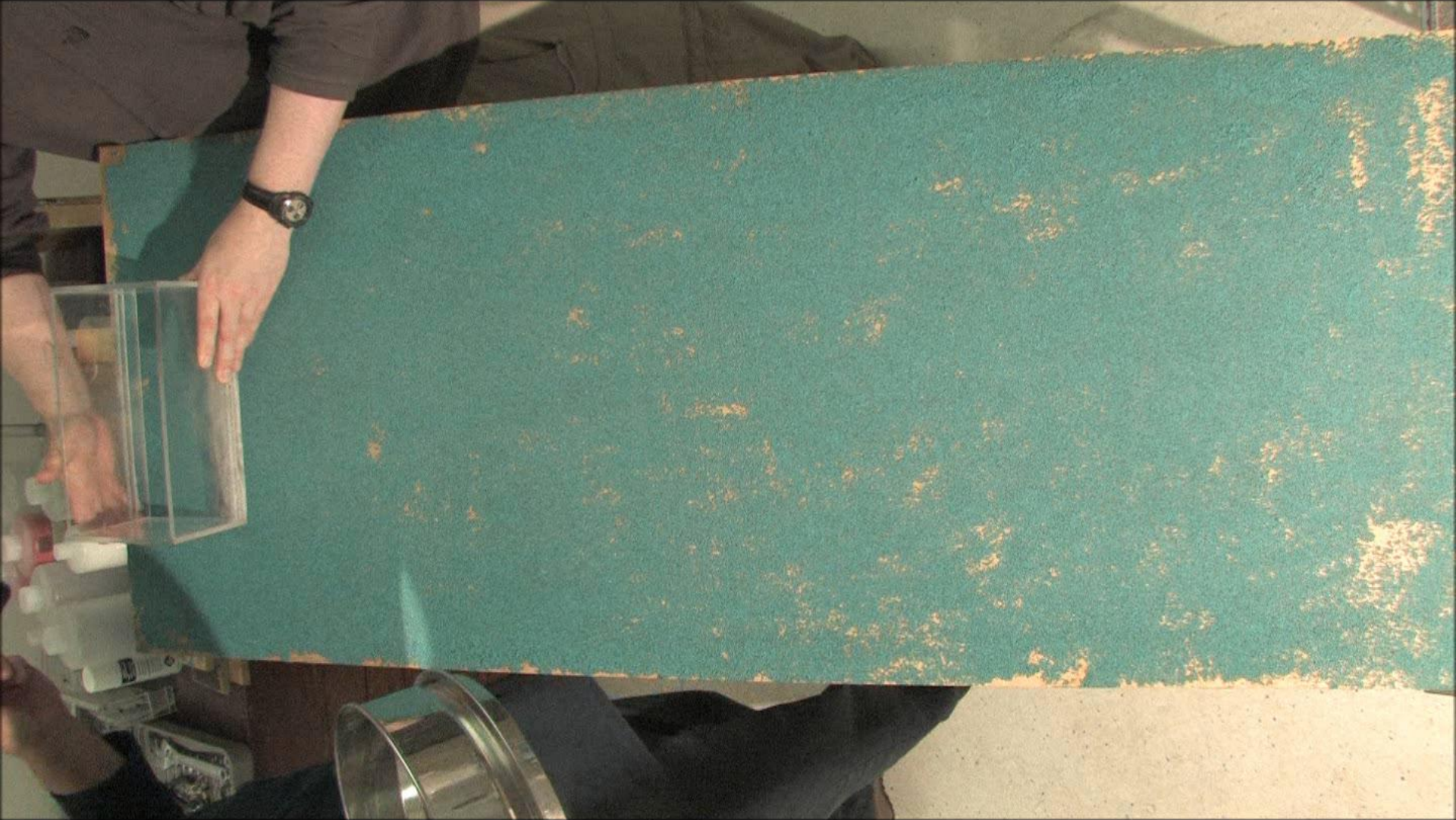
Coarse rich levee



- segregation occurs in many hazardous natural flows
 - debris-flows, pyroclastic flows & snow avalanches
- and leads to spontaneous flow organization and longer run-out

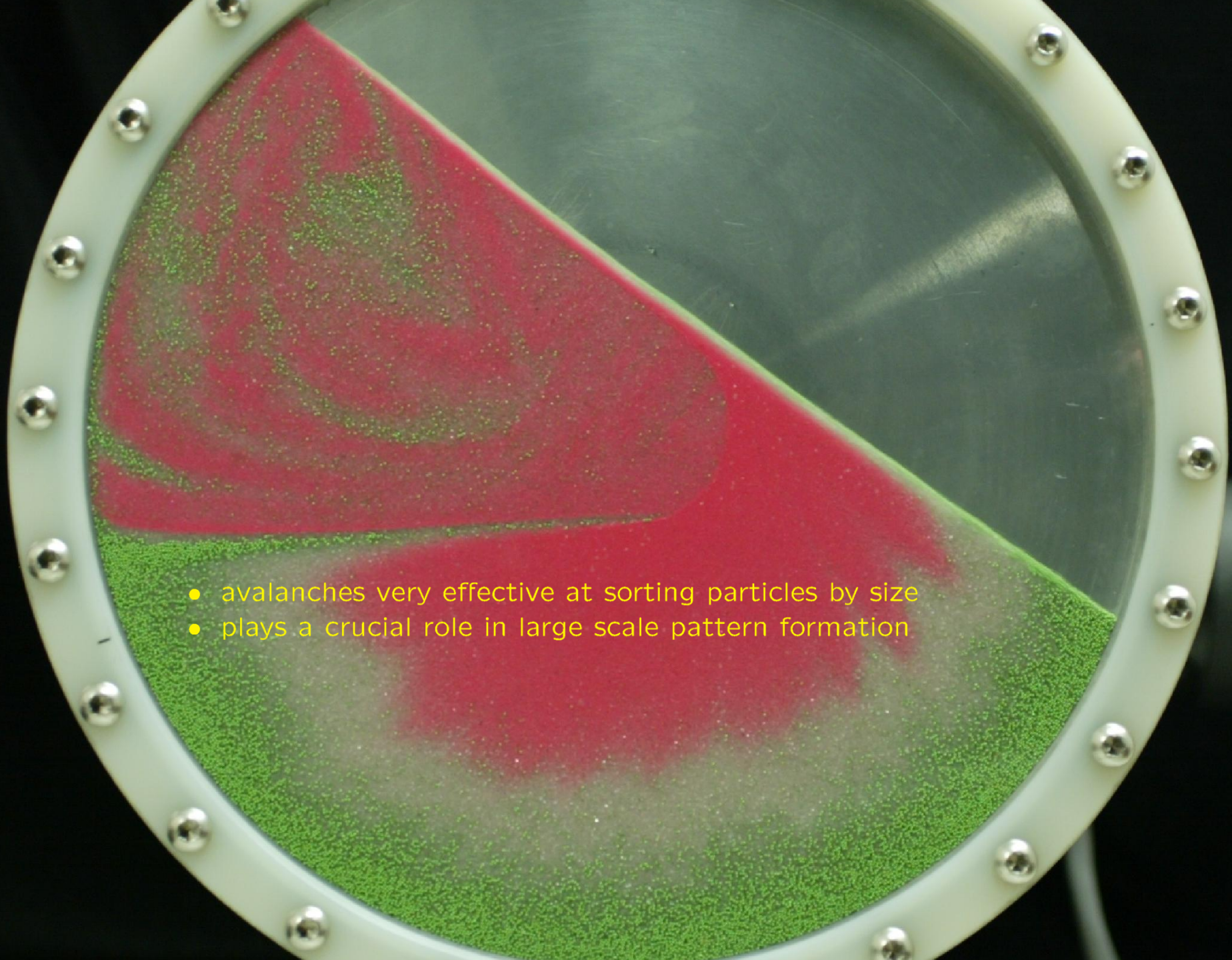


Segregation induced finger formation ...

