

GRAIN SEGREGATION AND LEVEE FORMATION IN GEOPHYSICAL MASS FLOWS

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Dick Iverson (US GEOLOGICAL SURVEY)

REFERENCES

Further details of the material in the talk are given in:

C. G. Johnson, B. P. Kokelaar, R. M. Iverson, R. G. LaHusen, M. Logan and J. M. N. T. Gray (2012)
Grain-size segregation and levee formation in geophysical mass flows,
J. Geophys. Res. **117**, F01032, doi:10.1029/2011JF002185

Videos are at <http://onlinelibrary.wiley.com/doi/10.1029/2011JF002185/full>



The model coupling particle size to basal friction, shown in the final slide, is described by:

M. J. Woodhouse, A. R. Thornton, C. G. Johnson, B. P. Kokelaar and J. M. N. T. Gray (2012)
Segregation-induced fingering instabilities in granular free-surface flows,
J. Fluid Mech. **709**, p.543–580, doi:10.1017/jfm.2012.348



D. Miller, US Geological Survey



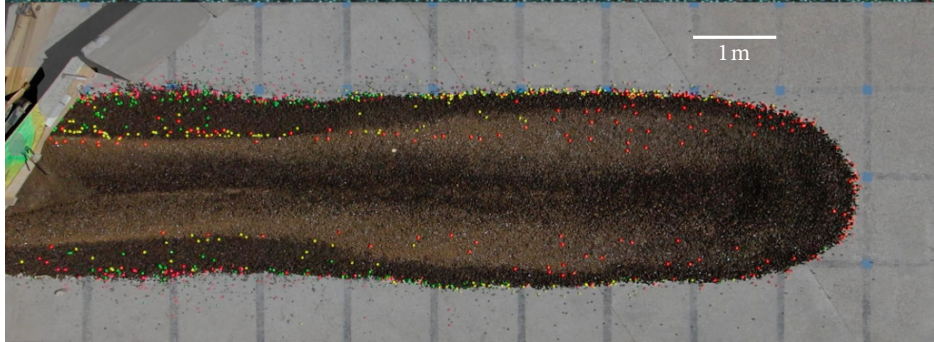
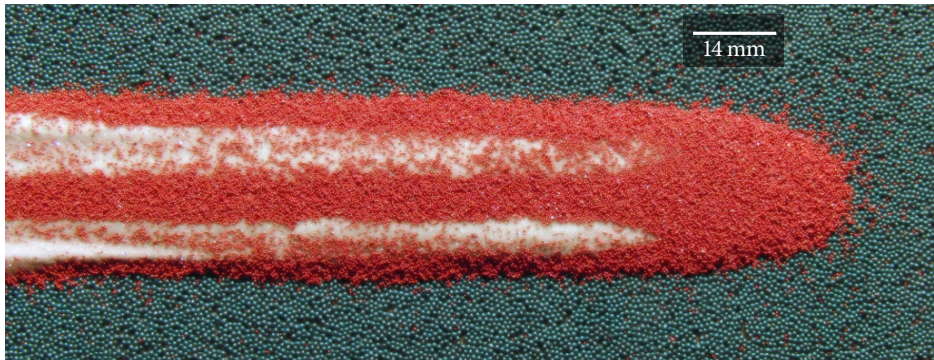
D. Miller, US Geological Survey



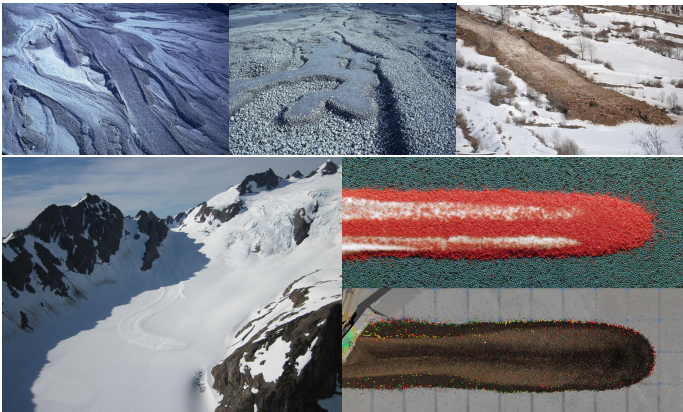
P. Bartelt, WSL/SLF, Switzerland



Blue Glacier, Olympic National Park, WA



LEVEED DEPOSITS



A wide variety of deposits that:

- Are on smooth topography at a shallow incline
- Have an elongated, levee-bounded, (possibly fingered?) morphology
- Have a steep, rounded distal termination

By what mechanism are these deposits emplaced?

USGS DEBRIS FLOW FLUME

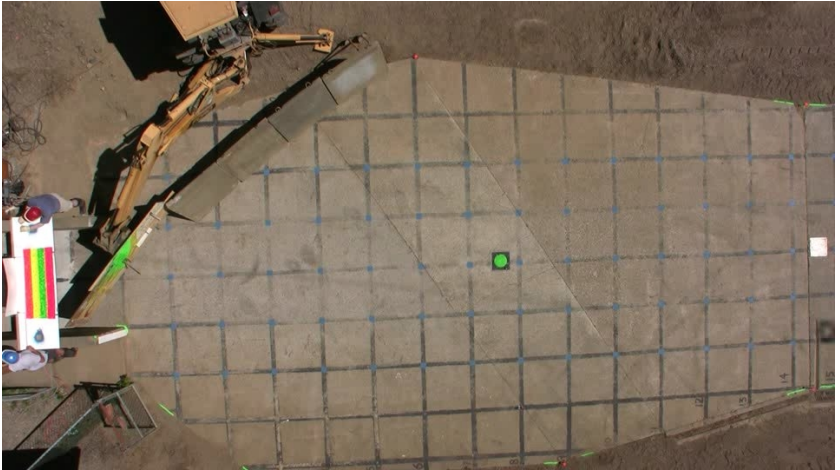


- Flume is 95 m long, 2 m wide
- Inclined at 31°
- Release of 10 m³ of water-saturated sediment
- Sediment is water-saturated sand and gravel, size-range predominantly 0.0625–32mm
- Flume opens onto flat runout area

FLOW TRACKING

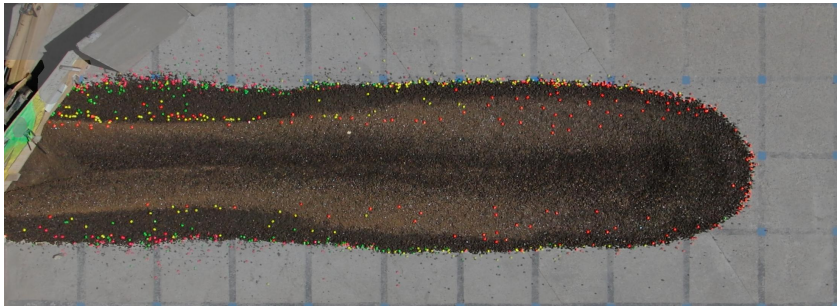


DEBRIS FLOW RUNOUT



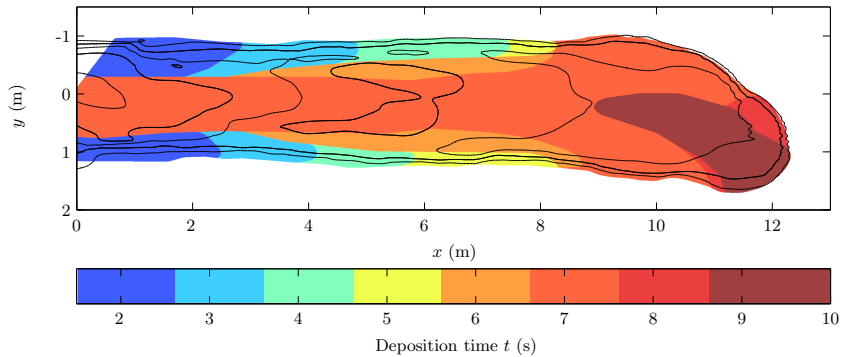
- Flow runs out onto near-horizontal surface ($\sim 2.4^\circ$ incline)
- Forms elongated deposit

DEBRIS FLOW RUNOUT

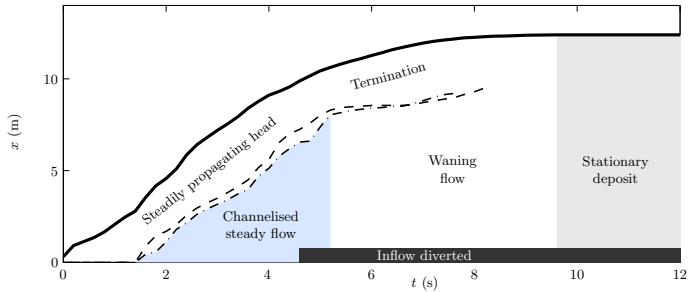
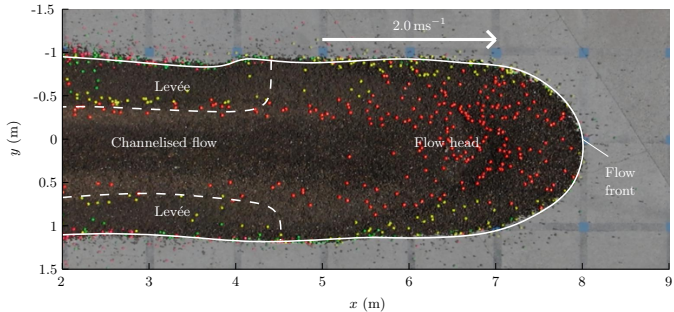