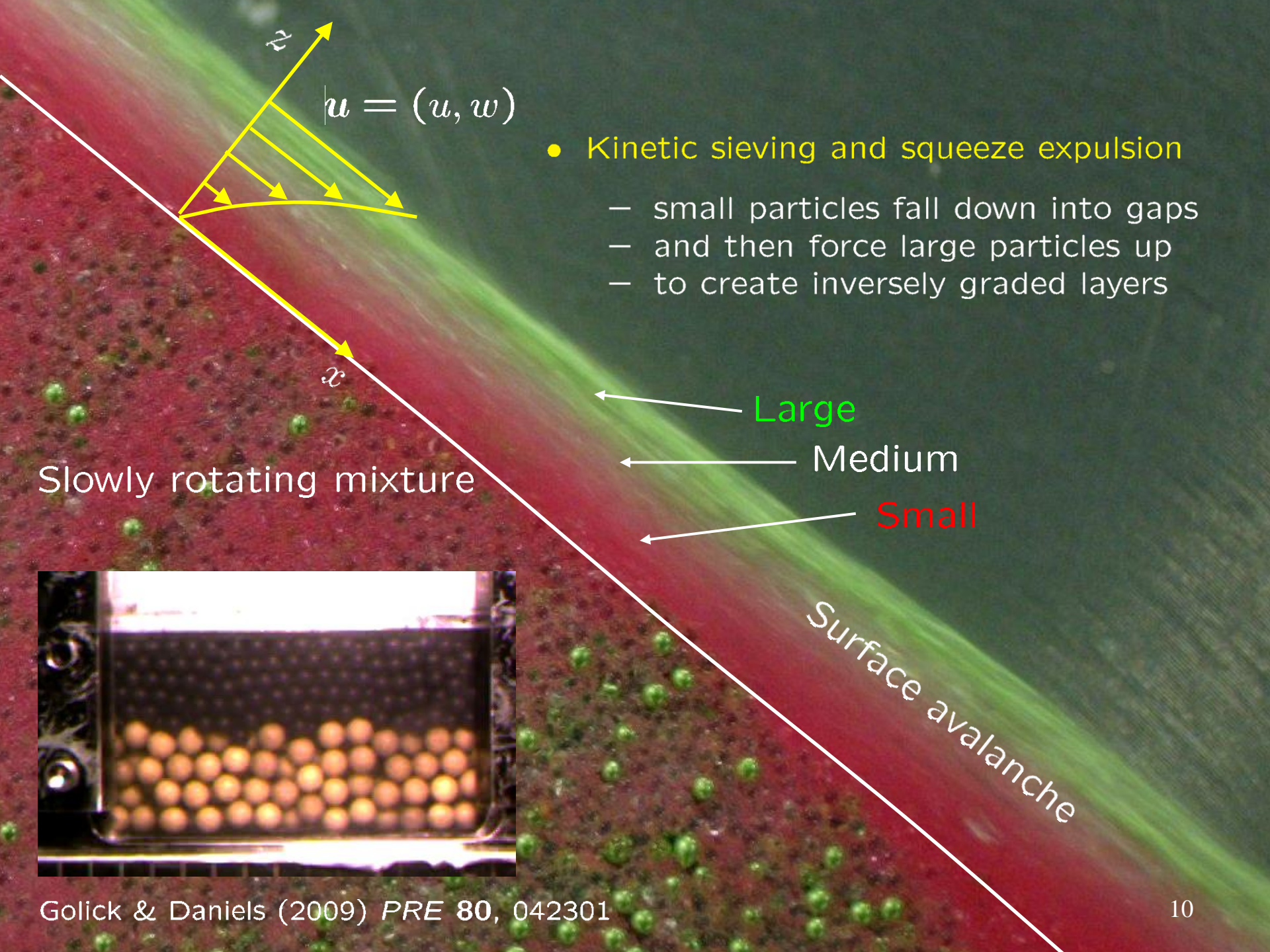


Pouliquen, Delours & Savage (1997), *Nature*. **386**, 816-817.
Woodhouse *et al.* (2012), *J. Fluid Mech.* **709**, 543-580.





- avalanches very effective at sorting particles by size
- plays a crucial role in large scale pattern formation



z

$$|\mathbf{u} = (u, w)$$

x

Slowly rotating mixture

- Kinetic sieving and squeeze expulsion
 - small particles fall down into gaps
 - and then force large particles up
 - to create inversely graded layers

Large

Medium

Small

Surface avalanche



Mixture framework

- the volume fraction ϕ^ν of constituent ν , per unit volume of mixture, lies in the range

$$0 \leq \phi^\nu \leq 1.$$

and their sum

$$\sum_{\forall \nu} \phi^\nu = 1.$$

- In standard mixture theory the partial and intrinsic density, stress, pressure and velocity fields satisfy

$$\rho^\nu = \phi^\nu \rho^{\nu*}, \quad \boldsymbol{\sigma}^\nu = \phi^\nu \boldsymbol{\sigma}^{\nu*}, \quad p^\nu = \phi^\nu p^{\nu*}, \quad \mathbf{u}^\nu = \mathbf{u}^{\nu*}$$

- The bulk density, pressure and velocity are

$$\rho = \sum_{\forall \nu} \rho^\nu, \quad p = \sum_{\forall \nu} p^\nu, \quad \rho \mathbf{u} = \sum_{\forall \nu} \rho^\nu \mathbf{u}^\nu$$

Mass and momentum balances for each constituent

- Each constituent satisfies individual mass

$$\frac{\partial \rho^\nu}{\partial t} + \nabla \cdot (\rho^\nu \mathbf{u}^\nu) = 0,$$

and momentum balances

$$\frac{\partial}{\partial t}(\rho^\nu \mathbf{u}^\nu) + \nabla \cdot (\rho^\nu \mathbf{u}^\nu \otimes \mathbf{u}^\nu) = \nabla \cdot \boldsymbol{\sigma}^\nu + \rho^\nu \mathbf{g} + \boldsymbol{\beta}^\nu,$$

where \otimes is the dyadic product and \mathbf{g} is the gravitational acceleration vector.

- The interaction drag $\boldsymbol{\beta}^\nu$ consists of three terms

$$\boldsymbol{\beta}^\nu = p \nabla f^\nu - \rho^\nu c (\mathbf{u}^\nu - \mathbf{u}) - \rho d \nabla \phi^\nu,$$

c is the coefficient of inter-particle drag,

d is the coefficient of diffusive remixing

- Acceleration negligible in the z direction.
- Bulk momentum balance \Rightarrow pressure lithostatic

$$p = \rho g(s - z) \cos \zeta.$$

- If pressure is shared in proportion

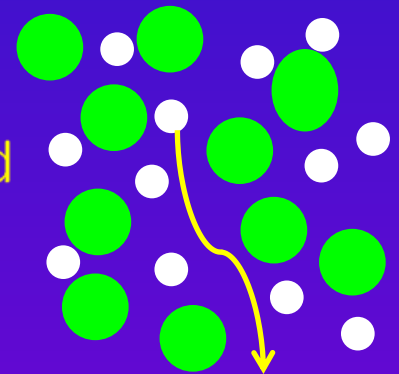
$$p^\nu = \phi^\nu p, \quad \Rightarrow \quad \text{NO SEGREGATION}$$

- Use non-standard partial/intrinsic pressure relation

$$p^\nu = f^\nu p,$$

where f^ν determines the proportion load

$$0 \leq f^\nu \leq 1, \quad \sum_{\forall \nu} f^\nu = 1.$$



- Assuming normal accelerations are negligible

$$\phi^\nu w^\nu = \phi^\nu w + (f^\nu - \phi^\nu)(g/c) \cos \zeta - (d/c) \frac{\partial \phi^\nu}{\partial z},$$

$f^\nu - \phi^\nu > 0$ particles rise

$f^\nu - \phi^\nu = 0$ no relative motion

$f^\nu - \phi^\nu < 0$ particles percolate downwards

- Particles in a pure phase carry all of the load

$$f^\nu = 1, \quad \text{when} \quad \phi^\nu = 1,$$

- When no particles, they cannot carry any load

$$f^\nu = 0, \quad \text{when} \quad \phi^\nu = 0.$$

- Large particles carry more load than small

$$f^l = \phi^l + B_{ls} \phi^l \phi^s \quad (\text{bi-disperse})$$

The multi-component segregation remixing equation

- f^ν must reduce to bidisperse case in any submixture
- suggests additive decomposition

$$f^\nu = \phi^\nu + \sum_{\forall \mu} B_{\nu\mu} \phi^\nu \phi^\mu, \quad (\text{polydisperse case})$$

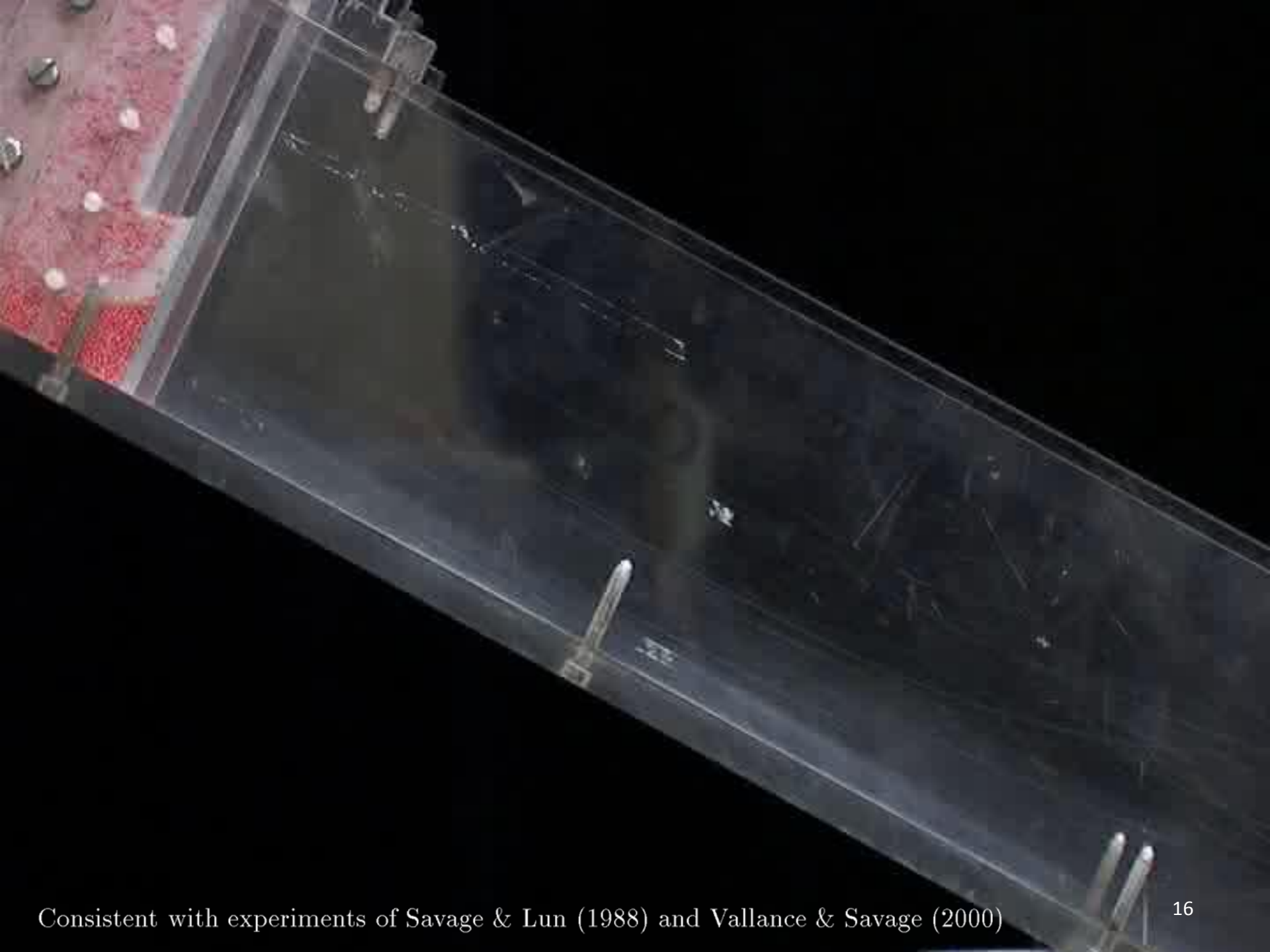
where $B_{\nu\mu} = -B_{\mu\nu}$ and $B_{(\nu\nu)} = 0$.