- Flow deposits continuously into stationary levees
- Levees restrict flow width, except close to the flow front

DEBRIS FLOW RUNOUT



DEBRIS FLOW RUNOUT



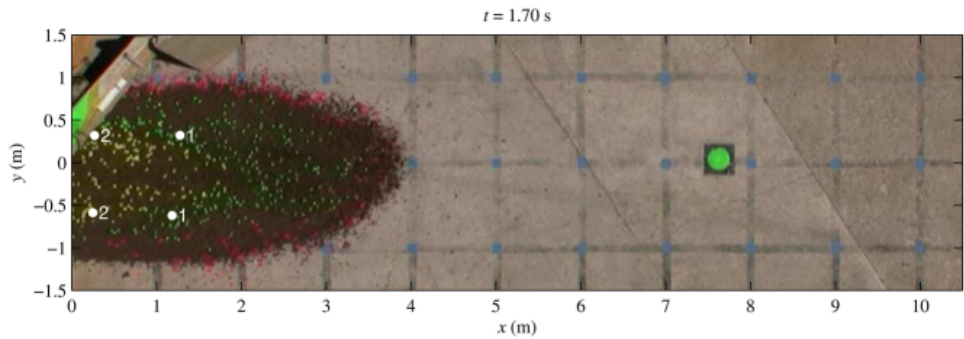
SURFACE TRACERS



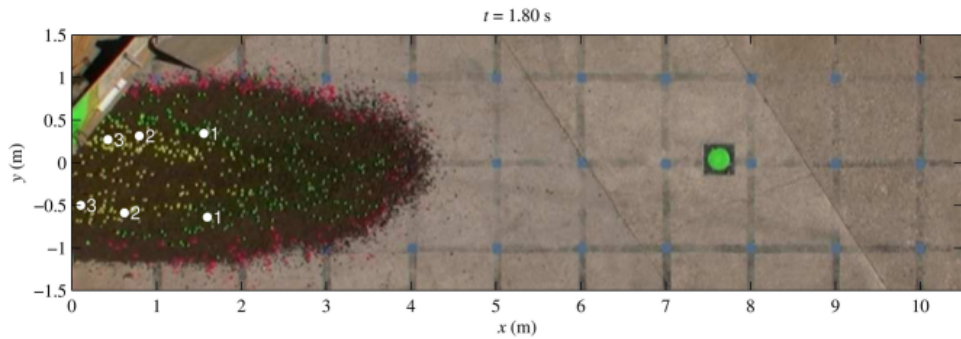
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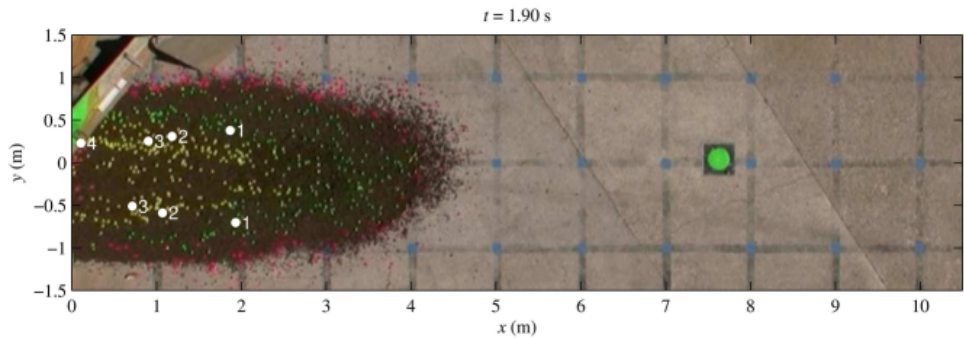
SURFACE PARTICLE PATHS



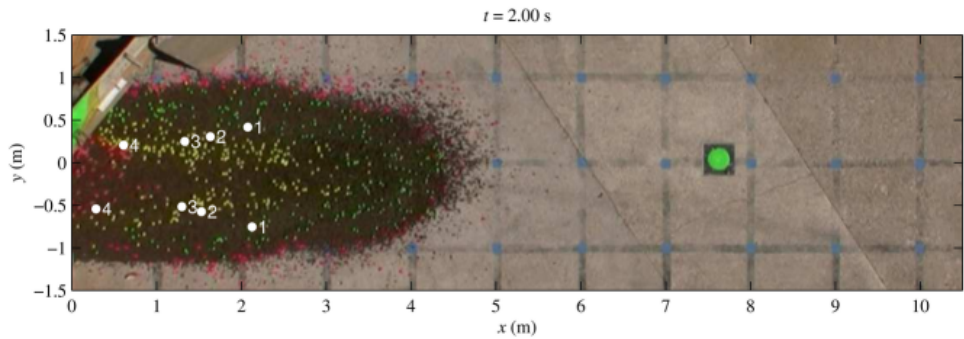
SURFACE PARTICLE PATHS



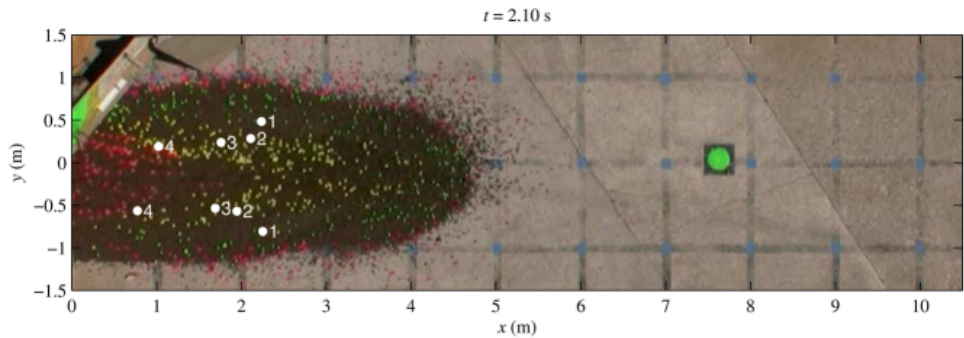
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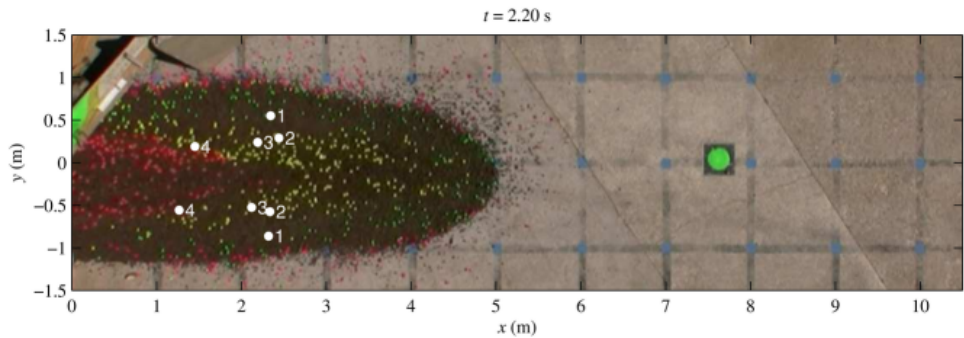
SURFACE PARTICLE PATHS