- Scaling on thickness H , length L and velocity U implies segregation remixing equation (phase ν) is

$$\frac{\partial \phi^\nu}{\partial t} + \nabla \cdot (\phi^\nu \mathbf{u}) + \frac{\partial}{\partial z} \left(\sum_{\forall \mu} S_{\nu\mu} \phi^\nu \phi^\mu \right) = \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^\nu}{\partial z} \right),$$

where

$$S_{\nu\mu} = \frac{Lg \cos \zeta}{HUc} B_{\nu\mu}, \quad D_r = \frac{Ld}{H^2Uc}.$$



Consistent with experiments of Savage & Lun (1988) and Vallance & Savage (2000)

Bi-disperse mixtures

- Yields two equations for large and small particles

$$\begin{aligned}\frac{\partial \phi^l}{\partial t} + \nabla \cdot (\phi^l \mathbf{u}) + \frac{\partial}{\partial z}(S_{ls} \phi^l \phi^s) &= \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^l}{\partial z} \right), \\ \frac{\partial \phi^s}{\partial t} + \nabla \cdot (\phi^s \mathbf{u}) - \frac{\partial}{\partial z}(S_{ls} \phi^s \phi^l) &= \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^s}{\partial z} \right).\end{aligned}$$

- The summation condition $\sum \phi^\nu = 1$ implies

$$\phi^l + \phi^s = 1,$$

- Large particle concentration can be eliminated

$$\frac{\partial \phi^s}{\partial t} + \nabla \cdot (\phi^s \mathbf{u}) - \frac{\partial}{\partial z}(S_{ls} \phi^s (1 - \phi^s)) = \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^s}{\partial z} \right).$$

Bridgwater, Foo & Stephens (1985), *Powder Technol.* **41**, 147-158

Savage & Lun (1988) *J. Fluid Mech.* **189**, 311-335

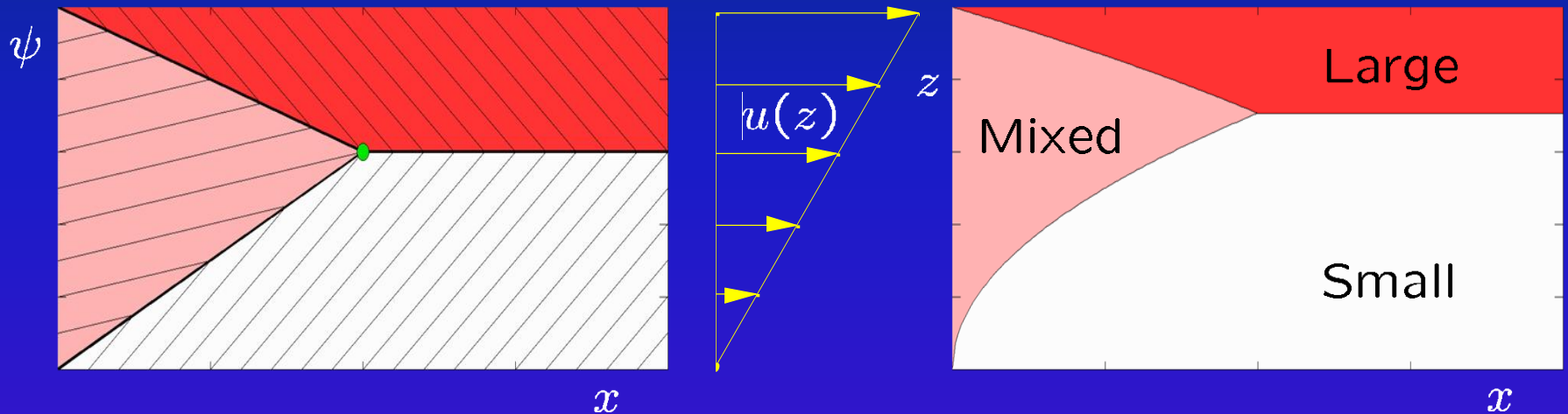
Dolgunin & Ukolov (1995) *Powder Technol.* **83**, 95-103

Gray & Thornton (2005) *Proc. Roy. Soc. A.* **461**, 1447-1473.

Thornton, Gray & Hogg (2006) *J. Fluid Mech.* **550**, 1-25.

Gray & Chugunov (2006) *J. Fluid Mech.* **569**, 365-398.

Steady-state concentration shocks in absence of diffusive-remixing



- shock height $s(x)$ satisfies the jump condition

$$\left[\phi u \frac{ds}{dx} + S_{ls} \phi (1 - \phi) \right] = 0 \Rightarrow u \frac{ds}{dx} = S_{ls} (\phi^+ + \phi^- - 1) = \frac{d\psi}{dx}$$

- Using depth-integrated velocity coordinates

$$\psi = \int_0^z u(z') dz'$$

- this can be integrated to show there are three intersecting shocks for a homogeneous inflow with $\phi = \phi_0$

$$\psi_1 = S_{ls} \phi_0 x, \quad \psi_2 = 1 - S_{ls} (1 - \phi_0) x, \quad \psi_3 = \phi_0$$

A ternary mixture of large medium and small particles

- Theory yields three equations, but one can be eliminated since

$$\phi^m = 1 - \phi^s - \phi^l$$

- To give two equations for the large and small particles

$$\begin{aligned} \frac{\partial \phi^l}{\partial t} + \nabla \cdot (\phi^l \mathbf{u}) + \frac{\partial}{\partial z} (S_{lm} \phi^l (1 - \phi^l - \phi^s) + S_{ls} \phi^l \phi^s) &= \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^l}{\partial z} \right) \\ \frac{\partial \phi^s}{\partial t} + \nabla \cdot (\phi^s \mathbf{u}) + \frac{\partial}{\partial z} (-S_{ls} \phi^s \phi^l - S_{ms} \phi^s (1 - \phi^l - \phi^s)) &= \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi^s}{\partial z} \right) \end{aligned}$$