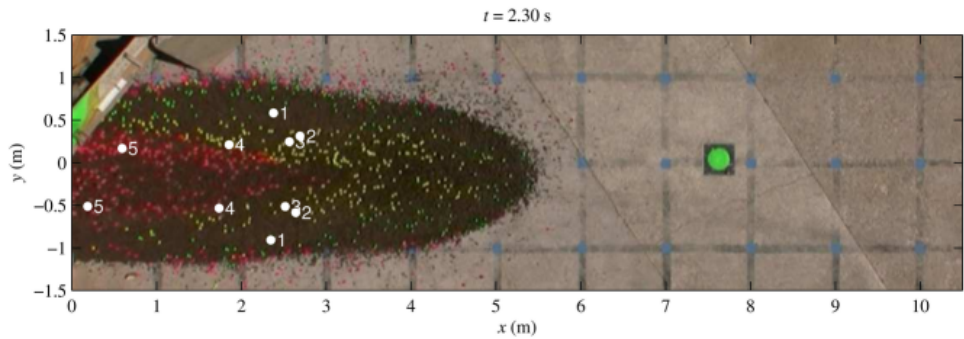
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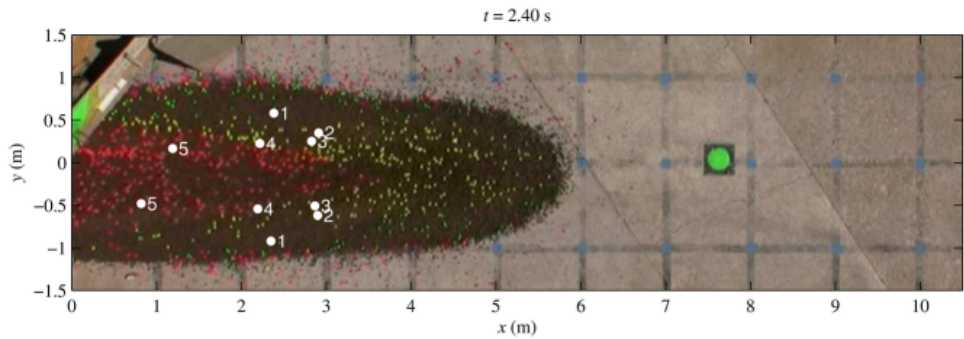
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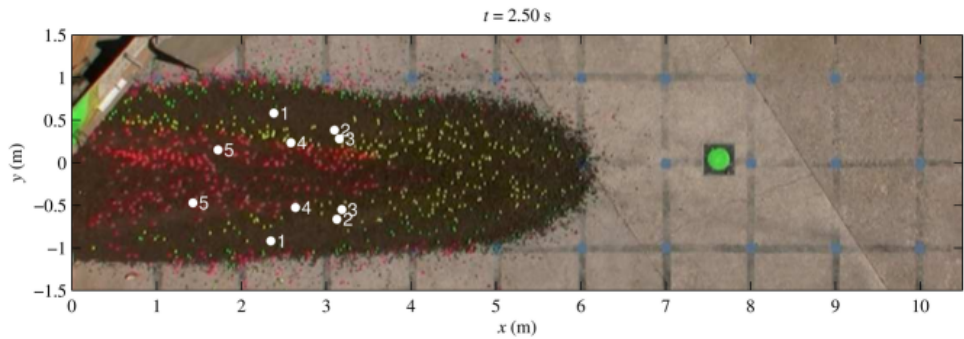
SURFACE PARTICLE PATHS