- A steady-state solution for a homogeneous inflow at $x = 0$

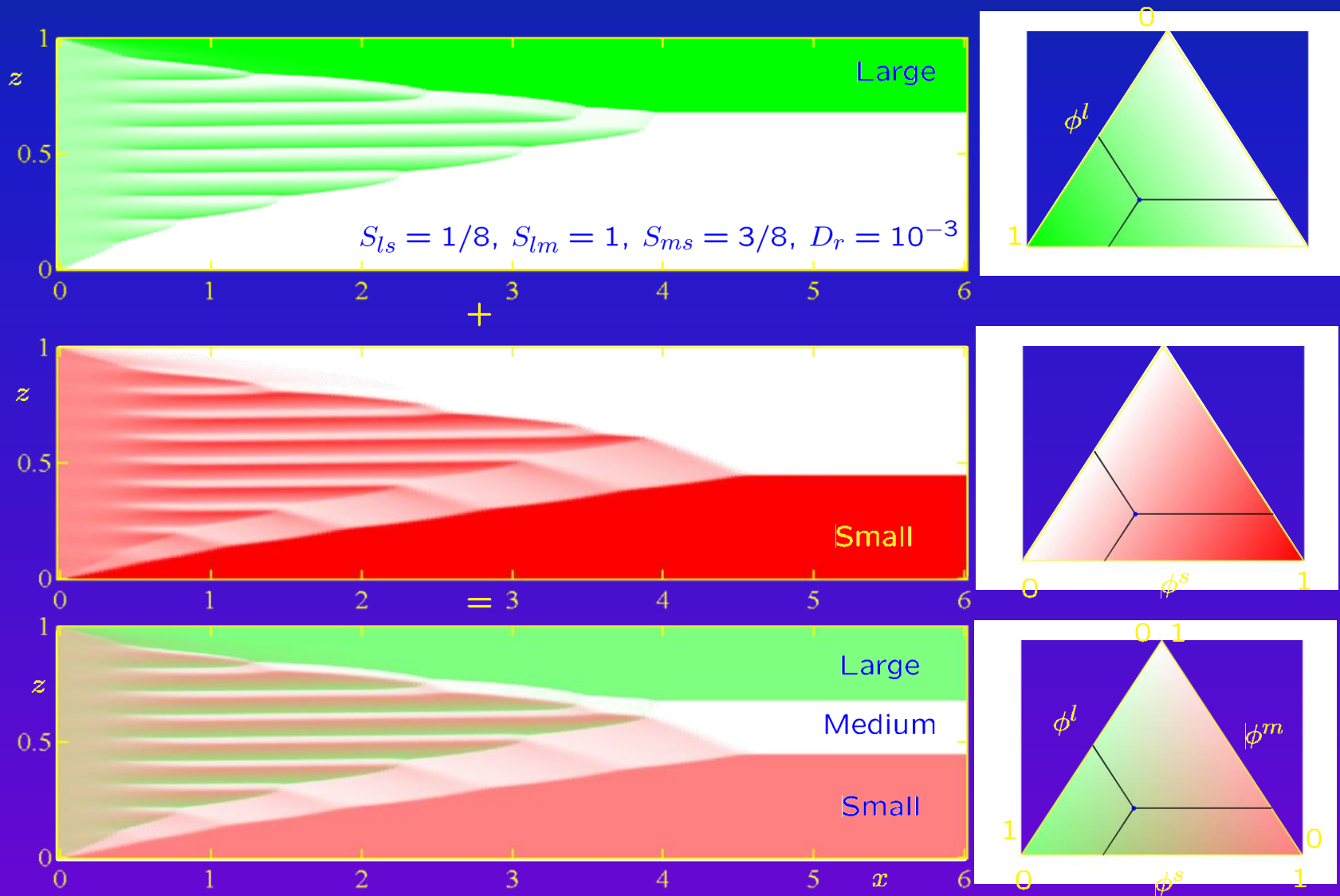
$$\phi^v(0, z) = \phi_0^v$$

- and with prescribed velocity field

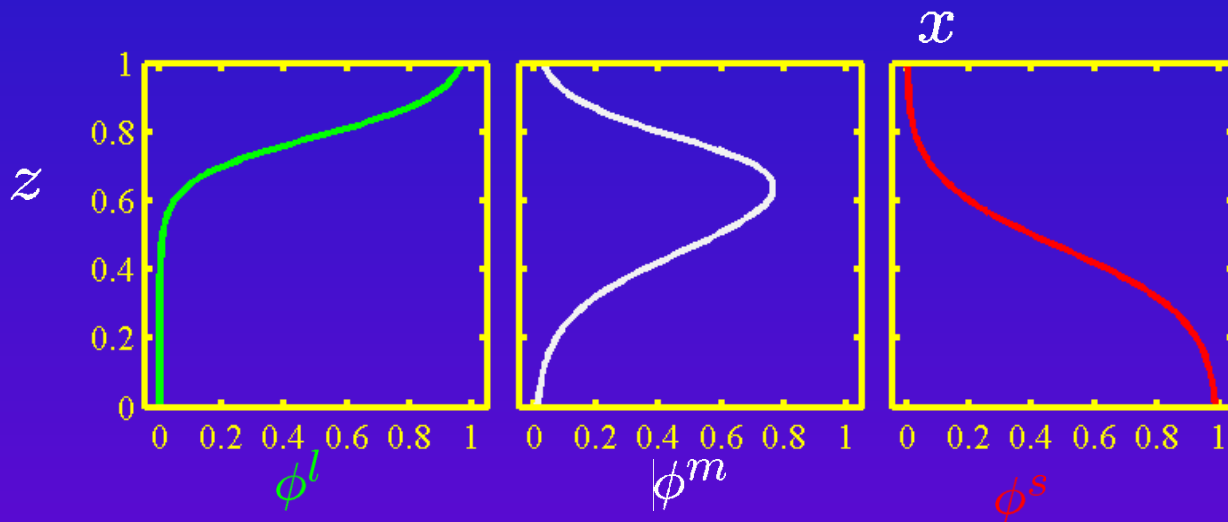
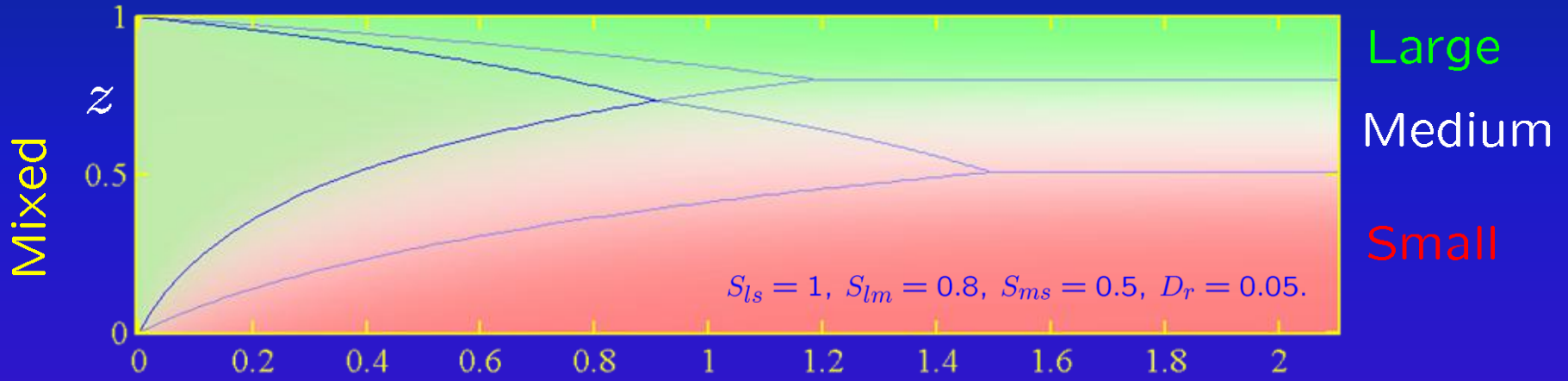
$$u = u(z), \quad w = 0$$

- subject to no-flux conditions at $z = 0, 1$
- can be computed using Matlab function pdepe (Galerkin Method)

For non-monotonic segregation rates there can be linear instabilities!



Reverse distribution grading



- comparable to experiments ...

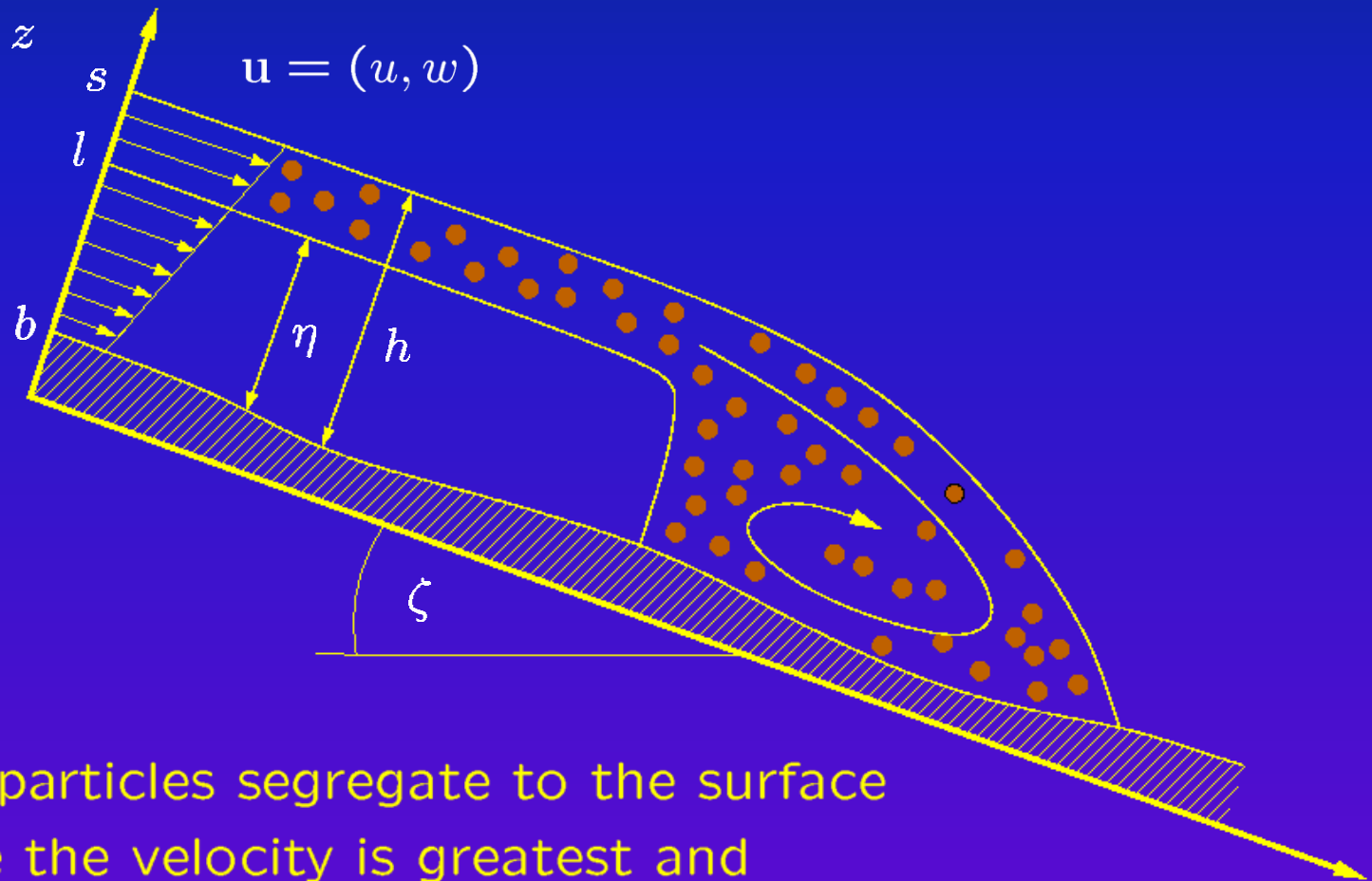
Gray & Ancy (2011) *J. Fluid Mech.* **678**, 535-558

Wiederseiner *et al.* (2011) *Phys. Fluids* **23**, 013301

- and DEM simulations (Jim McElwaine ...)