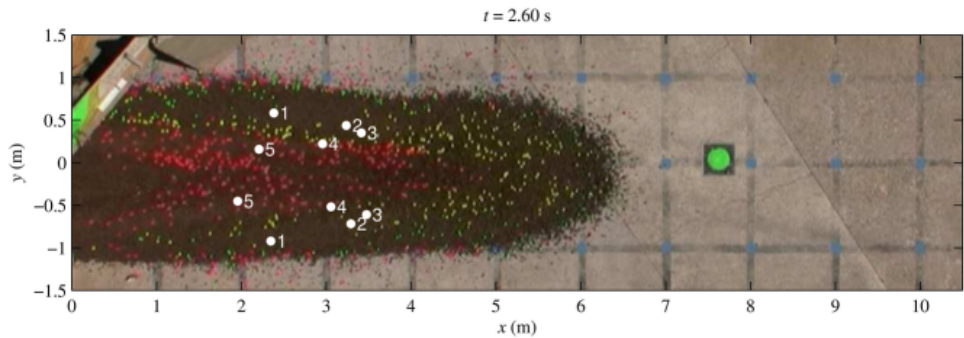
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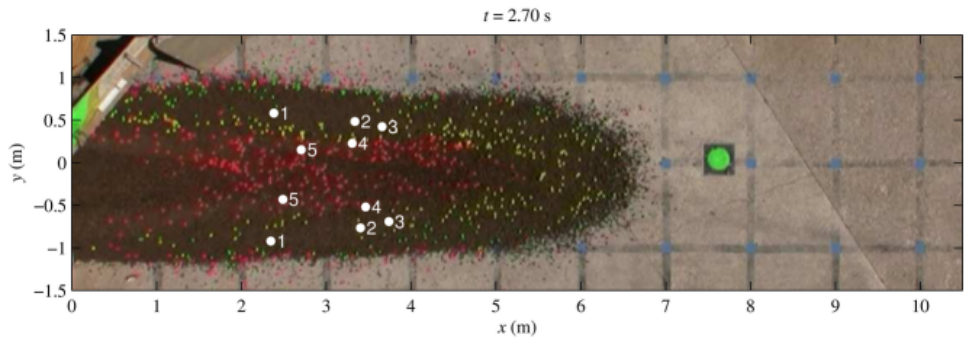
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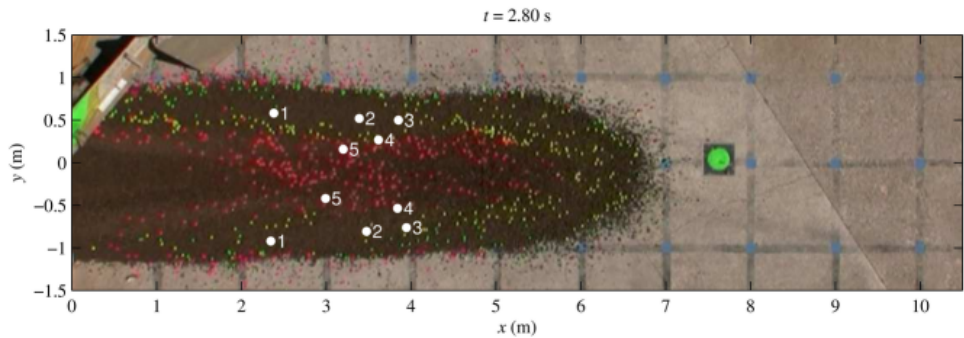
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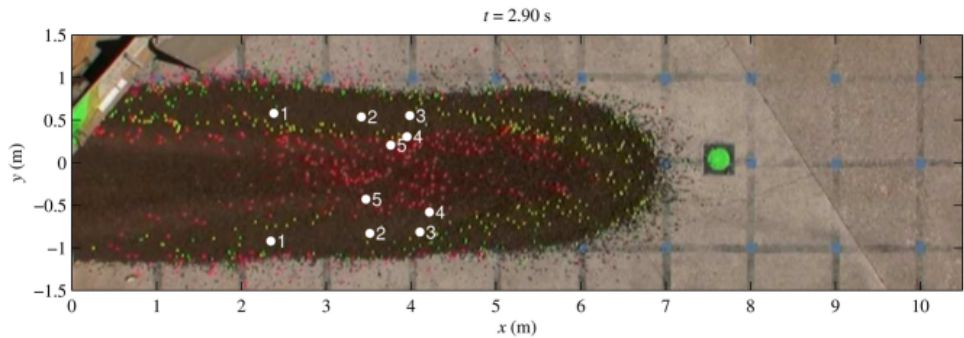
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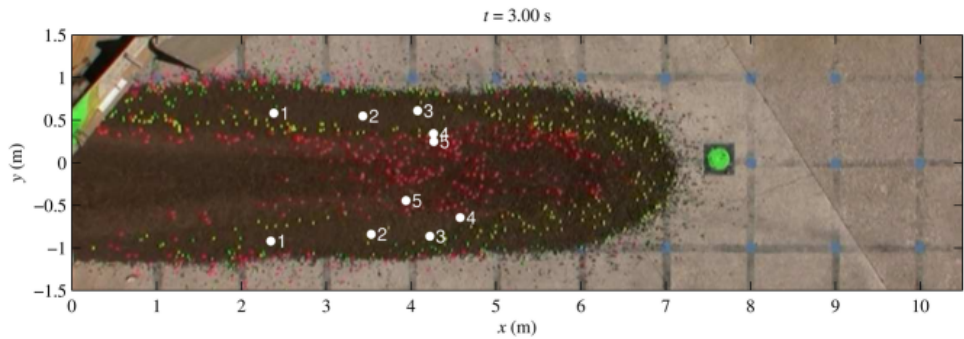
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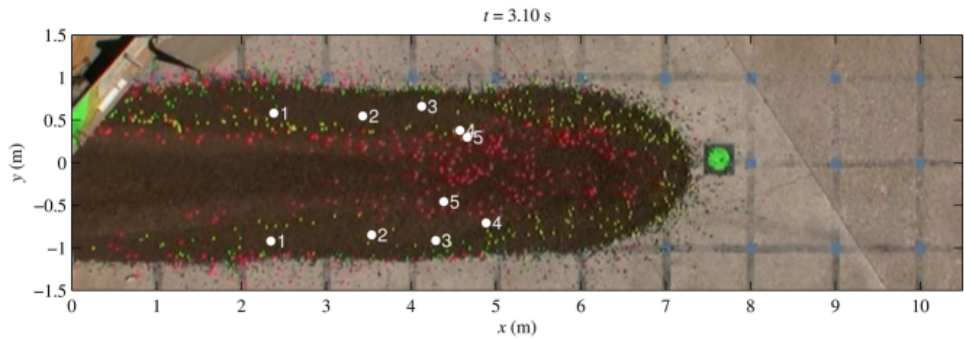
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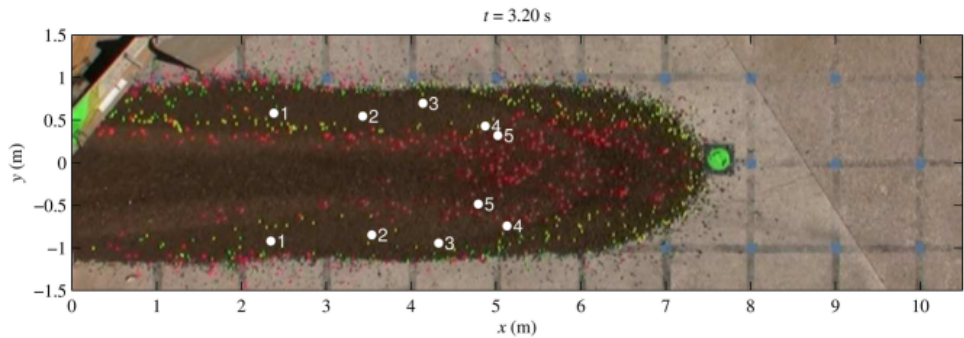
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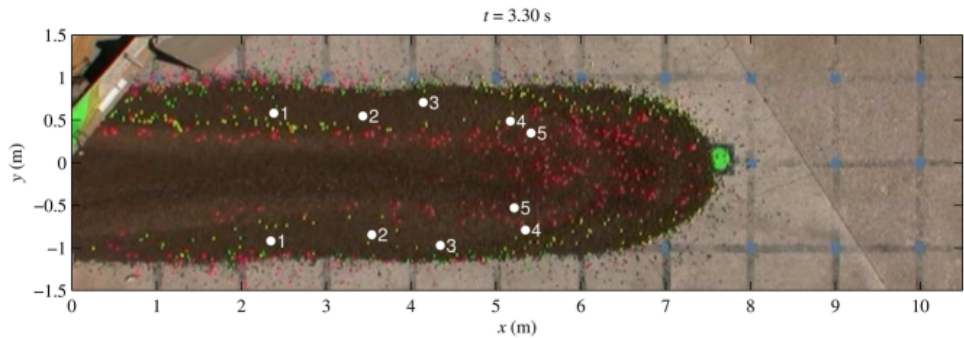
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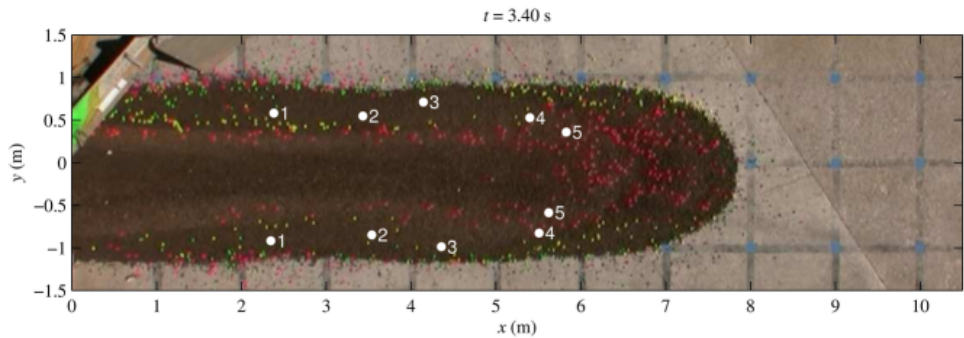
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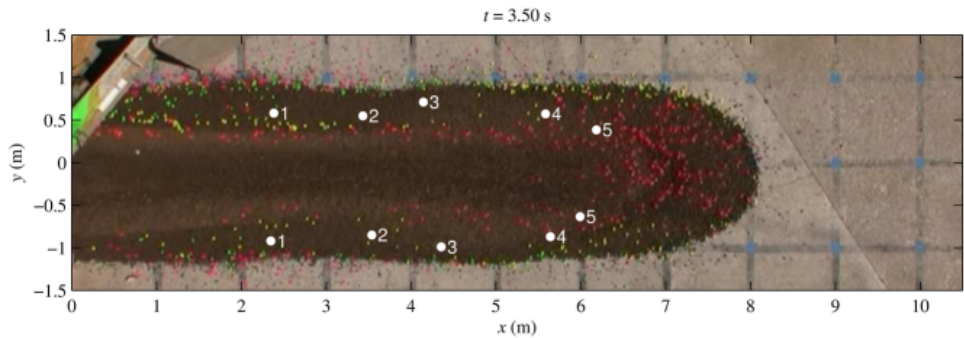
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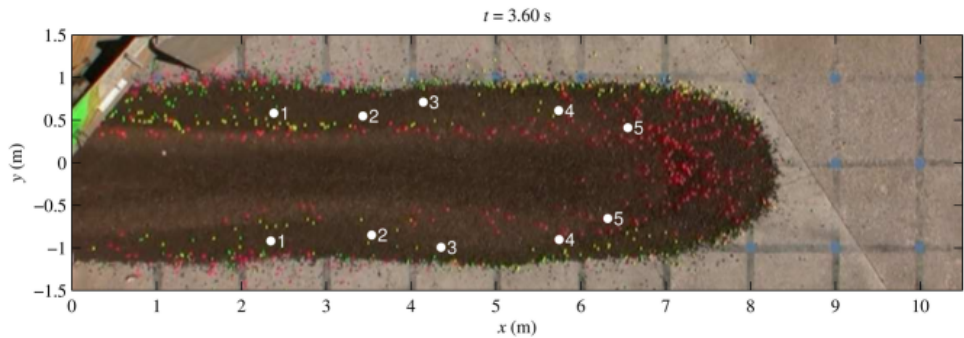
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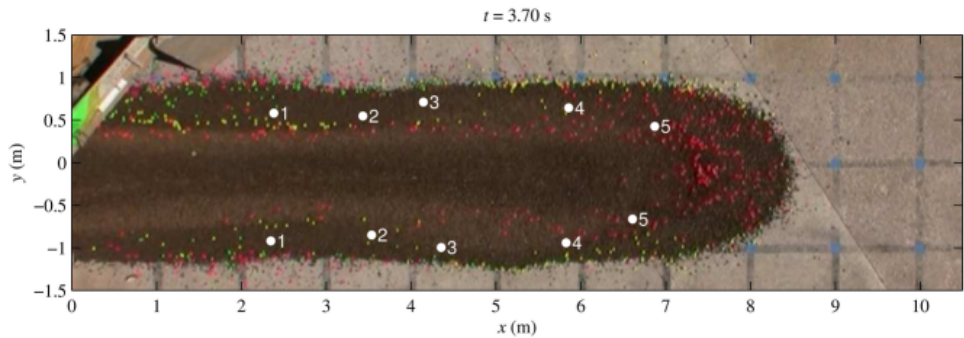
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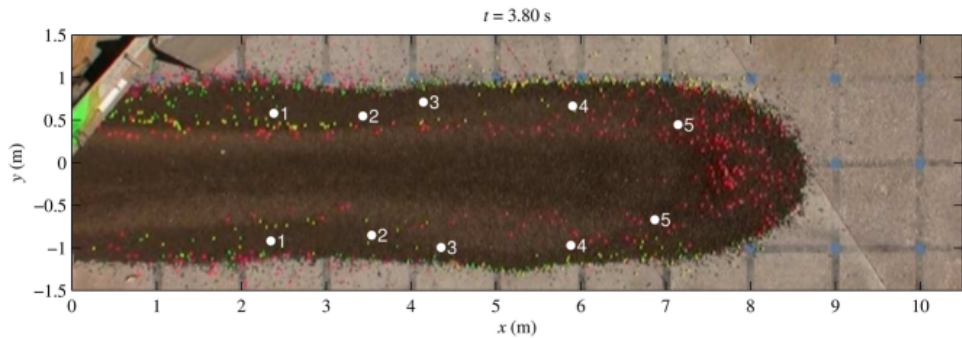
SURFACE PARTICLE PATHS