Transport and accumulation of large particles



- large particles segregate to the surface
- where the velocity is greatest and
- are transported to the flow front where they are
- over run and recirculated by particle size segregation

A depth averaged theory for particle size segregation

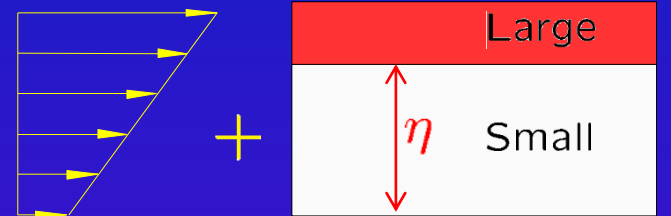
- Integrating the segregation-remixing equation w.r.t z
- subject to the no flux and kinematic boundary conditions gives

$$\frac{\partial}{\partial t}(h\bar{\phi}) + \frac{\partial}{\partial x}(h\bar{\phi u}) = 0$$

- where the integrals evaluated assuming

$$h\bar{\phi} = \int_b^s \phi^s dz = \eta$$

$$h\bar{\phi u} = \int_b^s \phi^s u dz = \eta \bar{u} - (1 - \alpha) \bar{u} \eta \left(1 - \frac{\eta}{h}\right)$$



i.e. linear velocity with basal slip and sharp segregation

- This yields the large particle transport equation

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(\eta \bar{u}) - \frac{\partial}{\partial x} \left((1 - \alpha) \bar{u} \eta \left(1 - \frac{\eta}{h}\right) \right) = 0.$$

- for the evolution of the inversely graded shock interface η .

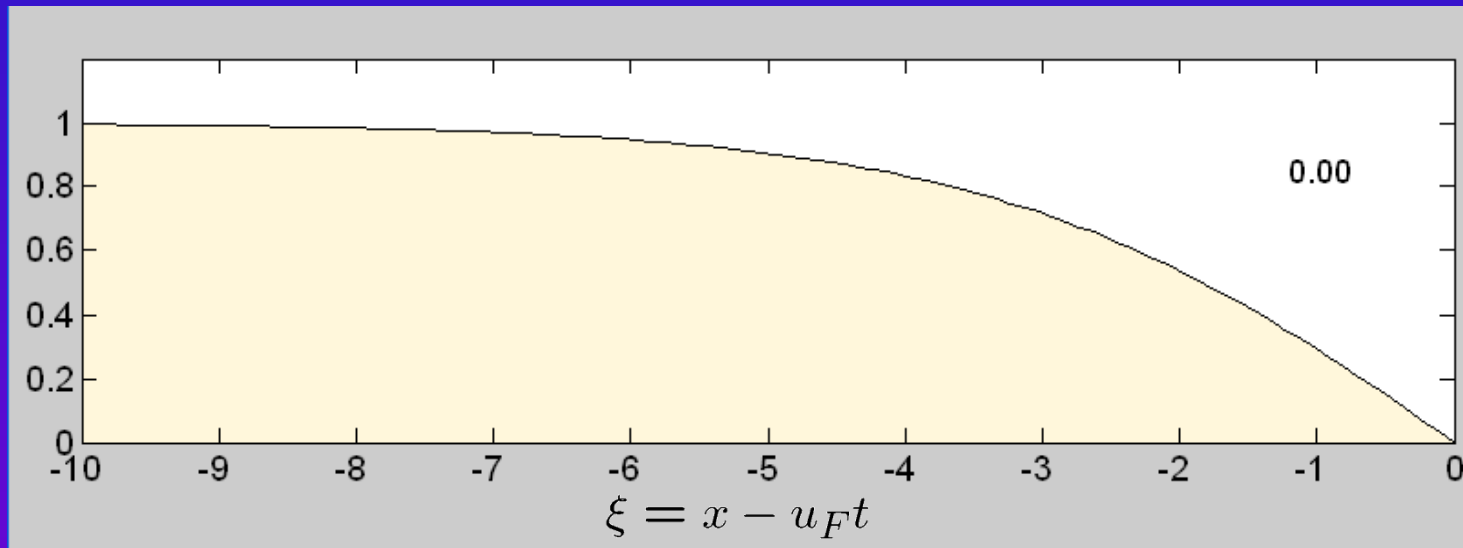
- Using $\eta = h\bar{\phi}$ this can also be rewritten as

$$\frac{\partial}{\partial t}(h\bar{\phi}) + \frac{\partial}{\partial x}(h\bar{\phi}\bar{u}) - \frac{\partial}{\partial x} \left((1 - \alpha)h\bar{u}\bar{\phi} (1 - \bar{\phi}) \right) = 0.$$

- Remarkably similar to the segregation equation ...

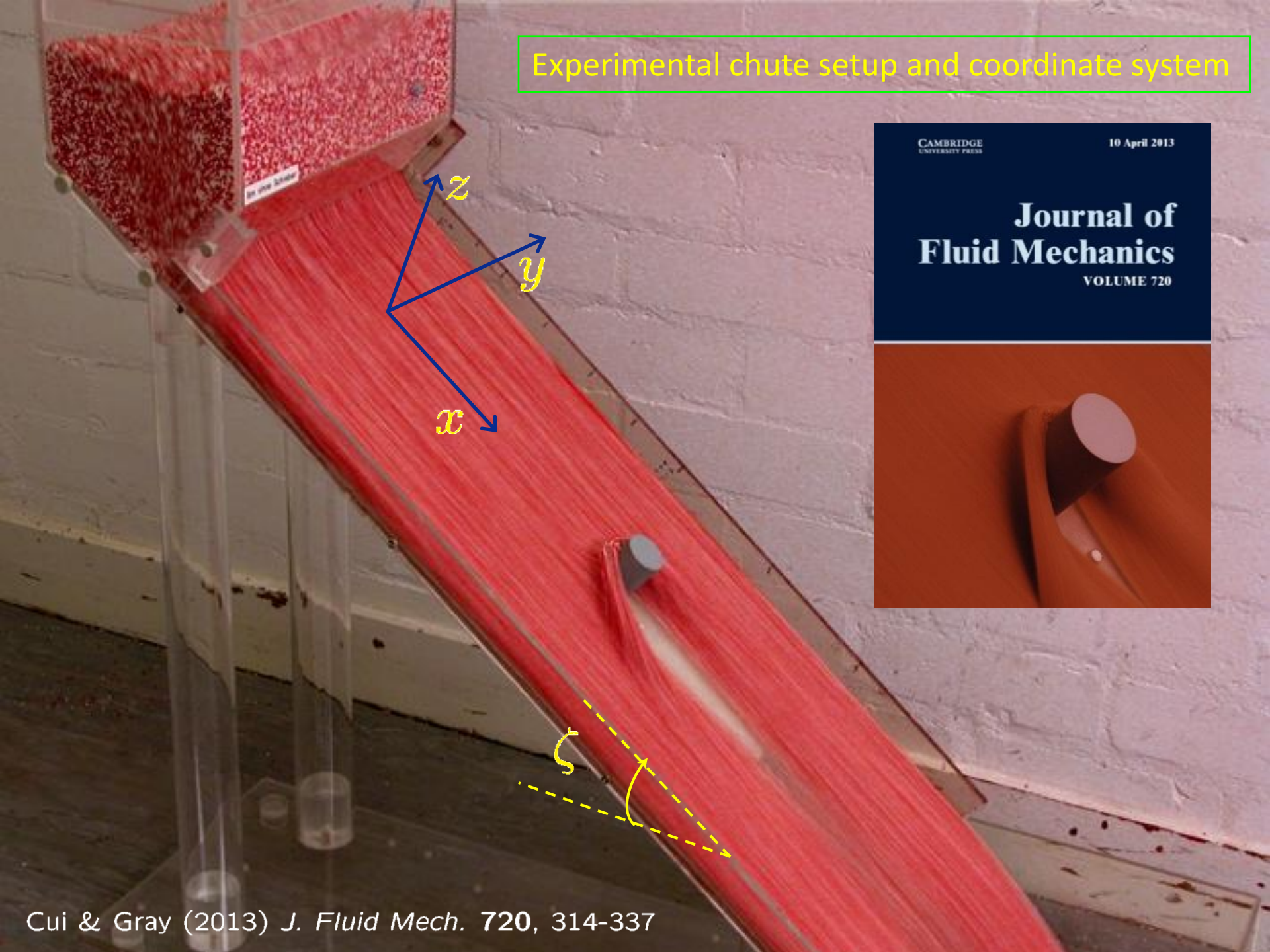
$$\frac{\partial \phi}{\partial t} + \frac{\partial}{\partial x}(\phi u) + \frac{\partial}{\partial z}(\phi w) - S_{ls} \frac{\partial}{\partial z}(\phi(1 - \phi)) = \frac{\partial}{\partial z} \left(D_r \frac{\partial \phi}{\partial z} \right)$$

- Large grains transported forwards to form bouldery flow front



- more RESISTIVE larger particles \Rightarrow feedback on bulk flow

Experimental chute setup and coordinate system



- For avalanche thickness h and mean velocity $\bar{\mathbf{u}} = (\bar{u}, \bar{v})$ in the downslope x and cross-slope y directions.

$$\frac{\partial h}{\partial t} + \frac{\partial}{\partial x}(h\bar{u}) + \frac{\partial}{\partial y}(h\bar{v}) = 0,$$

$$\frac{\partial}{\partial t}(h\bar{u}) + \frac{\partial}{\partial x}(h\bar{u}^2) + \frac{\partial}{\partial y}(h\bar{u}\bar{v}) + \frac{\partial}{\partial x} \left(\frac{1}{2}gh^2 \cos \zeta \right) = hgS_{(x)},$$

$$\frac{\partial}{\partial t}(h\bar{v}) + \frac{\partial}{\partial x}(h\bar{u}\bar{v}) + \frac{\partial}{\partial y}(h\bar{v}^2) + \frac{\partial}{\partial y} \left(\frac{1}{2}gh^2 \cos \zeta \right) = hgS_{(y)},$$

- source terms composed of gravity, basal friction μ and gradients of the basal topography b

$$S_{(x)} = \sin \zeta - \mu(\bar{u}/|\bar{\mathbf{u}}|) \cos \zeta - \frac{\partial b}{\partial x} \cos \zeta,$$

$$S_{(y)} = -\mu(\bar{v}/|\bar{\mathbf{u}}|) \cos \zeta - \frac{\partial b}{\partial y} \cos \zeta,$$

- system is hyperbolic, Froude number $Fr = |\bar{\mathbf{u}}|/\sqrt{gh \cos \zeta}$

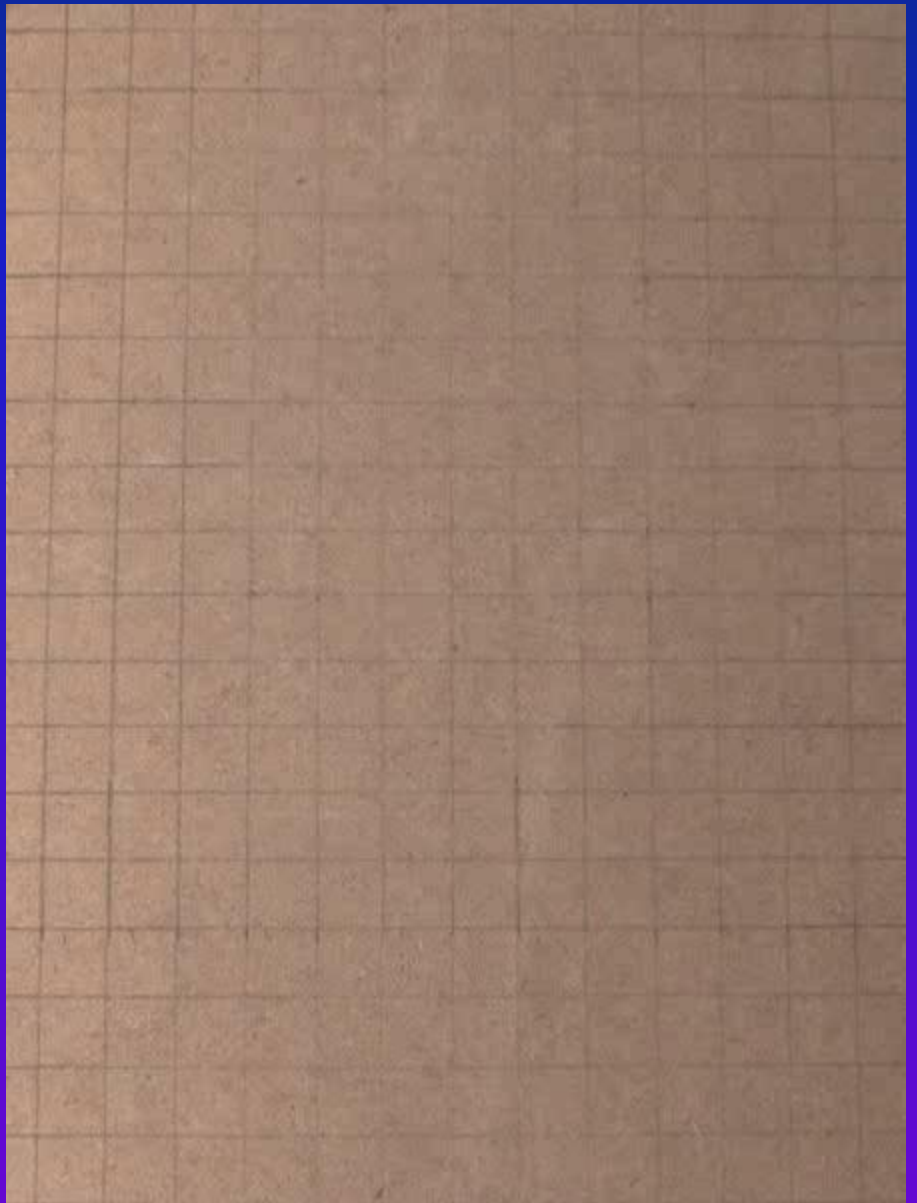
Granular jets and hydraulic jumps on an inclined plane

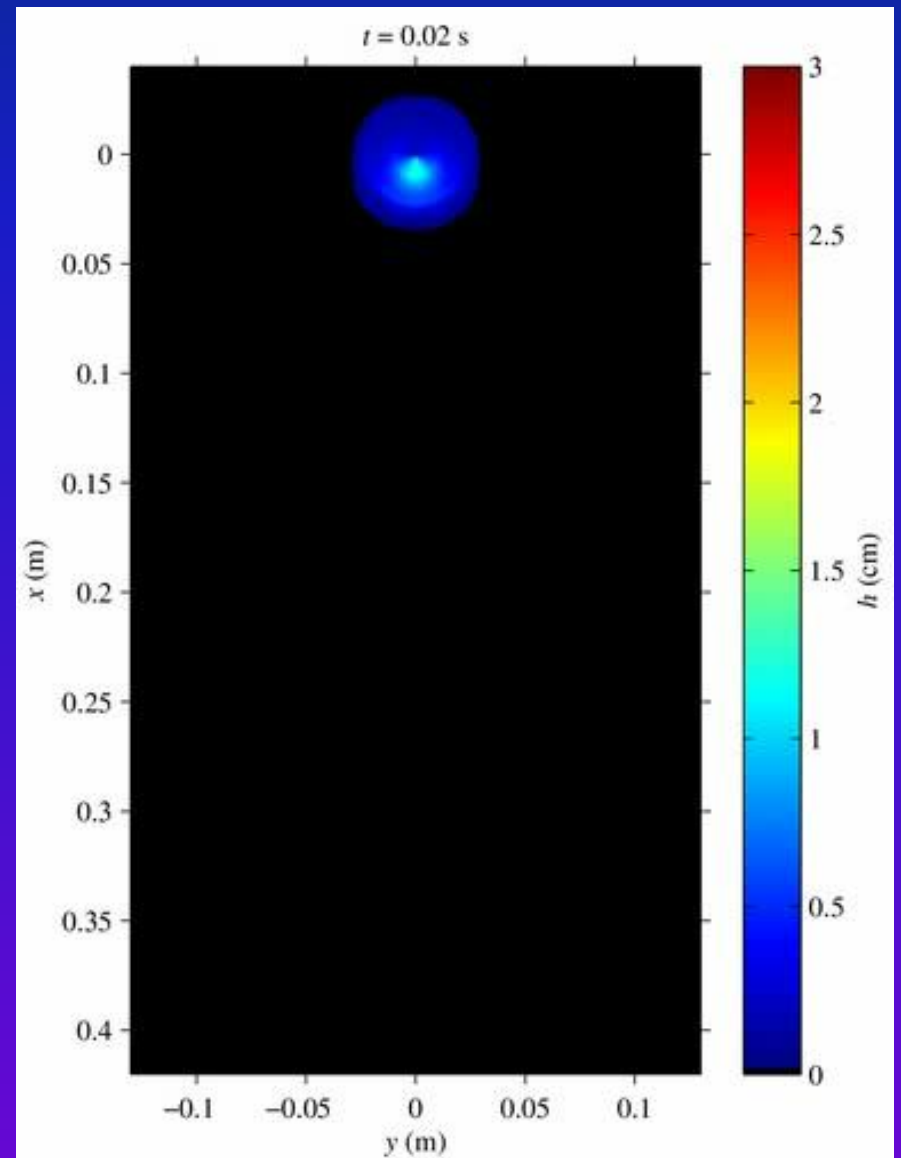
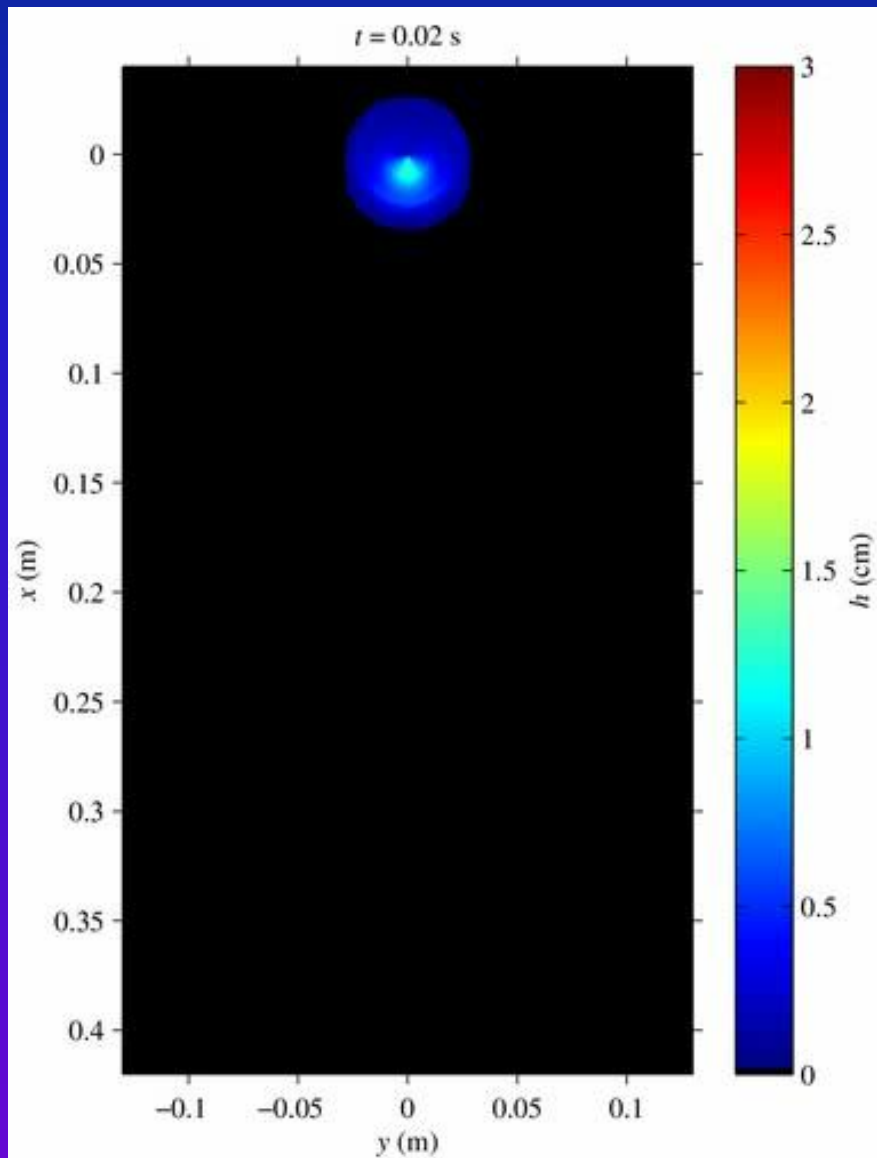