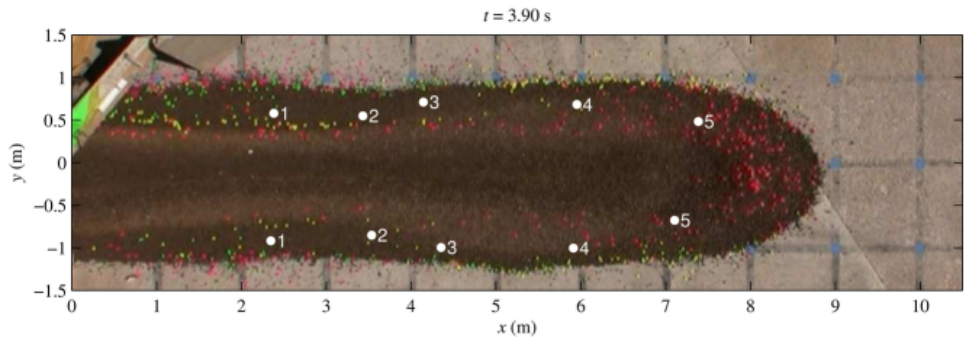
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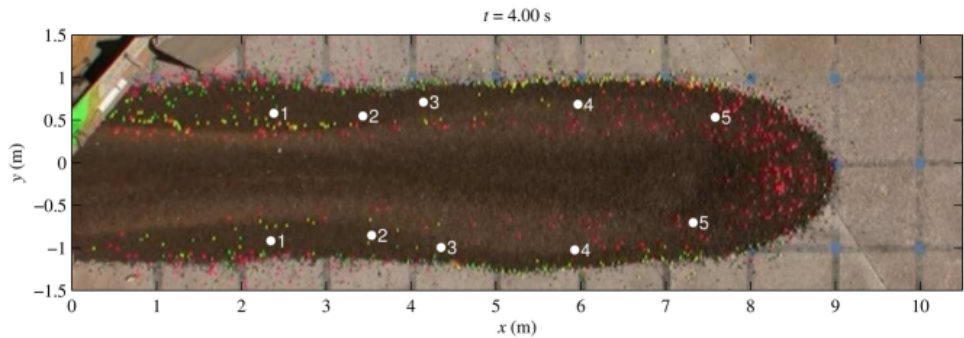
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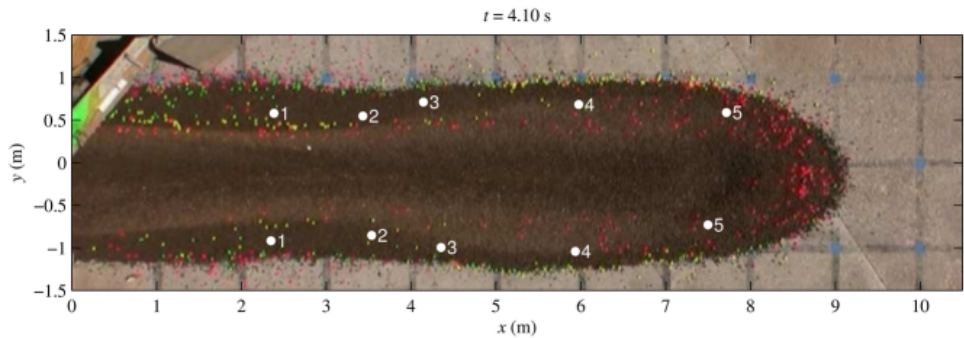
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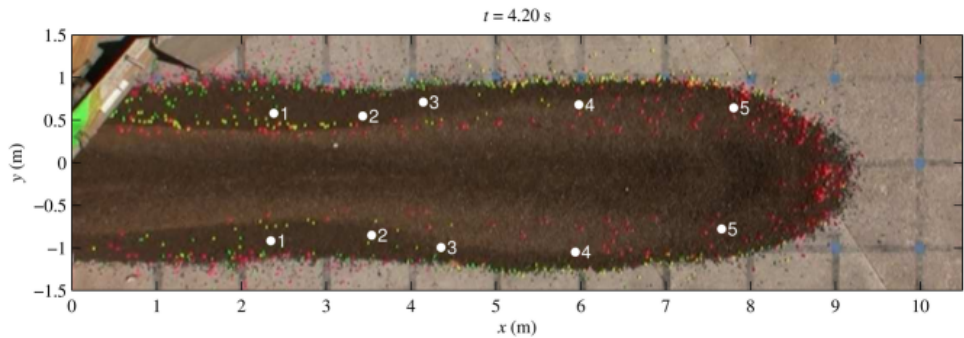
SURFACE PARTICLE PATHS



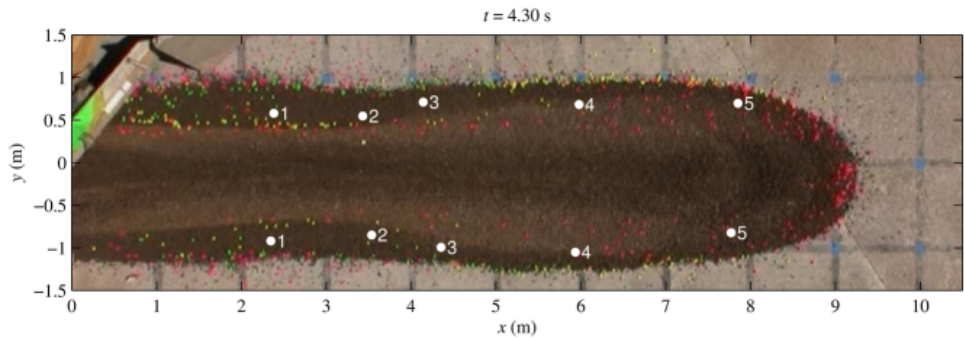
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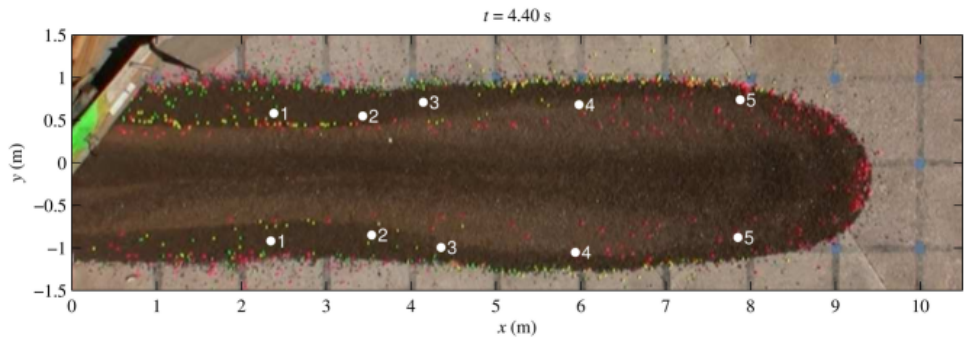
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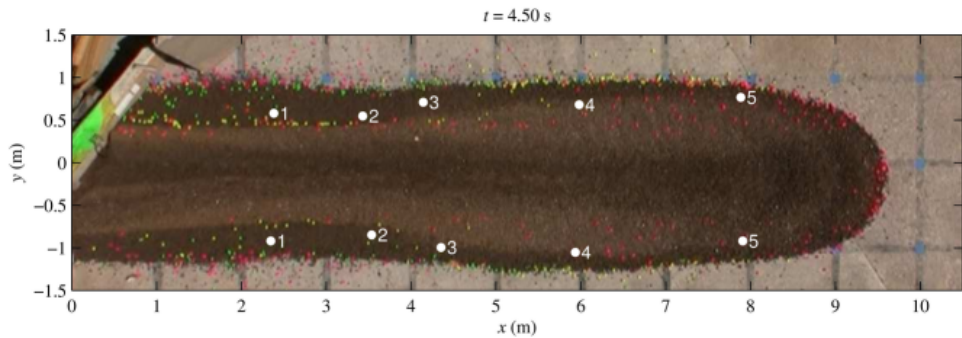
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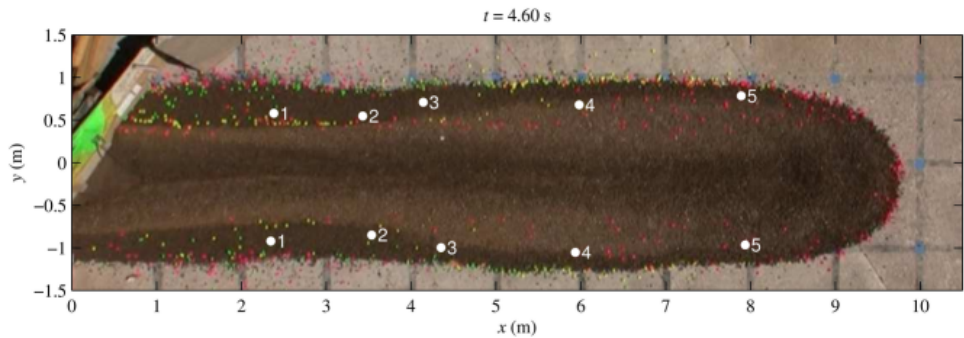
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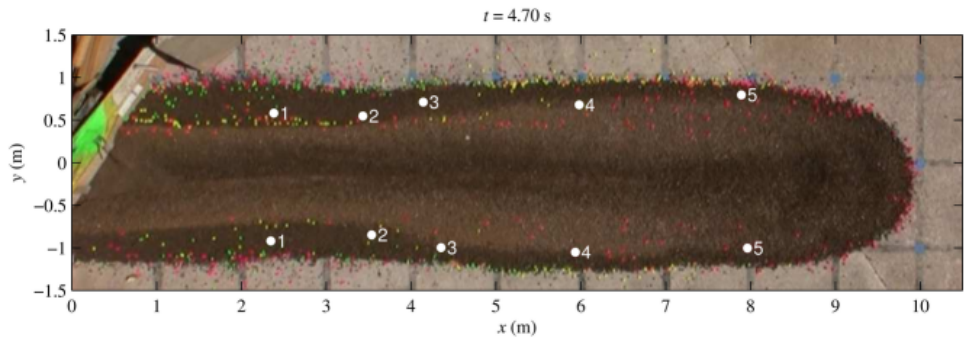
SURFACE PARTICLE PATHS



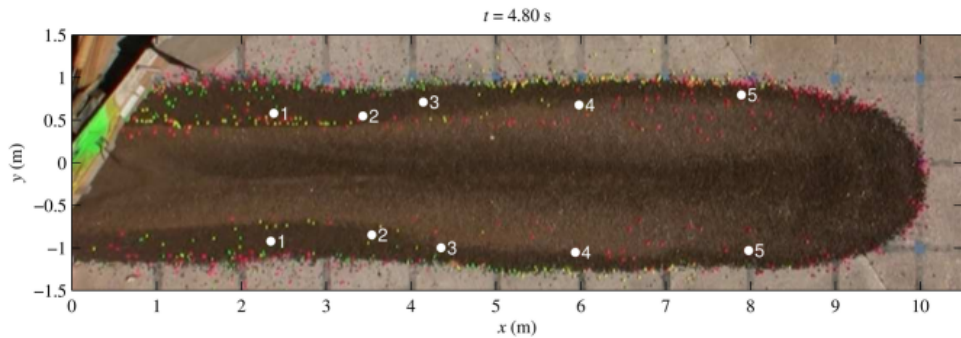
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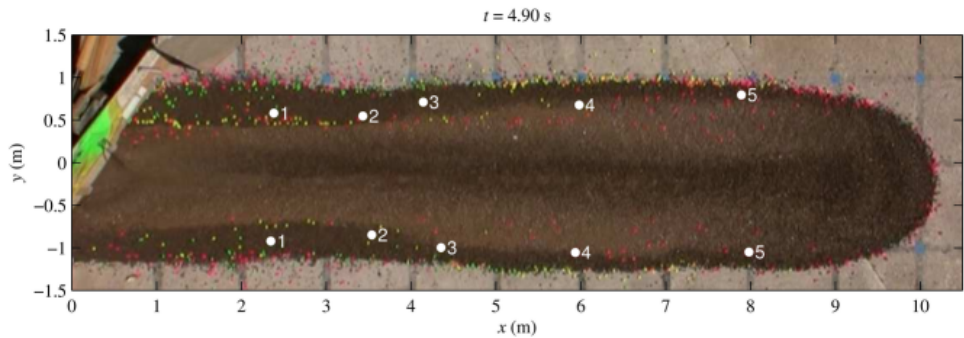
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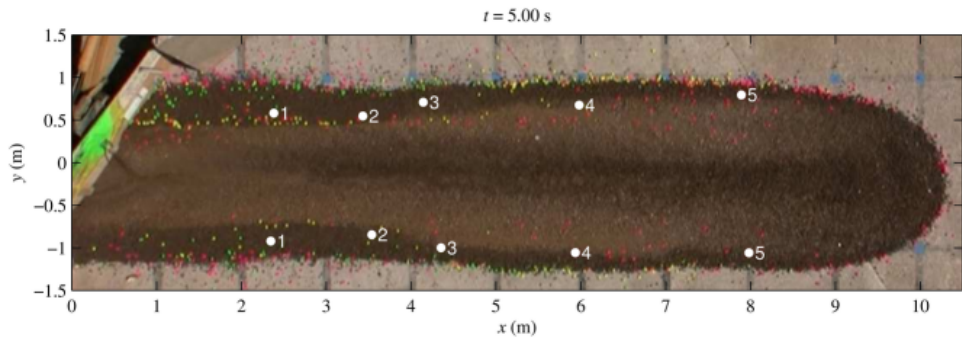
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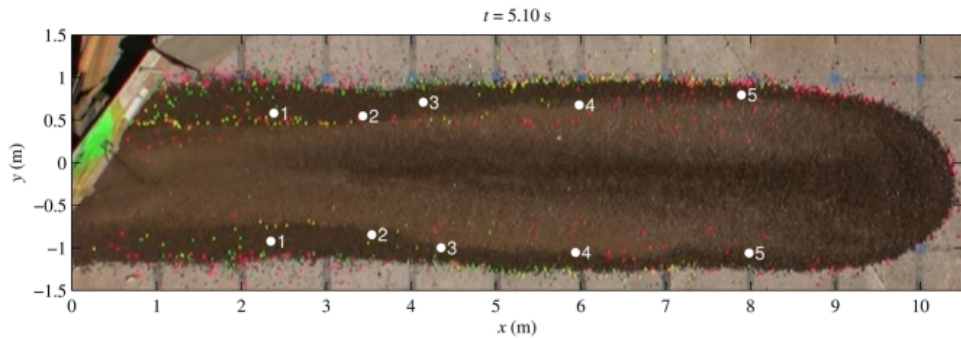
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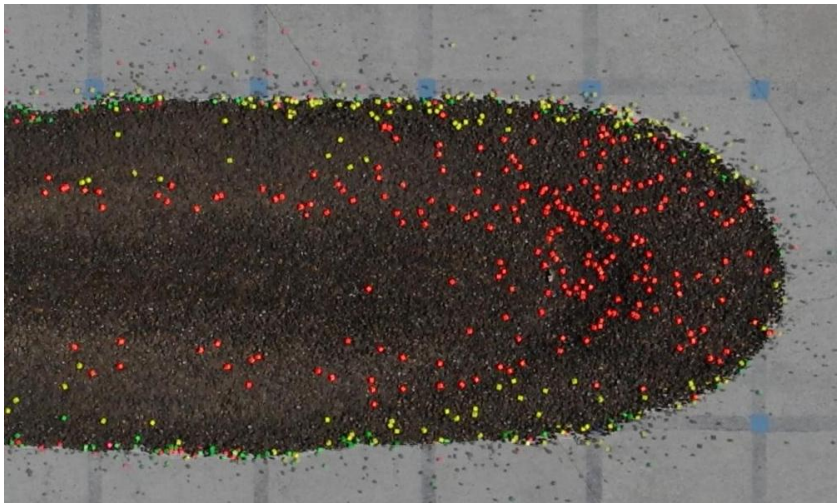
SURFACE PARTICLE PATHS



SURFACE PARTICLE PATHS

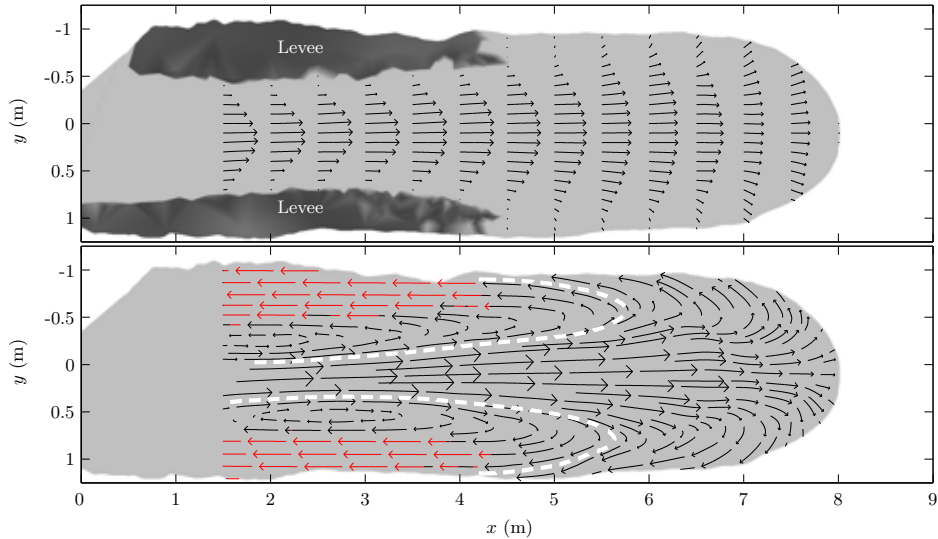


USGS DEBRIS FLOW FLUME



- Flow steady in a frame moving with the front

USGS DEBRIS FLOW FLUME



NATURAL DEBRIS FLOW



Source: Costa & Williams, USGS Open File Report 84-606 (1984)

NATURAL DEBRIS FLOW



Source: Costa & Williams, USGS Open File Report 84-606 (1984)

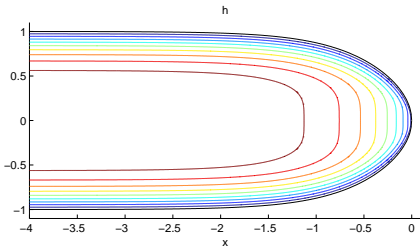
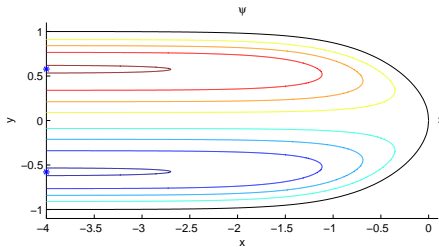
MOVING-FRAME MODELLING

For modelling, construct a flow depth and depth-integrated velocity field that resemble the experimental flows:

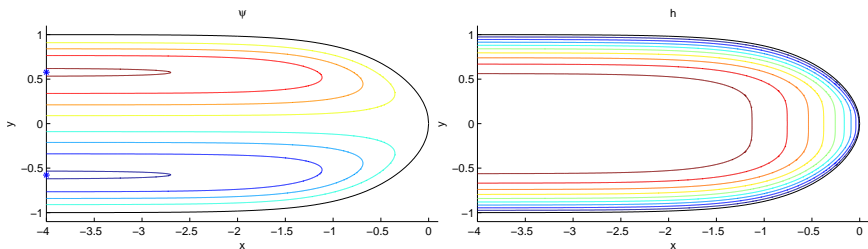
$$h(\xi, y) = y_0 (1 - \hat{y}^{2n})$$
$$h\bar{\mathbf{u}} = (\psi_y, -\psi_\xi) \quad \text{with}$$

$$\psi(\xi, y) = y_0^3 (A\hat{y}^{-1} - B\hat{y}^{2n+1} - C\hat{y}^{2m+1} + D\hat{y}^{2n+2m+1})$$

$$y_0(\xi) = \sqrt{\tanh(-\xi)}, \quad \hat{y} = y/y_0.$$

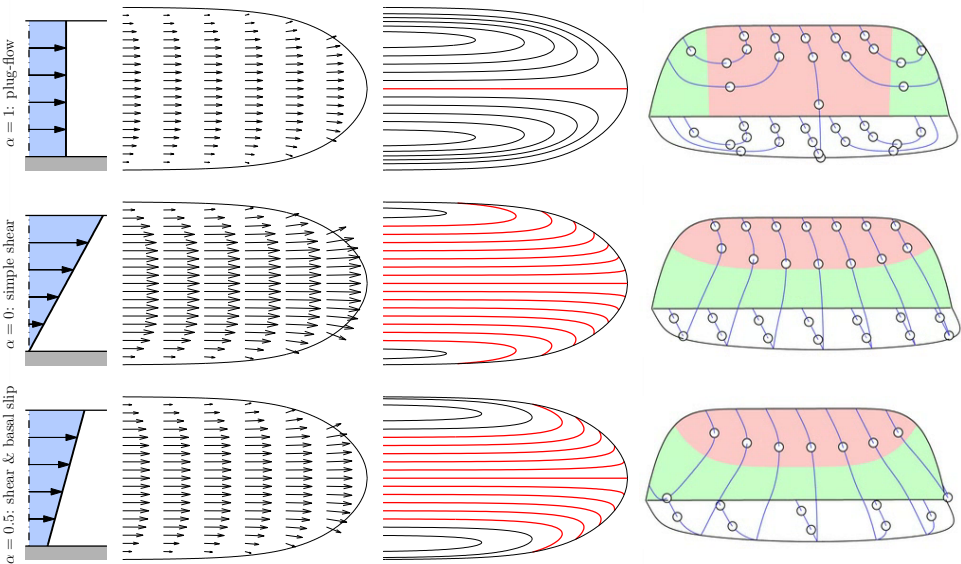


MOVING-FRAME MODELLING

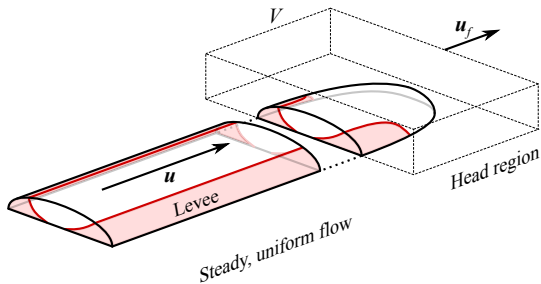


- Assume a constant vertical velocity profile: $(u, v)(\xi, y, z) = \bar{\mathbf{u}}(\xi, y)f(z/h)$
- For example, linear velocity profiles with basal slip: $(u, v) = \bar{\mathbf{u}}\left(\alpha + (1 - \alpha)\frac{2z}{h}\right)$
- Vertical velocity w determined by mass conservation: $\nabla \cdot \mathbf{u} = 0$

MOVING-FRAME MODELLING



MOMENTUM IN THE FLOW HEAD



Momentum conservation: $\rho \left(\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} \right) = \nabla \cdot \boldsymbol{\sigma} + \mathbf{f}$

Assume steady in a moving frame, integrate over volume V moving with the front:

$$\underbrace{- \int_{\partial V} [(\mathbf{u} - \mathbf{u}_f) \cdot \mathbf{n}] u \, dA}_{\text{Momentum advected into head}} + \underbrace{\int_{x=u_f t} (\boldsymbol{\sigma} \cdot \mathbf{n})_x \, dA}_{\text{Pressure force}} + \underbrace{g \sin \theta \int_V \rho \, dV}_{\text{Gravity}} = \underbrace{- \int_{z=0} (\boldsymbol{\sigma} \cdot \mathbf{n})_x \, dA}_{\text{Basal friction / viscous resistance}}$$

Extra momentum advected into head due to lateral shape profile.

Must be balanced by additional basal friction / viscous resistance in head.

PARTICLE SIZE SEGREGATION



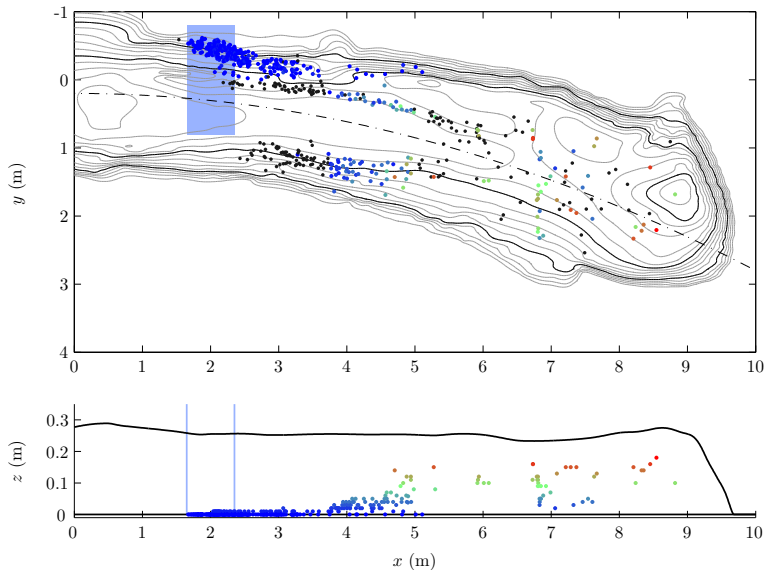
PARTICLE SIZE SEGREGATION



PARTICLE SIZE SEGREGATION



PARTICLE SIZE SEGREGATION



- Infer a typical rise rate of $\sim 3.5\text{cm s}^{-1}$

SEGREGATION MODELLING

In flume experiments, continuum particle size distribution forms three classes:

- Gravel (> 8 mm), segregates upwards
- Sand (0.0625 – 8 mm), segregates downwards
- Mud and fine sand (< 0.0625 mm) ($\sim 2\%$), advected with pore fluid

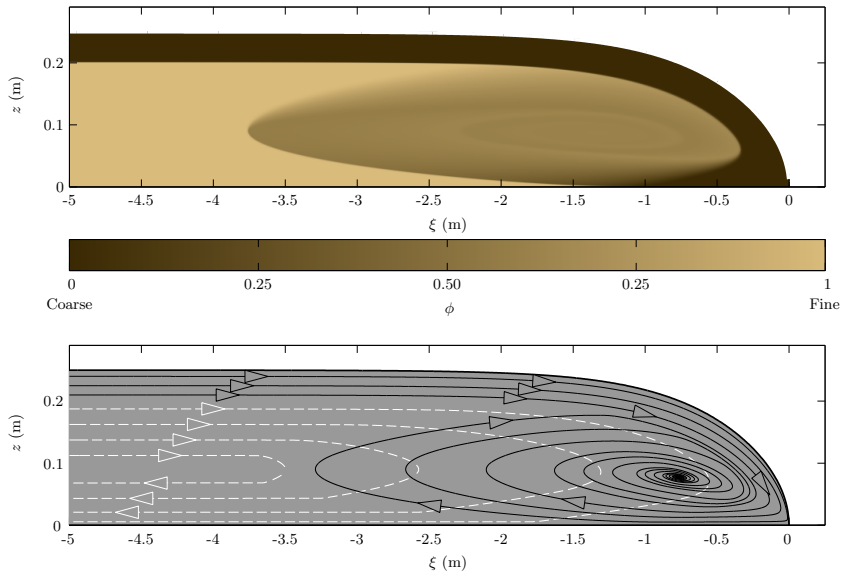
Bidisperse particle segregation model of Gray & Thornton (Proc. R. Soc. **461**)

- Large (gravel) and small (sand) particle volume fractions ϕ_l and ϕ_s
- Assume incompressible flow $\phi_l + \phi_s = 1$
- Large particles move with velocity $\mathbf{v}_l = \mathbf{u} + q\phi_s\mathbf{k}$, concentration governed by