- Oblique impingement of an inviscid jet (Hasson & Peck 1964)
- Friction law for rough beds

$$\mu = \tan \zeta_1 + \frac{\tan \zeta_2 - \tan \zeta_1}{1 + \beta h / (\mathcal{L} Fr)},$$

- including treatment of static material for $0 < Fr < \beta$ (Pouliquen & Forterre 2002)







A simple two-dimensional fully coupled segregation model

- For avalanche thickness h , small particle thickness η and depth-averaged velocity $\bar{\mathbf{u}}$ the 2D coupled model is

$$\frac{\partial h}{\partial t} + \text{div}(h\bar{\mathbf{u}}) = 0,$$

$$\frac{\partial \eta}{\partial t} + \text{div} \left(\eta\bar{\mathbf{u}} - (1 - \alpha)\eta \left(1 - \frac{\eta}{h} \right) \bar{\mathbf{u}} \right) = 0,$$

$$\frac{\partial}{\partial t}(h\bar{\mathbf{u}}) + \text{div}(h\bar{\mathbf{u}} \otimes \bar{\mathbf{u}}) + \text{grad} \left(\frac{1}{2}h^2 \cos \zeta \right) = h\mathbf{S},$$

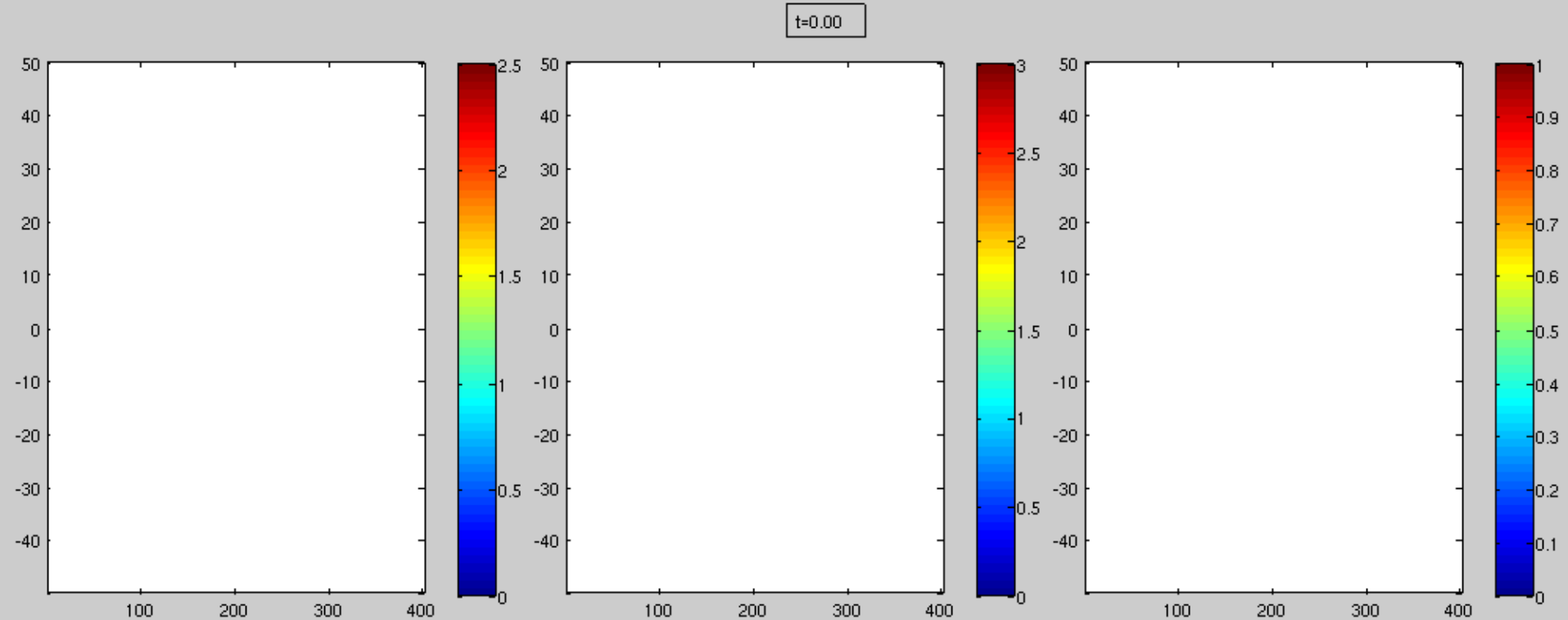
- source terms composed of gravity and basal friction

$$\mathbf{S} = \begin{pmatrix} \sin \zeta - \mu(\bar{u}/|\bar{\mathbf{u}}|) \cos \zeta, \\ -\mu(\bar{v}/|\bar{\mathbf{u}}|) \cos \zeta, \end{pmatrix}$$

- coupling through $\bar{\phi} = \eta/h$ dependent friction coefficient

$$\mu = (1 - \bar{\phi}) \mu^L + \bar{\phi} \mu^S, \quad \mu^L > \mu^S$$

Woodhouse *et al.* (2012), J. Fluid Mech. 709, 543-580.



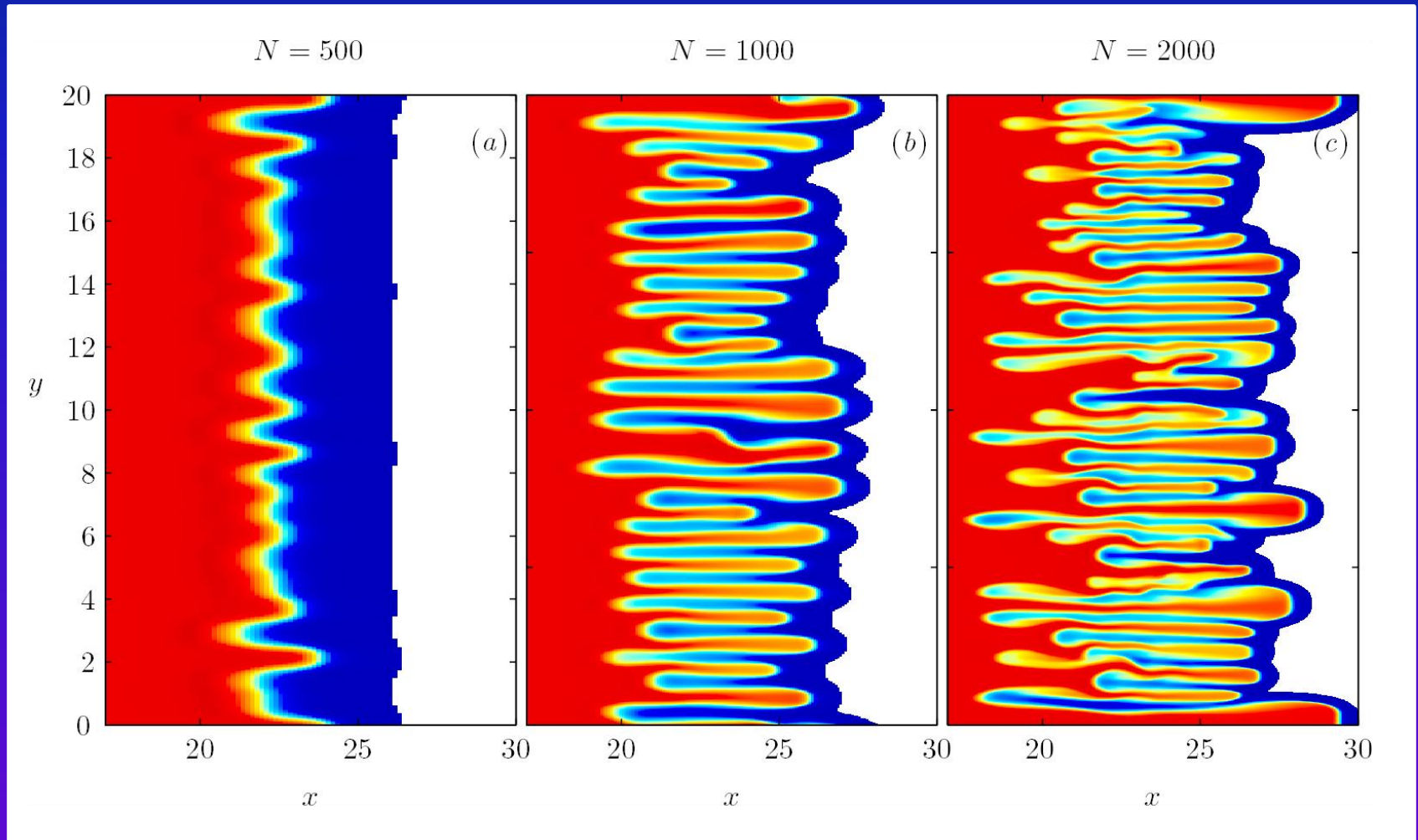
h

$|\bar{u}|$

$\bar{\phi}$

- The model is hyperbolic
- captures the instability mechanism
- and forms large rich lateral levees, BUT