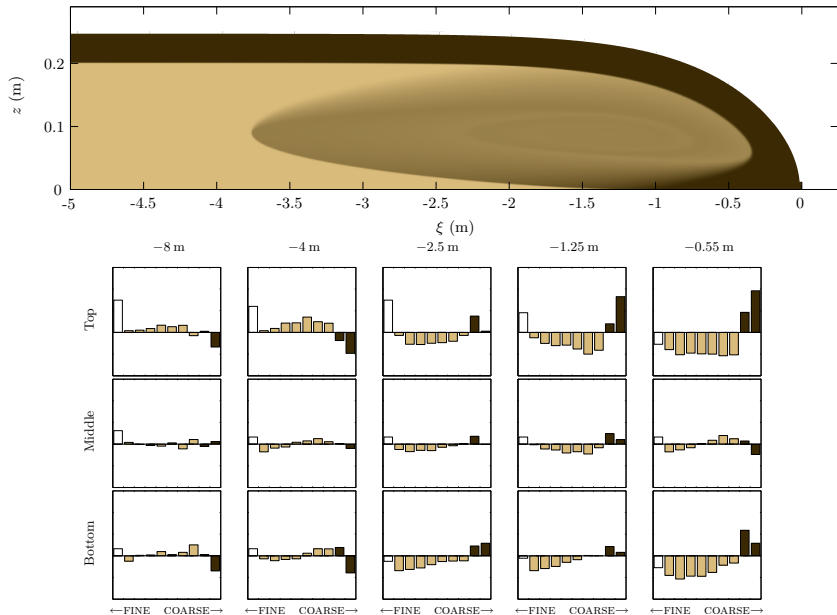
$$\frac{\partial\phi_l}{\partial t} + \nabla \cdot (\phi_l\mathbf{v}_l) = 0, \text{ where}$$

- \mathbf{u} is the prescribed bulk velocity field
- q is a segregation speed ($\approx 3.5 \text{ cm s}^{-1}$)

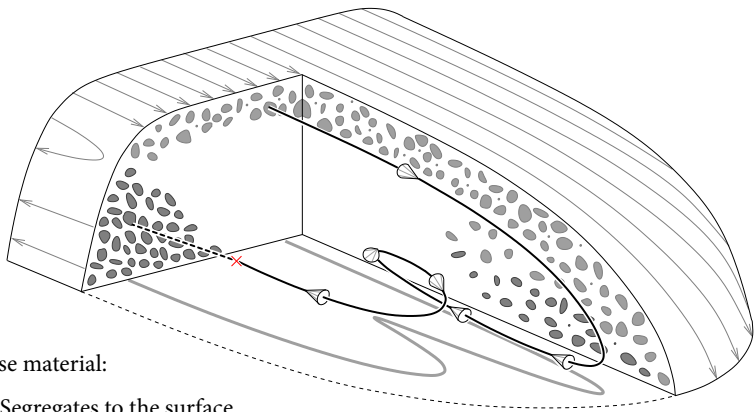
SEGREGATION MODELLING



SEGREGATION MODELLING



PARTICLE PATHS IN A SEGREGATING FLOW FRONT

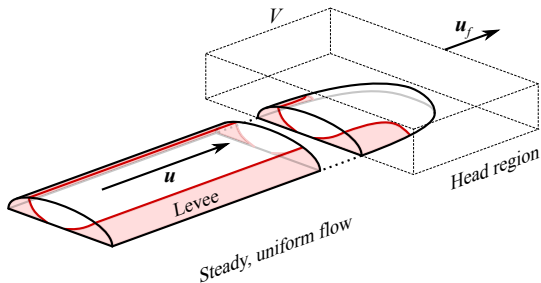


Coarse material:

- Segregates to the surface
- Is transported to the front
- Recirculates with a spiral path within the flow head
- Is transported laterally, into the levees

(Johnson *et al.*, JGR Earth Surface **117** (2012))

MASS BALANCE IN THE FLOW HEAD



Large particle conservation: $\frac{\partial \phi_l}{\partial t} + \nabla \cdot (\phi_l \mathbf{v}_l) = 0$

Assume steady in a moving frame, integrate over volume V moving with the front:

$$\int_{\partial V} [\phi_l (\mathbf{u} - \mathbf{u}_f)] \cdot \mathbf{n} dA = 0$$

Flux of large particles advected into flow head at the surface of the channelised flow must balance flux of large particles leaving head in levees.

DEPTH-INTEGRATED MODELLING

Wish to encapsulate vertical structure into a two-dimensional depth-averaged model.

- Integrate three-dimensional equation

$$\frac{\partial \phi_l}{\partial t} + \nabla \cdot (\phi_l (\mathbf{u} + q\phi_s \mathbf{k})) = 0$$

between $z = 0$ and $z = h$.

- Project $\phi_l(x, y, z, t)$ onto a 'first moment' function $\phi_l(z) = f(z, \bar{\phi}_l(x, y, t))$.

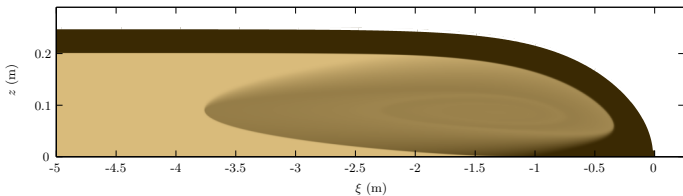
Assuming instantaneous segregation (Gray & Kokelaar, *JFM* **652**), obtain

$$\frac{\partial \eta}{\partial t} + \nabla_h \cdot \left[\eta \bar{\mathbf{u}} \left(\alpha + (1 - \alpha) \frac{\eta}{h} \right) \right] = 0$$

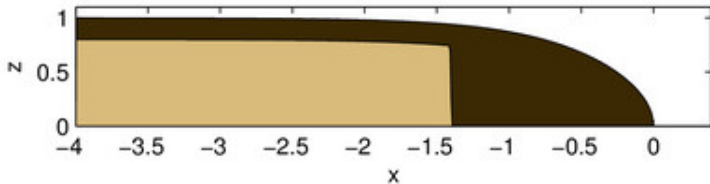
- $\eta(x, y) = h\bar{\phi}_l$ is the height of the interface between segregated small and large particles
- The depth-integrated model captures just the enhanced transport of large particles at the surface

DEPTH-INTEGRATED MODELLING

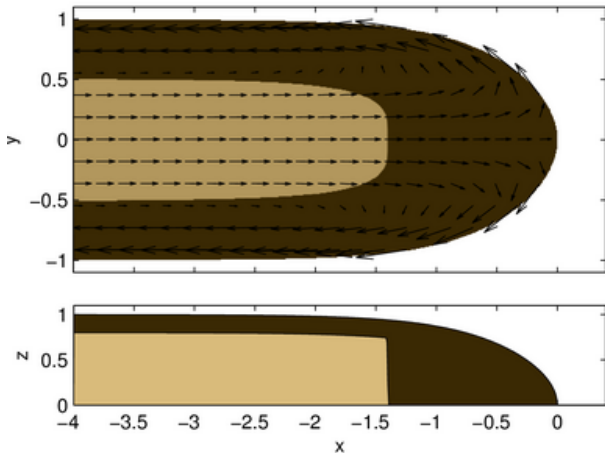
$$\frac{\partial \phi_l}{\partial t} + \nabla \cdot (\phi_l (\mathbf{u} + q\phi_s \mathbf{k})) = 0$$



$$\frac{\partial \eta}{\partial t} + \nabla_h \cdot \left[\eta \bar{\mathbf{u}} \left(\alpha + (1 - \alpha) \frac{\eta}{h} \right) \right] = 0$$

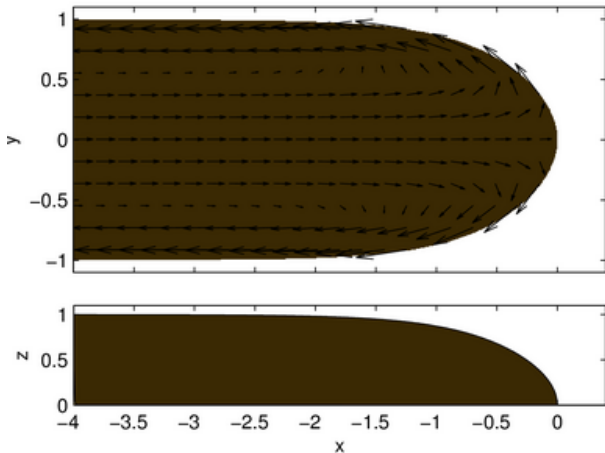


DEPTH-INTEGRATED MODELLING



$\bar{\phi}_l = 0.2$: Large particles transported away from flow front in levees

DEPTH-INTEGRATED MODELLING



$\bar{\phi}_l = 0.3$: Large particles accumulate in flow head





SHALLOW-WATER & DEPTH-INTEGRATED SEGREGATION MODEL

Model depth-integrated:

- Mass conservation of each species
- Momentum conservation
- Effect of large particle concentration on friction

$$\frac{\partial h}{\partial t} + \nabla_h \cdot [h\bar{\mathbf{u}}] = 0$$

$$\frac{\partial \eta}{\partial t} + \nabla_h \cdot \left[\eta \bar{\mathbf{u}} \left(\alpha + (1 - \alpha) \frac{\eta}{h} \right) \right] = 0$$

$$\frac{\partial}{\partial t} (h\bar{\mathbf{u}}) + \nabla_h \cdot \left[h\bar{\mathbf{u}}\bar{\mathbf{u}} + \frac{gh^2}{2} \right] = \mathbf{i} \sin \theta + \mu \frac{\bar{\mathbf{u}}}{|\bar{\mathbf{u}}|}$$

$$\mu = \mu(h, |\bar{\mathbf{u}}|, \eta/h)$$

(Woodhouse et al., JFM 709 (2012))