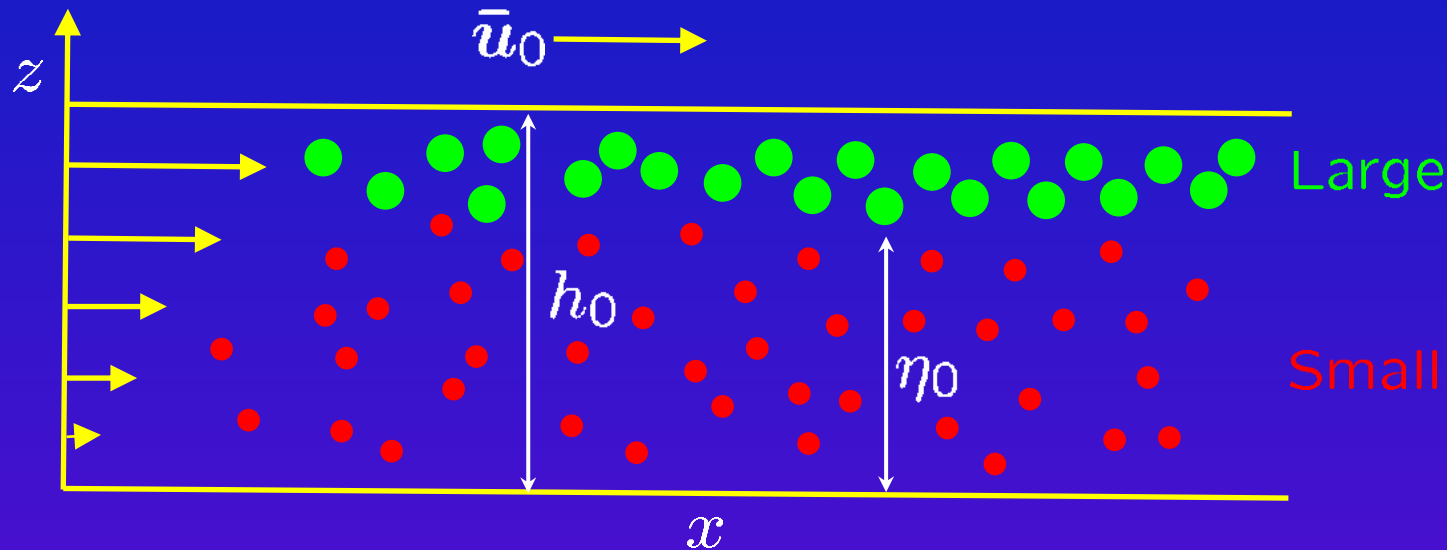
Numerical solutions are grid dependent ...!



- this indicates that there is some important physics missing

- Linear stability about a steady uniform base state

$$\bar{u} = \bar{u}_0, \quad \bar{v} = 0, \quad h = h_0, \quad \eta = \eta_0$$

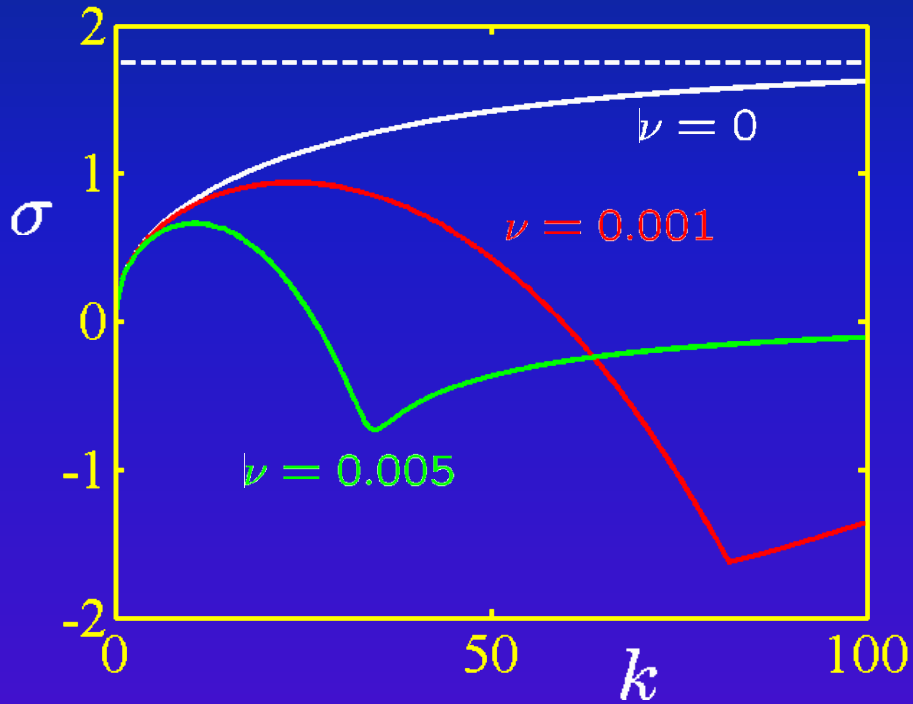


- Predicts unbounded short wavelength instability when

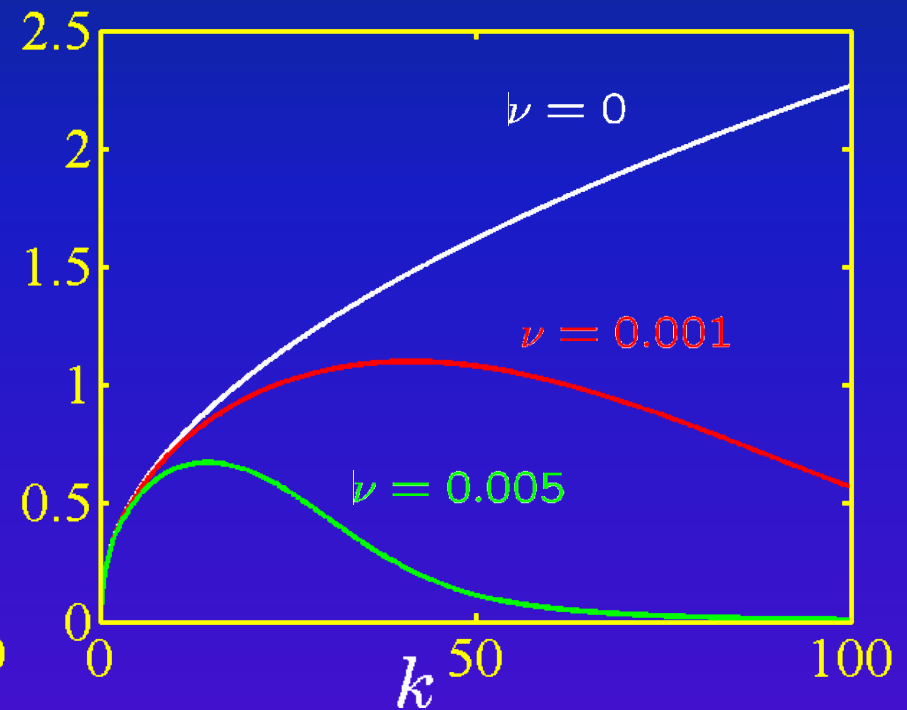
$$Fr = Fr_c = \frac{1}{(1 - \alpha)|2\eta_0 - 1|}$$

- This is when the characteristics coincide

$Fr \neq Fr_c$



$Fr = Fr_c$



- Depth averaging the $\mu(I)$ rheology suggests
- adding a diffusive term to the righthand side of the form

$$\frac{\partial}{\partial x} \left(\nu h^{\frac{3}{2}} \frac{\partial \bar{u}}{\partial x} \right)$$

- which gives cut-off ($Fr = Fr_c$) and boundedness ($Fr \neq Fr_c$)

A two-dimensional fully coupled model including rheology

- When the depth-averaged $\mu(I)$ -rheology is generalized to 2D it suggests a system of conservation laws of the form

$$\frac{\partial h}{\partial t} + \text{div}(h\bar{\mathbf{u}}) = 0,$$

$$\frac{\partial \eta}{\partial t} + \text{div} \left(\eta\bar{\mathbf{u}} - (1 - \alpha)\eta \left(1 - \frac{\eta}{h} \right) \bar{\mathbf{u}} \right) = 0,$$

$$\frac{\partial}{\partial t}(h\bar{\mathbf{u}}) + \text{div}(h\bar{\mathbf{u}} \otimes \bar{\mathbf{u}}) + \text{grad} \left(\frac{1}{2}h^2 \cos \zeta \right) = h\mathbf{S} + \text{div} \left(\nu h^{\frac{3}{2}} \mathbf{D} \right),$$

- where the two-dimensional strain-rate tensor is

$$\mathbf{D} = \frac{1}{2} (\mathbf{L} + \mathbf{L}^T)$$

- and $\mathbf{L} = \text{grad}(\bar{\mathbf{u}})$ is the depth-averaged velocity gradient
- Numerics converges ... (Baker, Johnson & Gray in prep)

